

Calcium distribution in the world ocean waters

Calcium/chlorinity ratio
Salinity
Calcium carbonate
Gulf of Elat
Rapport calcium/chlorinité
Salinité
Carbonate de calcium
Golfe d'Eilat

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ABSTRACT

The calcium/chlorinity ratio is known to be a very useful tool for studying some phenomena occurring in natural marine systems, such as water mass circulation, mixing and diffusion through coastal and open sea water, geochemical and biological precipitation of calcium carbonate, etc. Therefore numerous determinations of calcium/chlorinity ratios have often been carried out on a regional basis. However, a complete analysis of all the existing data has been almost entirely lacking.

A consecutive statistical analysis of all existing literature data from 1884 to 1980 concerning calcium/chlorinity ratios in different locations all over the world showed that this ratio for open seas and oceans is equal to $0.02128 \pm 0.00015 \text{ g. kg}^{-1} (\text{Cl}\text{‰})^{-1}$. The value obtained practically coincides with the value earlier suggested by Culkin and Cox in 1966.

In addition, in this work, the calcium distribution in the nearshore waters of the Gulf of Elat (Aqaba) has been determined. The seawater samples were collected along the shore of the Sinai Peninsula at 13 stations in the vicinity of a coral reef. The calcium/chlorinity ratios of the nearshore waters of the Gulf of Elat (0.02083 ± 0.00005) were found to be considerably less than those for water masses of open seas and oceans. This fact was explained by the biological precipitation of calcium carbonate from the adjacent surface waters of the gulf.

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RÉSUMÉ

La distribution du calcium dans les mers et océans du monde

Il est bien connu que le rapport calcium/chlorinité est un excellent instrument pour l'étude de certains phénomènes rencontrés dans le milieu marin, tels que la circulation des masses d'eau, le mélange et la diffusion entre les eaux côtières et les eaux de haute mer et la précipitation géochimique et biologique du carbonate de calcium. En conséquence, de nombreuses déterminations du rapport calcium/chlorinité ont été effectuées sur une base régionale. Cependant une interprétation générale de toutes les données existantes manque presque complètement.

L'analyse statistique des données publiées de 1884 à 1980, concernant les rapports calcium/chlorinité en différents endroits à travers le monde a montré que, pour les mers et océans, ce rapport est égal à $0.02128 \pm 0.00015 \text{ g. kg}^{-1} (\text{Cl}\text{‰})^{-1}$, valeur qui coïncide pratiquement avec celle suggérée antérieurement par Culkin et Cox en 1966.

Le présent travail examine la distribution du calcium dans les eaux, proches du rivage, du golfe d'Eilat. Les échantillons d'eau ont été prélevés à 13 stations situées sur la côte du Sinai au voisinage d'un récif corallien. Les rapports calcium/chlorinité des eaux du golfe d'Eilat sont nettement inférieurs à ceux des mers et océans. Ce fait s'explique par la précipitation biologique du carbonate de calcium des eaux superficielles venant des régions avoisinantes du golfe.

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INTRODUCTION

Marcet (1819) was the first to suggest that the major components of seawater occur together in constant proportions. This hypothesis has been corroborated by Forchhammer (1865), who analyzed hundreds of water samples, and by Dittmar (1884), who analyzed 77 samples of seawater collected by the Challenger Expedition from oceans around the world. Dittmar concluded that the ratio of alkaline earth elements to chlorinity is nearly constant in all open oceans. It is now commonly accepted that in most of the seas and oceans, the ratios of all major ions to salinity (or chlorinity) are constant within the analytical precision of proper ion determination (Culkin, 1965; Culkin, Cox, 1966; Riley, Tongudai, 1967; Riley, Chester, 1971; Whitfield, Turner, 1979). However, modern precise analytical procedures have shown that although the hypothesis of constant ionic composition is largely correct, some variations occur in the composition of sea water bodies from different regions. Some conflicting evidence concerning variability in the ratios of some major ions to chlorinity (salinity) has recently appeared. These phenomena have been observed especially in some closed lagoons (Krumgalz *et al.*, 1980), in the Baltic Sea region with its large inflow of river waters (Zarins, Ozolins, 1935; Gripenberg, 1937; Anderson, Dyrssen, 1981), in the Red Sea with its hot brines (Brewer *et al.*, 1965; Miller *et al.*, 1966; Brewer, Spencer, 1969; Danielsson *et al.*, 1980), etc. As the analytical precision of different ion determinations in seawater is improved, new information concerning some chemical processes in the water bodies of oceans can be obtained.

Seawater (up to saturation depth) is known to be practically saturated with respect to calcium carbonate in the form of calcite or aragonite (Sillen, 1961; Schmalz, Chave, 1963; MacIntyre, Platford, 1964; Pytkowicz, 1969, 1970; Lyakhin, 1971; Tsunogai *et al.*, 1971; Ben-Yaakov *et al.*, 1974; Bezborodov *et al.*, 1978; Morse *et al.*, 1979). This fact is in agreement with some observations on the supersaturation with calcium carbonate of surface seawater (Cloud, 1962; MacIntyre, Platford, 1964; Wells, Illing, 1964; Groot, 1965; Broecker, Takahashi, 1966; Pytkowicz, Fowler, 1967; Lyakhin, 1968; Chave, Suess, 1970; Edmond, Gieskes, 1970; Alekin, Katunin, 1974). However, the delay of calcium carbonate precipitation from sea water can probably be explained by some kinetic factors: the effects of some inorganic ions (Simkiss, 1964; Pytkowicz, 1965; 1973; Weyl, 1965; Bischoff, 1968; Lippman, 1973; Berner, 1975), phosphates (Raistrick, 1949; Reddy, Nancollas, 1973; Berner *et al.*, 1978) or naturally occurring organic compounds (Chave, 1965; Kitano, Hood, 1965; Chave, Suess, 1967; 1970; Meyers, Quinn, 1971; Berner *et al.*, 1978) on the rate of CaCO_3 nucleation from natural saline waters. Therefore calcium ions can be considered as a quasi-conservative constituent of seawater, but sometimes, under certain conditions, calcium carbonate precipitation does occur. Recently such a phenomenon has been observed for some natural systems: the Persian Gulf (Wells, Illing, 1964), small hard-water lakes (Wetzel,

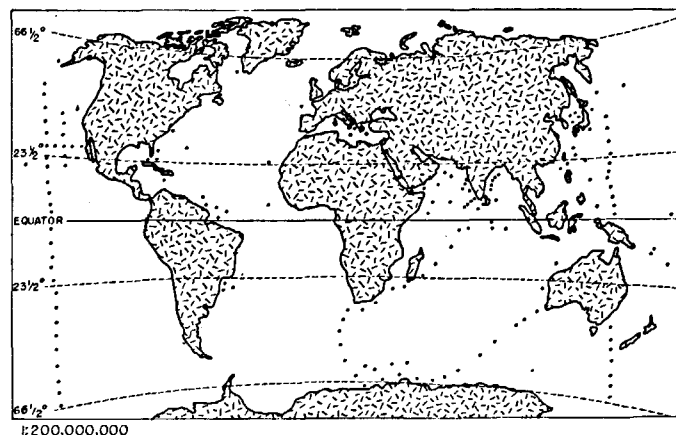


Figure 1
Map of the station location for calcium determination in seawater.

1966), shallow carbonate marine environments (Broecker, Takahashi, 1966) and the Great Lakes (Strong, Eadie, 1978). Thus, changes in the calcium/chlorinity ratio can characterize some natural processes occurring in seawater, such as calcium carbonate precipitation, the mixing of river water with sea water (since the calcium/chlorinity ratio for river water is much higher than that of seawater—Hutchinson, 1957), and the assimilation of calcium from seawater by various kinds of marine organisms to form calcareous shells, etc. Freemantle *et al.* (1978) tried to use Ca^{2+} ion distribution in the Gulf of Elat (Aqaba) to study seawater pollution by phosphate rock. Therefore numerous determinations of calcium/chlorinity (salinity) ratios have often been carried out on a regional basis. Figure 1 presents the location of the sampling station where seawater samples were taken during different expeditions for calcium determination. However, a complete analysis of all data concerning the calcium distribution in the world ocean waters has been almost entirely lacking. In this article we try to conduct such an analysis.

RESULTS AND DISCUSSION

All available data concerning calcium content for different seas and oceans from 1884 to 1980 are presented in Table 1. It should be emphasized that in many cases, the experimental literature data were recalculated by us to common units $\text{g} \cdot \text{kg}^{-1} (\text{Cl} \%)^{-1}$ with seawater density computed from the equation of state of seawater (Millero *et al.*, 1976).

The results shown in Table 1 can be divided into three groups: 1) the first group, with a calcium/chlorinity ratio equal to $0.02586 \pm 0.0013 \text{ g} \cdot \text{kg}^{-1} (\text{Cl} \%)^{-1}$, includes the Baltic Sea, the Gulf of Finland and the Gulf of Bothnia; 2) the second group includes hot brines in the Red Sea, with calcium/chlorinity ratios equal to $0.0308 \pm 0.0018 \text{ g} \cdot \text{kg}^{-1} (\text{Cl} \%)^{-1}$; 3) the third group includes all other data, ranging from 0.02033 to $0.02258 \text{ g} \cdot \text{kg}^{-1} (\text{Cl} \%)^{-1}$, with mean value 0.02137. The question is whether the difference between these three groups is significant, or whether they can be considered as coming from the same population. The *t*-

Table 1
Calcium/Chlorinity ratios throughout the world oceans and seas.

Location	$\text{Ca}^{2+}/\text{Cl}^{\circ}/_{\text{oo}}^{-1}$ $\text{g. kg}^{-1} (\text{Cl}^{\circ}/_{\text{oo}})^{-1}$	References	Remarks and coordinates
<i>Group 1</i>			
Baltic Sea:		Zarins, Ozolins, 1935	
surface water	0.026 48 ± 0.000 06		
bottom water	0.024 26 ± 0.000 07		
Baltic Sea:		Gripenberg, 1937	57°22'-59°46'N; 18°26'-20°24'E.
surface water	0.026 19 ± 0.000 24		
bottom water	0.024 20 ± 0.000 34		
Gulf of Bothnia	0.026 49 ± 0.000 17	Gripenberg, 1937	60°11.5'-65°38'N; 17°52'-23°47'E. The $\text{Ca}^{2+}/\text{Cl}^{\circ}$ ratios were identical for surface and bottom waters (distances from bottom were 1-27 m and maximum depth was 175 m).
Gulf of Finland		Gripenberg, 1937	59°21'-60°31'N; 22°59'-28°48'E.
surface water	0.027 7 ± 0.001 3		
bottom water	0.025 7 ± 0.000 8		
Baltic Sea	-	Anderson, Dyrssen, 1981	The authors obtained the following relation, which should be valid for the Danish Sounds as well as the Baltic Sea: $\text{Ca} (\text{g/kg}) = 0.020 4 \text{ Cl}^{\circ}_{\text{oo}} + 0.022 5$. This equation is identical to the formula presented by Gripenberg (1937).
<i>Group 2</i>			
Hot brines in the Red Sea:			
a) Discovery Deep	0.030 3 0.033 0 0.029 8	Brewer <i>et al.</i> , 1965 Brewer, Spencer, 1969 Danielsson <i>et al.</i> , 1980	The Ca/Cl ‰ ratio was calculated by us from their graphical data.
b) Atlantis II Deep	0.033 0 0.033 0 0.030 9 0.028 9 0.030 2	Miller <i>et al.</i> , 1966 Brewer, Spencer, 1969 Brewer, Spencer, 1969 Danielsson <i>et al.</i> , 1980 Danielsson <i>et al.</i> , 1980	A layer with temperature 56°C. A layer with temperature 44°C. An intermediate brine. A deep brine.
c) Chain Deep	0.028 2	Brewer, Spencer, 1969	
<i>Group 3</i>			
All oceans	0.021 28	Dittmar, 1884	Only four samples.
North Pacific Ocean	0.021 15	Thompson, Wright, 1930	
Pacific Ocean	0.021 35	Kirk, Moberg, 1933	
Western Pacific Ocean	0.021 20	Miyake, 1939	
Tokyo Bay	0.021 30	Matida, 1950	
Pacific Ocean (Gulf of Alaska);		Chow, Thompson, 1955	
surface water	0.021 96		
5000 m depth	0.021 88		
Atlantic Ocean	0.021 22	Carpenter, 1957	Samples were from open ocean water.
Bahamas Bank	0.020 63	Carpenter, 1957	Samples were from a region of active precipitation of CaCO_3 on the Bahamas Bank. 48°12'N, 127°17'W.
Pacific Ocean:		Pate, Robinson, 1958	
surface water	0.021 41		
1559 m depth	0.021 33		
Pacific Ocean	0.021 8 ± 0.001 0	Sugawara, Kawasaki, 1958	
Indian Ocean	0.022 0 ± 0.000 4	Sugawara, Kawasaki, 1958	
Antarctic seas	0.022 2 ± 0.000 2	Sugawara, Kawasaki, 1958	
Labrador Sea	0.021 7	MacIntyre, Platford, 1964	The results obtained were characteristic of total Ca^{2+} and Sr^{2+} content in seawater. The $\text{Ca}^{2+}/\text{Cl}^{\circ}$ ratios were calculated by us. Considerable CaCO_3 precipitation was observed by the authors. Approximately 70 samples from major oceans and seas were carefully analyzed. A year later, Riley and Tongudai (1967) applied other methods for the analysis of the same seawater samples. The results obtained by these two groups of scientists were in excellent agreement. Therefore, for statistical analysis, only the 1966 values were taken.
Great Bahama Banks	0.020 68 ± 0.000 25	Broecker, Takahashi, 1966	
North Pacific Ocean	0.021 28	Culkin, Cox, 1966	
South Pacific Ocean	0.021 28	Culkin, Cox, 1966	
North Atlantic Ocean	0.021 28	Culkin, Cox, 1966	
South Atlantic Ocean	0.021 20	Culkin, Cox, 1966	
Northern seas	0.021 21	Culkin, Cox, 1966	
Southern seas	0.021 30	Culkin, Cox, 1966	
Indian Ocean	0.021 24	Culkin, Cox, 1966	
Mediterranean Sea	0.021 31	Culkin, Cox, 1966	
Red Sea	0.021 15	Culkin, Cox, 1966	
Persian Gulf	0.021 23	Culkin, Cox, 1966	
North Sea	0.021 18	Culkin, Cox, 1966	
Baltic Sea	0.021 27	Culkin, Cox, 1966	
Indian Ocean	0.021 52	Dietrich <i>et al.</i> , 1966	
Atlantic off South Coast of Brazil	0.021 56	Kato, 1966 a	
Lagoon of Cananea (Brazil)	0.020 6	Kato, 1966 b	The author suggested that a biogeochemical precipitation of dissolved calcium took place in the lagoon water.

Table 1 (continuation)

Location	$\text{Ca}^{2+}/\text{Cl}^{\circ}/_{\text{oo}} \text{g. kg}^{-1} (\text{Cl}^{\circ}/_{\text{oo}})^{-1}$	References	Remarks and coordinates
Coastal water of Cananeia (Brazil)	0.02168	Kato, 1966 <i>b</i>	
Gulf of Trieste	0.0218 ± 0.0005	Mameli, Mosetti, 1966	
Atlantic Ocean:		Fabricand <i>et al.</i> , 1967	
<i>a</i>) surface water	0.0218		18°59'N, 22°30.5'W.
3 250 m depth	0.0210		
<i>b</i>) surface water	0.0209		10°56'N, 49°30'W.
4 750 m depth	0.0211		
Straits of Florida:		Traganza, Szabo, 1967	
<i>a</i>) at the edge of Great Bahama Bank	0.02113		
<i>b</i>) on the Great Bahama Bank	0.02033		
Northern Adriatic Sea	0.02124	Mameli, Mosetti, 1968	44°56.5'N, 13°25.4'E to 44°58.2'N, 12°38.0'E.
Otranto Channel	0.02127	Mameli, Mosetti, 1968	39°52.1'N, 18°34.2'E to 39°51.2'N, 19°18.1'E.
A region between Greece and the Island of Crete	0.02114	Mameli, Mosetti, 1968	36°17.2'N, 22°35.5'E to 35°30.2'N, 23°19.6'E.
Mediterranean Sea	0.02116	Mameli, Mosetti, 1968	
Western North Pacific Ocean:			
<i>a</i>) the Oyashio region			
surface water	0.02121 ± 0.00004	Tsunogai <i>et al.</i> , 1968 <i>a</i>	
depth > 1 000 m	0.02132 ± 0.00003		
<i>b</i>) the Kuroshio region			
surface water	0.02108 ± 0.00003		
depth 500-1 000 m	0.02120 ± 0.00005		
Pacific Ocean		Tsunogai <i>et al.</i> , 1968 <i>b</i>	38°01'N, 143°58'E.
surface water	0.02217		The author noted that the values obtained characterized the total amount of Ca^{2+} and Sr^{2+} in seawater.
Sargasso Sea	0.02068 ± 2.79%	Billings <i>et al.</i> , 1969	32°10'N, 64°30'W. A strong temporal variation throughout the water column (up to 2600 m) was discovered.
Western Pacific Ocean	0.0214	Sagi, 1969	The $\text{Ca}^{2+}/\text{Cl}^{\circ}/_{\text{oo}}$ ratios ranged from 0.02107 to 0.02174.
Western Atlantic Ocean (near South America)		Lyakhin, 1971	
surface water	0.02155 ± 0.00024		
5 000 m depth	0.0220		
Indian Ocean			
surface water	0.02104	Tsunogai <i>et al.</i> , 1971	
Antarctic Ocean			
surface water	0.02116	Tsunogai <i>et al.</i> , 1971	
deep water	0.02116		
South China Sea	0.0215	Chen, Jeng, 1972	The $\text{Ca}^{2+}/\text{Cl}^{\circ}/_{\text{oo}}$ ratios ranged from 0.02053 to 0.02229. The precision of calcium determination was ± 2.2%.
Caribbean Sea (south of Puerto Rico)		Atwood <i>et al.</i> , 1973	17°38'N, 67°00'W.
surface water	0.0208		
depth 1 500 m	0.0210		
Pacific Ocean		Tsunogai <i>et al.</i> , 1973	
surface water	0.02104		
deep water	0.02126 ± 0.0006		
mean of the whole			
Pacific Ocean	0.02123		
South Pacific Ocean		Horibe <i>et al.</i> , 1974	There is depth distribution. The stations occupied were in the western boundary region of the South Pacific between American Samoa and Tonga.
surface water			
(up to 500 m)	0.02134 ± 0.00003		
deep water			
(from 500 to 5 200 m)	0.02149 ± 0.00004		
Central Ionian Sea		Magazzu, Corigliano, 1974	
surface water	0.02153		
depth 1 000-4 000 m	0.02145		
Eastern coasts of Sicily		Magazzu, Corigliano, 1974	
surface water	0.02258		
Tropical North Atlantic Ocean	0.02120 ± 0.00004	Lebel, Poisson, 1976	
Bashi Channel (near Taiwan)	0.0205 ± 0.004	Li, Chen, 1976	The precision of calcium determination was ± 2.0%.
Equatorial Pacific Ocean		Almgren <i>et al.</i> , 1977	The authors could explain the relation between $\text{Ca}/\text{Cl}^{\circ}/_{\text{oo}}$ and $\text{A}/\text{Cl}^{\circ}/_{\text{oo}}$ by including the acid produced by the oxidation of peptides and fatty acids.
surface water	0.02136		Calcium concentration of nearshore waters showed significant seasonal variations; therefore these data were not included in the statistical calculations.
depth 1 100 m	0.02154		
Northern Indian Ocean		Naik, 1978	
Arabian Sea (nearshore waters of Goa)	0.0213 ± 0.0019		

Table 1 (end)

Location	$\text{Ca}^{2+}/\text{Cl}^{\circ}/\text{‰}$ $\text{g.kg}^{-1} (\text{Cl}^{\circ}/\text{‰})^{-1}$	References	Remarks and coordinates
Northern Indian Ocean:			
a) Bay of Bengal			
surface water	0.02190	Sen Gupta <i>et al.</i> , 1978	In the Bay of Bengal this ratio is influenced by river runoff.
depth 3000 m	0.02196		
b) Gulf of Mannar			
surface water	0.02193		
c) Arabian Sea			
surface water	0.02193		
depth 2000 m	0.02196		
Laccadive Sea:	0.02168 ± 0.00002	Naqvi, Reddy, 1979	
Kavaratti Lagoon	0.02145 ± 0.00004		
Minicoy Lagoon	0.02142 ± 0.00005		
Red Sea	0.0223	Danielsson <i>et al.</i> , 1980	The Ca/Cl ‰ ratio was estimated by us from their graphical data.
Pacific Ocean:			
depth 4244 m	0.02154	Kanamori, Ikegami, 1980	26°11'N, 136°42'E.
Mediterranean Sea:			
surface water	0.02123	Krumgalz, Holzer, 1980	34°29'89"E, 32°59'04"N.
depth 1130 m	0.02126		
Mediterranean Sea:			
a) Haifa Beach	0.0211 ± 0.0002	Krumgalz <i>et al.</i> , 1980	
b) 1.5 km from Haifa Beach	0.0213 ± 0.0003		
Great Bitter Lake			
(30 m offshore)	0.02147	Krumgalz <i>et al.</i> , 1980	
North Pacific	0.02116 ± 0.00012	Shiller, Gieskes, 1980	There are depth and longitudinal distributions.

test can be used to compare the difference between two averages, both of which have a degree of uncertainty:

$$t = \frac{\bar{X}_1 - \bar{X}_2}{s \sqrt{1/n_1 + 1/n_2}} \quad (1)$$

where:

$$s = \sqrt{\frac{(n_1 - 1)s_1^2 + (n_2 - 1)s_2^2}{n_1 + n_2 - 2}} \quad (2)$$

In these equations, \bar{X}_i is the average value of a number of observations, n_i is the number of observations in a group, and s_i is the standard deviation in the proper group. When comparing the calcium/chlorinity ratios for the first and third groups ($\bar{X}_1 = 0.02586$, $\bar{X}_3 = 0.02137$, $n_1 = 7$, $n_3 = 86$, $s_1 = 0.0013$, $s_3 = 0.00040$) and for the second and third groups ($\bar{X}_2 = 0.0308$, $n_2 = 9$, $s_2 = 0.0018$), the t values are equal to 8.89 and 14.42, respectively. These values are larger than the critical t value equal to $t_{0.001}(n_i + n_j - 2) = 3.40$. This means that the probability is greater than 0.999 that the average values for the calcium/chlorinity ratios of the above-mentioned groups are significantly different according to the Student t -test for the difference between two means.

Some physical explanations could be suggested for such significant differences in calcium/chlorinity ratios. If there is a non-marine water source in any ocean or sea having ionic concentration significantly different from that of sea water, then the major element/chlorinity ratios after the mixing of these water bodies would be considerably different from those in sea water. The Baltic Sea region and the hot hypersaline brines found in the central region of the Red Sea provide excellent examples of this phenomenon. For the Baltic Sea region, the calcium/chlorinity ratio is larger than that of other seas and oceans. This is due to the inflow of a large amount of continental river freshwater enriched by Ca^{2+} ions. The

abnormally high calcium/chlorinity ratios in the hot brines of the Red Sea can be explained by the nature of the water masses taking part in their formation.

These brines are known to be formed by hot, highly saline water with temperatures of 45-62°C (Riley, 1967; Pugh, 1969; Ross, 1969; Danielsson *et al.*, 1980) and salinities of 255-326‰ (Brewer *et al.*, 1965; Riley, 1967; Brewer, Spencer, 1969) entering the sea through fissures in the sea floor. The mixture of these waters, enriched by some elements including calcium, with ordinary seawater leads to the unusually high calcium/chlorinity ratios characterizing the hot brines in the Red Sea. Brewer *et al.* (1965) observed, in addition, a mixed water layer where calcium/chlorinity ratios gradually changed from anomalous values characteristic of these brines to values representative of ordinary sea water.

In an analysis of the third group of calcium/chlorinity ratio values using the t -test, it can easily be shown that not all values of the calcium/chlorinity ratios in this group belong to the same population. Variation of the calcium/chlorinity ratio in seawater is expected to occur with the depth and other hydrological conditions. The reasons for this are: (1) the effect of river waters enriched by calcium; and (2) uptake of calcium from the surface seawater by calcareous organisms and re-dissolution of the calcareous shells of dead organisms sinking to the deep water. We have found by consecutive t -tests of all cited literature data that the calcium/chlorinity ratios for waters of open seas and oceans is equal to $0.02128 \pm 0.00015 \text{ g.kg}^{-1} (\text{Cl}^{\circ}/\text{‰})^{-1}$. The value obtained practically coincides with the value 0.02126 obtained by Culkin and Cox (1966). As more new data concerning calcium/chlorinity ratios in seawater for different locations become available, more precise calcium/chlorinity ratios for waters of open seas and oceans can be obtained. For all 86 locations analyzed, 59 locations have calcium/chlorinity values characteristic of open seas and oceans. The other 27 locations have

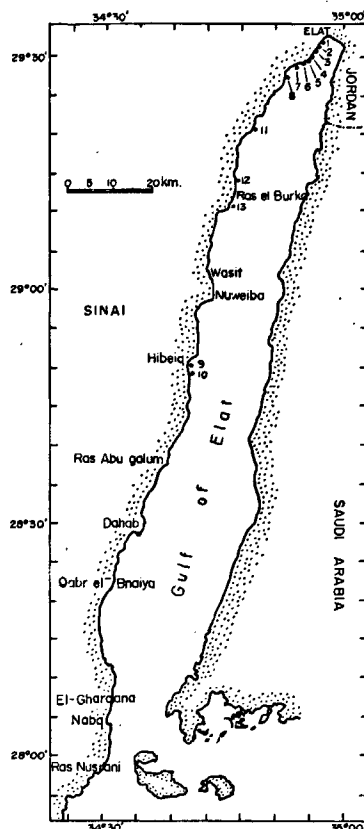


Figure 2
Map of the Gulf of Elat
with station locations.

calcium/chlorinity values either smaller or larger than those for open seas and oceans.

The smaller values in some cases can be explained by geochemical or biological precipitation of calcium carbonate (Carpenter, 1957; Broecker, Takahashi, 1966; Kato, 1966b). On the other hand, the larger values in some cases can be explained by the influence of a strong river inflow (Sen Gupta *et al.*, 1978). Since almost every different water body has a characteristic calcium/chlorinity ratio which depends on its hydrological conditions, the variations in calcium/chlorinity ratios can provide information on some natural processes occurring in the particular water body, such as water mass circulation; mixing and diffusion through coastal and open sea water, geochemical and biological

precipitation of calcium carbonate, etc. In water bodies with intensive coral production, interesting information can be obtained by the study of calcium distribution. The Gulf of Elat (Aqaba), with its numerous coral reefs, could be a natural model for studying the influence of coral growth on the calcium/salinity ratio. This study has been conducted on the calcium distribution in the Gulf of Elat water body in the vicinity of a coral reef during September-October 1979.

Samples of seawater were collected along the shore of the Sinai Peninsula at 13 stations indicated in Figure 2. The samples were collected from the surface layer at a distance of 20-50 m from shore. The calcium content of the seawater samples was determined by potentiometric titration with ethyleneglycol bis-(β -aminoethyl ether)-N-N'-tetraacetic acid (EGTA). For the potentiometric titration, a digital titration system - DTS-833 (precision ± 0.1 mv) with a calcium electrode (F2110 Ca) and a saturated KCl-calomel reference electrode (K 401) was used. A detailed description of the procedure used has been given previously (Krumgalz, Holzer, 1980). Salinity values were determined by a salinometer, Autosal Model 8400 having an accuracy of $\pm 0.003\%$ after instrument calibration with standard seawater of 35.002‰ salinity. The results obtained are presented in Table 2. The chlorinities of these samples were calculated from experimentally measured salinities using the equation:

$$\text{Cl}\% = \frac{\text{S}\%}{1.80655} \quad (3)$$

As can be seen from Table 2, calcium/chlorinity ratios are practically constant and are equal to $0.02083 \pm 0.00005 \text{ g.kg}^{-1}(\text{Cl}\%)^{-1}$.

It is very interesting to compare the calcium distribution in the Gulf of Elat with the calcium content in other world oceans. The average calcium/chlorinity ratio found for the samples of seawater from the Gulf of Elat, $0.02083 \text{ g.kg}^{-1}(\text{Cl}\%)^{-1}$, is much smaller than those found for open sea waters. This is the result of biological precipitation of calcium carbonate by some marine organisms, including corals, which deplete calcium in surface waters by the deposition of CaCO_3 in their skeletons. D. Dyrssen, in his review of this article, noted

Table 2
Calcium ion distribution in the Gulf of Elat (Aqaba).

Station (*)	Calcium concentration, $\text{g Ca}^{2+}/\text{kg seawater}$	Salinity, (‰)	Calcium/chlorinity ratio, $\text{g.kg}^{-1}(\text{Cl}\%)^{-1}$
1	0.4680 ± 0.0017	40.626	0.02081 ± 0.00007
2	0.4684 ± 0.0002	40.677	0.02081 ± 0.00001
3	0.4676 ± 0.0011	40.652	0.02078 ± 0.00005
4	0.4693 ± 0.0012	40.700	0.02083 ± 0.00005
5	0.4689 ± 0.0012	40.666	0.02083 ± 0.00005
6	0.4693 ± 0.0006	40.678	0.02085 ± 0.00002
7	0.4662 ± 0.0014	40.658	0.02072 ± 0.00007
8	0.4696 ± 0.0010	40.656	0.02087 ± 0.00004
9	0.4698 ± 0.0010	40.736	0.02083 ± 0.00004
10	0.4710 ± 0.0038	40.793	0.02087 ± 0.00016
11	0.4698 ± 0.0010	40.766	0.02081 ± 0.00004
12	0.4722 ± 0.0025	40.787	0.02092 ± 0.00011
13	0.4712 ± 0.0008	40.801	0.02087 ± 0.00004

(*) Stations 1-10, September 1979; Stations 11-13, October 1979.

that his data concerning the Red Sea and the Gulf of Aden also confirm that the low values of the Ca/Cl‰ ratio are due to the activity of coral reefs. There now exists a significant body of literature alluding to calcium depletion in coral reef waters as inferred from alkalinity depletions. However, in our article we did not review this field of research, and therefore the studies using the indirect alkalinity approaches were omitted.

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