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# The importance of turbulent kinetic energy on transport of juvenile clams (*Mya arenaria*)

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## Abstract:

Soft-shell clam, *Mya arenaria*, culture on the east coast of Canada is characterized by high loss following seeding. To evaluate the importance of passive transport due to currents, an experimental flume study was designed. The purpose was to measure the effects of hydrodynamic conditions, substrate, and clam size on dispersal in controlled laboratory condition and to interpret these results in relation to field measurements. Unidirectional currents with gradual increasing velocities (0 to 60 cm s<sup>-1</sup>) were applied to three substrates (muddy sand, medium sand, and coarse sand) in which clams from one of three size classes (10, 15, and 20 mm) had burrowed. We also examined the resulting effects of turbulent kinetic energy on the erosion of medium-grain sand and clams from the three size classes. Turbulent energy was created with a homemade device acting on the unidirectional currents. Nearly 95% of buried clams (all substrates and size classes together) were unaffected by unidirectional currents of up to 60 cm s<sup>-1</sup>, but only 10% withstood turbulent kinetic energy of 10.1 J m<sup>-3</sup>, a level that is lower than that measured in the field during an autumnal storm. The transport of clams was found to be directly related to substrate erosion-levels.

**Keywords:** Clam transport; Substratum erosion; Unidirectional currents; Turbulence; ADV; Flume

## 1. Introduction

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The soft-shell clam *Mya arenaria*, Linnaeus, 1758 is a burrowing suspension-feeding bivalve mollusc that is widely distributed over the North American coast. Clam fishing and culture is an important socio-economic activity in many small communities; however, as is the case along the US East coast, the biomass has generally declined, mainly as a result of overfishing, environmental degradation and diseases (Arnold et al., 2002; Beal and Vencile, 2001; McGladdery et al., 2001). In Canada, mollusc culture is an important industry, and the soft-shell clam is considered as an important species for diversification purposes (Anonymous, 2009).

Very little is known about the cultivation of soft-shell clams in eastern Canada as it is relatively new and is still mostly at a R&D level (Calderon, 2007). These clams reach commercial size four to five years after seeding (Beal, 2005). During this period, they are subjected to losses mainly attributable to diseases, like haemic neoplasia, predation and passive transport due to currents (Beal, 2006a, b; Beal and Kraus, 2002; Bourque et al., 2002; Hunt, 2004b; McGladdery et al., 2001; Miron et al., 2005). In lagoons of the Îles-de-la-Madeleine (southern Gulf of St. Lawrence, eastern Canada), sometimes less than 30% of clams seeded during summer at 15 mm shell-length (SL) are retrieved the following spring (Chevarie et al., 2006). We suggest that mechanisms associated with clam losses will be related to predation and/or passive transport. In eastern Canada, only the population of *M. arenaria* from Prince Edward Island was unequivocally affected by haemic neoplasia (Delaporte et al., 2008) that was reported to cause massive mortality (McGladdery et al., 2001). In Îles-de-la-Madeleine lagoons, a neoplasia diagnosed by hematocytology showed that clams were only slightly affected (Delaporte et al., 2008). Indeed, a five-year histopathological survey showed that this disease did not cause the losses observed during the first autumn and winter following seeding among clams in the Îles-de-la-Madeleine (Réjean Tremblay, personal observation).

Survival of juvenile clams is correlated with the burrowing depth of individuals, which is also linked with SL (Zaklan and Ydenberg, 1997). Larger individuals burrow deeper than smaller ones (Zaklan and Ydenberg, 1997). Since small clams are nearer the sediment surface, they are susceptible to transport by substrate erosion at the same shear velocities as inert particles (Matthiessen, 1960; Roegner et al., 1995; St-Onge and Miron, 2007; St-Onge et al., 2007). *Mya arenaria* lives in sediments with a large range of grain sizes (LeBlanc and Miron, 2006; Roegner et al., 1995). Relationships between shear velocity, sediment grain size and SL length on sediment erosion and transport of juvenile clams have been established at shear velocities as low as  $1.6 \text{ cm s}^{-1}$  (Hunt, 2004b). Turbulence (measured as turbulent kinetic energy, TKE) are major factors influencing bedload erosion and could have a strong influence on clam transport (Hendriks et al., 2006; Pope et al., 2006). Therefore, experiments manipulating TKE are likely more representative of conditions encountered by soft-shell clams in the wild than those investigating only the effect of unidirectional current velocities.

Our study was designed to determine the effect of unidirectional current and combined current-waves on clam transport as a function of SL and sediment type. Our main hypothesis is that TKE has a higher impact on clam transport than shear velocity. Because critical shear velocity ( $u^*_c$ ) increases with increasing grain-size and/or sediment cohesiveness (Sobral and Widdows, 2000), we predicted that clam transport varies according to sediment type.

## 2. Material and methods

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### 2.1 Aquatron flume description

Experiments were carried out in a racetrack recirculating flume (Aquatron) at the Aquaculture Station of ISMER in Pointe-au-Père (Rimouski, Qc, Canada). The Aquatron flume (Fig. 1) is

a smaller version of the HYCOBENTHOS flume located in the Marine Station of Dinard, France (Muséum National d'Histoire Naturelle, MNHN) (Jonsson et al., 2006; Olivier, 1997). The main differences concern the flume dimensions (total length = 4.3 m, channel width = 0.45 m), the experimental section characteristics (length = 0.91 m, depth = 0.15 m), and the flow induction device (12 PVC disks [0.5 m diameter] connected via a pulley system to a direct-power 1 hp variable-speed motor). Flow through the flume channel is driven by the friction of the rotating disks; these disks are immersed in the seawater and their speed is set by digital command. When needed, a wave-current benthic boundary layer (BBL) can be produced using a wave generator located at the beginning of the straight flume section: an immersed triangular piece of wood moves up and down at a fixed frequency to create waves.

## 2.2 Aquatron flow characterization

The flume was filled to a depth of 15 cm with 10- $\mu$ m-filtered natural seawater (salinity = 28.5  $\pm$  0.2) at room temperature (20.2  $\pm$  0.4°C). Flow measurements were acquired with a 3-axis 16 MHz SonTek MicroADV (San Diego, USA; Acoustic Doppler Velocimeter) at a data rate of 25 Hz, and 7500 data points were averaged for each height. For each flow profile, velocities were measured at 13 heights between 0.5 and 9.0 cm above the bottom, and shear velocity ( $u_*$ ) was computed from the vertical velocity profile in the logarithmic layer according to the “law of the wall” using the Kármán-Prandtl equation (Bergeron and Abrahams, 1992). To maximize the BBL development, a 5 mm thick PVC sand-coated plate was fixed on the flume bed in the whole straight section.

The relationship between shear velocity and velocity at 5 cm above bottom ( $U_{5cm}$ ) was explored to simplify measurement during following erosion experiments. Nine flow profiles for velocities of  $U_{5cm}$  between 2 and 50 cm s<sup>-1</sup> were measured and their shear velocity computed. A linear regression gave following relationship between  $U_{5cm}$  and shear velocity:

$$u_* = 0.0338 \times U_{5cm} + 0.0251 \quad (R^2 = 0.949) \quad (1)$$

This equation was used subsequently to convert  $U_{5cm}$  for each flow speed treatment to shear velocity  $u_*$ .

To determine potential longitudinal and transversal velocity gradients in the experimental section, we measured  $U_{5cm}$  along three transects across the channel (A, B and C) of eight regularly distributed points (intervals of 5.1 cm; Fig. 1) for a free stream velocity of 15 cm s<sup>-1</sup>. Transects were positioned 15 cm after the start (C), at the center (B), and 15 cm before the end (A) of the experimental area. For current-wave experiments, the mean turbulent kinetic energy (TKE, J m<sup>-3</sup>; N=5) levels were acquired along the B transect using the MicroADV at a data rate of 25 Hz, with 1500 data points per measurement (T = 15.6  $\pm$  0.5°C; salinity = 27.6  $\pm$  0.6). TKE was then calculated using the following formula (Pope et al., 2006):

$$TKE = \frac{1}{2} \rho (u'^2_{mean} + v'^2_{mean} + w'^2_{mean}) \quad (2)$$

where  $\rho$  is the water density (kg m<sup>-3</sup>) and  $u'$ ,  $v'$ ,  $w'$  are the turbulent components of flow velocity (m s<sup>-1</sup>) on x, y and z axis respectively.

When sediment was introduced into the experimental section (see below), we included rugosity parameters to calculate shear stress ( $u_*$ ) from the  $U_{5cm}$  measurements using the Kármán-Prandtl formula:

$$u_* = \kappa \times U_z / \ln(z/z_0)$$

(3)

where  $\kappa$  is the Karman constant ( $\kappa=0.4$ ),  $z = 5$  cm,  $z_0 = k_s / 30$ ,  $k_s = 2.5 D_{50}$ , and  $D_{50}$  is the median grain size diameter of the sediment, varying from 0.3 to 0.44 cm.

### 2.3 Unidirectional current experiments

The three different substrates with distinct mean grain sizes used for this set of flume experiments originated from the Îles-de-la-Madeleine clam seeding site (47° 26' N 61° 50' W; *IM treatment*), and the maritime St. Lawrence Estuary in Métis (48° 40.5' N 68° 2.0' W; *ME treatment*) and Pointe Pouliot near Rimouski (48° 29.4' N 68° 29.6' W; *PP treatment*). These intertidal sediments were not sieved and only very large objects were removed before filling the recess of the flume.

Two surficial sediment samples were collected before and after each experiment to measure particle size. Samples for grain size determinations were homogenized, sieved, and dried at 60°C to determine percentage weights of different fractions using a sieve column with a ratio of successive mesh sizes of 2<sup>0.25</sup> (McManus, 1988). Grain-size values were determined by the geometric method of moments using the Gradistat program (Blott and Pye, 2001).

Before each experiment involving clams, we assessed the shear stress ranges in the experimental section for each substratum by measuring  $U_{5\text{cm}}$  on transects A, B, and C (Fig. 1). Values of  $u_*$  corresponding to  $U_{5\text{cm}}$  equal to 15, 30, 45 and 60 cm s<sup>-1</sup> were also calculated from equation (2). All substrates were kept in a 5°C controlled temperature room during the experiments until their use.

Juveniles were collected, in early November 2006, from HAM lagoon by sieving sediment on a 5 mm mesh. Clams were sent directly on sea ice to the Station Aquicole de Pointe-au-Père, which is associated to Institut des Sciences de la Mer de Rimouski. Prior to experiments, 600 clams (shell lengths between 5 and 30 mm) were maintained for one month before experiment in 200 l tanks filled with 12 cm deep ME-like sediment and 10 µm filtered natural seawater at constant temperature (16.7 ± 0.8°C) and salinity (26.0 ± 0.7). They were fed with a standard hatchery diet consisting of equal proportions of the microalgae *Pavlova lutheri*, *Isochrysis aff. Galbana* and *Nannochloropsis* sp. The diet was supplied continuously to maintain a total concentration of ~30 cells µl<sup>-1</sup> in the tanks, allowing clams to be fed to satiety.

Clams were sorted into three size classes based on their shell length (SL): 10.8 ± 1.0, 15.6 ± 1.0, and 19.8 ± 1.1 mm (here after referred to as 10, 15 and 20 mm SL). For each trial, the experimental section was filled with a randomly chosen substrate and the flume was then filled with 10 µm filtered seawater (T = 16.7 ± 0.8°C; salinity = 26.0 ± 0.7) to a height of 15 cm. Ten juvenile clams from one size class, also selected randomly, were placed on the sediment surface in the middle of the experimental area (B transect, Fig. 1). The clams were placed in parallel and separated by 5 cm from each other so that they were all facing the incoming flow. A unidirectional current of 10 cm s<sup>-1</sup> ( $U_{5\text{cm}}$ ) was applied for 16 h to induce individuals to burrow into the substrate: this velocity was found to significantly increase both the mean proportion of juvenile clams to burrow (St-Onge and Miron, 2007) and the burial rate (St-Onge and Miron, pers. com.). At the start of the experiment, the clams' burial state (buried or not) was noted and the current velocity was then increased with a maximal acceleration of 1 cm s<sup>-1</sup> min<sup>-1</sup> to avoid burst effects. Only clams that were initially completely buried with visible siphon in or out of the sediment were considered in the data analysis. Nine increasing flow regimes were applied over 30 minutes each, corresponding to free stream velocities ( $U_\infty$ ) equal to 20, 25, 30, 35, 40, 45, 50, 55, and 60 cm s<sup>-1</sup>; clam positions were noted at the beginning and at the end of each flow step. At the end of each test, the flume was completely cleaned and washed before the next trial. Five replicates (n=5) for each treatment (sediment × SL) were run (N=45) at a rate of one treatment per day. Each day, combination of treatment (sediment × SL) was selected randomly.

### 2.4 Turbulence experiments

The experimental protocol adopted for this set of experiments was similar to the previous one except for i) the water depth, which was increased to 21 cm, and ii) the flow

regimes applied to the buried clams, which lasted 30 minutes, corresponded to  $U_{\infty}$  equal to 20, 40, 45, 50 and 55  $\text{cm s}^{-1}$ , and integrated a constant turbulent energy produced by 120 cycles per minute (CPM) of the wave generator. A five-minute stabilization period was applied after each flow increase to allow the observation of clam positions. Temperature and salinity remained stable during the whole set of experiments ( $17.5 \pm 0.3^{\circ}\text{C}$  and  $24.5 \pm 0.5$ , respectively).

## **2.5 Field measurements**

To characterize the hydrodynamic regime in the field, an 6 MHz Vector velocimeter, (Nortek, Rud, Norway) was deployed from 5 to 13 October 2006 near the main clam seeding site ( $47^{\circ} 26' \text{ N } 61^{\circ} 50' \text{ W}$ ) in the high subtidal zone of Havre-aux-Maisons (HAM) lagoon in the Îles-de-la-Madeleine (southern Gulf of St. Lawrence, eastern Canada). The velocimeter sensor was located 27.7 cm above the sediment surface and was sampling at 8 Hz for 6.83 min (3280 data points) every 15 min. The sampling volume of the velocimeter is positioned 15.7 cm from the sensor; it was therefore located 12 cm above sediment level to avoid interference due to sand accumulation.

On the same site, bed elevation change was measured twice a week from 18 September to 28 November 2007 using nine horizontal rods initially anchored 30 cm above the sediment level and scattered over a total area of 110  $\text{m}^2$ . The first data recorded were considered as controls and further positive or negative values correspond to sediment deposition or erosion, respectively.

## **2.6 Data analysis**

Field data from the ADV were filtered to remove low tide and aberrant data. Maps of the shear velocities in the flume were produced using the linear interpolation gridding method in Surfer software (ver. 7.0). Two-way ANOVAs were performed on mean shear stress data to assess i) the transversal (four treatments from the internal to the external side of the flume) and longitudinal (three treatments from the beginning to the end of the experimental section) gradients (7126 data points per treatment) and ii) their relationship with both free stream velocity (four treatments) and sediment type (three treatments) corresponding to 24 data points per treatment (measurement grid in Fig. 1).

Three-way split plot ANOVAs were performed on the unidirectional current experiment data to determine differences in the clam transport rate as a function of substrate type, clam size and free stream velocities. The units of replication were the five runs performed for each of the nine substrate type  $\times$  clam size treatments; the main plots were substrate (MI, ME and PP) and clam size classes (10, 15 and 25 mm), and sub-plots were free stream velocities (10 values from 10 to 60  $\text{cm s}^{-1}$ ). Two-way split plot ANOVAs were performed on data from the turbulent kinetic energy experiments, with size classes as the main plot and TKE (four levels from 0 to 10.1  $\text{J m}^{-3}$ ) as subplots. These setups of effect of flow velocity include cumulative effect of increasing flow, even if a stabilization period have been applied between each flow increase. Thus, we consider flow velocity as a subplot factor with limiting interpretation, as the losses of clams could be influenced by earlier velocity. Advantage of this setup is the realization of each flow velocity measure in the same day and limit temporal variability of measures inherent to use of one flume.

For all these analysis, normality and homogeneity of variance was directly checked by observations of residuals; (Log+1) transformations were applied when required. When needed, LSD post-hoc comparisons tests were performed to identify groups differing from the others. Data in text, table and figures are presented with mean  $\pm$  SD.

### 3. Results

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#### 3.1 Aquatron flow characterization

In general, vertical velocity profiles were fully developed and the regression coefficients related to  $u_*$  calculated using the “law of the wall” method were always above 0.96 (see examples in Fig. 2). In the experimental area, the shear stress velocities associated with a  $15 \text{ cm s}^{-1}$  free stream velocity varied according to the longitudinal and transversal gradients (significant longitudinal  $\times$  transversal interaction,  $df=6$ ,  $MS=0.04$ ,  $F=288.9$ ,  $p<0.001$ ; Fig. 3). Post-hoc comparisons show significant differences between all treatments ( $p<0.001$ ), with minimum mean shear stresses found on the internal side of the flume, especially at the beginning of the experimental section (mean  $u_* = 0.679 \text{ cm s}^{-1}$ ). Maximum shear stress values were observed in the inner side of the experimental area, particularly at the end of the experimental area (mean  $u_* = 0.750 \text{ cm s}^{-1}$ , Fig. 3). The gradient between the external and the internal sides is related to the gradual increase in the benthic boundary layer from the leading edge, which is characteristic of flume flows (Muschenheim et al., 1986; Nowell and Jumars, 1984), as well as to the lateral gradient associated with the flume walls. However, in the experimental section of the Aquatron flume, the spatial variability of the shear stress is low for a mean  $U_\infty$  of  $15 \text{ cm s}^{-1}$ , always less than 10% (in the experimental section:  $u_* = 0.72 \pm 0.03 \text{ cm s}^{-1}$ ).

In the unidirectional current experiments, mean shear stresses were related to both mean  $U_{5\text{cm}}$  and substrate type (significant interaction,  $df=6$ ,  $MS=0.032$ ,  $F=3.447$ ,  $p=0.0027$ ). For a given  $U_{5\text{cm}}$ , the mean  $u_*$  was generally similar for the three sediments but was lowest for the IM– $15 \text{ cm s}^{-1}$  treatment. When the coarser ME substrate was exposed to high flow regimes, stronger mean shear stresses were generated (Table 1). In general, the BBL flow regime was always turbulent ( $Re > 9000$ ); it was most frequently characterized by a transitory rugosity on the experimental section ( $3.5 < Re^* < 49$ ), but also rarely showed a smooth flow (for  $U_{5\text{cm}} = 15 \text{ cm s}^{-1}$  on IM and ME sediments). In these flow conditions, TKE was very low (between  $0.07$  to  $1.78 \text{ J m}^{-3}$ ) with the lowest values measured above PP sediment.

In turbulence experiments, TKE was positively related to increasing  $U_{5\text{cm}}$  (Fig. 4), with maximum values ( $10.71 \text{ J m}^{-3}$ ) obtained for  $U_{5\text{cm}} = 55 \text{ cm s}^{-1}$ . This maximum was slightly below the mean TKE values measured in the field ( $\sim 12 \text{ J m}^{-3}$ ; see below).

#### 3.2 Flume experiments

Grain sizes of the three sediments are presented in Fig. 5. The samples comprised 1) a clean well-sorted medium sand with a mean grain size of  $0.297 \text{ mm}$  (IM), 2) a clean well sorted slightly gravelly to gravelly sand with a mean grain size of  $0.421 \text{ mm}$  (ME), and 3) a heterogeneous muddy gravelly sand with a mean grain size of  $0.332 \text{ mm}$  (PP).

At the beginning of the unidirectional current experiments ( $n=5$ ), the percentage of buried clams was always higher than 82%. We found only a significant size effect ( $df = 2$ ,  $MS=0.38$ ,  $F=3.62$ ,  $p=0.047$ ). Post-hoc LSD tests showed a weak significant decrease of the mean burial proportion of  $15 \text{ mm SL}$  clams ( $p<0.05$ ). In fact, we observed that  $15 \text{ mm SL}$  clams were dislodged slightly more in the IM treatment (10%) than the others sediment treatment (Fig. 6). For the other trials, we were unable to dislodge the clams even at the highest flow regime (for the ME– $60 \text{ cm s}^{-1}$  treatment, mean  $u_* = 2.10 \text{ cm s}^{-1} \pm 0.07 \text{ cm s}^{-1} \text{ SE}$ ).

The percentage of all size classes of buried clams was significantly affected by the wave-current flow regime ( $df=1$ ,  $MS=4.48$ ,  $F=85.9$ ,  $p<0.001$ ), with lower values corresponding to the most dynamic regime (LSD tests,  $p < 0.001$ ;  $TKE = 10.1 \text{ J m}^{-3}$ ). In fact, more than 80% of juvenile clams were dislodged when TKE reached  $10.1 \text{ J m}^{-3}$  (Fig. 7). However, lower percentages of the  $15$  and the  $20 \text{ mm}$  clams were buried at the beginning of the experiment (70% and 60%, respectively) than the smaller individuals (100%).

### 3.3 Field measurements

Wind, mean current speed and turbulent kinetic energy (TKE) measured in October 2006 at the HAM field site are shown in Fig. 8. Very high TKE values were observed when winds and currents were high ( $15 \text{ J m}^{-3}$  on 5 October, and up to  $20 \text{ J m}^{-3}$  on 13 October; mean TKE =  $12 \text{ J m}^{-3}$ ), corresponding to stormy conditions. The average temperature measured by the velocimeter was  $14.1 \pm 3.18^\circ\text{C}$ . In the field, instantaneous currents, generally co-occurring with windy conditions (wind velocities temporarily reaching  $15 \text{ m s}^{-1}$ ), were always below  $24 \text{ cm s}^{-1}$ ; this velocity would correspond to a maximum shear stress induced by currents of only  $1.16 \text{ cm s}^{-1}$ . Bed elevation measurements showed several erosion and deposition periods (Fig. 9). Stormy conditions generate sediment erosion during two main periods: between 2 and 18 October ( $\Delta z = -1.6 \text{ cm}$ ) and between 26 October and 5 November ( $\Delta z = -2.0 \text{ cm}$ ). Sediment deposition was observed between these periods at rates of  $0.1$  to  $0.2 \text{ cm d}^{-1}$ .

## 4. Discussion

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### 4.1 Effect of unidirectional currents

We were mostly unable to erode juvenile (10-20 mm) *Mya arenaria* in our flume, even with shear velocities reaching  $2.10 \text{ cm s}^{-1}$  (= unidirectional flow of  $60 \text{ cm s}^{-1}$ ). Only 15 mm SL clams were slightly eroded in the medium sand treatment at  $u_* > 1.5 \text{ cm s}^{-1}$  (= unidirectional flow  $> 45 \text{ cm s}^{-1}$ ), corresponding to the erosion of the top few millimetres of sediment. Our data are in accordance with most of the flume studies focusing on the bedload transport of small juvenile clams (SL  $< 8 \text{ mm}$ ) (Gulmann et al., 2001; Hunt, 2004b, 2005; Jennings and Hunt, 2009; Roegner et al., 1995): the transport rate of small juvenile *M. arenaria* increases with increasing shear velocity and sediment fluxes. In our study, we used 1- to 2-year-old juveniles (SL  $> 8 \text{ mm}$ ), which are typically the age seeded by shellfish farmers in the Îles-de-la-Madeleine. Hunt (2004b) showed that the transport rate of 4.9 mm SL *M. arenaria* was negligible, and she suggested that these juveniles had reached a refuge size where they were not subjected to erosion by shear velocities  $< 1.6 \text{ cm s}^{-1}$ . Our unidirectional flume studies confirm this hypothesis. By contrast, St-Onge and Miron (2007) observed high clam dislodgment rates (up to 50%) for larger juvenile clams ( $10 \text{ mm} < \text{SL} < 20 \text{ mm}$ ), corresponding to a clean sand, heavy bedload transport for free stream velocities above  $29 \text{ cm s}^{-1}$  ( $u_* = 2.01 \text{ cm s}^{-1}$ ) applied for one hour. The mean shear velocity in our study was  $2.1 \text{ cm s}^{-1}$  while each trial lasted more than 5 hours, and no bulk sediment transport was observed for any of the sediments used. Further, the clams in their study were buried manually just below the surface at the onset of the experiment while those in the present study were allowed to bury by themselves for about 16 h. Such differing results may also be due to differences in the BBL flume flows, since St-Onge and Miron (2007) used a small flume (straight channel length =  $0.75 \text{ m}$ , width =  $0.2 \text{ m}$ ), which does not allow good BBL flow development and may be affected by strong lateral velocity gradients and high turbulence levels (Muschenheim et al., 1986; Vogel, 1994). The Aquatron flume produces a well-developed BBL with low lateral gradients (less than 10% in our trials), characteristics that are typical of recirculating flumes (Jonsson et al., 2006). In our experiment, although sediment erosion occurred for shear stresses above  $1.5 \text{ cm s}^{-1}$  (= unidirectional flow  $> 45 \text{ cm s}^{-1}$ ) and was highest on the clean medium sand treatment, no clam dislodgment was observed. Such results imply active burrowing by clams to reduce or avoid transport, as suggested by Roegner et al. (1995) and Lundquist et al. (2004a). Burrowing behaviour and burial depth varies according to water temperature, sediment type, clam size, shear velocities, and algal mats (Auffrey et al., 2004; Lundquist et al., 2004a; Pariseau et al., 2007; St-Onge et al., 2007; Zaklan and Ydenberg, 1997). With the exception of the high shear velocities and the size classes tested, our experimental conditions were favourable for burrowing, i.e., water temperature of  $18^\circ\text{C}$ , mean grain diameter  $< 500 \mu\text{m}$ , no algal mats. In the absence of

turbulence, we can conclude that juvenile clams (SL > 8 mm) are large enough not to be eroded by shear velocities exceeding 2 cm s<sup>-1</sup>.

#### **4.2 Effect of turbulence**

Our turbulence experiments revealed that juveniles with a mean shell length as large as 20 mm are still subject to displacement when they are submitted to high interacting wave–current flows. For TKE exceeding 10 J m<sup>-3</sup>, the transport rate of clams was above 80% for all the tested size classes, and this was mainly caused by the bulk transport of the sediment. In laboratory studies, Roegner et al. (1995) and Hunt (2005) clearly showed the positive relationship between sediment and clam transport at high flow speeds through bedload transport. Several studies (Hunt, 2005; Hunt et al., 2007; Hunt and Mullineaux, 2002) have confirmed these results on soft-shell clam recruits in the highly stratified Navesink estuary, where turbulence is dampened by pycnoclines (Fugate and Chant, 2005), as have other field studies conducted on different clam species on tidal flats (Beukema and Devlas, 1989; Norkko et al., 2001) or in subtidal habitats (Olivier et al., 1996; Olivier and Retiere, 1998). The effect of turbulence through wind stress and associated waves has been observed in the field on the post-larvae and juveniles of many invertebrate species (Olivier et al., 1996), including bivalves (Emerson, 1991; Lundquist et al., 2006). Concerning *Mya arenaria* juveniles, Emerson (1991) and Emerson and Grant (1991) have clearly demonstrated, using sediment traps, that bivalve dispersal was correlated to wind-driven sediment transport, especially when wind speeds exceed 5–7 m s<sup>-1</sup>. In their experiments, bedload transport of juveniles (mean = 9 mm SL; up to 28 mm) was important, with maximum values in the traps of 2000 ind. m<sup>-1</sup> d<sup>-1</sup>. In the Îles-de-la-Madeleine, periods of high turbulence corresponding to autumn storms could induce bedload transport of seeded clam juveniles. We measured TKE exceeded 10 J m<sup>-3</sup> (up to 20 J m<sup>-3</sup>) in field conditions and demonstrated that this was a threshold value for eroding both sediment and seeded clams.

#### **4.3 Expected distance of dispersal**

Some recent studies have focused on the expected spatial scales for bedload transport through field or modelling experiments. Hydrodynamic models predict that the average distance travelled by bedload-dispersing organisms in shallow estuaries is on the order of 100 m d<sup>-1</sup> (Lundquist et al., 2004b). Field observations by Lundquist et al. (2006) revealed that rates of passive bedload transport, as well as resuspension higher into the water column, increased with wave activity, suggesting that storm-related transport is an important process for long-distance recolonization. Norkko et al. (2001) measured the dispersal distances of *Maconoma liliiana* juveniles in the field and calculated that individuals could travel over several metres within a tidal cycle on an intertidal sandflat. On the same species, Petuha et al. (2006) estimated that bedload transport of juveniles could account for a dispersal distance of 80 m on a spring tide. For *Mya arenaria* juveniles, recent work by Jennings and Hunt (2009) assessed bedload dispersal rates of 2 to 40 cm h<sup>-1</sup> for small individuals (SL < 1.5 mm) in relatively high flume flows (1.13 < u\* < 1.30 cm s<sup>-1</sup>) and of 10 cm h<sup>-1</sup> in the field. Based on results from a 1D numerical bedload transport model of juvenile clams by tidal currents, including *M. arenaria* recruits, Hunt et al. (2009) concluded that dispersal distances could reach several kilometres in one month (100 m d<sup>-1</sup>) and would be responsible for modifying patterns of distribution and abundance. Such processes could explain in part the high recruitment variability observed for soft-shell clams in the field at scales of metres to kilometres (Bowen and Hunt, 2009; Emerson and Grant, 1991; Hunt et al., 2003; LeBlanc and Miron, 2006). In HAM lagoon, residual tidal currents rapidly decrease away from the inlets, with residual flows almost vanishing near the shellfish farming concessions (Guyondet and Koutitonsky, 2008). According to Koutitonsky and Tita (2006), wind events during storms would probably alter this residual circulation, but their influence



would be limited to a few tidal cycles. In HAM lagoon, strong local winds are predominantly blowing from the western sector (Drapeau, 1988), and we should thus expect bedload transport of juveniles from the tidal flats to the deeper parts of the lagoon, which would contribute to a loss of stock for shellfish farmers.

#### **4.4 Other potential causes of juvenile *Mya* loss**

The loss of juvenile clams during autumn could be attributed to post-settlement processes such as predation (Beal and Kraus, 2002; Hunt et al., 2003; Hunt and Mullineaux, 2002). In HAM lagoon, potential predators are the rock crab *Cancer irroratus*, the green crab *Carcinus maenas* (newly arrived in the Îles-de-la-Madeleine) and the nemertean *Cerebratulus lacteus*, which could drastically affect populations of soft-shell clam adults and juveniles (Bourque et al., 2002; Miron et al., 2005). The risk of predation by epibenthic crabs generally decreases with increasing size/age (Sousa, 1993). However, Miron et al. (2005) showed that the native rock crab is a potential threat for commercial bivalves, particularly juveniles (SL < 25 mm).

Predator foraging may increase sediment erosion and thus displace juveniles (Hunt, 2004a) by decreasing the  $u_{*c}$  for bedload transport. More generally, bedload transport of both sediment and juveniles depends on a variety of biological factors that modify critical shear velocities to erode the substratum, including stabilization or destabilization due to infauna (Olivier et al., 1996; Volkenborn et al., 2009; Widdows and Brinsley, 2002)

## **Conclusion**

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Our flume studies have clearly shown that high levels of turbulence could explain the loss of commercial juvenile soft-shell clams in the Îles-de-la-Madeleine by wind generated waves during storms. We now have to quantify these processes in the field using bedload traps to assess scales of dispersal and turnover of the seeded population, and to test the hypothesis of predation by rock crabs, green crabs, and nemerteans, which were rejected until now to explain high losses of seed.

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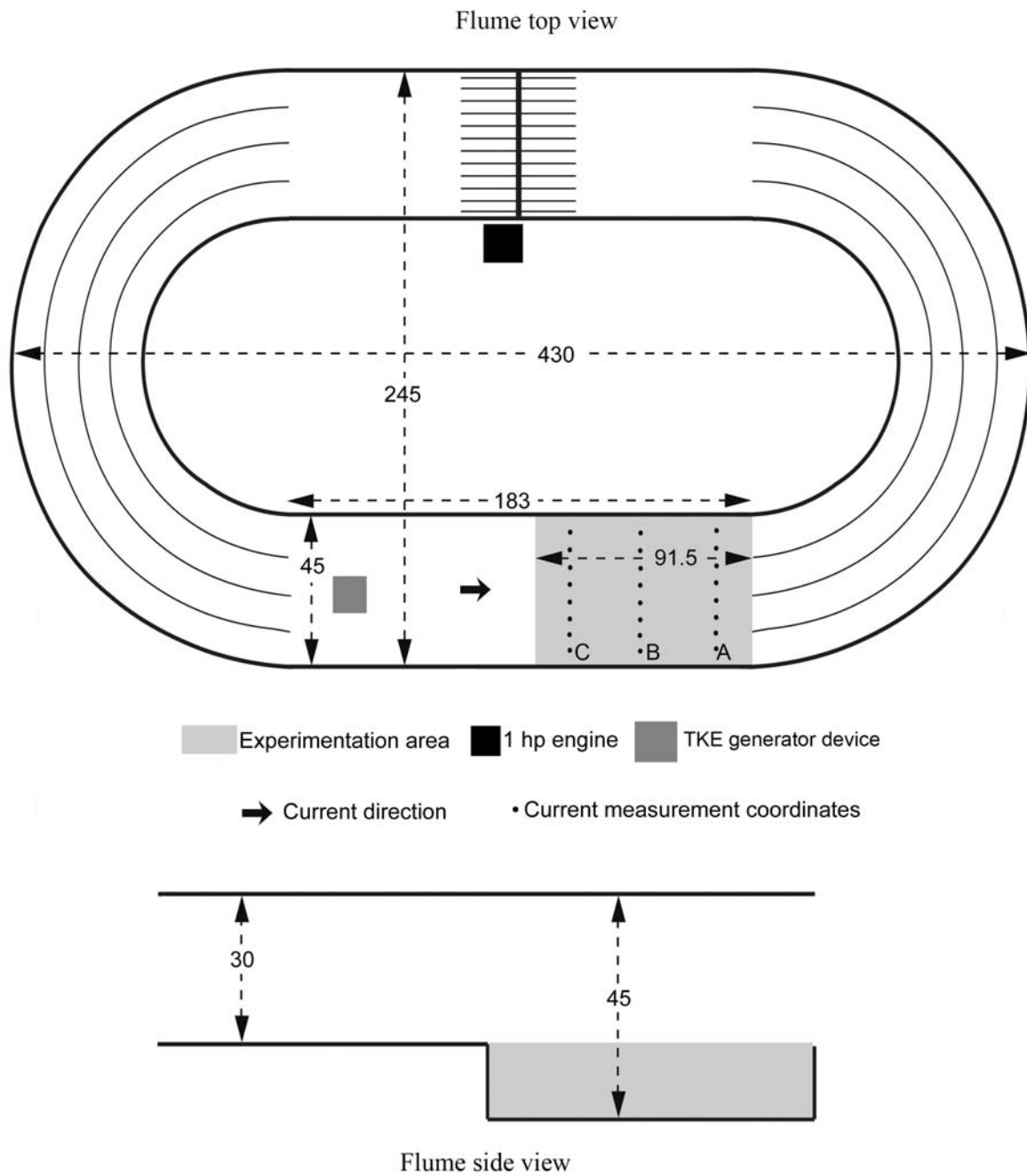
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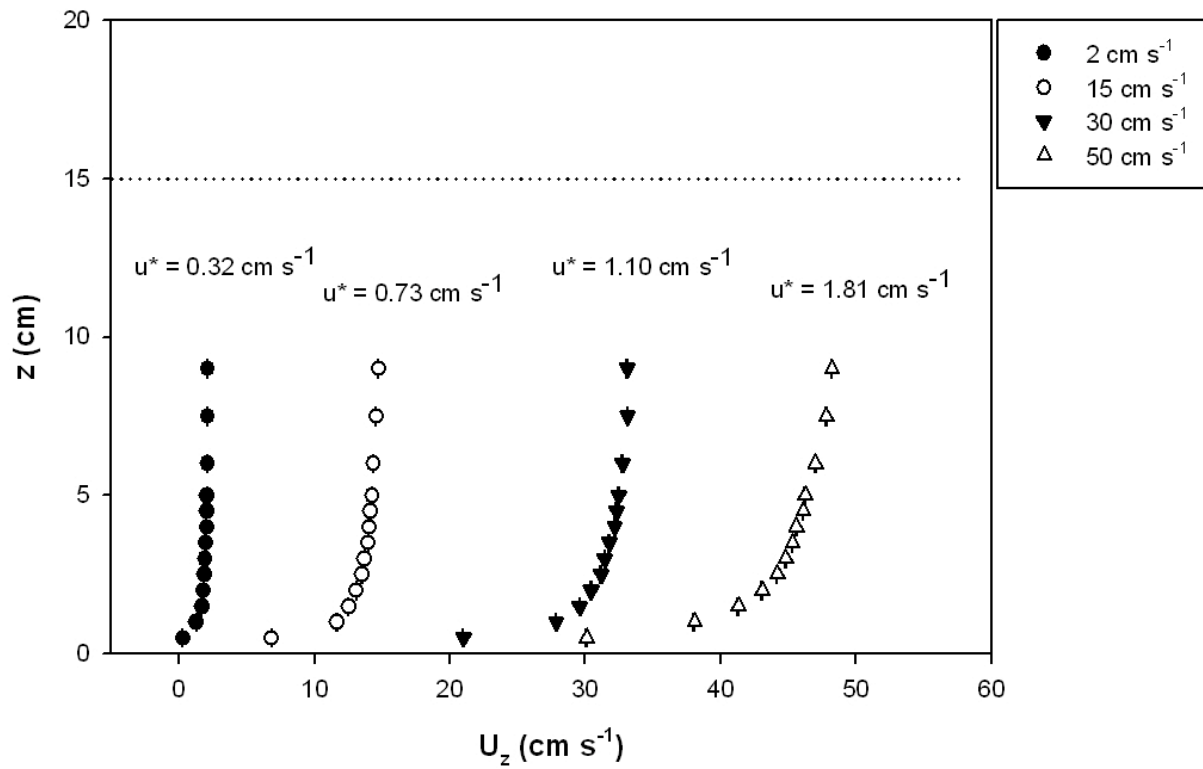
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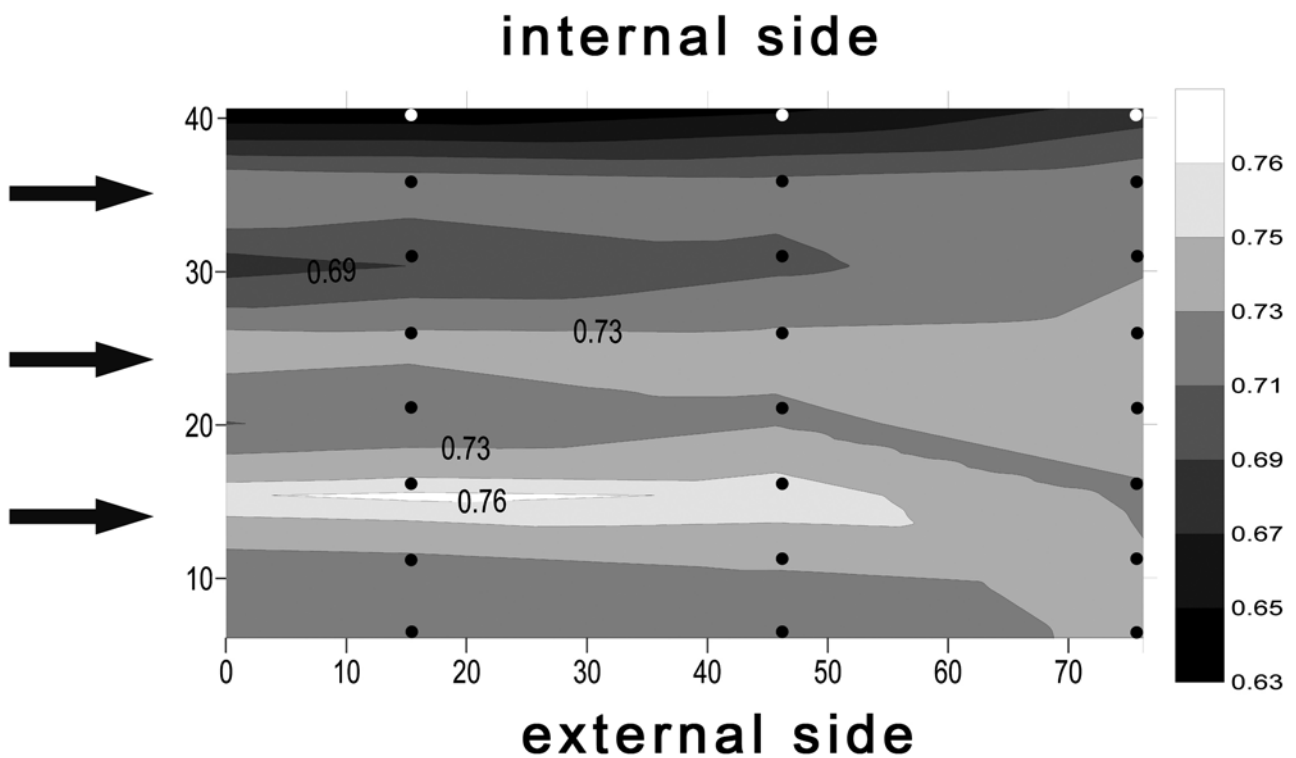
## Figures



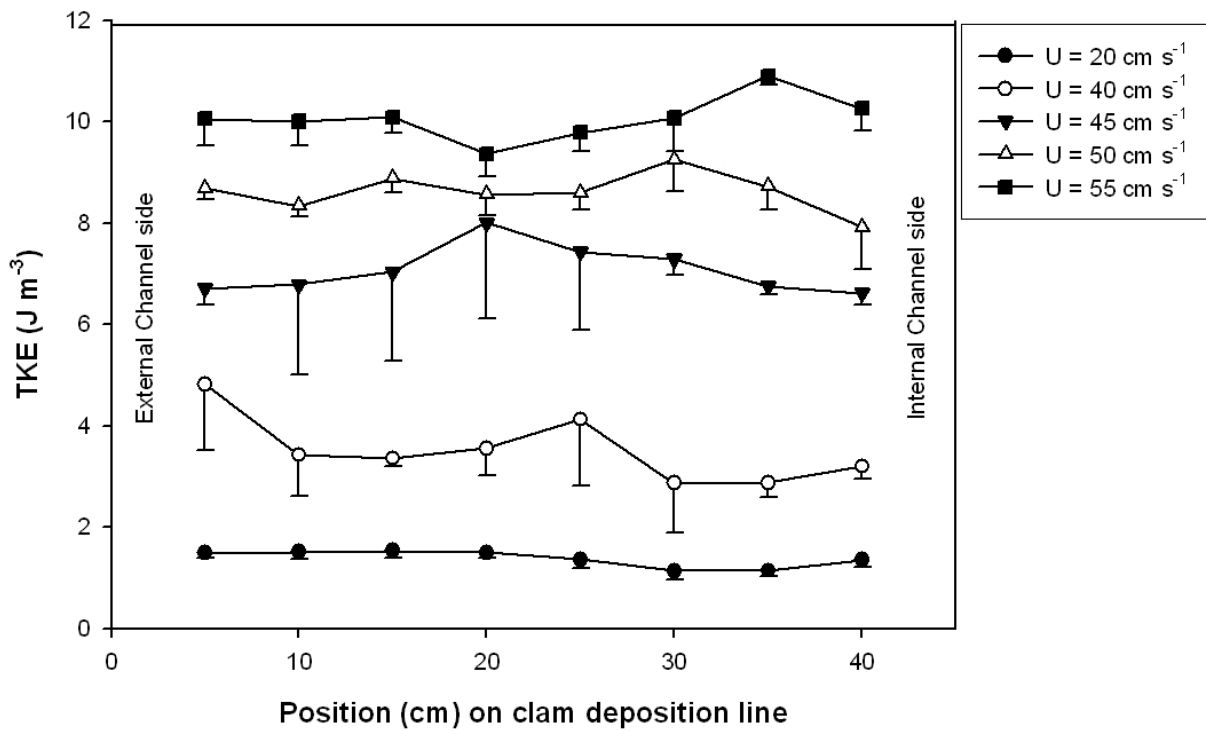
**Figure 1.** Schematic drawing of the benthic biological flume Aquatron, with positions of flow-measurement transects A, B (also clam position line), and C. Dimensions are in cm.



**Figure 2.** Vertical profiles for increasing velocities ( $U_\infty$ ) (mean  $\pm$  SD) in the benthic flume at different height ( $z$ ) values. The horizontal dotted line indicates the water level in the flume (15 cm).



**Figure 3.** Horizontal variations of  $u_*$  ( $\text{cm s}^{-1}$ ) across the experimental zone for a  $15 \text{ cm s}^{-1}$  free stream velocity (dimensions are given in cm).

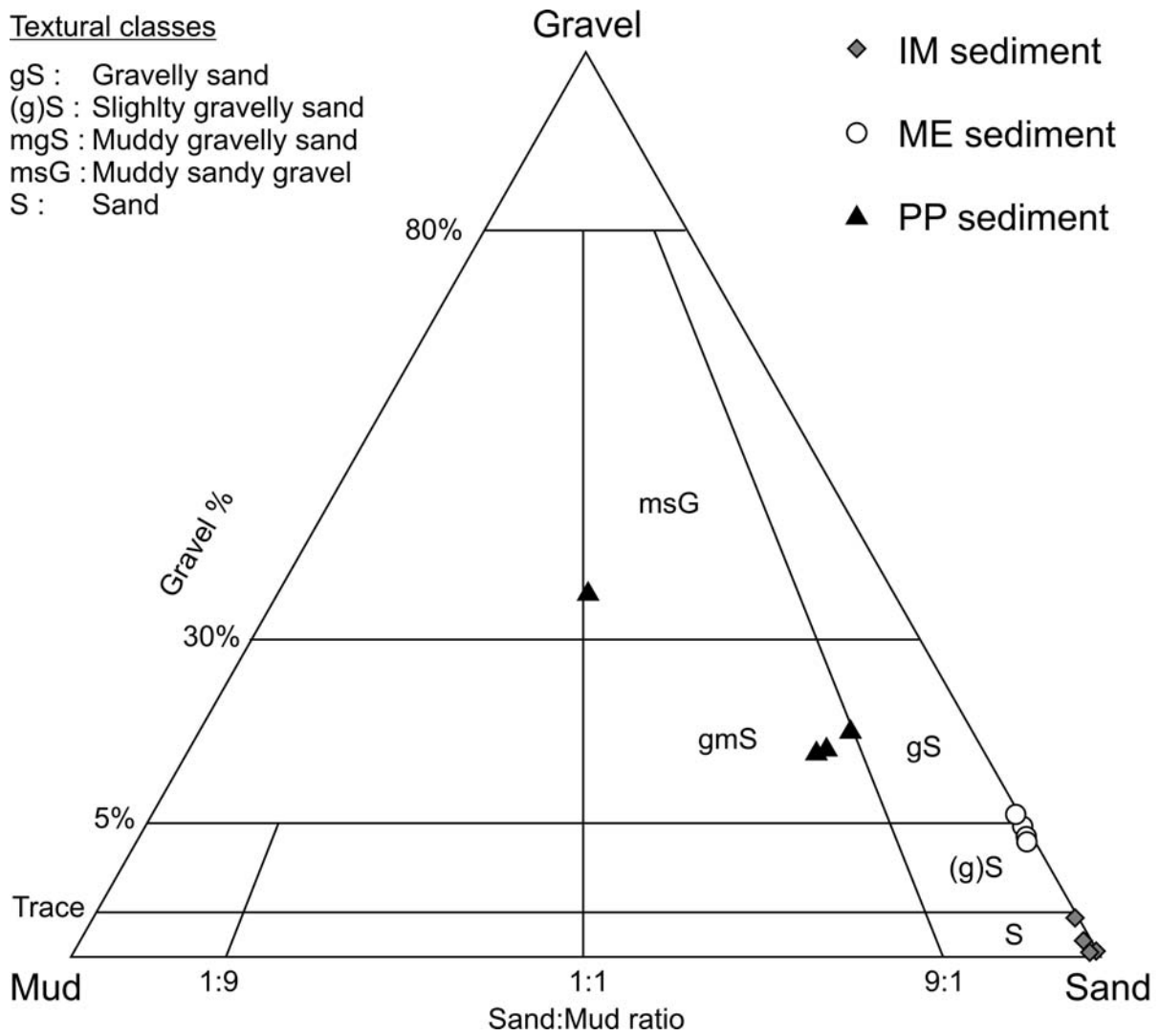


**Figure 4.** Turbulent kinetic energy (TKE) (mean  $\pm$  SD) during the wave-current experiments measured along the clam position line across the channel (B in Fig. 1), five current steps with constant wave-generator frequency of 120 CPM.

Textural classes

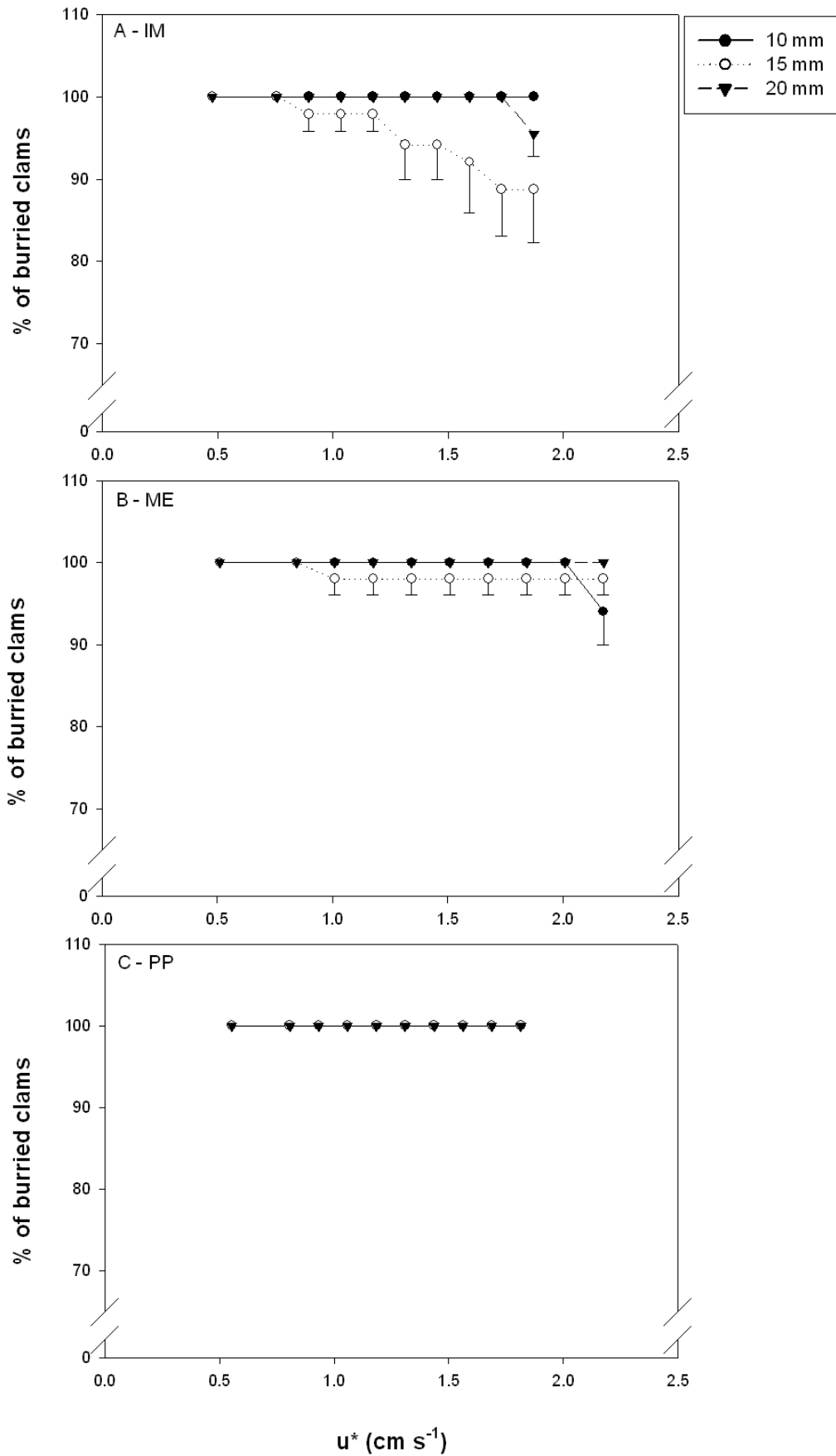
gS : Gravelly sand  
 (g)S : Slightly gravelly sand  
 mgS : Muddy gravelly sand  
 msG : Muddy sandy gravel  
 S : Sand

◆ IM sediment  
 ○ ME sediment  
 ▲ PP sediment



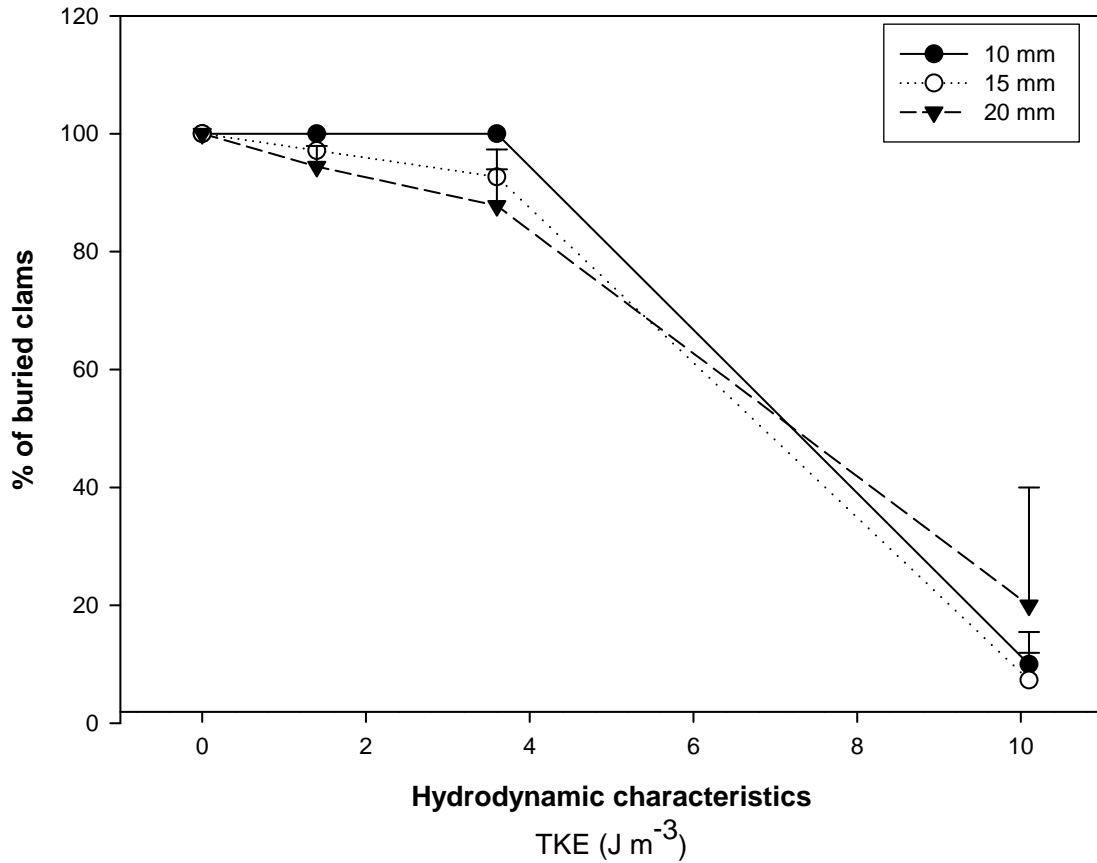
**Figure 5.** Grain size analyses of the three experimental substrates presented in the ternary gravel–sand–mud diagram with textural classes according to Folk (1954).



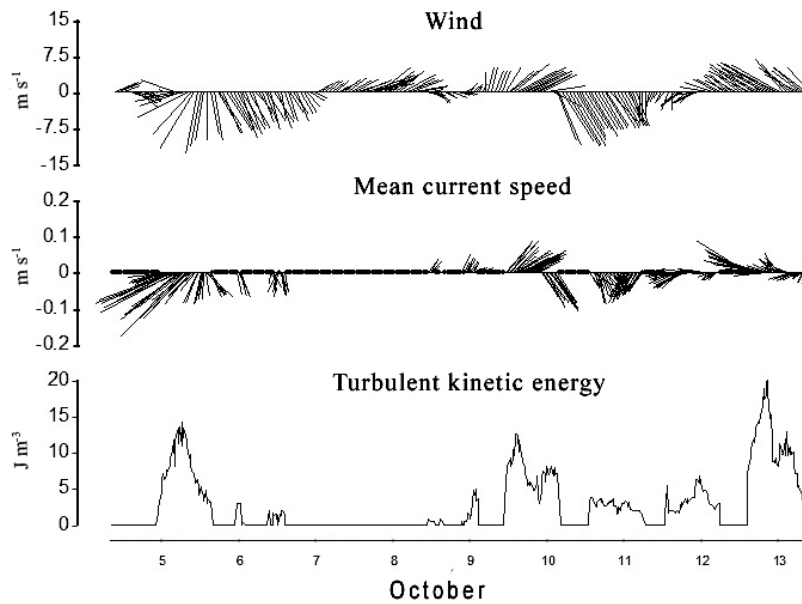


**Figure 6.** Percentage of buried clams (mean  $\pm$  SD) for the three length classes in different substrates after exposition to unidirectional currents of different velocities in the experimental

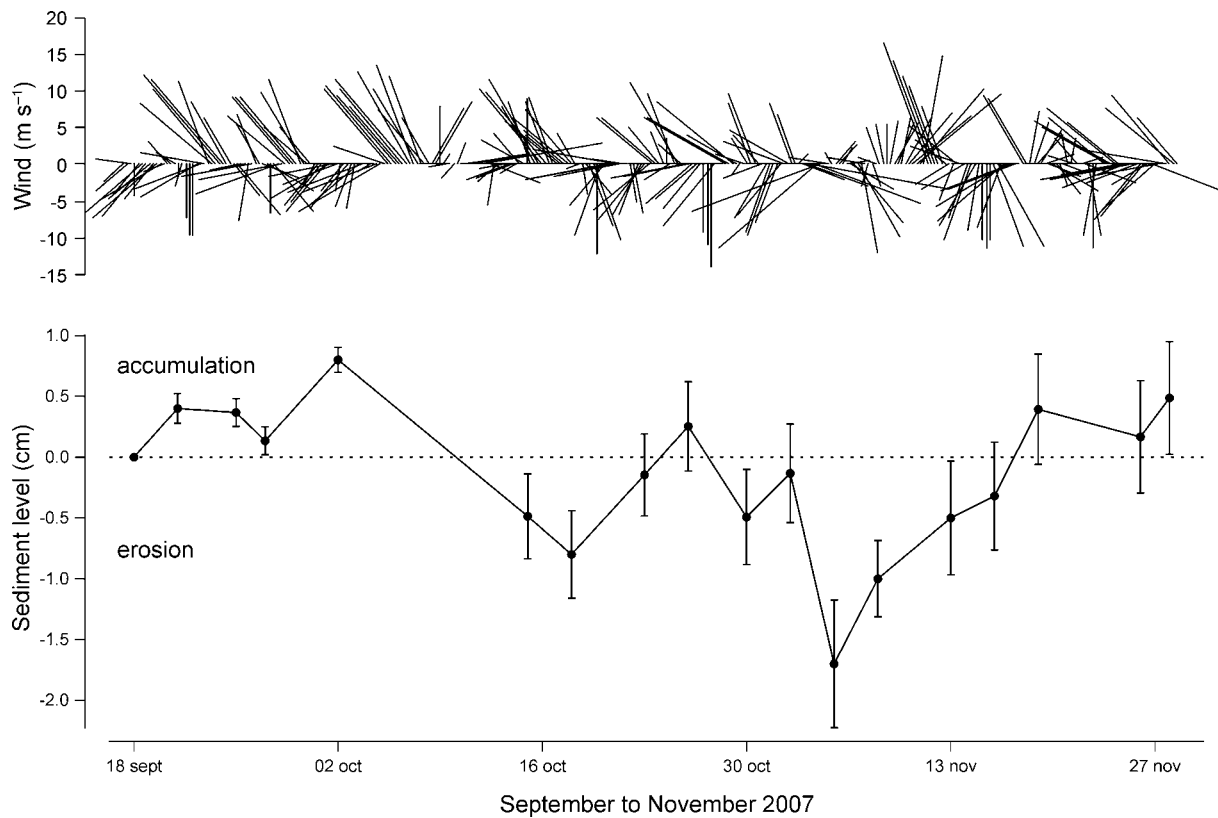
benthic flume: a) medium sand from Îles-de-la-Madeleine (IM), b) coarse sand from Métis (ME), c) muddy sand from Pointe-au-Père (PP).



**Figure 7.** Percentage of buried clams (mean  $\pm$  SD) in medium sand (IM) when exposed to wave generator action (120 CPM) with an increasing current velocity ( $cm s^{-1}$ ) in the experimental benthic flume.



**Figure 8.** Environmental conditions measured at the Havre-aux-Maisons field site from 5 to 14 October 2006.



**Figure 9.** Changes in bed levels (mean  $\pm$  SD) measured at the Havre-aux-Maisons field site from September to November 2007.

## Tables

Free stream Velocity (cm s <sup>-1</sup> )	15	30	45	60
<b>Mean U* ± SD (cm s<sup>-1</sup>)</b>				
<b>IM</b>	0.44 ± 0.05	1.25 ± 0.13	1.55 ± 0.22	1.73 ± 0.44
<b>ME</b>	0.59 ± 0.06	1.28 ± 0.08	1.73 ± 0.16	2.10 ± 0.33
<b>PP</b>	0.56 ± 0.10	1.20 ± 0.15	1.51 ± 0.31	1.72 ± 0.41
<b>Re</b>				
<b>IM</b>	10343.75	29483.33	36596.35	40808.85
<b>ME</b>	13340.63	28942.19	39245.31	47589.58
<b>PP</b>	9052.60	19389.58	24445.83	27723.96
<b>Re*</b>				
<b>IM</b>	1.17	3.33	4.13	4.61
<b>ME</b>	2.12	4.59	6.22	7.55
<b>PP</b>	15.90	34.05	42.93	48.69
<b>Mean TKE ± SD (J m<sup>-3</sup>)</b>				
<b>IM</b>	0.14 ± 0.11	0.63 ± 0.35	0.89 ± 0.61	1.78 ± 1.09
<b>ME</b>	0.14 ± 0.02	0.56 ± 0.09	0.99 ± 0.15	1.13 ± 0.89
<b>PP</b>	0.07 ± 0.05	0.15 ± 0.15	0.24 ± 0.19	0.28 ± 0.24

**Table 1.** Hydrodynamic characteristics (mean ± SD) above three different sediments (Medium sand - IM, coarse sand - ME and muddy sand - PP). Re is measured for a 15 cm water column.