

An edited version of this paper was published by [AGU](#).

Importance of the sea surface curvature to interpret the normalized radar cross section

A. A. Mouche^{1,3}, B. Chapron¹, N. Reul^{1,*}, D. Hauser², Y. Quilfen¹

¹Laboratoire d'Océanographie Spatiale, Ifremer, Plouzané, France

²Centre d'Etude des Environnements Terrestre et Planétaires, IPSL, CNRS/UVSQ, Vélizy, France

³Also at Centre National d'Etudes Spatiales, Paris, France.

*: Corresponding author : N. Reul, email address : nreul@ifremer.fr

Abstract:

Asymptotic models (small perturbation and small slope approximation at first-order, Kirchhoff approximation or two-scale model) used to predict the normalized radar cross section of the sea surface generally fail to reproduce in detail backscatter radar measurements. In particular, the predicted polarization ratio versus incidence and azimuth angles is not in agreement with experimental data. This denotes the inability of these standard models to fully take into account the roughness properties with respect to the sensor's configuration of measurement (frequency, incidence, and polarization). On the basis of particular assumptions, to decompose the scattered electromagnetic field between zones covered with freely propagating waves and others where roughness and slopes are enhanced, recent works were able to match observations. In this paper, we do not assume such a decomposition but study the latest improvements obtained in the field of approximate scattering theories of random rough surfaces using the local and resonant curvature approximations. These models are based on an extension of the Kirchhoff Approximation up to first order to relate explicitly the curvature properties of the sea surface to the polarization strength of the scattered electromagnetic field. Consistency with previous approaches is discussed. As shown, dynamically taking into account the sea surface curvature properties of the surface is crucial to better interpret normalized radar cross-section and polarization ratio sensitivities to both sensor characteristics and geophysical environment conditions. The proposed developments, termed the Resonant Curvature Approximation (RCA), are found to reproduce experimental data versus incidence angle and azimuth direction. The polarization sensitivity to the wind direction and incidence angle is largely improved. Finally, Gaussian statistical assumption adopted to derive the analytical expression of the normalized radar cross section is also discussed. In particular, the third-order cumulant function is shown to better reproduce the second-order up-/down-wind azimuth modulation. The proposed developments appear very promising for improvement of our understanding and analysis of both sea surface radar backscatter and Doppler signals.

Keywords: Resonant Curvature Approximation; sea surface radar; backscatter.

1. Introduction

36 As the capabilities of remote sensing instruments ever-increase, new opportunities to
37 develop consistent inversion schemes of the sea surface geometry and kinematic appear.
38 For instance, the high resolution SAR images are now commonly used to retrieve wind
39 fields thanks to the backscattered intensity power [*Monaldo and Kerbaol, 2003*] or the
40 sea surface velocity using the Doppler anomaly analysis [*Chapron et al., 2003, 2005*].
41 Consequently, one could think about consistent inversion merging these two sources of
42 information. To resolve the remaining ambiguities in the normalized radar cross-section
43 (NRCS) interpretation. In particular, it will help to better decipher between wind effects
44 and current impacts on the apparent surface roughness.

45 However, to date, large discrepancies between observations and model predictions still
46 remain. Asymptotic theories as the Small Slope Approximation (SSA) [*Voronovich, 1994*;
47 *Plant, 2002*], Kirchhoff Approximation (KA), or more standard approaches as the Two-
48 Scale Model (TSM) [*Plant, 1986; Thompson, 1988; Romeiser et al., 1997*] and the Small
49 Perturbation Method at first order (SPM-1) [*Valenzuela, 1978*] fail to correctly predict
50 the NRCS in both HH and VV co-polarizations under all environmental conditions and
51 sensor configurations and characteristics. Data comparisons by several authors system-
52 atically recall these weaknesses. The above cited models cannot reproduce the NRCS
53 in both VV and HH polarizations with respect of incidence angle (see e.g. [*Voronovich*
54 *and Zavarotny, 2001; Kudryavtsev et al., 2003; Plant, 2003; Mouche et al., 2006b*]). Data
55 acquired simultaneously for both co-polarizations help to study more precisely the polar-
56 ization ratio (PR) defined as the ratio of the NRCS in VV over the NRCS in HH. *Mouche*

57 *et al.* [2005] show that in C-band the PR is azimuth dependent with respect to the wind
58 direction. In C-band, this dependency was clearly evidenced for incidence larger than 30° .
59 *Mouche et al.* [2006b] show that the first order expansion of the Small Slope Approxima-
60 tion (SSA-1) or TSM fail to reproduce this azimuth dependency. By construction SSA-1,
61 SPM-1 and KA cannot reproduce any azimuthal variation for the PR. This is a strong
62 limitation of such models. In opposite, TSM can lead to an azimuth dependency, but
63 generally not in agreement with the data. This difference between TSM and other cited
64 models clearly results from the efforts made in the TSM formalism to take into account
65 the depolarization effects of the largest waves on the Bragg resonant waves (e.g. [*Plant*,
66 1986; *Thompson*, 1988; *Romeiser et al.*, 1997; *Valenzuela*, 1978]). However, as often dis-
67 cussed, TSM's formalism is somehow arbitrary to the choice of the parameter separating
68 large modulating and small modulated waves. To overcome these issues and to propose
69 consistent inversion schemes, one certainly need to advance in the field of approximate
70 theories of scattering from random rough surface [*Elfouhaily and Guérin*, 2004].

71 In particular, *Elfouhaily et al.* [2003a] proposed a new asymptotic theory for wave
72 scattering from rough surface taking into account the curvature effect of the surface on
73 the scattered field. This curvature effect is a first order correction term of the zeroth
74 order expression of the scattered field given by the Kirchhoff Approximation (KA). The
75 Local Curvature Approximation (LCA-1) formalism has the advantage to dynamically
76 reach both KA and the Small Perturbation Method at first order (SPM-1) depending
77 upon the surface properties. *Mouche et al.* [2006a] applied LCA-1 to the problem of
78 microwave scattering from ocean surface. Based on analytical comparisons with the Two-
79 Scale Model but also data comparisons, it was shown that LCA-1 polarization sensitivity

80 was very close to TSM and thus somehow inadequate to reproduce the NRCS of the ocean
81 surface. But, in comparison with the TSM, LCA-1 solution is more general as it unifies
82 dynamically SPM and KA asymptotic solutions and removes the issue concerning the
83 dividing scale of the surface. Based on this analysis, a new asymptotic solution which
84 provides a more realistic polarization sensitivity than LCA-1 has been proposed to restrict
85 the curvature correction to the so-called resonant Bragg waves. This model, namely the
86 Resonant Curvature Approximation (RCA), conserves the dynamical properties of LCA-
87 1 to reach KA and SPM-1 asymptotic solution. Accordingly, the polarization ratio at a
88 given incidence angle will be sea surface roughness dependant.

89 In this paper, we first expose the remaining issues in the field of the sea surface NRCS
90 prediction. Data and model comparisons are presented to illustrate our comments. Then,
91 we briefly present the asymptotic solutions based on the extension of the KA to take into
92 account the surface curvature effect of depolarization for the incident electromagnetic
93 waves. Comparisons between the model and radar data help to discuss the importance of
94 the sea curvature for backscattered signal interpretation. The issue about the statistical
95 representation of the sea surface is also commented to resolve the up/down-wind asymme-
96 try of the observed NRCS. Our conclusions and perspectives for the use of these models
97 in the field of ocean remote sensing ends this work.

2. Position of the problem

To date, there is no electromagnetic model able to reproduce the NRCS in both VV and HH polarizations for all incidence angles, radar wavelength and wind conditions (speed, direction). In particular, the polarization sensitivity is not correctly reproduced. Figure 1 illustrates this point with two examples of PR in Ku and C band. The PR is defined

as the ratio of the NRCS in VV over the NRCS in HH polarization. On figure 1 (a), we present the PR versus incidence angle calculated from Ku-band NRCS in both VV and HH polarizations measured by NSCAT. The data acquired in Ku-band were already presented by *Quilfen et al.* [1999]. We consider incidence angles from 20° to 50°. a_0^{pp} stands for the coefficient of the standard three-term Fourier model used for empirical formulations of the NRCS versus viewing angle with respect to wind direction:

$$\sigma_0^{pp}(U, \theta, \Phi) = a_0^{pp}(U, \theta) + a_1^{pp}(U, \theta) \cos(\Phi) + a_2^{pp}(U, \theta) \cos(2\Phi), \quad (1)$$

98 where U is the near-surface wind speed, θ the radar's incidence angle and Φ the wind
 99 direction relative to the radar's azimuth look direction. pp denotes the co-polarization
 100 considered. With radar data, model predictions given by SPM-1, TSM, KA and SSA-1
 101 are presented. The sea surface description is given by the unified spectrum for short
 102 and long wind driven waves proposed by *Elfouhaily et al.* [1997]. To be consistent with
 103 these observations, we consider only the isotropic part of the spectrum. Analytical NRCS
 104 expressions of this models are recalled in the appendix. SPM-1 and TSM predictions
 105 are only presented for incidence angles greater than 20° as they are not valid for lower
 106 incidences. As already reported (see e.g., review by *Valenzuela* [1978]; *Kudryavtsev et al.*
 107 [2003]), SPM-1 underestimates the NRCS in HH polarization whereas it is rather good
 108 for the VV polarization. The direct consequence is an overestimation of the PR. Adding
 109 a modulation from the longer waves to the resonant Bragg waves (which provide SPM-1
 110 backscattering) through a TSM enables to better reproduce the depolarization effect on
 111 the predicted NRCS. Yet, NRCS in HH polarization is found lower than the measurements
 112 and TSM's PR is not in agreement with the data. As KA does not provide any polarization
 113 sensitivity, it is clear that it cannot match the data in VV or HH polarizations for the

114 largest incidence angles where there is a significant polarization difference between the
115 backscattered signal in co-polarizations. SSA-1, by construction, imposes a polarization
116 sensitivity too strong. This comes from the fact that as the incidence angle increases,
117 SSA-1 tends very quickly to the SPM-1 asymptotic solution. To lower this effect, higher
118 orders of SSA must be considered. However, *Voronovich and Zavarotny* [2001] showed
119 that the addition of the second order is not sufficient to reproduce data. The conclusion of
120 these comparisons is that none of the model presented above is able to predict the correct
121 polarization sensitivity versus incidence angle in term of mean level.

122 On figure 1 (b), we present the PR versus azimuth angle obtained from the NRCS in
123 both VV and HH polarizations measured by the STORM radar. This data acquired in
124 C-band were already presented by *Mouche et al.* [2006b]. A complete presentation of the
125 radar could be found in [*Hauser et al.*, 2003]. We consider a 40° incidence angle and
126 a 11m/s wind speed. Obviously, none of the model is able to reproduce the observed
127 azimuthal PR modulation. As observed, the PR is dependent on geophysical parameters
128 such as the wind direction. Necessarily, the PR expression given by an asymptotic model
129 must be thus sensitive to the sea surface roughness description. This points out the
130 limitations of the above cited models for the understanding of the scattering processes at
131 the sea surface and the development of consistent inversion scheme to retrieve geophysical
132 parameters.

133 Taking into account these limitations in existing models, we propose to apply two
134 extended versions of the Kirchhoff Approximation model based on a first order correction
135 attributed to the surface curvature.

3. Extended Kirchhoff model for the Normalized Radar Cross Section

3.1. Coordinates system and definitions

136 To expose the general scattering problem, we adopt the same vectorial conventions than
 137 used by *Elfouhaily and Guérin* [2004] in their review on electromagnetic scattering theo-
 138 ries. The right cartesian coordinate system is defined by the triplet of normalized vectors
 139 $(\hat{x}, \hat{y}, \hat{z})$, where the z-axis is directed upward. Σ is the rough surface which separates the
 140 upper medium and the lower medium (respectively air and water in our specific case).
 141 The (sea) surface elevation is represented by $z = \eta(x, y) = \eta(\mathbf{r})$, where \mathbf{r} is the horizontal
 142 component of the three-dimensional position wave vector $\mathbf{R} = (\mathbf{r}, z)$. According to these
 143 conventions, we consider a incident downward propagating electromagnetic plane wave
 144 with a wave-vector $\mathbf{K}_0 = (\mathbf{k}_0, -q_0)$. The up-going scattered waves is characterized by the
 145 wave-vector $\mathbf{K} = (\mathbf{k}, q_k)$. \mathbf{k}_0 and \mathbf{k} are the horizontal components of the incident and
 146 scattered waves whereas q_0 and q_k are the vertical ones. We define also \mathbf{Q}_h and Q_z related
 147 to the coordinates of the wave numbers \mathbf{K} and \mathbf{K}_0 : $\mathbf{Q}_h = \mathbf{k} - \mathbf{k}_0$ and $Q_z = q_0 + q_k$.

The scattered field above and far away ($R \rightarrow \infty$) from the sea surface is assumed to be related to the incident wave through the relation:

$$\mathbf{E}_s(\vec{R}) = -2i\pi \frac{e^{iKR}}{R} \mathbb{S}(\mathbf{k}, \mathbf{k}_0) \cdot \hat{E}_0. \quad (2)$$

$\mathbb{S}(\mathbf{k}, \mathbf{k}_0)$ is the so-called scattering operator. $\mathbf{E}_s(\mathbf{R})$ and $\mathbb{S}(\mathbf{k}, \mathbf{k}_0)$ can be decomposed on the fundamental polarization basis:

$$\mathbf{p}_v^\pm(\pm k) = \frac{k\hat{z} \mp q_k \hat{k}}{K} \quad \mathbf{p}_h^\pm(\pm k) = \hat{z} \times \hat{k}, \quad (3)$$

where the subscripts v and h indicate the vertical and horizontal polarizations, respectively. The minus superscript corresponds to the down-going plane waves while the plus

superscript to the up-going waves. In this vectors basis, the scattering operator is related to the scattering amplitude 2×2 matrix through:

$$\mathbb{S}(\mathbf{k}, \mathbf{k}_0) = \begin{bmatrix} \mathbf{p}_v^-(\mathbf{k}_0) \\ \mathbf{p}_h^-(\mathbf{k}_0) \end{bmatrix}^T \cdot \begin{bmatrix} S_{vv}(\mathbf{k}, \mathbf{k}_0) & S_{vh}(\mathbf{k}, \mathbf{k}_0) \\ S_{hv}(\mathbf{k}, \mathbf{k}_0) & S_{hh}(\mathbf{k}, \mathbf{k}_0) \end{bmatrix} \cdot \begin{bmatrix} \mathbf{p}_v^+(\mathbf{k}) \\ \mathbf{p}_h^+(\mathbf{k}) \end{bmatrix}, \quad (4)$$

148 where the superscript T stands for the transpose operator. In the 2×2 matrix, the
149 first subscript indicates the incident polarization whereas the second one indicates the
150 scattered polarization configuration considered.

For a given polarization configuration pq , $\mathbb{S}^{pq}(\mathbf{k}, \mathbf{k}_0)$ is further written as:

$$\mathbb{S}^{pq}(\mathbf{k}, \mathbf{k}_0) = \frac{1}{Q_z} \int_{\mathbf{r}} \mathbb{N}^{pq}(\mathbf{k}, \mathbf{k}_0; \eta(\mathbf{r})) e^{-iQ_z \eta(\mathbf{r})} e^{-i\mathbf{Q}_H \cdot \mathbf{r}} d\mathbf{r}, \quad (5)$$

151 where $\mathbb{N}^{pq}(\mathbf{k}, \mathbf{k}_0; \eta(\mathbf{r}))$ is a Kernel depending on the approach considered to establish the
152 solution.

The scattering cross-section is given by the incoherent second order statistical expression:

$$\sigma^{pq} = \langle |\mathbb{S}^{pq}(\mathbf{k}, \mathbf{k}_0)|^2 \rangle - |\langle \mathbb{S}^{pq}(\mathbf{k}, \mathbf{k}_0) \rangle|^2 \quad (6)$$

3.2. Local and Resonant Curvature Approximations

153 Based on the work performed by *Elfouhaily et al.* [2003b], *Elfouhaily et al.* [2003a]
154 expanded the scattering matrix up to the first order such as:

$$\begin{aligned} \mathbb{S}^{pq}(\mathbf{k}, \mathbf{k}_0) &= \frac{\mathbb{K}(\mathbf{k}, \mathbf{k}_0)}{Q_z} \int_{\mathbf{r}} e^{-iQ_z \eta(\mathbf{r})} e^{-i\mathbf{Q}_H \cdot \mathbf{r}} d\mathbf{r} \\ &- i \int_{\mathbf{r}} \int_{\boldsymbol{\xi}} T(\mathbf{k}, \mathbf{k}_0; \boldsymbol{\xi}) \hat{\eta}(\boldsymbol{\xi}) e^{-iQ_z \eta(\mathbf{r})} e^{-i(\mathbf{Q}_H - \boldsymbol{\xi}) \cdot \mathbf{r}} d\boldsymbol{\xi} d\mathbf{r}, \end{aligned} \quad (7)$$

155 where

$$T_{\text{lca}}(\mathbf{k}, \mathbf{k}_0; \boldsymbol{\xi}) = [\mathbb{B}(\mathbf{k}, \mathbf{k}_0; \boldsymbol{\xi}) - \mathbb{K}(\mathbf{k}, \mathbf{k}_0)], \quad (8)$$

156 is a kernel defined to take into account the surface curvature effects on the scattered field.
 157 \mathbb{B} is the Bragg Kernel and \mathbb{K} is the Kirchhoff Kernel (see e.g. *Elfouhaily et al.* [2003a] for
 158 their analytical expression). In Eq.(), the second term represents a first order correction
 159 to KA given by the first term. This first order curvature term has the property to reach
 160 dynamically both KA and SPM-1 limits with respect to the frequency and the properties
 161 of the surface considered. $\hat{\eta}(\boldsymbol{\xi})$ is the Fourier transform of the surface height function
 162 $\eta(\mathbf{r})$, $\boldsymbol{\xi}$ the wave-number of the surface in the spectral domain.

163 In [*Mouche et al.*, 2006a], we showed that we could choose a formulation for this Kernel
 164 which conserves all the dynamic properties of this proposed solution but with a weaker
 165 polarization sensitivity, considering only the curvature effect of the resonant Bragg waves.
 166 In this case, the Kernel expression is:

$$T_{\text{rca}}(\mathbf{k}, \mathbf{k}_0; \boldsymbol{\xi}) = [\mathbb{B}(\mathbf{k}, \mathbf{k}_0; \boldsymbol{\xi}) - \mathbb{K}(\mathbf{k}, \mathbf{k}_0)]\delta(\boldsymbol{\xi} - \mathbf{Q}_H). \quad (9)$$

167 As already discussed by *Mouche et al.* [2006a], this solution can be compared with the
 168 improved Green's function method proposed by *Shaw and Dougan* [1998] excepted that
 169 the formulation helps to preserve the required shift and tilt invariance properties due to
 170 the LCA-1-like formalism of the RCA solution.

Assuming Gaussian statistics for the sea surface description, the derivation of the NRCS using the scattering matrix expansion up to the first order for any expansion such as $N^{pq}(\mathbf{k}_0, \mathbf{k}) = N_0^{pq}(\mathbf{k}_0, \mathbf{k}) + \int_{\boldsymbol{\xi}} N_1^{pq}(\mathbf{k}_0, \mathbf{k}; \boldsymbol{\xi})\hat{\eta}(\boldsymbol{\xi})e^{i\boldsymbol{\xi}\cdot\mathbf{r}}d\boldsymbol{\xi}$ was already done and discussed in the context of LCA/RCA models by *Mouche et al.* [2006a]. In the case of microwave scattering from the sea surface sea surface, it was concluded that the predicted NRCS is very similar than the one using the phase perturbation method firstly proposed by *Berman and Dacol* [1990] and then applied by *Voronovich and Zavarotny* [2001] in the

context of SSA-2. Thus, in this paper, for RCA, we consider that the first order term of the Volterra series in the scattering matrix expansion is a small perturbation in the phase term of the zeroth order contribution in Eq. () such as:

$$\mathbb{S}(\mathbf{k}, \mathbf{k}_0) = \mathbb{K}_0(\mathbf{k}, \mathbf{k}_0) \int_{\mathbf{r}} e^{-iQ_z \eta(\mathbf{r})} e^{-iQ_z \delta_{\mathbf{k}, \mathbf{k}_0} \eta(\mathbf{r})} e^{-i\mathbf{Q}_H \cdot \mathbf{r}} d\mathbf{r}, \quad (10)$$

with

$$\delta_{\mathbf{k}, \mathbf{k}_0} \eta(\mathbf{r}) = \int_{\xi} \frac{T_{\text{lca/rca}}(\mathbf{k}, \mathbf{k}_0; \xi)}{\mathbb{K}(\mathbf{k}, \mathbf{k}_0)} \hat{\eta}(\xi) e^{i\xi \cdot \mathbf{r}} d\xi. \quad (11)$$

171 $\tilde{\eta}(\mathbf{r}) = \eta(\mathbf{r}) + \delta_{\mathbf{k}, \mathbf{k}_0} \eta(\mathbf{r})$ can be seen as a modified surface elevation. In the case of RCA,
 172 the first order curvature term is applied on the small resonant waves responsible of the
 173 Bragg scattering mechanism according to SPM-1 theory. Thus, the simplification of the
 174 scattering matrix expansion to describe the contribution of the curvature correction term
 175 through a phase modification is consistent with the small perturbation hypothesis.

Using the modified surface elevation for the statistical derivation of the NRCS, in the Kirchhoff integral, the characteristic function $\langle e^\eta \rangle$ is replaced by $\langle e^{\tilde{\eta}} \rangle$. Under Gaussian statistics, this formalism enables to have a tractable expression for the NRCS:

$$\sigma_0^{pq}(\theta, \phi) = \left| \frac{\mathbb{K}(\mathbf{k}, \mathbf{k}_0)}{Q_z} \right|^2 e^{-Q_z^2 \tilde{\rho}(0)} \int_{\mathbf{r}} [e^{-Q_z^2 \tilde{\rho}(\mathbf{r})} - 1] e^{-i\mathbf{Q}_H \cdot \mathbf{r}} d\mathbf{r}, \quad (12)$$

with:

$$\tilde{\rho}(\mathbf{r}) = \int_{\xi} \left| 1 + \frac{T_{\text{lca/rca}}(\mathbf{k}, \mathbf{k}_0; \xi)}{\mathbb{K}(\mathbf{k}, \mathbf{k}_0)} \right|^2 S(\xi) e^{i\xi \cdot \mathbf{r}} d\xi. \quad (13)$$

176 $\tilde{\rho}(\mathbf{r})$ is the so-called modified correlation function of a filtered spectrum and $S(\xi)$ the sea
 177 surface elevation spectrum. In the following, we use this formulation.

4. Results and Discussion

4.1. With Gaussian statistics for the sea surface description

178 As a first comparison between data and model, we present the PR versus incidence angle
179 given by RCA, LCA-1 and KA for a 10m/s wind speed in C- and Ku-band in the case
180 of an isotropic sea surface on figures 2 (a) and 2 (b). In both cases, we observe that the
181 curvature correction term in RCA or LCA-1 ensures to the extended KA to get polarization
182 sensitivity as the incidence increases. As expected, since the curvature correction term in
183 RCA is restricted to the resonant Bragg waves, the induced polarization sensitivity is less
184 than for LCA-1. From the data comparisons in Ku and C band presented here, we have a
185 better agreement with RCA than with LCA-1. Figure 2 (c) presents the PR predicted by
186 RCA and SSA-1 versus incidence angle for three frequencies. Focusing on the frequency
187 dependency, we observe that the PR decreases when the frequency increases with both
188 models. For SSA-1 (same comments can be done with KA or SPM-1) the only frequency
189 dependency comes through the Kernel definition and is too small. In LCA-1 or RCA, the
190 surface description controls the PR. This ensures, by construction, to reach dynamically
191 both SPM-1 and KA asymptotic solutions, lead to KA results when $k_0 \rightarrow \infty$ or when the
192 perception of the surface by the sensor is flat (no curvature). Numerical computations of
193 the PR show that the RCA solution is more frequency sensitive than SPM-1, SSA-1 or
194 KA. This explains why the model agrees well with the data in both Ku and C bands on
195 figures 2 (a) and 2 (b).

196 As already mentioned above, another important feature in the backscattered signal
197 for a given polarization, is the azimuth modulation with respect of the wind direction
198 relative to the radar's azimuth look direction. From this modulation, we can infer the

199 wind direction. Moreover, data analysis show that this modulation is incidence angle and
200 frequency dependent. Most of electromagnetic models are able to reproduce these two
201 dependencies. If this modulation is quite well reproduced in each co-polarization, it is
202 not sufficient as the predicted modulation is not polarization sensitive (or not sufficiently
203 for TSM). The comparison on figure 1 shows it. On figures 3 (a) and 4 (a), we present
204 PR measured with STORM data for two cases of different wind speeds (11 and 14 m/s).
205 These measurements exhibit an azimuth modulation dependent on the wind direction. To
206 show the impact of the curvature model family, we present the results given by LCA-1 and
207 RCA. As an example we also plot the SSA-1 results. LCA-1/RCA formalism enables to
208 reproduce an azimuth modulation for the PR due to the curvature correction term. More
209 precisely, following our hypothesis on Gaussian statistics for the sea surface representation,
210 such kind of model can only reproduce the first order harmonic of the PR. We will see in
211 the next section that this issue can be improved considering skewness effect. Comparisons
212 between LCA-1 and RCA confirms that the mean level of the PR is better reproduced
213 by RCA. The curvature effect attributed to the resonant waves provides also a better
214 trend for the azimuth modulation amplitude. Figures 3 (c-d) and 4 (c-d) presents the
215 NRCS in both VV and HH polarizations separately for these two wind speeds which are
216 the direct measurable quantities. As expected from our previous conclusion (see figures
217 2), RCA predicts a correct mean level for the NRCS in both HH and VV polarizations
218 whereas LCA and SSA-1 underestimate the NRCS in HH. As for the PR, due to Gaussian
219 statistics, the three models only reproduce a first order harmonic modulation. However,
220 with the curvature effect, this modulation is polarization sensitive. This fundamental
221 aspect enables to predict a PR sensitive to geophysical parameters such as wind speed or

222 wind direction relative to the radar's azimuth look direction. To be really complete with
223 these data comparisons in C-band, we present on figure 3 (b) and 4 (b) the difference
224 of NRCS $\sigma_0^{vv} - \sigma_0^{hh}$ in linear unit, DP hereafter, versus the wind direction relative to
225 the radar's azimuth look direction for the two wind speeds considered here. Once again,
226 RCA model is in better agreement with the data than other models. DP quantity is
227 interesting since in literature, authors proposed to decompose the measured NRCS in a
228 polarized and a scalar part (e.g. [Quilfen et al., 1999]). Using DP quantity, we remove
229 the scalar contribution to keep only the polarized part of the signal. In the case of
230 backscattering, this part is taken into account through the first order curvature correction
231 term in LCA/RCA formalism. As the comparisons with data are also satisfying for the
232 DP quantity, we are confident in the first order resonant curvature term of RCA.

233 To evaluate the ability of RCA to reproduce the data for different incident wavelengths,
234 we also present a set of comparisons in X-Band. These data were collected during POL-
235 RAD'96 experiment. The wind speed considered here was provided by buoys, ships and/or
236 model. The data set and the instrument were presented in details by *Hauser et al.* [1997].
237 On the figure 5, as for STORM data, we present the PR (a), DP (b) and the NRCS (c-d)
238 in both co-polarizations. In this case, from buoys and ship measurements, the wind speed
239 is approximatively 8 m/s (we consider the mean of the four available measurements).
240 As we already concluded for the C-Band, RCA/LCA formalism enables to reproduce an
241 azimuth modulation for the PR. RCA predictions give a better agreement with the data
242 than any of the other models presented in this paper.

243 These good agreement between RCA model and the data in C- and X- Band for the
244 different sets of four comparisons has to be seen as a cross-validation of the RCA model as

245 it would be easy to resolve only one aspect of the NRCS models but destabilizing an other.
246 The consistency between RCA with data on these four aspects, where SSA-1, LCA-1 or
247 KA fails (at least on one of these aspects), show the robustness of the model.

4.2. On the skewness effect on the NRCS

248 In all the data presented here, we observe a difference between the NRCS levels observed
249 upwind and downwind. This asymmetry (UDA hereafter) was already evidenced by many
250 authors thanks to radar data and is taken into account in the standard three-term Fourier
251 model used for empirical formulations of the NRCS versus viewing angle with respect
252 to wind direction (e.g. [*Stoffelen and Anderson, 1997; Bentamy et al., 1999; Herbasch,*
253 *2003*]). Measurements, reveal that the NRCS level in upwind direction is greater than
254 in downwind direction at high incidence angles (say $> 30^\circ$) and lower at small incidence
255 angles. In empirical models such as CMOD type models in C-Band or SASS in Ku-
256 Band, the a_1 coefficient (see Eq.) takes into account this second order effect in the
257 azimuth modulation. In physical models, a standard explanation for this asymmetry is
258 done through the hydrodynamic modulation of Bragg waves. However, data analysis
259 combining the dual co-polarization to remove the so-called scalar contribution thanks to
260 the DP quantity firstly performed by *Chapron et al. [1997]* reveal that the contribution
261 of the Bragg waves to the backscattered signal is more important downwind whereas the
262 total contribution of the total UDA predicts more signal upwind. This analysis proves that
263 if the hydrodynamic modulation of Bragg waves exists, its effect is dominated by an other
264 one. Indeed, the slopes of the longer waves are more steep leeward than windward. On
265 the longer waves slopes, the lower amplitude Bragg waves are predominantly windward.
266 On the windward side, the longer waves may be slightly steeper and rougher to decrease

267 the polarized contribution. This explains the observed UDA in the DP signal. Moreover,
 268 this implies that the opposite contribution of the UDA comes through the skewed form
 269 of the longer and the breaking waves which participate to the backscattering through
 270 the Kirchhoff mechanism. This idea consisting in associating the UDA of the breakers
 271 in a NRCS model was firstly applied by *Kudryavtsev et al.* [2003]. Same kind of dual
 272 co-polarizations analysis than the one performed by *Chapron et al.* [1997] or *Quilfen et al.*
 273 [1999] was proposed by *Mouche et al.* [2006b] in C-Band. This study based on STORM
 274 data for observations and *Kudryavtsev et al.* [2003] model also supports the idea of the
 275 breaking waves importance for the UDA through a scalar contribution to the NRCS.

276 Thus, we need to consider higher moments in the statistical derivation of the NRCS.
 277 The third order correction to the characteristic function enables to consider the skewness
 278 effect of the waves. In the local frame of RCA model, the modified characteristic function
 279 up to the third order is simply:

$$\langle e^{jQ_z(\tilde{\eta}_2 - \tilde{\eta}_1)} \rangle \approx e^{-Q_z^2(\tilde{\rho}(0) - \tilde{\rho}(\mathbf{r}))} e^{iQ_z^3 \tilde{S}_{\text{skew}}(\mathbf{r})}, \quad (14)$$

280 where \tilde{S}_{skew} is the skewness function associated to the modified surface height function.
 281 According to RCA formalism it is clear that any correcting order of the characteristic
 282 function can be polarization and frequency dependent. Following this idea, the NRCS
 283 when considering the skewness effect is:

$$\sigma_0(\theta, \phi) = \left| \frac{\mathbb{K}(\mathbf{k}, \mathbf{k}_0)}{Q_z} \right|^2 e^{-Q_z^2 \tilde{\rho}(0)} \int_{\mathbf{r}} [e^{-Q_z^2 \tilde{\rho}(\mathbf{r}) + iQ_z^3 \tilde{S}_{\text{skew}}(\mathbf{r})} - 1] e^{-i\mathbf{Q}_H \cdot \mathbf{r}} d\mathbf{r}. \quad (15)$$

284 In this work we choose for the skewness function a generic formulation as proposed by
 285 *Elfouhaily* [1997] adjusted on surface slope skewness when $\mathbf{r} \rightarrow 0$ [*Cox and Munk*, 1954]:

$$S_{\text{skew}}(\mathbf{r}) \underset{\mathbf{r} \rightarrow 0}{=} -\frac{1}{6}x\sigma_{sx}(x^2\sigma_{sx}^2C_{03} + 3y^2\sigma_{sy}^2C_{21}) \approx -\frac{r^3}{6}\sigma_{sx}^3C_{03}\cos(\phi), \quad (16)$$

$$S'_{\text{skew}}(0) = S''_{\text{skew}}(0) = S'''_{\text{skew}}(0) = 0 \quad (17)$$

286 where ϕ is the angle of the wind direction relative to the radar's azimuth look direction
 287 and C_{03} , C_{21} two empirical coefficients given by *Cox and Munk's* measurements [*Cox and*
 288 *Munk*, 1954].

289 On the figure 6 (a-f), we present the UDA asymmetry of the NRCS in both-co-
 290 polarizations and of the PR as measured with STORM radar and compared with the
 291 prediction of RCA model when considering the skewness effect. To complete the data set,
 292 we also show the UDA given by two empirical models CMOD-IFREMER [*Bentamy et al.*,
 293 1999] and CMOD-5 [*Herbasch*, 2003]. Data indicates that the asymmetry is incidence,
 294 wind speed and polarization dependent. Results given by the RCA model give very good
 295 agreements with data. In particular, RCA is able to predict realistic and different UDA
 296 for VV and HH polarizations. As a direct consequence, the predicted UDA for the PR
 297 is also in agreement with the data. In Ku Band, we compare directly the $a1$ coefficient
 298 in linear scale for a 10 and 15m/s wind speed. We observe on figure 7 that the trend
 299 with incidence angle is rather well reproduced by the model thanks to the third order
 300 correction in the characteristic function.

301 Taking into account the skewness according to Eq. () also modifies the azimuth mod-
 302 ulation. To see the impact on the whole azimuth range, we present the same plot as on
 303 figure 8 where the model is compared to radar data acquired versus azimuth angle. The
 304 results from RCA are obtained with and without skewness effect. We observe that the

305 third order correction gives more realistic trends for the NRCS with azimuth angles, as
306 it adds a significant second order azimuthal modulation. As observed and predicted, this
307 effect is greater in HH than in VV polarization.

5. Conclusion

308 We used recent improvements obtained in the field of electromagnetic scattering wave
309 theories from random rough surfaces to consider the sea surface curvature influence on the
310 radar backscatter measurements. LCA-1 and RCA were applied to the case of scattering
311 from a 2-Dimensional sea surface and compared with other existing models. Comparisons
312 with data showed that RCA results are in better agreement with the data whereas LCA
313 results are very close to the TSM predictions. This difference between the two models
314 comes from the fact that the RCA only takes into account the curvature effect of the
315 resonant Bragg waves to the NRCS prediction.

316 The formalism of LCA/RCA has the advantage to take into account the depolarization
317 effect of the sea surface through a dynamical term which depends on both the configuration
318 of the instrument (incidence, frequency and polarization) and the sea surface curvature
319 properties. This is a key element for an improved understanding of the electromagnetic
320 and oceanic waves interactions. In the framework of RCA, the first order term enables
321 to reproduce the NRCS in both co-polarization versus incidence angle in the microwave
322 domain. Good agreements for each polarization allows the model to reproduce the mean
323 PR. This leads to conclude that the curvature correction impact on the polarization is
324 certainly necessary. From our knowledge, RCA is the only asymptotic electromagnetic
325 theory applied to a 2-dimensional sea surface able to reproduce these results. Satisfying
326 solutions, in term of results, are given by *Kudryavtsev et al.* [2003] and *Kudryavtsev et al.*

327 [2005] who proposed a semi-empirical model based on an explicit decomposition of the
328 sea surface. In this model, the sea surface is separated in two parts. First, a regular one
329 which is responsible for the specular reflection near nadir and for the Bragg scattering
330 of short modulated waves by longer one through a TSM (which means also a separation
331 of the scales for this regular part). Second, zones with enhanced roughness due to effect
332 of breaking waves on the sea surface which produce a scalar contribution to the NRCS
333 through specular reflection on these steep slopes. Obviously a parallel between these
334 approaches can be done as the zeroth order of RCA could be compared to the breaking
335 waves and longer waves contribution invoked by *Kudryavtsev et al.* [2003] and *Kudryavtsev*
336 *et al.* [2005] while the curvature correction term and the Bragg contribution could be
337 associated to the same scattering process of the short resonant waves. Advantages of
338 RCA are of course the absence of any dividing parameter to separate scales of the sea
339 surface. Moreover, as the enhanced roughness zones contribution may be hard to precisely
340 parameterize, it could be convenient to use a model such as RCA which consider both
341 contributions from regular and non-regular surface implicitly through the characteristic
342 function. In a future work, an explicit comparison of these two models will be done. But
343 as an important issue, it can be stated that PR modulations shall follow the roughness
344 distribution and disturbances.

345 Finally, we discussed the implication of the non Gaussian statistics on the NRCS for the
346 sea surface description. Interestingly, it appears from RCA formalism that the charac-
347 teristic function of the scattered field is polarization, incidence and frequency dependent.
348 Moreover, RCA is in agreement with *Chapron et al.* [2003] conclusion about the UDA of
349 the Bragg waves and breaking waves at large incidence angle. Taking into account a third

350 order correction term in the characteristic function as a signature of breaking waves, it ap-
351 pears that extraction of the third moment in the backscattered signal would be interesting
352 to improve our understanding of the impact of breaking events on the NRCS.

353 A model such as RCA can be used to improve our understanding about the backscatter
354 signal modulations. As an example, since at large incidence angles the NRCS in HH
355 is lower than the prediction of KA whereas it is the contrary for the VV polarization,
356 it can simply be concluded that the sensor is very sensitive to the small waves in VV
357 polarization (more backscattered signal) whereas for HH polarization it is the contrary
358 (less backscattered signal) which means that the sensor is more sensitive to the longer and
359 steeper waves in this configuration. At large incidence angles, such a large sensitivity to the
360 longer and steeper waves than to the resonant Bragg waves will induces a larger Doppler
361 shift associated to the remote sensed waves than in VV polarization. Next, this model will
362 be used to derive an imaging radar model based on the correlation function modulation to
363 interpret the variation of the NRCS due to changes in the geophysical parameters. This
364 will help to define a more consistent combined analysis between measured Doppler shifts
365 and backscatter signals.

366 **Acknowledgments.** (Text here)

Appendix A: NRCS expression of existing in the case of Gaussian statistics

367 For convenience, we recall here the expression of the NRCS for SPM-1, TSM, SSA-1
368 and KA models as they are used through the paper. The derivation is done in case of
369 Gaussian statistics. We consider the same coordinates system and definitions than those
370 used to express the RCA solution in section 3 but also in the review on approximated

371 wave scattering theories from random rough surfaces proposed by *Elfouhaily and Guérin*
 372 [2004]. $\mathbf{K} = (\mathbf{k}, q_k)$ and $\mathbf{K}_0 = (\mathbf{k}_0, -q_0)$ respectively denotes the wave numbers of the
 373 scattered and incident waves. $\mathbb{B}(\mathbf{k}, \mathbf{k}_0)$ and $\mathbb{K}(\mathbf{k}, \mathbf{k}_0)$ are the so-called kernels of Bragg and
 374 Kirchhoff Approximations. Their expression can be found in [*Elfouhaily et al.*, 2003b].
 375 $\mathbf{Q}_H = \mathbf{k} - \mathbf{k}_0$ and $Q_z = q_k + q_0$.

A1. Small Perturbation Method at first order

$$\sigma_0^{BR} = |\mathbb{B}(\mathbf{k}, \mathbf{k}_0)|^2 S(\mathbf{Q}_H) \quad (\text{A1})$$

A2. Two Scale Model

$$\sigma_0^{TSM} = \int_{-\infty}^{\infty} d(\tan \Psi) \int_{-\infty}^{\infty} d(\tan \delta) \sigma_0^{BR}(\theta_i) P(\tan \Psi, \tan \delta), \quad (\text{A2})$$

376 where $P(\tan \Psi, \tan \delta)$ is the joint probability density of slopes for the long waves, θ_i the
 377 local angle, and $\sigma_{0_{BR}}$ the NRCS given by the SPM-1 due to the small roughness elements
 378 modulated by the longer waves. In our calculation this probability density is assumed
 379 Gaussian. The calculation of $\sigma_{0_{BR}}$ is done considering the angles corrections given by
 380 *Elfouhaily et al.* [1999] instead of initial Valenzuela's results [*Valenzuela*, 1978]:

$$\theta_i = -\cos^{-1}[\cos(\theta + \Psi) \cos(\tan^{-1} \delta \cos \Psi)]$$

with $S_x = \tan \Psi$ and $S_y = \tan \delta$,

381 the slopes of longer waves in and perpendicular to the incident plane.

A3. Small Slope Approximation at first order

$$\sigma_0^{SSA-1} = \left| \frac{\mathbb{B}(\mathbf{k}, \mathbf{k}_0)}{Q_z} \right|^2 e^{-Q_z^2 \rho(0)} \int_{\mathbf{r}} [e^{-Q_z^2 \rho(\mathbf{r})} - 1] e^{-i\mathbf{Q}_H \cdot \mathbf{r}} d\mathbf{r}. \quad (\text{A3})$$

A4. Kirchhoff Approximation

$$\sigma_0^{KIR} = \left| \frac{\mathbb{K}(\mathbf{k}, \mathbf{k}_0)}{Q_z} \right|^2 e^{-Q_z^2 \rho(0)} \int_{\mathbf{r}} [e^{-Q_z^2 \rho(\mathbf{r})} - 1] e^{-i\mathbf{Q}_H \cdot \mathbf{r}} d\mathbf{r}. \quad (\text{A4})$$

References

- 382 Bentamy, A., P. Queffelec, Y. Quilfen, and K. Katsaros, Ocean surface wind fields esti-
383 mated from satellite active and passive microwave instruments, *IEEE Trans. on Geosc.*
384 *and Remote Sens.*, *37*, 2469–2486, 1999.
- 385 Berman, D., and D. Dacol, Manifestly reciprocal scattering amplitudes for rough surfaces
386 interface scattering, *J. Acous. Soc. of Am.*, *87(5)*, 1990.
- 387 Chapron, B., V. Kerbaol, and D. Vandermark, A note on relationships between sea-
388 surface roughness and microwave polarimetric backscatter measurements: results from
389 polrad'96, *Proc. Int. Workshop POLRAD'96*, pp. 55–64, 1997.
- 390 Chapron, B., F. Collard, and V. Kerbaol, Satellite synthetic aperture radar sea surface
391 doppler measurements, *Proc. of the 2nd Workshop on Coastal and Marine Applications*
392 *of SAR, ESA SP-565*, pp. 133–140, 2003.
- 393 Chapron, B., F. Collard, and F. Ardhuin, Direct measurements of ocean surface velocity
394 from space: Interpretation and validation, *J. Geophys. Res.*, *110(C07008)*, 2005.
- 395 Cox, C., and W. Munk, Measurements of the roughness of the sea surface from photographs
396 of the sun's glitter, *J. Opt. Soc.*, *44(11)*, 838–850, 1954.
- 397 Elfouhaily, T., *A consistent wind and wave model and its application to microwave remote*
398 *sensing of the ocean surface*, Thesis dissertation. Denis Diderot university - Paris 7,
399 France, 1997.
- 400 Elfouhaily, T., and C.-A. Guérin, A critical survey of approximate scattering wave theories
401 from random rough surfaces, *Waves In Rand. Media*, *14(4)*, R1–R40, 2004.
- 402 Elfouhaily, T., B. Chapron, K. Katsaros, and D. Vandermark, A unified directionnal wave
403 spectrum for long and short wind-driven waves, *J. Geophys. Res.*, *102*, 15,781–15,796,

- 404 1997.
- 405 Elfouhaily, T., D. Thompson, D. Vandemark, and B. Chapron, A new bistatic model for
406 electromagnetic scattering from perfectly conducting random surfaces, *Waves In Rand.*
407 *Media*, 9, 281–294, 1999.
- 408 Elfouhaily, T., S. Guignard, R. Awdallah, and D. Thompson, Local and non-local cur-
409 vature approximation: a new asymptotic theory for wave scattering, *Waves In Rand.*
410 *Media*, 13, 321–337, 2003a.
- 411 Elfouhaily, T., M. Joelson, S. Guignard, and D. Thompson, Analytical comparison be-
412 tween the surface current integral equation and the second-order small slope approxi-
413 mation, *Waves In Rand. Media*, 13, 165–176, 2003b.
- 414 Hauser, D., P. Dubois, and G. Caudal, Polarimetric wind-scatterometer measurements
415 during polrad’96, *Proc. of Int. Workshop POLRAD’96*, pp. 55–64, 1997.
- 416 Hauser, D., T. Podvin, M. Dechambre, R. Valentin, G. Caudal, and J.-F. Daloze, Storm:
417 A new polarimetric real aperture radar for earth observations, *Proc. of ESA POLinsar*
418 *Int. Workshop*, 2003.
- 419 Herbasch, H., Cmod5 an improved geophysical model function for ers scatterometry,
420 *ECMWF, Internal Report*, 2003.
- 421 Kudryavtsev, V., D. Hauser, G. Caudal, and B. Chapron, A semiempirical model of the
422 normalized radar cross-section of the sea surface: 1. background model, *J. Geophys.*
423 *Res.*, 108(C3), 2003.
- 424 Kudryavtsev, V., D. Akimov, J. Johannessen, and B. Chapron, On radar imaging
425 of current features: 1. model and comparison with observations, *J. Geophys. Res.*,
426 110(C07016), 2005.

- 427 Monaldo, F., and V. Kerbaol, The sar measurement of ocean surface winds: An overview,
428 *Proc. of the 2nd Workshop on Coastal and Marine Applications of SAR, ESA SP-565*,
429 pp. 15–32, 2003.
- 430 Mouche, A., D. Hauser, J.-F. Daloze, and C. Guérin, Dual-polarization measurements at
431 c-band over the ocean: Results from ariborne radar observations and comparison with
432 envisat asar data, *IEEE Trans. on Geosc. and Remote Sensing*, *43(4)*, 753–769, 2005.
- 433 Mouche, A., B. Chapron, and N. Reul, A simplified asymptotic theory for ocean surface
434 electromagnetic waves scattering, *Submitted to WRM*, 2006a.
- 435 Mouche, A., D. Hauser, and V. Kudryavtsev, Radar scattering of the ocean surface and
436 sea-roughness properties: A combined analysis from dual-polarizations airborne radar
437 observations and models in c band, *J. Geophys. Res.*, *111(C09004)*, 2006b.
- 438 Plant, W., A two-scale model of short wind-generated waves and scatterometry, *J. Geo-*
439 *phys. Res.*, *91(C9)*, 10,735–10,749, 1986.
- 440 Plant, W., A stochastic, multiscale model of microwave backscatter from the ocean, *J.*
441 *Geophys. Res.*, *107(C9)*, 2002.
- 442 Plant, W., Microwave sea return at moderate to high incidence angles, *Waves In Rand.*
443 *Media*, *9*, 339–354, 2003.
- 444 Quilfen, Y., B. Chapron, A. Bentamy, J. Gourrion, T. Elfouhaily, and D. Vandermark,
445 Global ers-1 and 2 nscat observations: Upwind/crosswind and upwind/downwind mea-
446 surements, *J. Geophys. Res.*, *104(C5)*, 11,459–11,469, 1999.
- 447 Romeiser, R., W. Alpers, and V. Wismann, An improved composite surface model for
448 the radar backscattering cross-section of the ocean surface: 1. theory of the model and
449 optimization/validation by scatterometer data, *J. Geophys. Res.*, *102(C11)*, 25,237–

- 450 25,250, 1997.
- 451 Shaw, W., and A. Dougan, Green's function refinement as an approach to radar backscat-
452 ter: General theory and application to lga scattering from the ocean, *IEEE Transaction*
453 *on Antenna Propagation*, *46(1)*, 57–66, 1998.
- 454 Stoffelen, A., and D. Anderson, Scatterometer data interpretation: Estimation and vali-
455 dation of the transfer function cmod4, *J. Geophys. Res.*, *102*, 5767–5780, 1997.
- 456 Thompson, D., Calculation of radar backscatter modulations from internal waves, *J.*
457 *Geophys. Res.*, *93(C10)*, 12,371–12,380, 1988.
- 458 Valenzuela, G., Theories for the interactions of electromagnetic and oceanic waves - a
459 review, *Boundary-Layer Met.*, *13*, 61–85, 1978.
- 460 Voronovich, A., Small-slope approximation for electromagnetic wave scattering at a rough
461 interface of two dielectric half-spaces, *Waves In Rand. Media*, *4*, 337–367, 1994.
- 462 Voronovich, A., and V. Zavarotny, Theoretical model for scattering of radar signals in ku-
463 and c-bands from a rough sea surface with breaking waves, *Waves In Rand. Media*, *11*,
464 247–269, 2001.

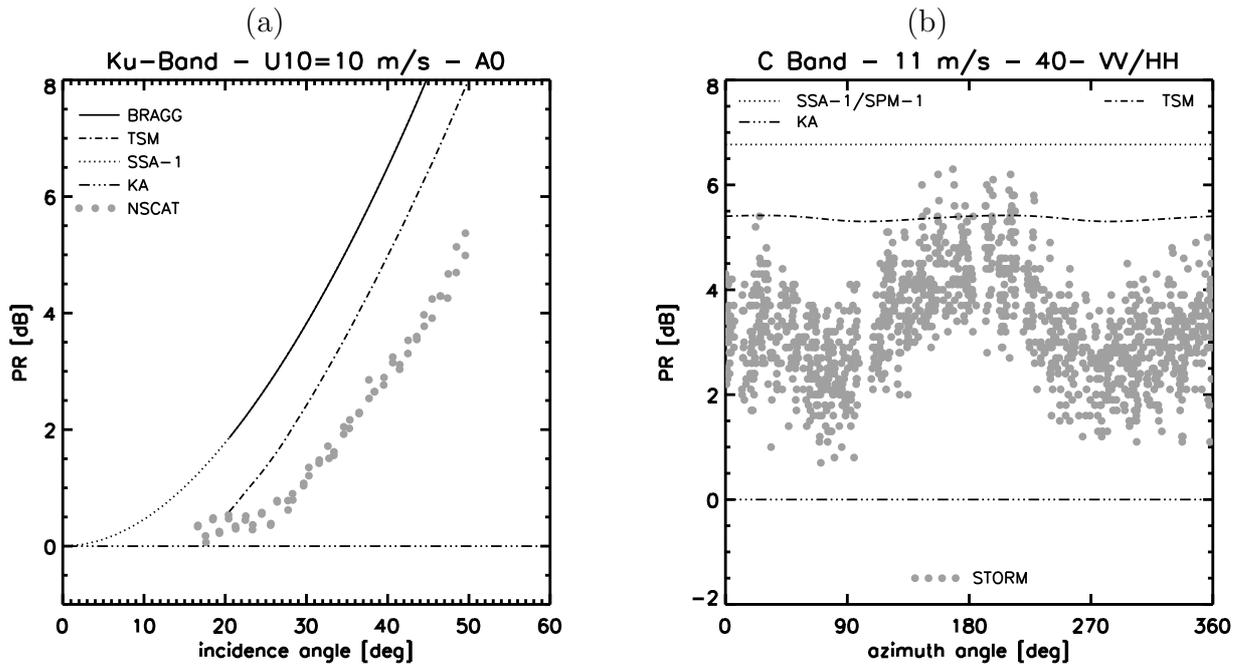


Figure 1. (a) Polarization ratio versus incidence angle in Ku band for a 10 m/s ten meters high wind speed in the case of an isotropic sea surface. (b) Polarization ratio versus wind direction relative to the radar's azimuth look direction for a 11 m/s ten meters high wind speed and a 40° incidence angle.

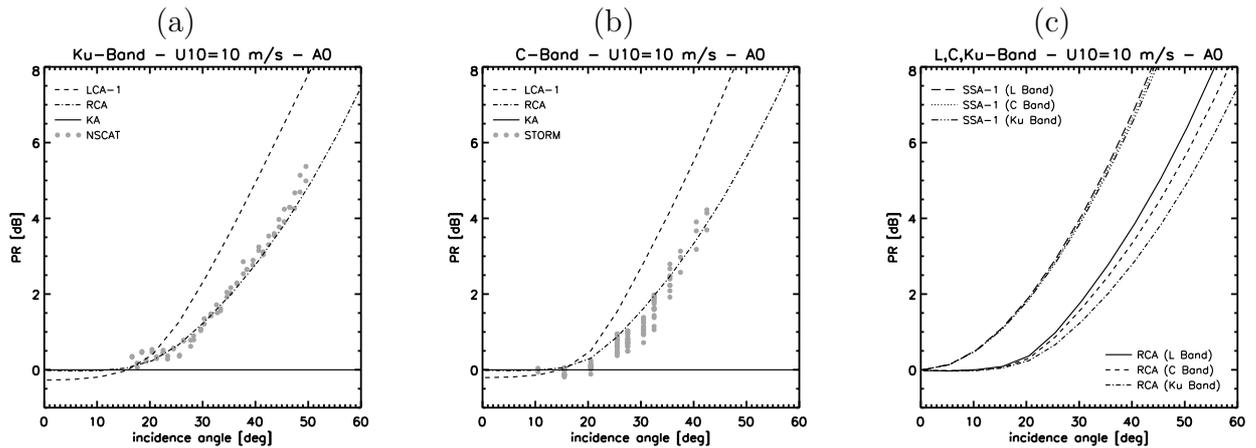


Figure 2. (a) Polarization ratio versus incidence angle in Ku band for a 10 m/s ten meters high wind speed in the case of an isotropic sea surface. Solid, dashed and dashed-dotted lines are respectively the predictions given by KA, LCA and RCA models. Data are from NSCAT. (b) Same than (a) but for C-Band. Data are from STORM radar. (c) Polarization ratio versus incidence angle for a 10 m/s ten meters high wind speed in the case of an isotropic sea surface in Ku, C and L Band given by RCA and SSA-1.

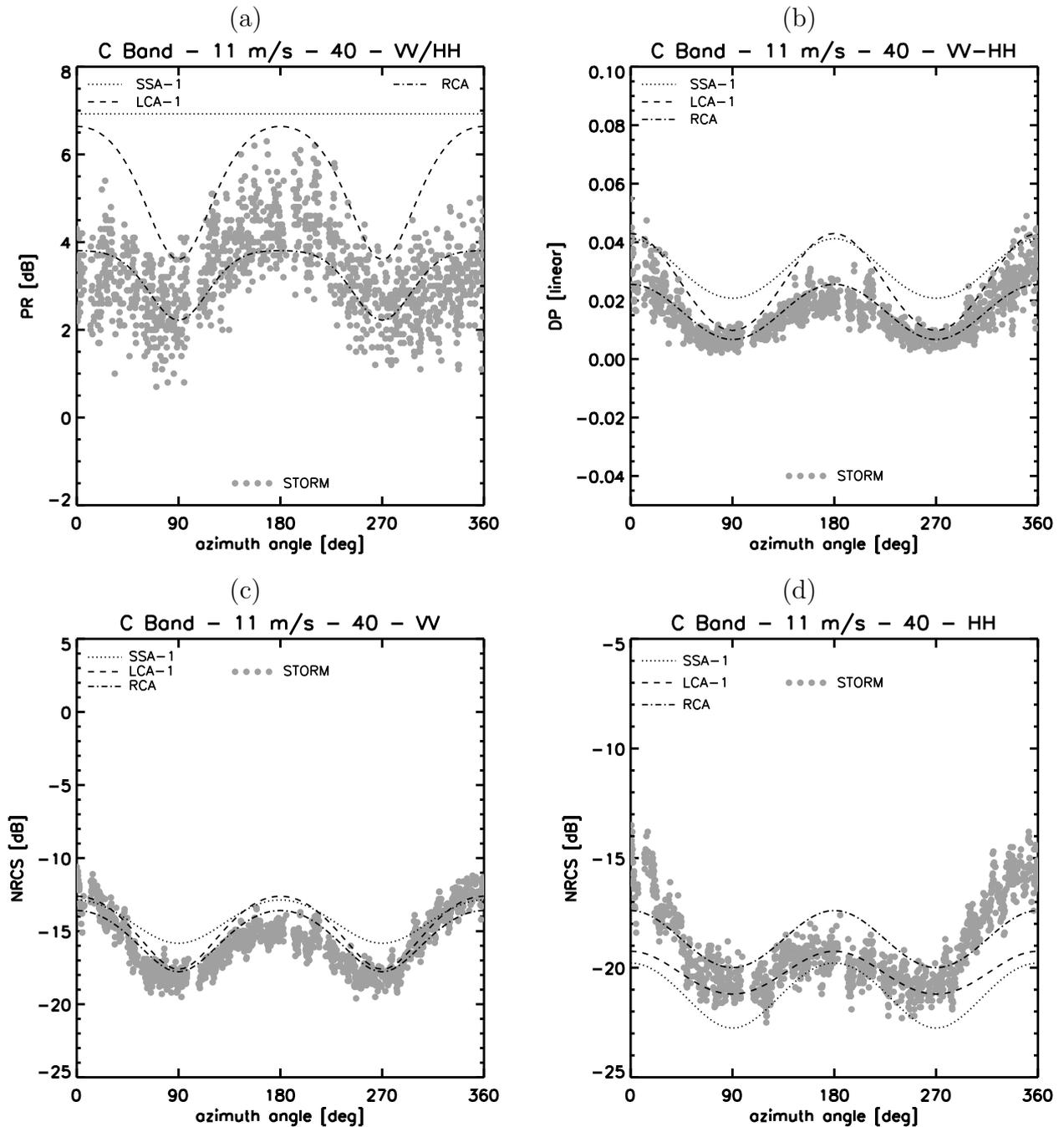


Figure 3. (a) Polarization ratio versus wind direction relative to the radar’s azimuth look direction for a 11 m/s ten meters high wind speed, a 40° incidence angle in C band. (b) Same but for the difference of NRCS. (c) Same but for the NRCS in VV polarization. (d) Same but for the NRCS in HH polarization.

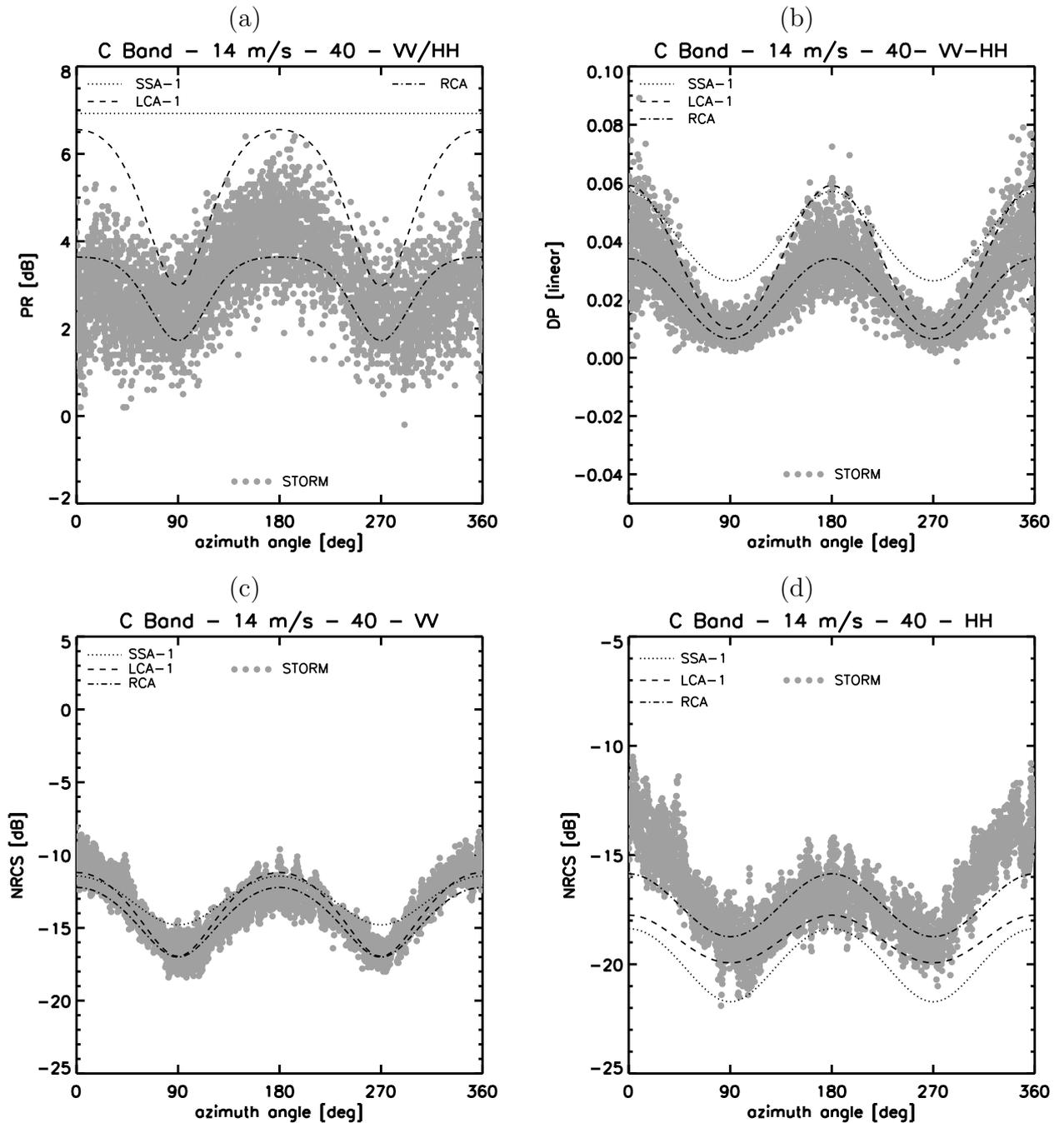


Figure 4. (a) Polarization ratio versus wind direction relative to the radar's azimuth look direction for a 14 m/s ten meters high wind speed, a 40° incidence angle in C band. (b) Same but for the difference of NRCS. (c) Same but for the NRCS in VV polarization. (d) Same but for the NRCS in HH polarization.

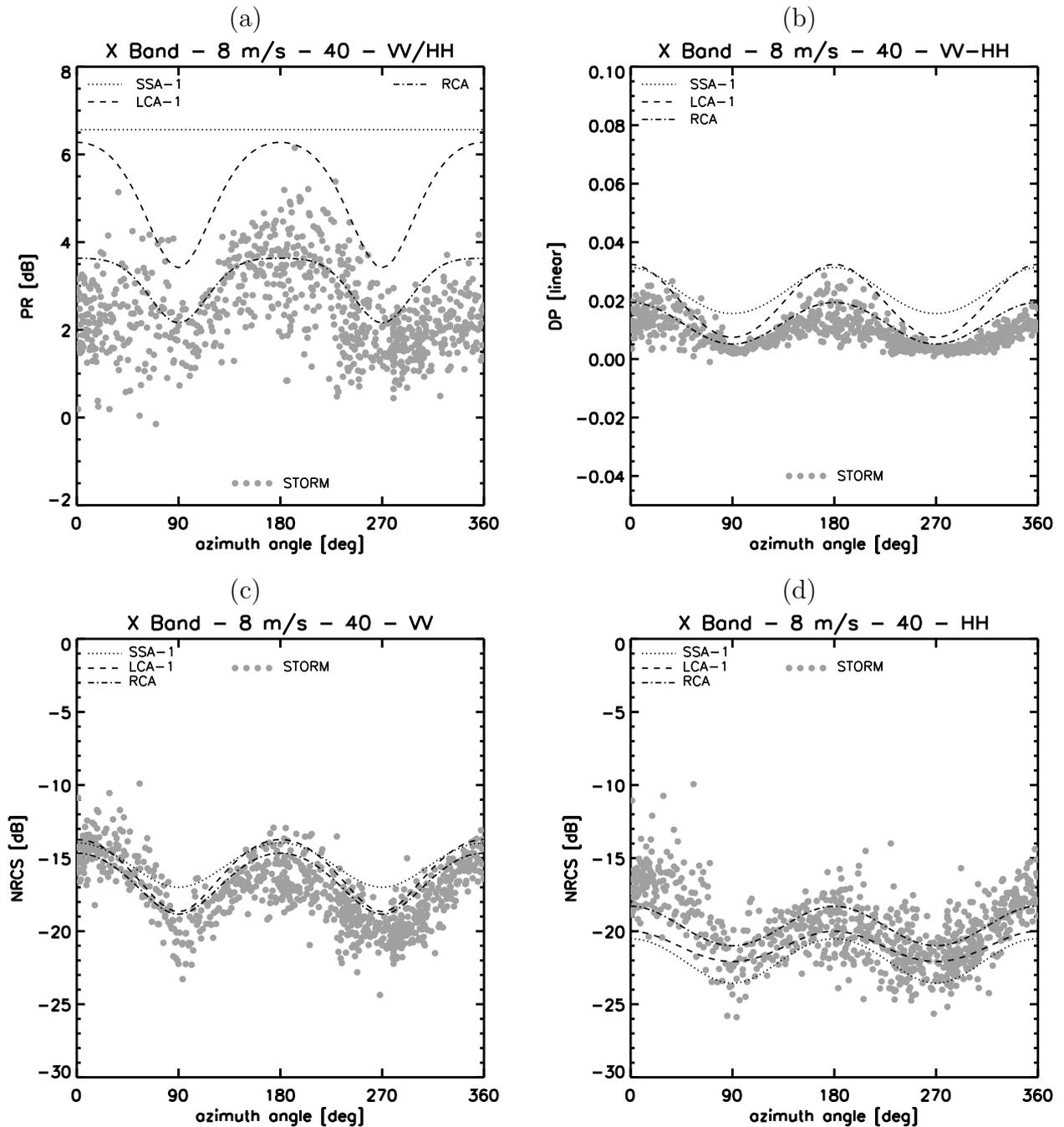


Figure 5. (a) Polarization ratio versus wind direction relative to the radar’s azimuth look direction for a 8 m/s ten meters high wind speed, a 40° incidence angle in X band. (b) Same but for the difference of NRCS. (c) Same but for the NRCS in VV polarization. (d) Same but for the NRCS in HH polarization.

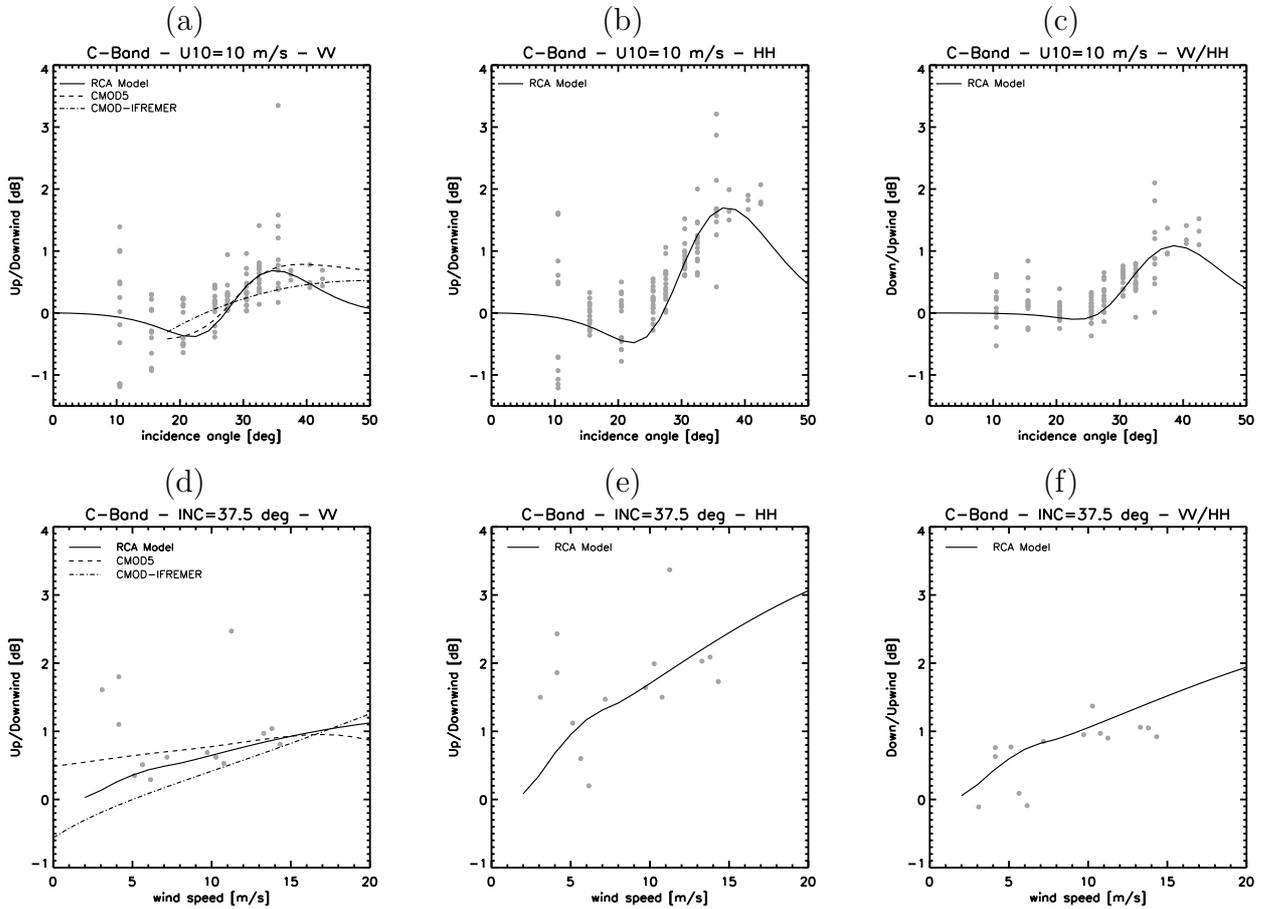


Figure 6. Top panel: UDA versus incidence angle in (a) VV and (b) HH polarizations for a 10m/s ten meters high wind speed. (c) Downwind to upwind asymmetry of the PR versus incidence angle for a 10m/s ten meters high wind speed. Bottom panel: UDA versus incidence angle for (a) VV and (b) HH polarizations for a 37.5° incidence angle. (c) Downwind to upwind asymmetry of the PR versus incidence angle for a 37.5° incidence angle

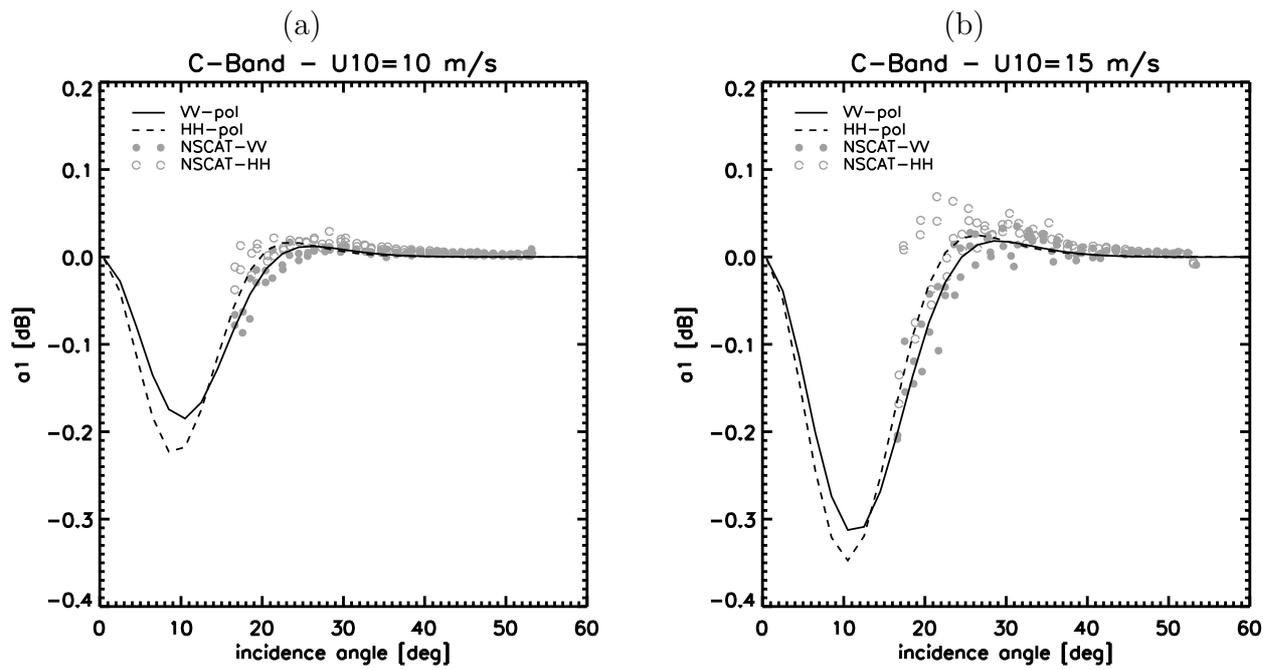


Figure 7. a_1^{vv} and a_1^{hh} coefficients as a function of the incidence angle in Ku-Band for two given ten meters high wind speeds: (a) 10m/s and (b) 10m/s.

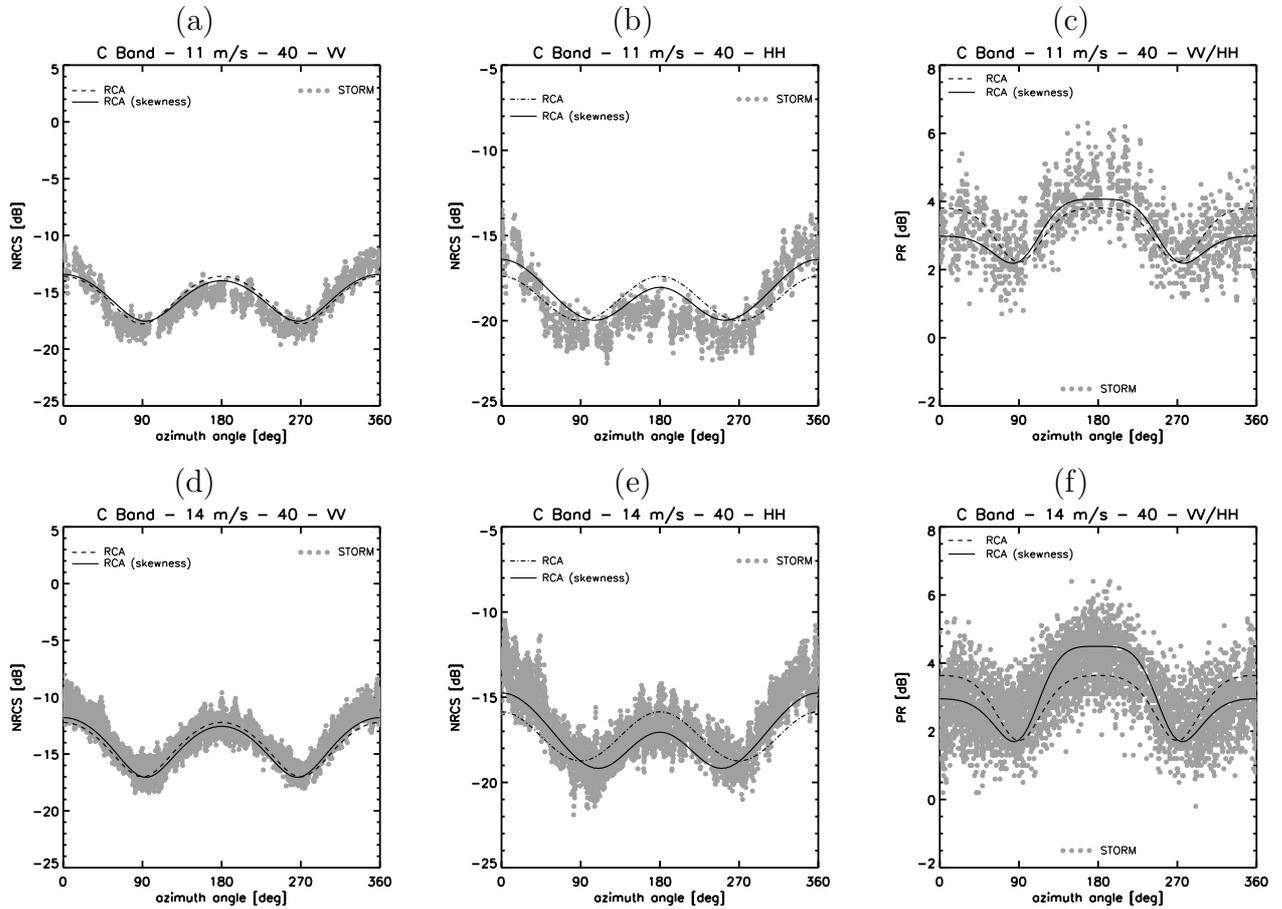


Figure 8. Top panel: NRCS versus wind direction relative to the radar’s azimuth look direction for a 11 m/s ten meters high wind speed, a 40° incidence angle in C band in (a) VV and (b) HH polarizations. (c) Same for the PR. Bottom panel: NRCS versus wind direction relative to the radar’s azimuth look direction for a 14 m/s ten meters high wind speed, a 40° incidence angle in C band in (a) VV and (b) HH polarizations. (c) Same for the PR.