
Temporal changes in the phytoplankton community along the French coast of the eastern English Channel and the southern Bight of the North Sea

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Abstract :

Long-term trends in the phytoplankton community along the French coast of the English Channel and southern Bight of the North Sea were studied and related to physico-chemical factors and large-climatic indices. Phytoplankton and hydrological data were acquired through three transects. Sampling took place between 1992 and 2011 as part of the French Phytoplankton Monitoring Network (REPHY) and the Regional Nutrients Monitoring Network (SRN). Trends in time-series were identified with dynamic linear models tailored to environmental monitoring data characteristics (e.g. irregular sampling frequency, missing data). Temporal and spatial patterns in the phytoplankton community were explained with a partial triadic analysis. Relationships between the phytoplankton community composition, environmental factors, and climatic indices were assessed using a redundancy analysis (RDA). The analyses revealed long-term changes in the community composition, characterized by a temporal structure that remained common to all transects. The abundance of some groups of taxa such as the one composed by *Gymnodinium* and *Gyrodinium*, as well as the group of *Pseudo-nitzschia* increased during the study period, whereas the abundance of other taxa as for example *Guinardia* and the group of *Coscinodiscus* and *Stellarima* globally decreased. More generally, the proportion of dinoflagellates relative to diatoms increased. Trends in environmental variables were also observed in most sites and related to decreases in nutrient concentrations and an increasing trend in salinity. The RDA indicated that the Atlantic Multidecadal Oscillation index and salinity were the main factors defining the temporal structure of the phytoplankton community. This suggests that variations observed in the phytoplankton community are linked to hydro-climatic changes in the coastal environment.

Keywords : climatic indicators, dynamic linear model, environmental drivers, multivariate analysis, phytoplankton

1. Introduction

In a context of climate change and concerns about anthropogenic pressures over coastal ecosystems, understanding the response of biological communities has become a necessity. In marine ecosystems, the phytoplankton is at the base of marine foodwebs, and its specific composition determines the pathways and efficiencies of the energy transfer through the upper trophic levels (Cloern and Dufford, 2005; Dickman *et al.*, 2008). The relative species abundance and the community composition contribute to the characterization of phytoplankton assemblage structure. This structure is governed by various environmental factors, allelopathic interactions and grazing, affecting cellular processes, and determining the population growth, community composition, and algal succession (e.g. Tilman *et al.*, 1982; Marshall *et al.*, 2006). These factors lead to recurrent patterns of phytoplankton composition observed at different temporal scales. However, these community properties may vary significantly with climate change and water-quality trends in coastal areas.

Long-term variations in phytoplankton communities have been recently observed and related to hydro-climatic changes, affecting biomass, community composition, and phenology (Edwards and Richardson, 2004; Suikkanen *et al.*, 2007). These modifications in phytoplankton communities are often studied at different taxonomic scales. Thereby, recent works show a modification of the relative proportion between diatoms and dinoflagellates (Klais *et al.*, 2011; Zhai *et al.*, 2013) and significant variations in the biomass of some taxonomic groups such as chrysophytes and chlorophytes (Suikkanen *et*

76 *al.*, 2007). Functional approaches have also been used to assess temporal changes in
77 phytoplankton communities related to physical factors and climate (David *et al.*, 2012). On
78 the French coast of the eastern English Channel and the southern Bight of the North Sea,
79 some studies highlighted the link between large-scale climatic features, local
80 meteorological and hydrographical conditions, modulating phytoplankton dynamics
81 (Breton *et al.*, 2006, Goberville *et al.*, 2010). These studies have proved the importance of
82 monitoring networks to assess long temporal changes in phytoplankton communities.

83 The Phytoplankton Monitoring Network (REPHY) is implemented and managed by
84 the French Research Institute for the Exploitation of the Sea (IFREMER). This program
85 aims to describe spatial and temporal dynamics of phytoplankton populations and to detect
86 and monitor phycotoxin-producing species involved in toxic shellfish outbreaks and
87 affecting human health (Amzil *et al.*, 2001; David *et al.*, 2012; Gailhard *et al.*, 2002).

88 Previous works with these data have greatly improved our knowledge on spatial patterns of
89 microphytoplankton and highlighted the influence of hydrodynamic features in the
90 geographical structure of microalgal populations (David *et al.*, 2012, Gailhard *et al.*, 2002).
91 Seasonal variability and temporal succession patterns have also been characterized through
92 the analyses of phytoplankton events, defined by the phases of sudden growth, reached
93 maxima and decline of abundance (Beliaeff *et al.*, 2001). In addition, Guarini *et al.* (2004)
94 showed the importance of the interaction between benthic and planktonic communities
95 through assessing the contribution of the microphytobenthos to the total phytoplankton
96 abundance. Regarding the eastern English Channel, Lefebvre *et al.* (2011) analyzed the
97 long-term variability of the major phytoplankton groups and gave a first insight of
98 establishing a link between environmental conditions and phytoplankton. However, even
99 though several studies have been conducted on the REPHY data sets, temporal tendencies
100 have been only partially addressed, either for the major phytoplankton components or the

101 specific community composition. At present, data on phytoplankton have been collected
102 over 20 years, and completed by a regional network which targets relations between
103 phytoplankton and nutrients.

104 Besides these ecological results, it should be noted that the analyses of these time
105 series requires methods able to identify structures in space and time with data that exhibit a
106 very large variability. Several methods have been used with data coming from monitoring
107 programs, specially multivariate and time-series analyses techniques (*e.g.* David *et al.*,
108 2012, Goberville *et al.*, 2010, Wasmund *et al.*, 2011).

109 The main objective of this paper is to characterize and quantify long temporal
110 changes in the phytoplankton community structure at three sites in the French coast of the
111 eastern English Channel and the southern Bight of the North Sea, and to explain these
112 changes in relation to physico-chemical factors and large-scale climatic indices. We
113 examine if variations of the phytoplankton community can be observed at different scales,
114 from specific community composition, to groups such as diatoms and dinoflagellates. A
115 statistical framework is developed to deal with common issues in the analysis of long-term
116 monitoring time series and depict the spatial structure of temporal patterns.

117

118 Material and Methods

119

120 Study area

121 The study of the phytoplankton community was conducted in three coast-to-offshore
122 transects, along the French coast: Bay of Somme (SOM), Boulogne (BL), and Dunkerque
123 (DK) (Figure 1). Each transect consists of three to four sampling stations. Boulogne and
124 the Bay of Somme are located in the eastern English Channel, and are characterized by a
125 coastal flow system often separated from the open sea by a frontal area (Brylinski *et al.*,

126 1996). The sampling area of Dunkerque, located in the southern Bight of the North Sea, is
127 characterized by shallow and well-mixed waters (Lefebvre *et al.*, 2011). In addition, the
128 sampling zone is influenced by a macrotidal regime and by freshwater inputs coming from
129 the Seine, Somme and other tributaries (such as Canche, Authie and Liane rivers).

130

131 Data sets

132 Data was collected as part of the REPHY and the Regional Nutrients Monitoring Network
133 (SRN). Water samples were acquired from fortnightly to a monthly frequency from 1992
134 to 2011 between 0 and 1 meter, along with physical measurements and completed with
135 chemical analyses. Finally, the environmental variables included in this study were water
136 temperature (°C), salinity (psu), dissolved inorganic nitrogen ($\text{DIN} = \text{NH}_4^+ + \text{NO}_3^- + \text{NO}_2^-$,
137 $\mu\text{mol l}^{-1}$), silicate ($\mu\text{mol l}^{-1}$) and phosphates ($\mu\text{mol l}^{-1}$). For quantitative phytoplankton
138 analyses, samples were fixed with Lugol's solution and counted according to the Utermöhl
139 method (Utermöhl, 1958). Organisms were identified to the lowest possible taxonomic
140 level. Taxa that are difficult to discriminate with optical microscopy were grouped. This
141 was the case of various groups representing a set of multiple species or even genus (*e.g.*
142 *Pseudo-nitzschia "seriata complex"*, *Gymnodinium-Gyrodinium*). These groups correspond
143 to the taxonomic units extracted from the database. In addition, within sample observation
144 conditions, phytoplankton experts identified and counted organisms whose size is greater
145 than 20 μm but also species whose size is smaller but have a chain structure or form a
146 colony. Smaller species were also counted if they are potentially toxic/noxious (*e.g.*
147 *Chrysochromulina*, *Phaeocystis*). Further details about sampling and processing of
148 phytoplankton and physico-chemical parameters are available in the literature (Aminot and
149 K  rouel, 2004, 2007; Belin and Neaud-Masson, 2012; Gossel, 2006; Lefebvre *et al.*,
150 2011).

151 Two large-scale climatic indices were used: the winter North Atlantic Oscillation
152 index (NAO_w, Hurrell, 1995) and the Atlantic Multidecadal Oscillation (AMO, Kerr,
153 2000). The NAO index can be defined as the difference in atmospheric pressure at sea
154 level between the high pressure centre near the Azores and the low pressure centre near
155 Iceland (Rodwell *et al.*, 1999). This oscillation has been related with factors such as
156 temperature, precipitation, wind, storms and mixed-layer depth (Hurrell and Deser, 2009;
157 Trigo *et al.*, 2002). The winter NAO index (December-March) was used for its more
158 pronounced correlation with temperature, wind and precipitation, and because its influence
159 persists in subsequent months (Ogi *et al.*, 2003; Stenseth *et al.*, 2003). This index was
160 obtained from the website <http://climatedataguide.ucar.edu/>. The AMO is a pronounced
161 signal of climate variability in the North Atlantic sea surface temperature (Dijkstra *et al.*,
162 2006). The AMO index was downloaded from
163 <http://www.esrl.noaa.gov/psd/data/timeseries/AMO/>.

164

165 Preliminary processing of phytoplankton data

166 Despite the use of taxonomic units as described previously, some confusions due to
167 changes in phytoplankton taxonomists or any other expertise variation during the period
168 may arise. These changes may lead to heterogeneous data regarding taxonomic
169 classification and hence to a misinterpretation of phytoplankton time series. To
170 homogenize the data, taxonomic units were grouped into the lowest common taxonomic
171 level. For example, *Coscinodiscus*, *Coscinodiscus radiatus- marginatus*, *Coscinodiscus-*
172 *Stellarima* are three different labels available in the database and were finally replaced by
173 *Coscinodiscus-Stellarima*. In total, 91 of such taxonomic units were identified over the
174 selected period. Furthermore, discussions with observers and phytoplankton experts were
175 undertaken to ensure the reliability of the analysis performed.

176 To normalize data and reduce the effect of dominant taxa, abundances were $\log_{10}(x$
177 $+ 1)$ -transformed. In addition, a simple phytoplankton community index was defined as the
178 proportion of dinoflagellates relative to diatoms. This proportion was calculated as the
179 abundance of total dinoflagellates divided by the sum of total abundance of diatoms and
180 dinoflagellates.

181

182 Data analyses

183 Univariate and multivariate analyses were performed to assess changes in phytoplankton
184 community, taking into account different scales of variability either taxonomic, temporal
185 or spatial. The Figure 2 shows the pathways of statistical analyses applied in this study. All
186 analyses and graphical representations were performed with the R software (R Core Team,
187 2012).

188

189 Univariate analyses of phytoplankton and environmental data

190 In order to characterize seasonal patterns of the main taxonomic groups, monthly median
191 abundances, less sensitive to extreme values, were calculated over 1992–2011.

192 Bacillariophyceae, Dinoflagellata, and Prymnesiophyceae were previously identified as the
193 most common classes in this zone (Lefebvre *et al.*, 2011).

194 Dynamic Linear Models (DLM, West and Harrison, 1997) were used to study long-
195 term trends in phytoplankton and environmental variables (Figure 2). DLM were
196 previously used in other studies for analysing phytoplankton time series (Cottingham and
197 Carpenter, 1998; Soudant *et al.*, 1997). They are indeed time series models, *i.e.* they take
198 into account the temporal structure. Moreover, the parameters of the model can evolve
199 with time, hence the "dynamic" designation. This class of models was chosen because they
200 are well tailored to environmental monitoring data characteristics (*e.g.* irregular sampling

201 frequency, missing data, outliers). Furthermore, as for all time series models, exogenous
202 information can be taken into account using an "intervention", that is, modifying some
203 model parameters. This was particularly useful for incorporating information such as
204 changes in measurement methods applied to physico-chemical variables. All these changes
205 in measurement methods were analysed. For example, from 1992 to 2011, two main
206 methods have been used for the measurement of salinity: Knudsen titration and
207 conductivity sensors. In such cases, using the intervention allows to correct the time series
208 and thus, ensures comparability and continuity of measurements through the time series.

209 To extract trends in time series, a common model structure was used for all
210 variables. The first component is a second order polynomial trend: it remains simple and
211 allows to capture complex evolutions. The second component is a 12-month seasonal
212 factor. This was chosen since sampling frequency was mainly monthly. Furthermore, this
213 temporal unit is fine enough to capture different annual periodic patterns. Finally, when
214 more than one observation was available within the month, the median was used to
215 aggregate the data. Outliers were detected by examination of standardized residuals. Under
216 the assumption of a normal distribution of standardized residuals, it is expected that about
217 95% of them are within the ± 1.96 interval. Values standing outside this interval were
218 considered as outliers. They were not removed but treated using interventions. Model
219 checking was assessed by residual diagnosis. Serial autocorrelation and normality were
220 tested by plotting the autocorrelation function of the standardized residuals as well as their
221 normal QQ-plot. Deviations from normality were rare. In the case where serial
222 autocorrelation was present, an AR(1) term was added to the model. The results showed
223 that adding this term does not affect trends. Thus, models have been kept in the original
224 structure as described above.

225 DLM were first used to study long-term trends in diatoms and dinoflagellates
226 functional groups, as well as trends of some taxonomic units, selected through multivariate
227 techniques (explained below). Physico-chemical trends were identified with DLM, and the
228 annual mean trend was used for multivariate analyses. Time series analyses with DLM
229 were conducted with the *dml* package for R (Petris, 2010).

230

231 Multivariate analyses of phytoplankton community composition

232 For multivariate analyses, the genus was used as the smallest taxonomic level. Aggregating
233 the data at this level allows to capture most of the information considering that little
234 information is lost using the genus level rather than the species one (Cottingham and
235 Carpenter, 1998). In order to focus on the main changes in phytoplankton community, rare
236 taxa were excluded from the analyses, based on the occurrence frequency over 1992–2011.
237 A taxa was discarded when its occurrence frequency was less or equal to three years in all
238 sites. This resulted in a loss of less than one percent of the total abundance for each site,
239 which represents 29 discarded taxonomic units. Sixty-two taxonomic units remained for
240 mutivariate analyses.

241 Spatial and temporal characteristics of the phytoplankton data led to obtain a three-
242 dimensional table (taxonomic units; sites; years). Partial Triadic Analysis (PTA), used with
243 such multi-table data, was computed to elucidate temporal patterns of the phytoplankton
244 community composition (Figure 2, package *ade4* for R, Dray *et al.*, 2007). This method
245 allows simultaneous principal component analyses of a set of tables by searching for an
246 average temporal structure called compromise, and studying the spatial stability of this
247 structure (Blanc and Beaudou, 1998). In addition to the analysis performed at genus level,
248 the PTA was also carried out at upper levels (*e.g.* family, order, class).

249

250 Relationships between phytoplankton, environmental factors and climatic indices
251 A Redundancy Analysis (RDA) was conducted with data from all sampling stations in
252 order to study the relationships between patterns in the phytoplankton community
253 composition, environmental variables and climatic indices (Figure 2, package *vegan* for R,
254 Oksanen *et al.*, 2012). The significance of the analysis and all variables was assessed using
255 permutation tests (499 permutations).

256

257 Results

258

259 From 1992 to 2011, Bacillariophyceae, Dinoflagellata, and Prymnesiophyceae were the
260 most abundant groups, averaging 81% of the total abundance. Among these groups,
261 diatoms and dinoflagellates represented over 84% of the total diversity. Principal diatom
262 taxa in terms of abundance were the genus *Guinardia*, *Chaetoceros*, *Rhizosolenia*, *Pseudo-*
263 *nitzschia*, *Paralia*, *Skeletonema* and *Leptocylindrus*. The group *Gymnodinium-Gyrodinium*
264 was the major dinoflagellate present over the period. *Phaeocystis* (Prymnesiophyceae) was
265 a major contributor to the phytoplankton total abundance between March and May. When
266 present, *Phaeocystis* median contribution varies from 74% to 90%, with the highest
267 concentrations generally observed in stations near the coast. The complete list of taxa
268 identified is given in the Supplementary Material (Table S1) with information on the
269 taxonomic groups they belong to, and highlighting those mainly discussed in this study.

270

271 Seasonal patterns in phytoplankton and environmental data

272 Although all sites within a same transect generally present the same patterns, three
273 sampling sites, each one belonging to a different transect were chosen as their results
274 present the most pronounced variations. Consequently, these sampling sites will be used

275 for graphical representations. Among the major phytoplankton groups, *Phaeocystis*
276 (*Prymnesiophyceae*) showed a clear seasonal pattern, with peaks that appear generally in
277 April (Figure 3). Diatoms were present throughout all the year with highest abundance
278 usually observed between June and July in Boulogne and the Bay of Somme (Figure 3). In
279 contrast, in Dunkerque, diatoms showed high abundances early in the year (March).
280 Seasonal patterns of diatoms, dinoflagellates and *Prymnesiophyceae* are showed for each
281 sampling station in Figure S1 in the Supplementary Material (SM).

282 Spring-summer diatoms generally included the genus *Guinardia*, *Chaetoceros*,
283 *Pseudo-nitzschia* and *Rhizosolenia* while *Leptocylindrus* was characteristic in summer. The
284 autumn-winter period was characterized by the presence of the diatom *Skeletonema*, the
285 groups *Thalassionema-Thalassiothrix-Lioloma*, *Thalassiosira-Porosira*, *Coscinodiscus-*
286 *Stellarima* and the genus *Ditylum*. Others groups present at this period of the year included
287 the small single-cell flagellate *Dictyocha*. Dinoflagellate abundances increased from the
288 beginning of the year peaking between July and September, then decreasing at the end of
289 the year (Figure 3). The dinoflagellates *Gymnodinium-Gyrodinium*, *Proto-peridinium-*
290 *Peridinium*, and *Prorocentrum* are common genera during this period.

291 Seasonal patterns of environmental variables are illustrated for each sampling
292 station in Figure S2 (SM). Nutrients data reflected typical temperate seasonal patterns,
293 remaining almost similar between sites. DIN, phosphate and silicates reached maxima
294 during the winter and decreased over the spring and early summer to levels close to the
295 detection limits (Figure 4). A further increase in nutrients was observed in the late
296 summer-autumn period. During winter, the Bay of Somme showed the highest
297 concentration of DIN and silicates, with average values between 22.7 $\mu\text{mol l}^{-1}$ and 44.3
298 $\mu\text{mol l}^{-1}$ for DIN and between 9.1 $\mu\text{mol l}^{-1}$ and 20.7 $\mu\text{mol l}^{-1}$ for silicates. Lower
299 concentrations of DIN and silicates were found at Boulogne with ranges of 11–23.8 $\mu\text{mol l}^{-1}$

300 ¹ and 5.1–8.7 $\mu\text{mol l}^{-1}$, respectively. At Dunkerque the concentration of these nutrients
301 varied from 14.5 to 31.5 $\mu\text{mol l}^{-1}$ for DIN and from 7.7 to 12.9 $\mu\text{mol l}^{-1}$ for silicates.

302 The phosphates concentration varied from 0.73 to 1.1 $\mu\text{mol l}^{-1}$ at Boulogne and
303 reached comparable values at Dunkerque and the Bay of Somme (0.83–1.27 $\mu\text{mol l}^{-1}$ and
304 0.82–1.76 $\mu\text{mol l}^{-1}$ respectively). Furthermore, a spatial heterogeneity of the nutrient
305 concentrations was observed in "coast-to-offshore" transects, with higher concentrations
306 near the coast. The temperature varied from an average minimum of 5.9°C between
307 January and February to a maximum of 18.9°C between July and August. The salinity
308 varied between 31.6 and 34.4 psu. In addition, the sampling station of Bif was
309 characterized by a lower salinity (minima 9.1 psu), related to a major influence of
310 freshwater inputs from the Somme river.

311

312 Long-term trend in diatoms and dinoflagellates groups and environmental data based on
313 univariate analysis

314 The long-term trends in the diatoms and dinoflagellates abundance are showed in Figure 5.
315 Minor changes in the abundance of diatoms are observed at Dunkerque and in the Bay of
316 Somme while a decrease in the abundance is observed in Boulogne between 1992 and
317 1997, followed by an increase between 1998 and 2005. The abundance of dinoflagellates
318 between 1992 and 2011 showed more pronounced variations. From 2000 to 2007, the
319 mean level of the dinoflagellates logarithmic abundance increased by 60% (from 2.5 to 4
320 \log_{10} cells l^{-1}). Spatial differences in the trend of dinoflagellates are observed among
321 transects between 1992 and 2000, but abundance patterns showed minor variations within
322 sampling stations of the same transect. The abundance of dinoflagellates decreased at
323 Dunkerque and Boulogne between 1992 and 2000, while minor changes were observed at
324 the Bay of Somme. The phytoplankton community showed an increasing trend in the

325 proportion of dinoflagellates relative to diatoms, with values generally below 0.1 between
326 1992 and 2001 and frequently above this value after 2002 (Figure 5). In addition, increases
327 in dinoflagellates abundance were observed throughout all the seasonal cycle after 2001.

328 Regarding physico-chemical variables, dissolved inorganic nitrogen decreased in
329 Dunkerque from the beginning of the period (Figure 6 a) and after 2000 at Boulogne. The
330 phosphate concentration decreased at Dunkerque and the Bay of Somme at the start of
331 period with a minor increase after 2000s (Figure 6 b). A decrease was observed at
332 Boulogne over the period. No trend was observed for silicates at Dunkerque while it seems
333 to slightly increase at Boulogne after 2000s (Figure 6 c). At the Bay of Somme, DIN and
334 silicates seem to vary cyclically, with periods of higher concentrations (1994–1995, 2000–
335 2001) followed by years with lower concentrations (1996–1998, after 2002). Time series of
336 salinity at Boulogne and the Bay of Somme showed large increases during the period
337 (Figure 6 d).

338
339 Long-term changes in the phytoplankton community based on multivariate analysis
340 As suggested by the PTA results, the community structure changed between 1992 and
341 2011 (Figure 7). The first two axes account for 58.7% of the total variability. PTA objects,
342 corresponding to (site;year) pairs, were projected on the first and second axis of the
343 compromise (Figure 7 a). Each dotted line represents a given sampling station. The first
344 axis globally distinguishes two periods: 1992–2001 and 2002–2011. In the latter period,
345 the second axis splits the period into two groups: 2002–2007 and 2008–2011. This
346 temporal trajectory was generally seen at all the sampling stations. Such result highlights a
347 temporal structure of the phytoplankton community composition and a trend in taxa
348 abundances that change largely during the sampling period while remaining almost
349 constant in space. Among diatoms, *Melosira*, and *Stephanopyxis* were usually common in

350 early years but rarely observed after 2002. Other taxa such as the dinoflagellates
351 *Amphidinium*, *Alexandrium*, and *Polykrikos* were characteristic after 2002 (Figure 7 b).
352 Finally, *Heterocapsa* (dinoflagellate), *Torodinium* (dinoflagellate), and *Eutreptiella*
353 (euglenid) were mostly common over the third period.

354 Changes in the abundance of some taxonomic units were observed over the period
355 (Figure 8 and Figure S3 in SM). The diatom *Guinardia* generally decreased at the Bay of
356 Somme and Dunkerque while no pronounced variation was detected at Boulogne. Other
357 taxa, mainly diatoms, showed a decline in their abundances (e.g. *Coscinodiscus-Stellarima*,
358 *Eucampia-Climacodium*, *Ceratulina*) while the abundance of *Pseudo-nitzschia* increased
359 during the period of study. Among dinoflagellates, the group of *Gymnodinium-Gyrodinium*
360 showed a strong increase in abundance, which has more than doubled (in logarithmic
361 scale) between 2002 and 2007. *Phaeocystis* showed no trend in the Bay of Somme, while a
362 minor increase was observed at Dunkerque and Boulogne.

363 The analyses of taxonomic groups from family to class indicate a similar temporal
364 structure. Nevertheless, with the increase in the level of aggregation, presence/absence
365 patterns of specific taxonomic units are more difficult to detect while changes in
366 abundances tend to become more apparent.

367

368 Relationships between phytoplankton community, environmental factors and climatic
369 indices

370 The redundancy analysis accounted for 21.8% of the total variance of the data (all
371 canonical axes). Of this, 68.8% was explained by the first two axes. These canonical axes
372 and the seven variables were significant ($p < 0.01$). On the first axis, the AMO index,
373 salinity, NAO_w index, and temperature were the main explanatory variables (Figure 9 a).
374 The first axis describes a temporal structure, with two different periods, before and after

375 2001 (Figure 9 a, c). The second axis is mainly defined by silicates, dissolved inorganic
376 nitrogen, but also salinity (Figure 9 a). It shows a spatial pattern between transects.
377 Sampling stations in the Bay of Somme are characterized by high concentrations of
378 silicates and dissolved inorganic nitrogen, and a low salinity (Figure 9 a). Temporal
379 patterns were observed in transects of Dunkerque and Boulogne in relation to a decreasing
380 trend in nutrients concentrations (mainly dissolved inorganic nitrogen) over 1992–2011
381 (Axis 2–RDA, Figure 9 a).

382 The distribution of phytoplankton taxonomic units along the first gradient was
383 positively related to the salinity, AMO index and temperature, and negatively to the NAO_w
384 index and phosphates concentration. The AMO index, salinity and temperature vectors
385 pointed to the second period (after 2001) when dinoflagellates increased (Figure 9 a, b).
386 Diatoms were mostly grouped in the lower-left quadrant and thus positively linked to
387 silicates, DIN and NAO_w index. However, some diatoms such as *Pseudo-nitzschia* showed
388 their highest abundance during the second period (positive score on the first axis),
389 suggesting that these diatoms were favoured under an increased salinity and positively
390 associated with the AMO index and temperature. Other genera show low scores in both
391 axes, thus suggest its presence along all transects and over mostly the whole period of
392 study (*e.g. Rhizosolenia, Skeletonema*).

393

394 Discussion

395

396 Patterns of phytoplankton community changes

397 During the past 20 years, the structure of the phytoplankton community of the eastern
398 English Channel and the southern Bight of the North Sea was subject to changes. We
399 observed an increasing trend of dinoflagellates during the last decade while minor changes

400 were observed for diatoms between 1992 and 2011. The PTA allowed us to understand
401 these changes at a finer taxonomic scale. Some diatoms taxa globally showed a declining
402 trend in their abundances (*e.g. Guinardia, Coscinodiscus-Stellarima*). In contrast, the
403 abundance of *Pseudo-nitzschia* appears to increase over the 20 years of study. Diatoms like
404 *Melosira* and *Stephanopyxis* were mainly found before 2001. Some dinoflagellates such as
405 *Gymnodinium-Gyrodinium* showed a strong increasing trend in their abundances after
406 2001. Other genera such as *Alexandrium*, *Amphidinium* and *Polykrikos* were mostly found
407 in the last decade (2002–2011).

408 In our study, the PTA enabled to describe temporal patterns and also facilitated
409 comparisons between multiple sampling stations, determining if the long-term trend is
410 spatially constant. This analysis was previously used to study the temporal and spatial
411 structures of phytoplankton communities (David *et al.*, 2012). Globally, three periods were
412 identified (1992–2001, 2002–2007 and 2008–2011), and such temporal pattern is globally
413 exhibited by all transects. In combination with RDA to further assess relationships between
414 phytoplankton communities, physical-chemical variables and climatic indices, we found a
415 clear separation between the periods before and after 2001, and some of the observed
416 trends in the phytoplankton community could be related with physico-chemical variations
417 and climatic indices. The time at which the changes in the phytoplankton community
418 structure occurred is consistent with a shift observed in the French coastal system after
419 2001, manifested through an increasing salinity and decreasing nutrients, impacting
420 chlorophyll *a* concentrations (Goberville *et al.*, 2010). Similar temporal patterns were
421 observed in the functional and specific diversity along the French Atlantic coast, although
422 spatial patterns were most apparent, mainly related to a latitudinal gradient and
423 hydrodynamic characteristics (David *et al.*, 2012). Moreover, during the last decades,
424 major modifications in the phytoplankton composition, leading to an increasing importance

425 of some groups such as dinoflagellates, have been documented in the central North
426 Atlantic Ocean, North Sea and Baltic Sea (Klais *et al.*, 2011; Leterme *et al.*, 2005; Zhai *et*
427 *al.*, 2013). The rather low percentage of the variance explained by the RDA in this study
428 has also been observed in previous analyses of phytoplankton data (Suikkanen *et al.*, 2007,
429 Wasmund *et al.*, 2011). This has been related to environmental factors usually not
430 considered in the analyses, but having influence in phytoplankton populations, such as
431 allelopathic interactions, interspecific competition and grazing.

432 It should be noted, however, that changes in staff involved in the identification of
433 microalgal species, or any other expertise variation, may lead to a misinterpretation of
434 phytoplankton time series. These issues have been already questioned, especially with long
435 term phytoplankton data sets (Wasmund and Uhlig, 2003; Wiltshire and Dürselen, 2004).
436 Obviously, with more than 20 years of REPHY monitoring network along the French
437 coast, such changes did occur. To deal with these biases, other studies have suggested a
438 regrouping of taxonomic units (Gailhard *et al.*, 2002, Guarini *et al.*, 2004). In our study,
439 clustering taxonomic units has been carried out with morphologically close taxa.
440 Furthermore, discussions with observers directly involved in the identification and
441 counting, as well as phytoplankton experts, suggest that our trends are reliable. In addition,
442 the results of PTA performed with different taxonomic grouping, from family to class,
443 show a similar temporal structure, which supports our interpretation.

444 Previous studies have highlighted temporal changes in the phytoplankton
445 community structure at different scales. In the English Channel, Widdicombe *et al.* (2010)
446 identified major changes in the community such as the decrease of the average abundance
447 of diatoms and the increase of some dinoflagellates abundance. For the study area, changes
448 have also been observed in chlorophyll *a* concentrations, used as an indicator of
449 phytoplankton biomass trends (Goberville *et al.*, 2010; Lefebvre *et al.*, 2011). Other

450 approaches can also be suitable with available species data sets. Modifications in the
451 ecological niche of species have been detected using long-term species data (Gebühr *et al.*,
452 2009; Grüner *et al.*, 2011). Moreover, a functional approach can be particularly relevant to
453 understand prey-predator processes and assess factors driving the structure of
454 phytoplankton communities (David *et al.*, 2012).

455

456 Links between phytoplankton changes and environmental variables

457 Several dinoflagellates genera have shown an increasing trend over the period and such
458 trend seems to be linked to AMO index, salinity and water temperature. In our study, the
459 AMO index was one of the principal factors defining the temporal structure observed in the
460 phytoplankton community, suggesting a possible link between large-scale climate
461 variability and phytoplankton. Similar temporal patterns were globally observed in all
462 transects, suggesting that the climate forcing may be stronger than local prevailing
463 conditions. The link between phytoplankton community structure and climate forcing has
464 been previously discussed in literature (Breton *et al.*, 2006; Irigoien *et al.*, 2000; Leterme
465 *et al.*, 2005). The AMO index reflects changes in the sea surface temperature, a main factor
466 governing the changes in the coastal environment (Beaugrand, 2009) and may favour
467 dinoflagellates growth. Although the AMO seems to be more related with the
468 phytoplankton community structure, the NAO index was also positively associated with
469 some diatoms and negatively related to dinoflagellates, a relation previously observed in
470 the North Atlantic Ocean and the English Channel (*e.g.* Irigoien *et al.*, 2000; Zhai *et al.*,
471 2013).

472 Regional climate conditions modify coastal environment, influencing physico-
473 chemical parameters (Goberville *et al.*, 2010). Along the French coast of the English
474 Channel, observed changes in environmental conditions, such as salinity increases and

475 decreasing nutrient concentrations were related with the sea surface temperature increase,
476 reduced rainfall and intensity of zonal winds (Goberville *et al.*, 2010). Oceanic-
477 atmospheric drivers such as AMO and NAO are likely to impact precipitation patterns and
478 may affect the river run-off as well as inputs on terrestrial-derived nutrients, salinity and
479 turbidity (*e.g.* Enfield *et al.*, 2001; Harley *et al.*, 2006, Trigo *et al.*, 2004). In a study of the
480 three main Iberian river basins, Trigo *et al.*, (2004) have shown an influence of the NAO
481 on precipitation patterns and river flow, but also highlighted the irregularity and the high
482 inter-decadal variability of this impact. The river flow patterns may also reflect human
483 impacts such as damming and channels. Concerning the Somme river, we found a weak
484 correlation between the NAO and the river flow (Pearson correlation coefficient $r = 0.15$, p
485 < 0.02) although some influence can be observed. These results may suggest more local
486 influencing conditions over the river flow. When analysing the influence of the Somme
487 river flow on the nutrient concentrations and salinity trends in the Bay of Somme, we
488 generally observed a significant ($p < 0.05$) and a high correlation coefficient (highest
489 values $r_{\text{DIN}} = 0.89$, $r_{\text{SIOH}} = 0.77$, $r_{\text{SALI}} = -0.52$, r_{PO4} not significant). Both, NAO and river flow
490 seem to modulate separately the Bay of Somme hydrologic conditions. In addition, the
491 frequency of important flow events in the Somme river seems to decrease after 2004. In
492 2001, the maxima river flow was about $96 \text{ m}^3 \text{ s}^{-1}$, while from 2004 to 2011, the maxima
493 river flows ranges between 30 and $49 \text{ m}^3 \text{ s}^{-1}$ (Data: Artois–Picardie Water Agency).

494 Generally, our results suggest an increasing trend in salinity and a decreasing trend
495 in dissolved inorganic nitrogen and phosphates, while a certain variability in trends
496 between sampling sites has to be noted. At Dunkerque, DIN concentrations decreased
497 throughout the study period and during the 2000's in Boulogne. In the Bay of Somme, a
498 period of low concentrations was observed from 2002. Our analysis suggests a decreasing
499 trend in phosphates in Boulogne over the whole period while at Dunkerque and the Bay of

500 Somme it decreases at the start of 1990's with minor increases in 2000's. The decreasing
501 trend in phosphates has been already highlighted in SRN data series and attributed to a
502 result of phosphates-reduction practice since 1970s (Lefebvre *et al.*, 2011). Similar results
503 were observed by Romero *et al.* (2013) while assessing nutrients export for a large number
504 of rivers in south-western Europe (including France). The authors highlighted a general
505 decrease in phosphorus, a general decrease in phytoplankton biomass during summer, as
506 well as an increase in the summer abundance of dinoflagellates over the past 20 years. It is
507 well-known that the nutrient availability is a major factor for algal growth (Tilman *et al.*,
508 1982). In general, diatoms dominate under high-nutrient and turbulent conditions whereas
509 dinoflagellates prevail under periods of stratified conditions and low nutrient
510 concentrations (Margalef, 1978). Under low-nutrient conditions, the use of mixotrophy in
511 algal flagellates may represent an advantage, making the growth not strictly dependent on
512 the inorganic nutrients supply (Cloern and Dufford, 2005; Smayda, 2002). Considering the
513 heterogeneity of phytoplankton taxonomic classification levels in the REPHY database, we
514 were not able to associate species, mainly dinoflagellates, to mixotrophic and heterotrophic
515 functional groups. However, potential groups where mixotrophic species are present, like
516 *Gymnodinium-Gyrodinium*, showed an increasing trend in abundance.

517 Phytoplankton changes are usually not disconnected from other ecosystem changes
518 and interactions between ecosystem components may amplify or attenuate changes at the
519 ecosystem level. Multiple studies have recently highlighted the response of marine
520 ecosystems to climate change (*e.g.* Beaugrand, 2009; Harley *et al.*, 2006; Hays *et al.*,
521 2005). These changes can be observed at several levels of the ecosystem such as primary
522 producers, zooplankton and fish communities (Beaugrand, 2004; Beaugrand, 2009;
523 Hawkins *et al.*, 2003). Thus, they may alter community properties like diversity,
524 dominance, community composition and lead to modifications in the marine ecosystems

525 functioning (Beaugrand, 2004; Wernberg *et al.*, 2011). However, the lack of long-term
526 data on other biological components makes difficult to separate ecosystem responses to
527 changes. There is a growing interest in analysing phytoplankton data sets in coastal areas,
528 establishing trends for phytoplankton as well as for associated variables such as inorganic
529 nutrients. These constitute the activities being addressed by the ICES Working Group on
530 Phytoplankton and Microbial Ecology (WGPME, Li *et al.*, 2011).

531

532 Conclusion and future works

533

534 The analysis of the phytoplankton time series allowed the identification of important
535 temporal changes in the phytoplankton community structure between 1992 and 2011 along
536 the French coast of the eastern English Channel and the southern Bight of the North Sea.
537 We found some evidence of changes in the community composition as well as significant
538 long-term trends in the abundance of some taxonomic units that seem to reflect climate-
539 driven changes in the coastal environment. Since phytoplankton plays a key role in the
540 pathways and efficiency of energy transfer through the upper trophic levels and
541 participates in the nutrients cycles, the observed long-term changes in the phytoplankton
542 community structure may have important ecological consequences (*e.g.* Edwards and
543 Richardson, 2004; Hays *et al.*, 2005). However, further studies are necessary to assess the
544 modifications of the phytoplankton community and the factors forcing these changes. In
545 particular, the causal relationships between climate and environmental drivers on one hand,
546 and phytoplankton communities and species phenology on the other hand, need to be
547 described to assess whether changes correspond to a major shift in the ecosystem structure
548 and functioning. Future works will therefore follow two lines of research: 1) apply the
549 same statistical framework to other REPHY monitoring sites, test for similar occurrences

550 of changes and relate changes to environmental drivers, 2) determine the ecological niche
551 of the most significant species and assess whether and how the phytoplankton phenology
552 responds to environmental changes.

553

554 Supplementary material

555

556 The following supplementary material is available at ICESJMS online version of the
557 manuscript:

558 Table S1 A list with assigned codes and taxonomic groups of phytoplankton taxa
559 identified.

560 Figure S1 Seasonal patterns of the three main phytoplankton groups at ten stations along
561 the French Coast of the English Channel and the southern Bight of the North Sea between
562 1992 and 2011.

563 Figure S2 Seasonal patterns of environmental variables at ten stations along the French
564 Coast of the English Channel and the southern Bight of the North Sea between 1992 and
565 2011.

566 Figure S3 Long-term trends in the abundance of six taxonomic units over the sampling
567 period at Dunkerque, Boulogne and the Bay of Somme.

568

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570

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580

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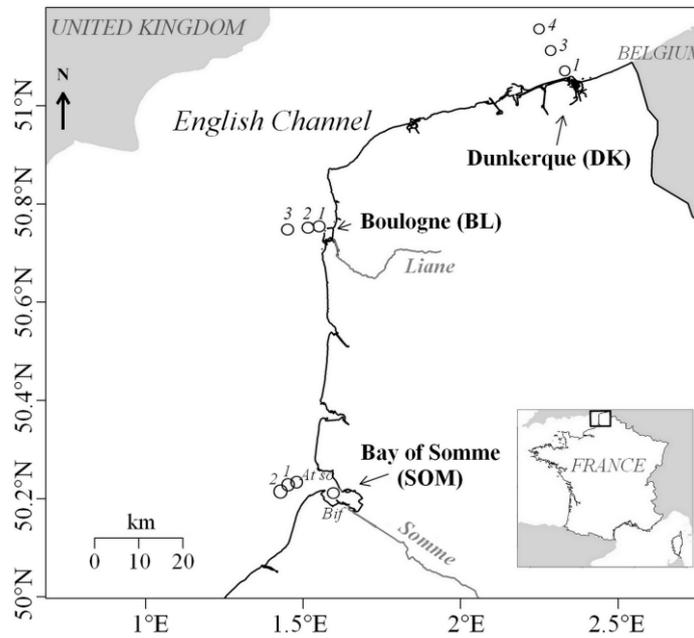
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768 Figures

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771 Figure 1. Location of the phytoplankton sampling stations in the Bay of Somme and
772 Boulogne transects (English Channel) and Dunkerque transect (southern Bight of the North
773 Sea).

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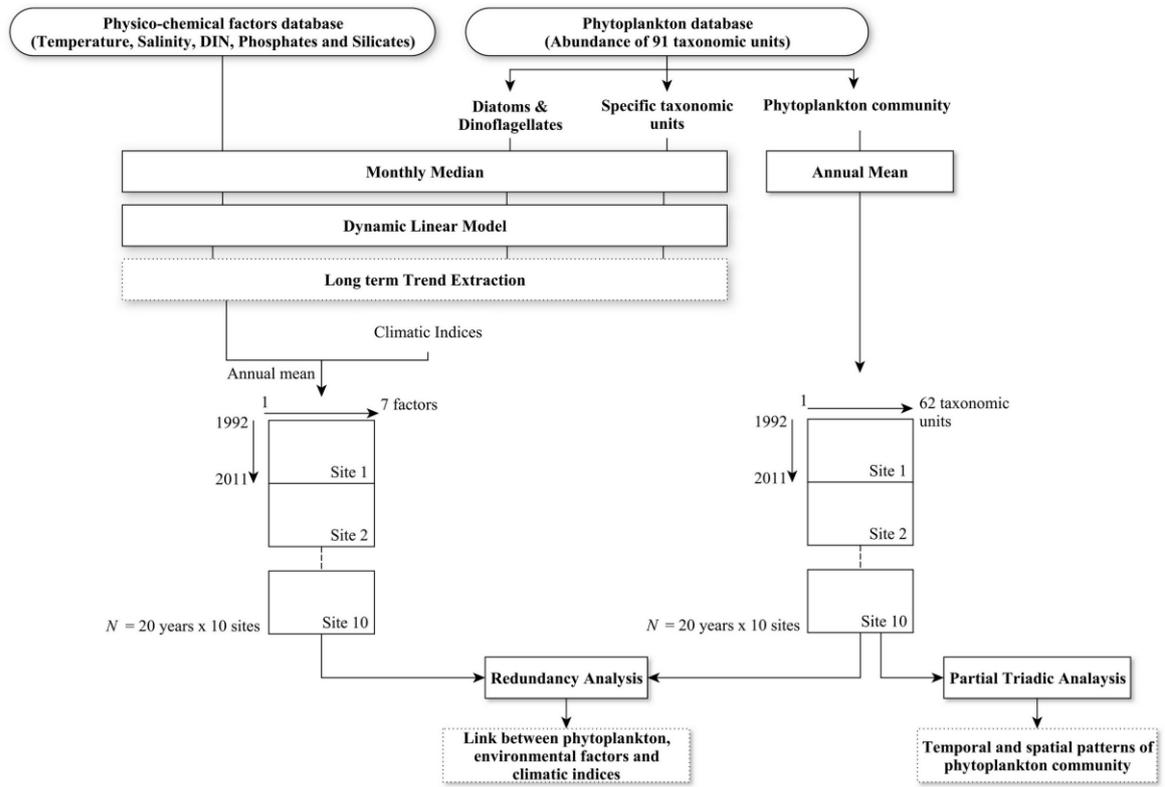
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787 Figure 2. Pathways of statistical analyses followed in this study.

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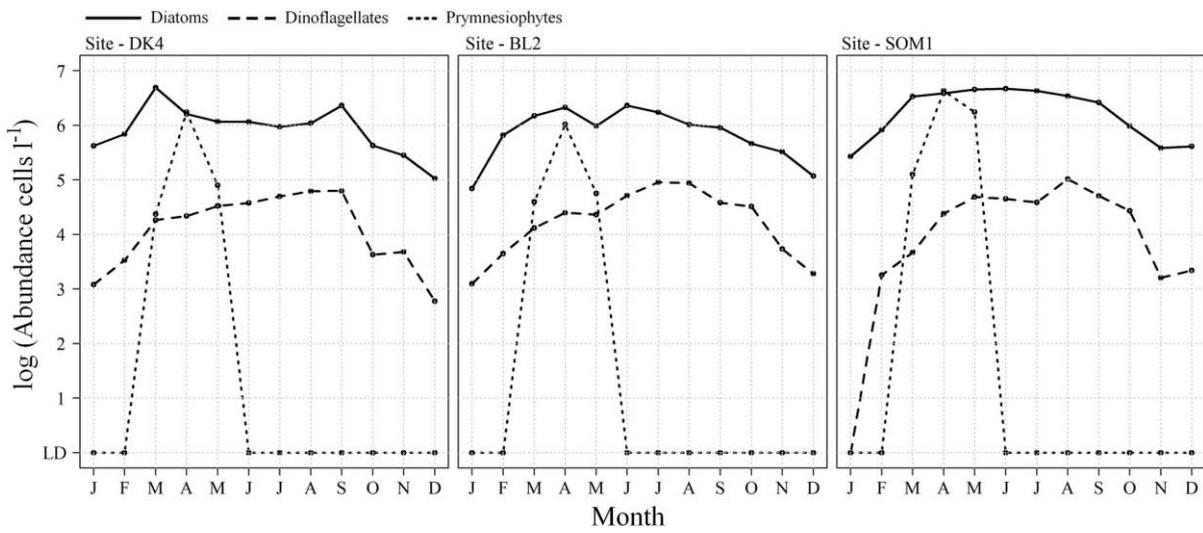
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802 Figure 3. Median monthly abundance of the main phytoplankton taxonomic groups in the
803 three transects between 1992 and 2011. LD (detection limit).

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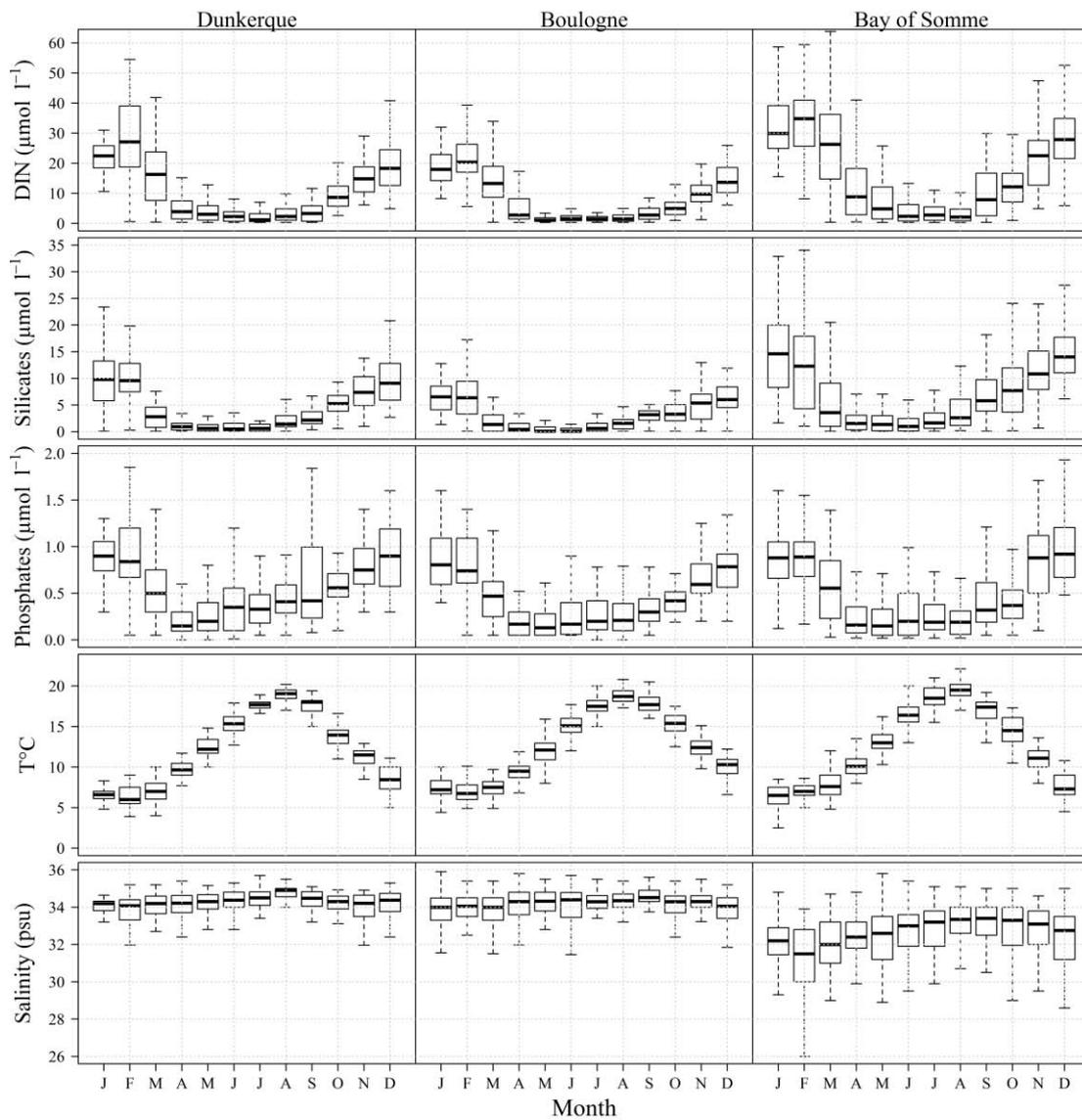
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821 Figure 4. Box-and-whisker plot of the annual variation of physico-chemical measurements

822 between 1992 and 2011 at Dunkerque, Boulogne and the Bay of Somme. Horizontal line

823 denotes the median value for each month, box represents first and third quartiles and

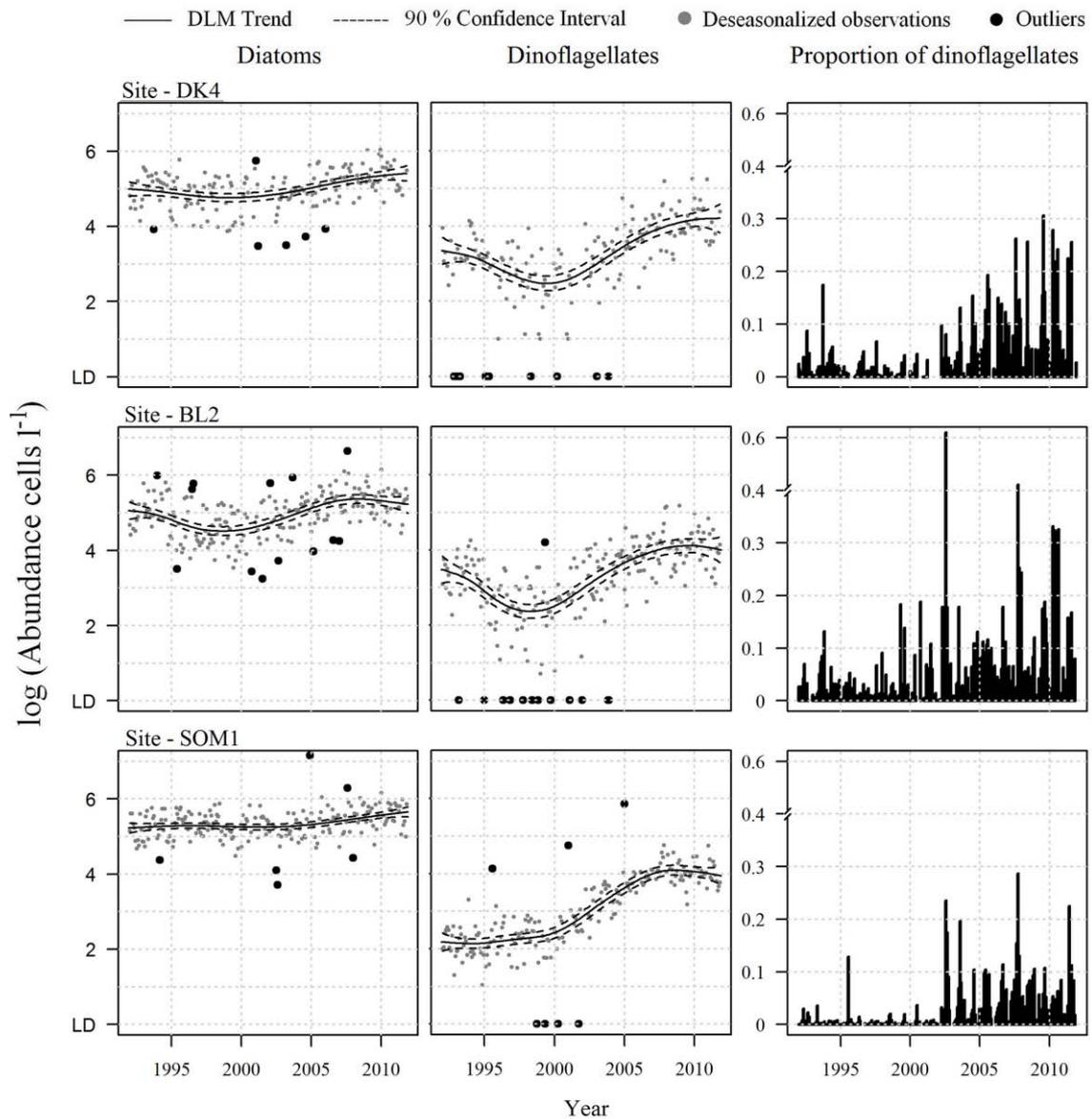
824 whisker represents last value within 1.5 times the interquartile distance.

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831 Figure 5. Long-term trends of diatoms (first column) and dinoflagellates (second column)

832 abundances, and proportion of dinoflagellates relative to diatoms (third column) between

833 1992 and 2011 in three sampling sites. Solid black lines and dotted lines represent trends

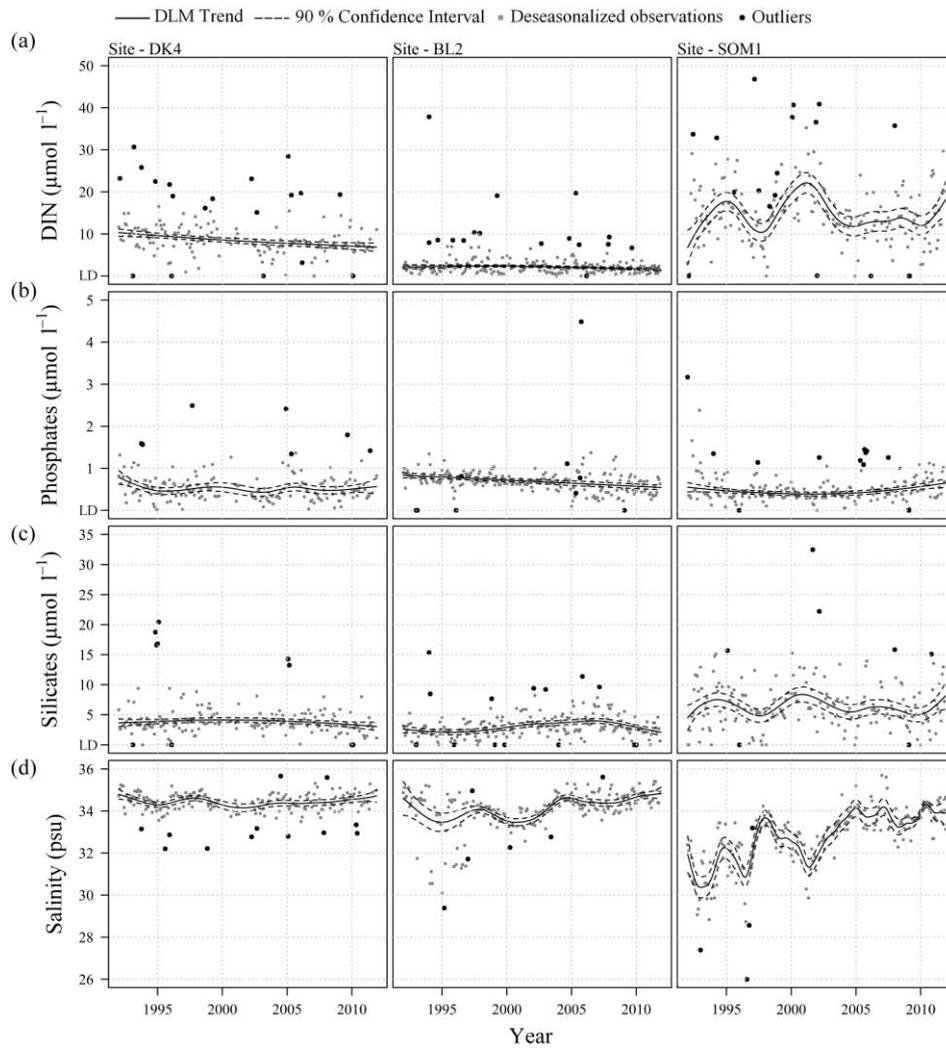
834 and 90% confidence interval respectively, as predicted by the Dynamic Linear Model.

835 Gray and dark dots correspond to deseasonalized observations and outliers respectively.

836 *LD* (detection limit).

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841 Figure 6. Long-term trends of (a) DIN, (b) phosphates, (c) silicates concentration and (d)
 842 salinity between 1992 and 2011 at Dunkerque (Site DK4), Boulogne (Site BL2) and Bay of
 843 Somme (SOM1). Solid black lines and dotted lines represent trends and 90% confidence
 844 interval respectively, as predicted by the Dynamic Linear Model. Gray and dark dots
 845 correspond to deseasonalized observations and outliers respectively. *LD* (*detection limit*).

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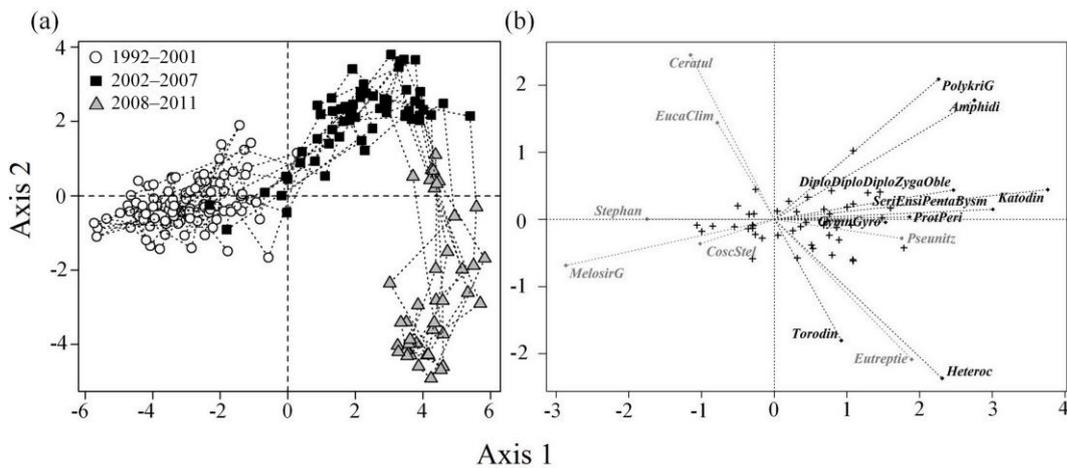
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853 Figure 7. Partial triadic analysis results conducted on the phytoplankton community data
854 from ten sampling stations with the periods 1992–2001, 2002–2007 and 2008–2011. (a)
855 First and second axis for (site; year) objects. Each dotted line represents a given sampling
856 station. (b) First and second axis for taxonomic units. For readability, only taxonomic units
857 with large coordinates in the first and second axis were represented, the other taxa were
858 represented with black crosses. The plot displays 43.3% and 15.4% of the variance in the
859 first and second axis respectively. For details in codes and taxonomic groups of
860 phytoplankton units, see Table S1 in SM.

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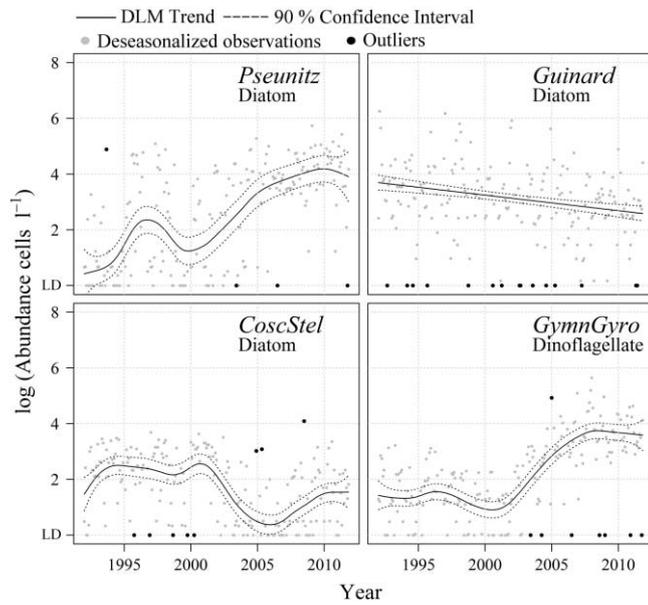
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873 Figure 8. Long term trend in the abundance of four selected taxonomic units over the
874 sampling period at the Bay of Somme, the transect showing the most pronounced
875 variations. Solid black lines and dotted lines represent trends and 90% confidence interval
876 respectively, as predicted by the Dynamic Linear Model. Gray and dark dots correspond to
877 deseasonalized observations and outliers respectively. *LD* (*detection limit*). *CoscStel*
878 *Coscinodiscus-Stellarima*, *Guinard Guinardia*, *GymnGyro Gymnodinium-Gyrodinium*,
879 *Pseunitz Pseudo-nitzschia*.

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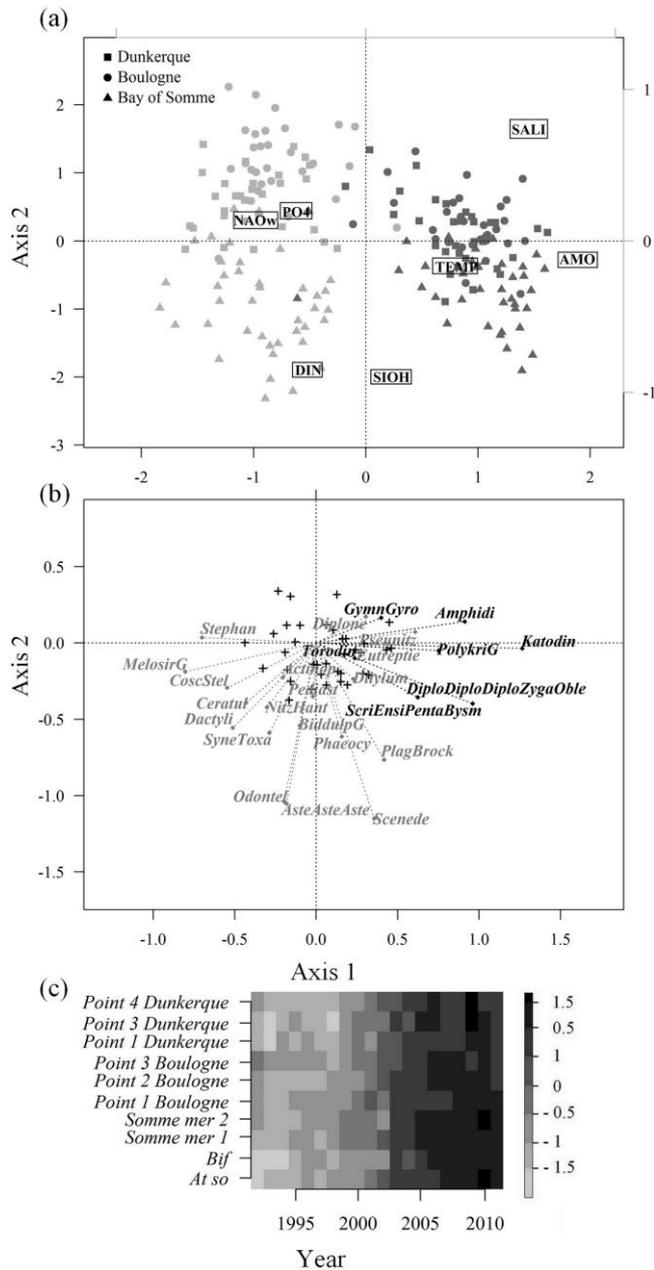
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887 Figure 9. Redundancy analysis of phytoplankton units between 1992 and 2011. (a) RDA
 888 biplot (site; year) score (filled symbols), and environmental variables (box). The bottom-
 889 left scale corresponds to site-year objects and top-right scales to environmental variables,
 890 respectively. Colours on filled symbols represent periods (light gray symbols for years
 891 before 2001, dark gray symbols after 2001). (b) Details for phytoplankton genera (black
 892 labels for dinoflagellates, gray labels for other taxa), genera which percentage of variance
 893 explained in RDA is lower than 20% are indicated by black crosses. (c) Sites-temporal
 894 trajectories represented for the first axis. Abbreviations of environmental variables: DIN

895 (dissolved inorganic nitrogen), SIOH (silicate), PO4 (phosphates), TEMP (temperature),
896 SALI (salinity), NAO_w (winter North Atlantic Oscillation index), AMO (Atlantic
897 Multidecadal Oscillation index). For details in codes and taxonomic groups of
898 phytoplankton units, see Table S1 in SM.
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