

Reprinted from Melson, W. G., Rabinowitz, P. D., et al., 1978.
Initial Reports of the Deep Sea Drilling Project, Volume XLV, Washington (U.S. Government Printing Office)

22. TRACE ELEMENTS IN BASALTS FROM 23°N and 36°N IN THE ATLANTIC OCEAN: FRACTIONAL CRYSTALLIZATION, PARTIAL MELTING, AND HETEROGENEITY OF THE UPPER MANTLE¹

H. Bougault, Centre Océanologique de Bretagne, Brest, France
M. Treuil, Laboratoire P. Sue, Saclay, France
and

J.L. Joron, Institut de Physique du Globe, Université Paris VI, Paris, France

INTRODUCTION

—From a study of first-transition series trace elements and low-partition-coefficient trace elements or so-called hygromagmatophile elements (Treuil and Varet, 1973; Treuil and Joron, 1976), we will try to define the relative importance of the three fundamental parameters—mantle source composition, extent of partial melting, and crystal fractionation—which bear on the formation of basalts recovered during Leg 45. We relate the presented data to the results of shipboard study (Chapters 7 and 8, this volume) and then give a tentative interpretation, taking into account our present knowledge of the behavior of trace elements during magmatic processes of FAMOUS and Leg 37 results. —

LEG 45 SAMPLES; DESCRIPTION OF RESULTS

Both major element and trace element data are presented in this section. The major oxides SiO₂, Al₂O₃, Fe₂O₃ total, MgO, CaO, K₂O, TiO₂, and the trace elements Sr, Zr, Ni, and Cr were analyzed on-board ship, in collaboration with J. M. Rhodes (see Introduction and Site Chapters, this volume). Analyses for MnO, P₂O₅, and Na₂O were completed on shore. FeO determinations were made by J. Norberg (Smithsonian Institution) according to the method outlined by L. C. Peck (1964).

All major element data presented were determined by X-ray fluorescence, except Na₂O, which was determined by atomic absorption. Trace element data were obtained by X-ray fluorescence (XRF), neutron activation (NA), or atomic absorption (AA). Co, Ni, and Zr were analyzed by both X-ray fluorescence and neutron activation; both results are given.

Trace element XRF determinations were corrected for matrix and instrumental effects by the procedures of Bougault et al. (1977); in addition, V and Cr measurements were corrected for the effects of enhancement and overlapping lines (Ti K β and V K β , respectively).

The shipboard values for Cr are only approximate, because it was not possible on board to calculate the

full matrix correction including the effect of iron enhancement and V K β interference.

The neutron activation method used was pure instrumental activation analysis (without chemical separation), using epithermal neutron irradiation (OSIRIS Reactor in Saclay—C.E.A. Groupe Pierre Sué). Because of the low order of magnitude of the concentrations of investigated elements in tholeiites (Tb, Hf, Ta, . . .), the use of epithermal neutrons is very important; this kind of irradiation allows the interaction of ⁴⁶Sc and ⁵⁵Fe to be strongly diminished. Irradiation was performed under Cd vials; then several measurements were made using a Ge-Li detector (resolution 2 KeV at 1.33 MeV) at different times from four days to one month after irradiation. The reference standards used are GSN and BCR1.

Tables 1 and 2 for Hole 395, Tables 3 and 4 for Hole 395A, and Tables 5 and 6 for Hole 396 show the concentrations of elements (major elements expressed in per cent and trace elements in ppm) for the samples investigated on-board ship.

Several different chemical units were defined on board (see Chapters 7 and 8, this volume), as shown by Figure 1, where concentrations of selected major and trace elements are plotted against depth. Units A₂ through A₅ are aphyric units, and P₂ through P₅ are porphyritic units. Some units were hard to define, because only one sample representing them was analyzed, or a given sample chemically seemed outside the range of analyses characterizing the unit. This is the case for Sample 395A-11-1, 105-107 cm; even though major elements resolve little difference between this sample and Unit A₂^{*}, additional trace element data appear to suggest a real difference between them. In Hole 395A, aphyric units are named P₂ to P₅, as a function of depth; In Hole 395, Samples 18-1, 37-41 cm and 18-2, 33-38 cm (Unit P₁^{*}) are chemically similar to Unit P₃ of Hole 395A.

Sample 395-20-1, 32-36 cm, within Unit P₂^{*}, nevertheless does not belong to that unit; it stands alone. Sample 395A-15-5, 0-11 cm, situated at the boundary of Units P₂ and P₃, remains difficult to classify; it is chemically similar to Unit P₅. The average composition given in Table 7 for Unit P₅ does not take into account breccia samples (which differ mainly in Cr and Ni). The deepest aphyric sample, 395A-67-2, 54-59 cm, is classified as belonging to Unit A₄; in fact, it is very similar to samples of Unit A₃.

¹Contribution 528 of the Département Scientifique, Centre Océanologique de Bretagne.

TABLE 1
Site 395, Hole 395, Major-Element Data (volatile-free concentrations, wt. %)

Unit Sample (Interval in cm)	10, CC (#1)	10, CC (#2)	10, CC (#3)	A2*				A2*				P1*		P1*		A2*		P2*
				11-1, 105-107	11-2, 62-64	12-2, 109-111	14-1, 112-114	14-1, 131-132	15-1, 71-73	15-1, 112-115	15-2, 130-133	17-1, 56-59	18-1, 37-41	18-1, 61-70	18-2, 33-38	18-2, 85-95	19-1, 18-20	20-1, 32-36
Depth (m)	93	93	93	101	102	112	131	131	139	139	140	159	166	167	168	169	177	185
Rock Type	s	G	s	VB	aphy. B	aphy. B	al B	aphy. B	aphy. B	aphy. B	aphy. B	G	pl. Pyr B	s	pl. pyr. B	s	aphy. B	pl. B
SiO ₂	43.6	45.7	44.8	48.9	49.3	49.8	47.01	49.7	49.81	49.1	49.0	50.2	49.4	43.1	49.4	43.4	49.8	49.6
Al ₂ O ₃	1.26	24.24	6.27	15.93	14.88	15.11	16.79	14.94	15.11	15.01	14.75	17.71	17.32	0.91	18.21	1.37	15.08	18.77
Fe ₂ O ₃ (t)	9.95	3.81	10.66	12.83	11.95	11.90	13.76	11.97	11.76	12.05	12.22	6.52	8.62	10.25	8.37	9.36	11.05	9.52
FeO	0.93		4.35	6.96		8.10		7.53				4.62	5.67	4.55		3.10	7.41	4.47
MnO	0.13	0.07	0.10	0.20	0.18	0.18	0.20	0.17	0.17	0.17	0.18	0.11	0.14	0.11	0.14	0.10	0.19	0.14
MgO	43.7	13.9	35.3	6.8	8.5	8.9	6.64	8.3	7.47	8.6	8.5	12.2	8.5	44.3	7.9	42.0	8.5	6.1
CaO	0.03	12.61	1.85	11.01	10.49	10.50	10.93	10.56	10.92	10.53	10.54	9.32	12.58	0.89	12.85	2.09	11.25	12.05
Na ₂ O	0.00	1.15	0.00	2.88	2.83	2.85	2.90	2.88	2.85	2.88	2.85	3.53	2.38	0.03	2.33	0.00	2.75	2.79
K ₂ O	0.02	0.08	0.02	0.22	0.09	0.09	0.28	0.13	0.25	0.10	0.13	0.20	0.07	0.00	0.13	0.00	0.31	0.30
TiO ₂	0.03	0.04	0.09	1.70	1.61	1.61	1.83	1.63	1.61	1.62	1.62	0.39	1.01	0.03	1.02	0.04	1.65	1.28
P ₂ O ₅	0.01	0.02	0.02	0.18	0.15	0.13	0.22	0.15	0.14	0.15	0.15	0.02	0.10	0.02	0.12	0.00	0.16	0.13
Total	98.73	101.62	99.11	100.65	99.98	101.07	100.36	100.43	100.09	100.21	99.94	100.20	100.12	99.64	100.47	98.36	100.84	100.68
LoI	-13.8	-4.2	-11.6	-0.9	-1.1	-1.3	-8.7	-1.2	-9.5	-1.5	-0.8	-4.9	-1.4	-7.6	-0.6	-9.1	-1.0	-1.5

Note: G = gabbro, s = serpentinite, V = variolitic, B = basalt, aphy. = aphyric, pl = plagioclase, al = altered.

In Hole 396, three units have been recognized, named P_a, P_b, and P_c, as a function of depth and are listed in Tables 5 and 6.

All the trace element results obtained confirm the classification and the definitions of chemical units made on-board ship. Since we bring no change or very little change to the definition of chemical units, we ask the reader to examine Chapters 7 and 8 in this volume for detailed information about particular units.

The alkali metals, Cs and Rb, together with Sr, show variations even within a given unit. Average values of other elements, which are more homogeneous within a given unit, are reported for each unit, together with associated σ (standard deviation) values, in Tables 7, 8, 9, and 10 for Sites 395 and 396, respectively. Examination of Tables 7 and 10 shows that concentrations of all elements investigated are typical of mid-ocean ridge tholeiites. Nevertheless, Th and Ta values are rather low, and some Sr values are high, especially in phyric units P₂* of Hole 395, P₂ and P₅ of Hole 395A. This problem of high Sr is still being investigated to determine if what we are seeing is enriched Sr because of plagioclase phenocrysts or is a result of alteration effects (since many vesicles filled with secondary minerals are present in these samples). Generally, trace element concentrations, except for Ni and Cr, are lower within phyric units than in aphyric units; this can be explained in part by the dilution effect of phenocrysts (mainly plagioclase), since many of these elements have low partition coefficients and low abundances in phenocryst phases, especially plagioclase.

Sc and Co vary within a narrow range, 30 to 38 ppm and 38 to 48 ppm, respectively, and show only a slight difference between aphyric and phyric units, as mentioned above. Taking into account the dilution effect of phenocrysts, one would obtain roughly similar concentrations for the matrix of phyric units as for aphyric units. Finding so little variation for those elements means that their behavior depends little on crystallization and melting processes; in addition, this means that the initial material (before melting) was homogeneous with respect to them. If the source had been heterogeneous, then melting and crystallization

would have had to occur in such a way as to compensate for the possible variations in the initial material, and this would be purely matter of chance.

Average values of Ni and Cr for Holes 395 and 395A are reported in Table 7, except for Units P₄ and P₄' ; Ni and Cr are more variable in these units, because of a variable content of olivine and clinopyroxene phenocrysts.

ALTERATION EFFECTS

Among the trace elements investigated, three can vary considerably because of alteration: Cs, Rb, and Sr (Hart, 1971). The problem of whether strontium is enriched in plagioclase or in the matrix, or is high because of alteration, is being investigated, as already mentioned.

As Figure 1 shows, the abundance of K₂O is somewhat random even within a given unit. Similar observations can be made for Cs and Rb. Cs is plotted versus K₂O in Figure 2, and versus Rb in Figure 3. Considering all points on Figure 2, no correlation between Cs and K₂O is evident. But within each unit, a positive correlation exists between Cs and K₂O, even taking into account the analytical precision (about 0.02 ppm for Cs and 0.2% for K₂O). Cs increases linearly with K₂O from low concentrations which could be considered as initial values before secondary alteration processes. In Figure 2 a general positive correlation between Rb and Cs can be observed, but there is no possibility of distinguishing between units, because of the analytical precision in determining Rb (~1 ppm).

The general correlation Cs = f(Rb) and the intra-unit correlation of Cs = f(K₂O) contrast with the homogeneity of other trace elements within units, allowing us to state that *these variations result from secondary processes (probably alteration)*.

GEOCHEMICAL METHOD OF INVESTIGATION

Properties of Trace Elements

It is assumed that trace element behavior follows the high-dilution law; as a consequence, a partition coef-

ficient between different phases can be defined directly in connection with the chemical potential of a given element in the different phases concerned. As a convention, the partition coefficient D of an element is defined as the ratio of its concentration in a mineral to its concentration in the liquid.

Among the investigated elements, one can find high-partition-coefficient (HPC) elements such as Cr and Ni ($D > 1$) and low-partition-coefficient (LPC) elements such as Th and Ta ($D \ll 1$). The bulk partition coefficients of some elements are variable, depending on the proportion of phases in equilibrium, and in this respect, their concentrations depend on the nature of magmatic processes, such as partial melting or fractional crystallization. This is the case, for instance, for V, Co, and Sr. The diverse partition coefficients are interpreted to reflect high stabilization energy in octahedral coordination for such elements as Cr and Ni (Burns, 1970; Curtis, 1964; Orgel, 1964; Allegre et al., 1968; Bougault and Hébinian, 1974, and the ability of other elements—Ta and Th, for instance—to form stable complexes in the liquid phase (Treuil and Varet, 1973; Treuil and Joron, 1975).

Behavior of Trace Elements During Crystallization and Melting Processes

As crystallization proceeds in a given magma, we assume that that Rayleigh law is valid:

$$\frac{C_L}{C_{LO}} = F^{(D-1)}$$

where

- C_L = concentration of the element in the liquid,
- C_{LO} = initial concentration of the element in the liquid,
- F = fraction of remaining liquid,
- D = bulk partition coefficient of the element.

Considering two elements (subscripts 1 and 2), the relationship between elements 1 and 2 is:

$$\log C_{L2} = \frac{D_2-1}{D_1-1} \log C_{L1} + \frac{(D_1-1) \log C_{LO2} - (D_2-1) \log C_{LO1}}{D_1-1} \quad (\text{Eq. 1})$$

1) If elements 2 and 1 have low partition coefficients, Equation 1 is simplified to:

$$C_{L2} = A C_{L1} \quad A = \frac{C_{LO1}}{C_{LO2}} \quad (\text{Eq. 2})$$

The slope A depends on the initial concentrations in the liquid C_{LO1} and C_{LO2} (Treuil and Varet, 1973; Treuil and Joron, 1975).

2) If elements 2 and 1 have high partition coefficients and if the proportion of phases crystallizing is

constant, the plot of $\log C_{L2}$ as a function of $\log C_{L1}$ is a straight line, the slope of which is

$$\frac{D_2-1}{D_1-1}$$

3) If element 2 is a HPC element and element 1 a LPC element, the slope of the line $\log C_{L2} = f(\log C_{L1})$ is approximated to (D_2-1) ; this can be considered as a way of determining the bulk partition coefficient D_2 of element 2.

In fact, these assumptions are valid insofar as we start from given values of C_{LO2} and C_{LO1} , which means starting from a given liquid with magmatic differentiation resulting only from crystallization processes. C_{LO2} and C_{LO1} depend on partial melting; it is necessary to determine to what extent they can vary (if at all) when partial melting occurs, using the elements investigated.

When partial melting occurs, let us assume that C_{LO} is given by the formula of Shaw (1970):

$$\frac{C_{LO}}{C_{SO}} = \frac{1}{D_O + F(1-D)}$$

where

- C_{LO} = concentration of the element in the liquid (which is the initial value with respect to the crystallization process),
- C_{SO} = initial concentration in the solid,
- D_O = initial bulk partition coefficient,
- D = bulk partition coefficient corresponding to melting phases,
- F = fraction of liquid produced.

Considering very low-partition-coefficient elements 1 and 2, it is easy to see that

$$\frac{C_{LO1}}{C_{LO2}} \approx \frac{C_{SO1}}{C_{SO2}}, \text{ almost independent of } F.$$

In this case, the slope A mentioned in Equation 2 is not dependent on F , and gives direct information about the initial solid (before melting) (Treuil and Joron, 1976),

$$A = \frac{C_{LO1}}{C_{LO2}} = \frac{C_{SO1}}{C_{SO2}}$$

If elements 1 and 2 are LPC elements but have partition coefficients with different orders of magnitudes, A depends on F ; liquids with different values of C_{LO1}/C_{LO2} can be produced by different partial melting. Basalts experiencing further crystal fractionation derived from these different liquids plot on lines with different slopes A (Treuil and Varet, 1973; Treuil and Joron, 1975).

Considering HPC elements, it can be shown that for Ni the ratio C_{LO}/C_{SO} is almost independent of the extent of partial melting (F) and of the composition of

TABLE 2
Site 395, Hole 395, Trace-Element Data (expressed in ppm)

Element		Sc	Ti	V	Cr	Mn	Fe	Co	Ni	Cu	Zn		
Analytical Method	Unit	NA	FX	FX	FX	FX	FX	FX	NA	FX	NA	AA	AA
Sample (Interval in cm)													
10, CC #1		8.25	180	49	2805	1007	69650	111	99.6	2160	1716	12	45
10, CC #2		1.92	240	3.6	323	542	26670	39	41.6	554	563	50	17
10, CC #3			540	34	1071	774	74620	104		1128		91	39
11-1, 105-107	A2*	38.4	10200	301	279	1548	80810	48	49.3	161	160	69	92
11-2, 62-64		36.6	9660	263.7	271	1394	83650	48	40.4	173	190	68	81
12-2, 109-111		37.3	9660	266	269	1394	83300	47	49.6	172	189	68	85
14-1, 112-114		40.8	10980	323	274	1548	96320	49	48.8		183	65	101
11-1, 131-132	A2*	37.3	9780	269	269	1316	83790	49	49.9	177	186	70	87
15-1, 71-73		38	9660	274	279	1316	82320	51	49.9		181	64	84
15-2, 130-133		37.9	9720	266	265	1316	85540	48	50.4	182	186	68	89
16-2, 104-105		36.9	9780	270	267	1394	85820	49	48.9	176	183	69	87
17-1, 50-69		35.2	2340	153	165	852	45640	40	43	155	170	0	19
18-1, 37-41	P1*	31.3	6060	217	361	1084	60340	38	41.3	139	152	68	53
18-1, 61-70		6.58	180	29	2677	852	71750	127	123	2370	2543	11	44
18-2, 33-38	P1*	30	6120	208	341	1084	58590	37	37.9	130	137	60	57
18-2, 85-95			240	52	4788	774	65520	116		2110		33	42
19-1, 18-20	A2*	38	9900	200	293	1472	77350	41	44	141	137	61	79
20-1, 32-36	P2*	30.6	7680	219	222	1084	66640	34	35	186	90	55	58

the presumed initial material. For Cr, the ratio C_{LO}/C_{SO} is somewhat dependent on F, increasing when partial melting increases, and is dependent on the composition of the presumed initial material (spinel content, for instance).

DISCUSSION AND INTERPRETATION OF RESULTS

From the shipboard data, we see that Ti and Zr behave similarly (Figure 4). It has been already shown that in mid-ocean ridge basalts, Ti behaves as an LPC trace element (Bougault, 1977). This is confirmed by the correlation shown on Figure 4. Leg 37 samples from Hole 332B plot on the same line as Leg 45 samples. The only difference between Leg 37 and Leg 45 samples is a shift along the line, both for phyruc samples on one hand and aphyric samples on the other. This can be attributed either to a difference of range of the extent of partial melting or to a different initial mantle material. In the remainder of this section, we shall examine to what extent a difference can be determined for the extent of partial melting and/or the composition of initial mantle material between 23°N (Leg 45) and 36°N (FAMOUS area).

The hygromagmatophile elements with low partition coefficients—Ta, Hf, Tb, Zr, and Ti—have been plotted on Figures 5 through 9 as a function of Th. Th is thought to have the lowest partition coefficient among these elements. On the same figures, the field where the FAMOUS samples plot is shown and the dispersion indicated. Both Leg 45 and FAMOUS samples plot on straight lines passing through the origin, with dispersions more or less significant. For Leg 45 sam-

ples straight lines can be drawn because in fact each point represents an average value of many samples analyzed in the same unit.

The first striking feature is the different slopes between Leg 45 samples on the one hand and FAMOUS samples on the other. Leg 37 data are not included in the FAMOUS range, because they were determined by other laboratories and there may be a problem of reproducibility of trace elements among them. Even so, Leg 37 data for Th and Ta from other laboratories fall in the field of FAMOUS samples and not at all in the field of Leg 45 samples.

From the theoretical discussion presented earlier, *this has to be attributed to a difference between FAMOUS area and Leg 45 initial mantle materials.* Samples from Hole 396, which is roughly symmetric to Site 395 on the other (eastern) side of the Mid-Atlantic Ridge, plot very near the line defined for Site 395.

Of all the fields of FAMOUS samples presented on Figures 5 to 9, Ta versus Th (Figure 5) has the lowest dispersion. This is because Ta and Th have the lowest partition coefficients among the studied elements and, from our earlier discussion, the slope of the line relating these elements is not dependent on partial melting. The other elements plotted versus Th are LPC elements, but their partition coefficients are higher than those of Th and Ta. The slopes corresponding to given series of crystallization differentiation depend on the initial concentration in the liquid, that is to say on the extent of partial melting. This is why the dispersion for these elements is higher around "average lines."

Table 11A presents the values of slopes of a given LPC element i , plotted against other LPC elements j ,

TABLE 2 - Continued

Rb	Sr		Zr	Sb		Cs	La	Eu	Tb	Hf	Ta	Th
NA	FX	FX	NA	NA	NA	NA	NA	NA	NA	NA	NA	NA
	6	3	<27	0.189	<0.027	<0.84	≤0.057	<0.014	<0.14	<0.008		0.05
	99	8	<14		<0.015	0.11	0.15	0.008	<0.067	<0.004		0.015
	4	4										
3.3	127	121	112	0.50	0.099	3.25	1.5	0.87	2.84	0.200		0.137
<1.4	119	112	126	0.020	0.024	3.2	1.33	0.82	2.71	0.183		0.132
<0.9	120	119	120	0.016	0.037	2.12	1.42	0.83	2.73	0.194		0.137
4.1			137	0.98	0.027	3.6	1.49	0.93	3.18	0.207		0.141
2.4	120	114	136	0.017	0.08	3.1	1.36	0.84	2.86	0.193		0.134
<1.1			108	0.018	0.05	4.1	1.40	0.83	2.83	0.192		0.138
1.9	118	116	112	0.015	0.020	3.4	1.45	0.85	2.88	0.187		0.130
<1.8	118	120	108	0.033	0.036	3.3	1.40	0.83	2.83	0.190		0.130
3.8	324	16	<27	0.017	0.34	0.44	0.41	0.193	0.19	<0.057		<0.019
2.1	112	67	63		0.038	2.04	0.95	0.52	1.65	0.103		0.079
<1.8	6	4	<25	0.018	<0.02	0.12	<0.05	0.011	<0.13	<0.007		<0.027
2.3	114	67	68	0.033	0.066	1.9	0.97	0.54	1.66	0.109		0.083
	10		1									
<1.5	130	127	131	0.026	0.040	3.8	1.42	0.84	3.00	0.201		0.154
8.1	162	94	98	0.101	0.40	2.77	1.18	0.64	2.21	0.149		0.113

both for FAMOUS and Leg 45 samples. The major differences between Leg 45 and FAMOUS samples are shown by comparing slopes (i versus j) when i or j corresponds to Ta or Th (the lowest partition coefficient elements). Table 11B also shows absolute values of LPC elements for both Leg 45 and FAMOUS aphyric basalts (Unit A₃ and Sample ARP 7-7, respectively) and plagioclase phyrlic basalts (Unit P₂ and Sample ARP 31-36, respectively). One can observe that for both types of basalts, Th and Ta values are lower, and other LPC elements are higher for Leg 45 samples than they are for FAMOUS samples. *This confirms that there is a difference between the initial mantle compositions of these areas*; whatever the partial melting or crystallization processes involved, it would not be possible after partial melting and crystallization to have some LPC elements lower and others higher if the mantle had the same initial composition in both cases.

For Leg 45 samples, it is not possible to determine differences in the extent of partial melting on the basis of LPC elements, because of the steep slopes of the lines and the analytical precision that would be required. Nevertheless, this can partly be shown through Cr and Ni data.

Log Cr versus log Ni is plotted on Figure 10. Leg 37 and FAMOUS samples define two common fields: one for aphyric samples, the other for plagioclase-phyric samples. Since the Ni concentration in the liquid is almost independent of partial melting and the Cr concentration is somewhat dependent, these trends are mainly representative of crystallization history. The FAMOUS aphyric field is narrow, but cannot be resolved to a straight line because of variations in

the proportions of crystallizing minerals. The FAMOUS plagioclase-phyric field is above and larger than the aphyric field, probably because of the variable content of clinopyroxene phenocrysts (the partition coefficient of Cr in clinopyroxenes is higher than 10). But the clinopyroxene content of FAMOUS plagioclase-phyric basalts is not at all sufficient to account for their high Cr concentrations relative to FAMOUS aphyric basalts for a given value of Ni. This is because the Cr concentration in the liquid increases with partial melting, as discussed theoretically earlier. In this respect, plagioclase-phyric basalts appear to derive from liquids produced by a higher degree of partial melting than aphyric basalts. This is in agreement with LPC element abundances, which are generally higher for plagioclase-phyric basalts than for aphyric basalts.

On Figure 10, both the fields of plagioclase-phyric basalts of Site 395 and 396 and of aphyric basalts, except A₃ and A₄ Sample 67-2, 54-59 cm, are shifted compared with FAMOUS and Leg 37 data. The different relative positions of the various Leg 45 aphyric basalts and the plagioclase phyrlic basalts is a consequence of different degrees of partial melting and of different proportions of crystallizing minerals. A plot of Ni or Cr versus an LPC element would result in a similar conclusion about partial melting; at least two and probably three different degrees of partial melting are represented.

It is somewhat difficult to draw conclusions about crystallization processes with Leg 45 samples, because of the low number of different rock types we have, compared with FAMOUS-Leg 37 data, and because of the different degrees of partial melting. Nevertheless, it

TABLE 5
Site 396, Hole 396, Major-Element Data (volatile-free concentrations, wt. %)

Unit	Pa					Pb				Pc			
	14-6, 20-45	15-4, 48-53	16-3, 137-142	18-1, 130-136	19-2, 138-141	22-4, 144-150	23-1, 143-147	24-3, 73-78	25-1, 96-100				
Sample (Interval in cm)													
Depth (m)	125	135	144	160	173	201	209	219	221				
Rock Type	pl. ol. B	pl. ol. B	pl. ol. B	pl. ol. B	pl. ol. B	pl. ol. B	pl. ol. B	pl. ol. B	pl. ol. B				
SiO ₂	49.3	49.5	49.8	49.7	49.7	49.8	49.2	49.1	49.7				
Al ₂ O ₃	16.25	16.75	16.73	16.93	16.70	16.93	16.68	16.58	16.71				
Fe ₂ O ₃ (t)	9.96	9.17	9.24	9.68	9.23	9.47	9.29	9.68	9.49				
FeO	5.36	4.36	4.45			4.64		5.23					
MnO	0.16	0.15	0.15	0.17	0.17	0.15	0.15	0.16	0.15				
MgO	7	8.5	8.1	7.9	7.7	7.6	7.3	7.9	7.4				
CaO	11.57	11.88	11.95	12.09	11.98	12.11	11.96	12.07	11.99				
Na ₂ O	2.62	2.59	2.57	2.62	2.73	2.69	2.72	2.65	2.65				
K ₂ O	0.31	0.21	0.24	0.29	0.21	0.25	0.21	0.23	0.13				
TiO ₂	1.50	1.24	1.27	1.27	1.29	1.29	1.29	1.29	1.29				
P ₂ O ₅	0.15	0.12	0.12	0.13	0.13	0.11	0.11	0.12	0.11				
Total	98.82	100.11	100.17	100.78	99.84	100.30	98.91	99.78	99.62				
Lol	-2.0	-3.1	-2.4	-2.6	-3.0	-2.3	-2.5	-2.0	-2.2				

Note: pl = plagioclase, ol = olivine, B = basalt.

From low-partition-coefficient hygromagmatophile elements, it can be stated, for Sites 395 and 396, that the initial mantle materials are similar, if not exactly the same. The initial mantle material at 23°N (Leg 45) is different from the initial mantle material at 36°N (FAMOUS area). This is the most important result of this study, which demonstrates mantle heterogeneity.

REFERENCES

Allegre, C. J., Javoy, M., and Michard, G., 1968. Etude de l'abondance des éléments de transition dans l'écorce terrestre, comparée à celle des terres rares. In Ahrens, L. H. (Ed.), *Origine et Distribution des Eléments*. London. (Pergamon Press), p. 914-928.

Bougault, H., 1977. First transition series elements fractional crystallization and partial melting. In Aumento, F., Melson, W. G., et al., *Initial Reports of the Deep Sea Drilling Project*, v. 37, Washington (U.S. Government Printing Office), p. 539-546.

Bougault, H. and Hekinian, R., 1974. Rift valley in the Atlantic Ocean near 36°50'N: petrology and geochemistry of basaltic rocks, *Earth Planet. Sci. Lett.*, v. 24, p. 249-261.

Bougault, H., Cambon, P., and Toulhoat, H. (1977). X-ray spectrometric analysis of trace elements in rocks; correction for instrumental interferences, *X-Ray Spectrometry*, v. 6, no. 2, p. 66-72.

Burns, R. G., 1970. *Mineralogical applications of crystal field theory*. Cambridge. (Cambridge Earth Sciences Series).

Curtis, C. D., 1964. Application of the crystal field theory: the inclusion of trace transition elements in minerals during magmatic differentiation, *Geochim. Cosmochim. Acta*, v. 28, p. 389-403.

Hart, S. R., 1971. Rb, Cs, Sr, and Ba contents, and Sr isotopic values of ocean floor basalts, *Phil. Trans. Roy. Soc. London*, ser. A., v. 268, p. 573-587.

Orgel, L. E., 1964. *Chimie des éléments de transition*. Dunod Ed., Paris.

Peck, L. C., 1964. Systematic analysis of silicates, *Geological Survey Bulletin 1170*. Washington (U.S. Government Printing Office).

Shaw, D. M., 1970. Trace elements fractionation during anatexis, *Geochim. Cosmochim. Acta*, v. 34, p. 237-243.

Treuil, M. and Joron, M., 1975. Utilisation des éléments hygromagmatophiles pour la simplification de la modélisation quantitative des processus magmatiques: exemple de l'AFAR et de la dorsale medio-atlantique. *Societa Italiana Mineralogia et Petrologia*, v. 31, p. 125-174.

_____, 1976. Etude géochimique des éléments en trace dans le magmatisme de l'Afar. Implication petrogénétique et comparaison avec le magmatisme de l'Islande et de la dorsale medio-atlantique. Inter-union Commission on Geodynamics: *Scientific Report 16*. Proceedings of an international symposium on Afar region and related rift problems, Bad Bergraben F. R. Germany, April 1-6 1974.

Treuil, M. and Varet, J., 1973. Critères volcanologiques pétrologiques et géochimiques de la genèse et de la différenciation des magmas basaltiques: exemple de l'AFAR, *Bull. Soc. Geol. France*, v. 15, p. 506-540.

TABLE 6
Site 396, Hole 396, Trace-Element Data (expressed in ppm)

Element	Sc	Ti	V	Cr	Mn	Fe	Co	Ni	Cu	Zn	Rb	Sr	Zr	Sb	Cs	La	Eu	Tb	Hf	Ta	Th			
Analytical Method	NA	FX	FX	FX	FX	FX	FX	NA	FX	NA	AA	AA	NA	FX	FX	NA	NA	NA	NA	NA	NA	NA		
Sample (Interval in cm)																								
Unit P _a 14-6, 20-25	36.8	9000	312	288		69720	40	40.7	111	114	58	73	4.7	148	109	118	0.021	0.260	3.5	1.38	0.79	2.64	0.207	0.161
Unit P _b 15-4, 48-53		7440	256	415		64190	41		171		65	66		155	86									
16-3, 137-142	34.8	7620	268	397		64680	40	40.8	143	142	64	67	4.4	155	90	90	0.020	0.250	3.0	1.14	0.64	2.14	0.167	1.22
18-1, 130-136		7620	262	371		67760	40		142		67	66		146	87									
19-2, 138-141	35.4	7740	269	391		64610	39	41.4	142	144	67	65	3.9	168	86	83	0.025	0.255	2.9	1.23	0.66	2.24	0.173	0.132
21-1, 87-89		7140	261	431		62650	41		189		67	65		150	85									
Unit P _c 22-4, 144-150	34.5	7740	254	284		66290	40	41.5	123	121	65	70	5.1	135	86	103	0.014	0.289	2.65	1.23	0.67	2.16	0.153	0.116
23-1, 143-147	34.2	7740	255	208		65030	41	42.4	128	125	65	72	4.2	133	93	89	0.011	0.228	2.7	1.19	0.67	2.16	0.152	0.096
24-3, 73-78	35.1	7740	254	293		67760	41	42.8	124	1340	66	73	4.3	128	86	90	0.025	0.234	2.7	1.26	0.68	2.31	0.156	0.116
25-1, 96-100	34.9	7740	259	288		66430	40	40.9	109	116	64	71	3.0	130	81	107	0.016	0.056	2.7	1.20	0.68	2.18	0.153	0.126

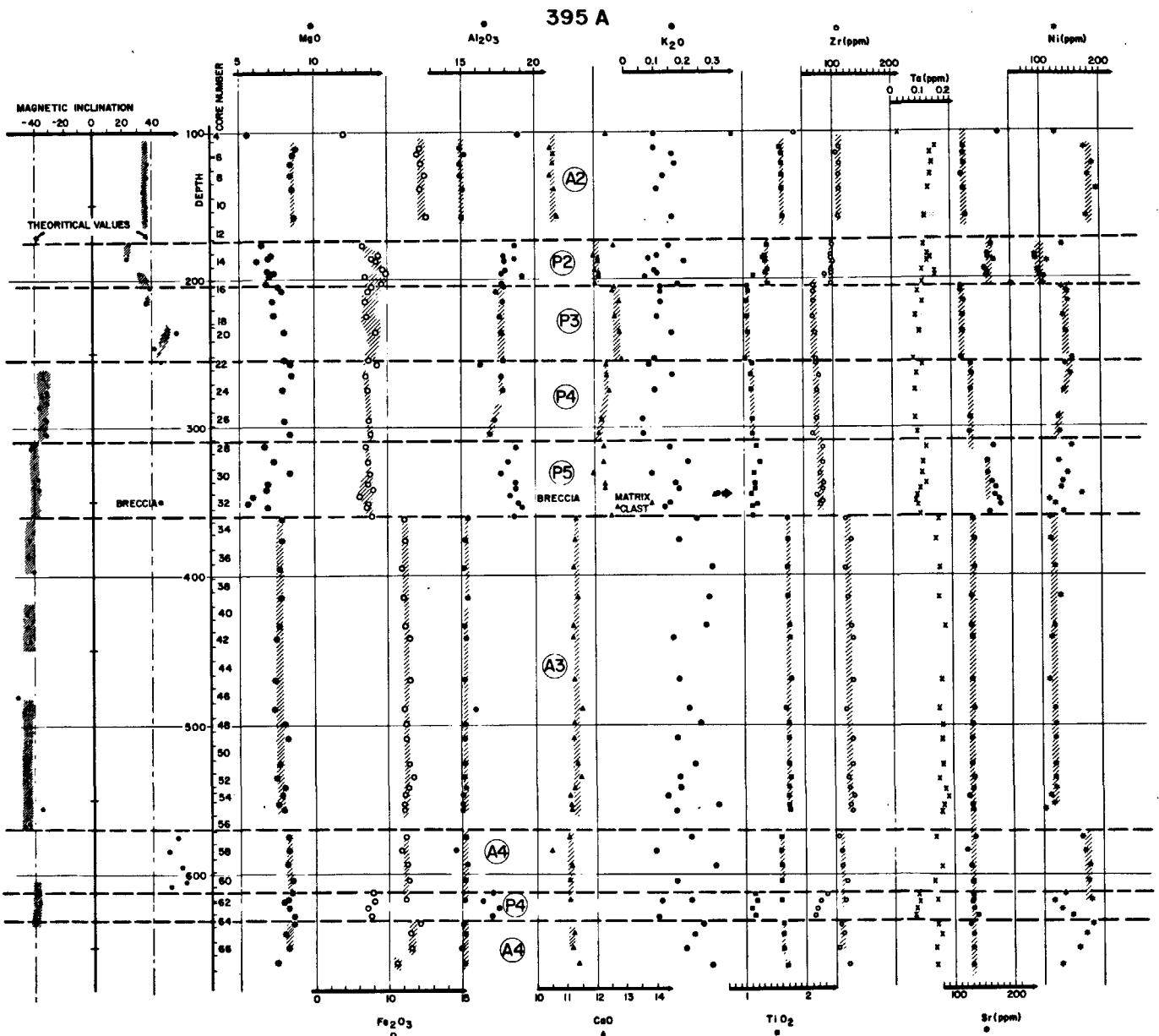


Figure 1. Selected major and trace elements concentrations versus depth: definition of chemical units.

TABLE 7
Site 395, Trace-Element Average Values for Each Unit (expressed in ppm)

Unit	A2*	A2*	A2	A3	A4	A4	A4	P1*	P2*	P2	P3	P4	P4'	P5	
Hole	395 11-1, 105-107 cm	395	395A	395A	upper Core 61 395A	lower Core 61 395A	67-2, 54-59 cm	395	395	395A	395A	395A	395A		
Sc	NA	38.4 1.4	37.8 1.2	36.6 0.7	38.7 0.4	36.8 0.4	37.8	30.6	30.6	32.6 0.4	31.5 0.9	32.1 0.9	31.7 1.5	30.12 0.8	
Ti	FX	10200	9710	9820	10337	9528	9800	10140	6090	7680	8177	6300	6694	6750	7045
V	FX	301	268	270	306	281	273	297	212	219	235	221	226	229	206
Cr	FX	279	271	270	251	303	296	263	350	222	232	351	- see text -		295
Co	FX	48	48.6	49.5	42.7	46.8	46.3	44	37.5	34	37.6	38.8	39.4	38.5	37
Co	NA	49	49.7	49.6	43.4	47.3	47.3	44	39.5	35	38.1	40.5	40.8	40.2	38.2
Ni	FX	161	176	178	117	172	169	129	135	86	94.8	140	1.6	1.5	136.1
Ni	NA	160	185	182	118	176	176	131	145	90	101	143.6	- see text -		145.3
Cu	AA	69	67.8	72.8	57.8	65.5	60.3	54	64	55	60.6	65.8	69	64.7	59.5
Zn	AA	92	85.5	84.2	84.6	81.5	84.5	83	55	58	64.3	59	63.8	65	64
Sr	FX	127	119	121	131.2	127.5	129	132	113	162	162	117	135.6	132.2	163.8
Zr	FX	121	116	110	125.6	111.2	109	124	67	94	97.6	67	72.1	73.5	78.2
Zr	NA	112	118	115	131.8	114.7	116	112	65	98	104.8	73	82.3	74.2	88.6
La	NA	3.25	3.39	3.25	4.1	3.5	3.6	4.0	2.0	2.8	3.0	2.1	2.2	2.2	2.45
Eu	NA	1.5	1.39	1.39	1.52	1.41	1.4	1.46	0.96	1.18	1.24	0.97	1.02	1.03	1.07
Tb	NA	0.87	0.83	0.82	0.89	0.81	0.82	0.87	0.53	0.64	0.68	0.55	0.56	0.57	0.58
Hf	NA	2.84	2.82	2.79	3.16	2.8	2.93	3.07	1.65	2.21	2.37	1.73	1.88	1.83	1.97
Ta	NA	0.2	0.190	0.189	0.228	0.193	0.198	0.225	0.106	0.149	0.161	0.112	0.110	0.106	0.125
Th	NA	0.137	0.134	0.130	0.164	0.144	0.147	0.142	0.081	0.113	0.120	0.084	0.083	0.074	0.096
		0.004	0.013	0.017	0.014	0.010	0.010				0.012	0.013	0.014	0.009	0.013

TABLE 8
Site 395, Major-Element Average Values for Each Unit (volatile-free concentrations, wt. %)

Unit	A2*	A2*	A2	A3	A4	A4	A4	P1*	P2*	P2	P3	P4	P4'	P5
Hole	395 11-1, 105-107 cm	395	395 A	395 A	395 A upper Core 61	395 A lower Core 61	395 A 67-2, 54-59 cm	395	395	395 A	395 A	395 A	395 A	395 A
No. Samples	1	6	6	16	4	4	1	2	1	8	6	6	4	6
SiO ₂	48.9	49.37	49.45	49.73	49.37	48.90	49.90	49.40	49.60	49.71	49.47	49.82	49.47	49.54
		0.32	0.23	0.14	0.77	0.29	0.46	0.63	0.46	0.32	0.37	0.26	0.13	0.35
Al ₂ O ₃	15.93	14.93	15.01	15.14	15.00	15.17	15.10	17.77	18.77	18.04	17.69	17.15	16.90	18.29
		0.12	0.14	0.22	0.36	0.10	0.10	0.63	0.32	0.32	0.19	0.61	0.47	0.39
Fe ₂ O ₃ (t)	12.83	12.06	12.31	11.20	11.17	11.46	10.56	8.49	9.52	9.33	8.77	8.82	8.87	8.73
		0.15	0.23	0.19	0.23	0.23	0.18	0.18	0.13	0.49	0.26	0.30	0.20	0.13
MnO	0.20	0.18	0.18	0.18	0.18	0.20	0.17	0.14	0.13	0.14	0.14	0.14	0.14	0.14
		0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01
MgO	6.8	8.58	8.53	7.61	8.20	8.40	7.30	8.20	6.10	6.84	7.63	7.78	8.35	7.24
		0.20	0.15	0.25	0.14	0.18	0.42	0.42	0.41	0.41	0.35	0.83	0.19	0.66
CaO	11.01	10.53	10.60	11.29	10.97	11.09	11.40	12.71	12.05	12.13	12.75	12.25	12.03	12.26
		0.03	0.09	0.11	0.32	0.16	0.19	0.19	0.21	0.21	0.13	0.19	0.10	0.22
Na ₂ O	2.68	2.66	2.64	2.46	2.40	2.5	2.5	2.15	2.59	2.59	2.16	2.29	2.22	2.48
		0.02	0.04	0.06	0.04	0.04	0.04	0.08	0.08	0.08	0.05	0.09	0.01	0.06
K ₂ O	0.22	0.11	0.14	0.21	0.20	0.23	0.29	0.10	0.30	0.13	0.12	0.14	0.09	0.16
		0.02	0.03	0.05	0.08	0.02	0.04	0.04	0.04	0.04	0.02	0.05	0.03	0.04
TiO ₂	1.70	1.62	1.64	1.72	1.59	1.62	1.69	1.01	1.28	1.37	1.05	1.12	1.13	1.18
		0.01	0.01	0.02	0.01	0.02	0.01	0.01	0.01	0.03	0.02	0.01	0.04	0.05
P ₂ O ₅	0.18	0.15	0.17	0.17	0.14	0.15	0.15	0.11	0.13	0.14	0.11	0.11	0.10	0.11
		0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.02
Total	100.45	100.01	100.67	99.71	99.22	99.72	99.06	100.08	100.47	100.42	99.89	99.62	99.30	100.13
Loi		-0.93	-0.98	-1.34	-1.9	-2.87	-2.60	-1.0	-1.5	-1.84	-1.2	-1.3	-1.85	-1.88

TABLE 9
Site 396, Major-Element Average Values
for Each Unit (volatile-free
concentrations, wt. %)

No. Samples	Unit		
	P _a	P _b	P _c
	1	4	4
SiO ₂	49.3	49.68	49.45
		0.13	0.35
Al ₂ O ₃	16.25	16.78	16.73
		0.10	0.15
Fe ₂ O ₃ (t)	9.96	9.33	9.48
		0.24	0.16
FeO	5.36		
MnO	0.16	0.16	0.16
MgO	7.0	8.17	7.55
		0.31	0.26
CaO	11.57	11.98	12.03
		0.09	0.07
Na ₂ O	2.40	2.41	2.46
		0.07	0.03
K ₂ O	0.31	0.24	0.21
		0.04	0.05
TiO ₂	1.50	1.27	1.29
		0.02	
P ₂ O ₅	0.15	0.12	0.11
		0.01	0.01
Total	98.60	100.14	99.47
Lol	-2.0	-2.8	2.25
		0.3	0.21

Note: Averages and standard deviations listed for Units P_b and P_c

TABLE 10
Site 396, Trace-Element Average Values for
Each Unit (expressed in ppm)

396		Unit		
		P _a	P _b	P _c
No. Samples		1	4	4
Sc	NA	36.8	35	34.7
Ti	FX	9000	7600	7740
V	FX	312	264	255
Cr	FX	288	393	290
Co	FX	40	40	40.5
Co	NA	40.7	41	42
Ni	FX	111	142	121
Ni	NA	114	143	124
Cu	AA	58	66	65
Zn	AA	73	66	72
Sr	FX	148	156	131
Zr	FX	109	87	86
Zr	NA	118	87	97
La	NA	3.5	2.9	2.7
Eu	NA	1.36	1.2	1.22
Tb	NA	.79	0.65	0.68
Hf	NA	2.64	2.2	2.2
Ta	NA	0.21	0.17	0.156
Th	NA	0.16	0.13	0.114

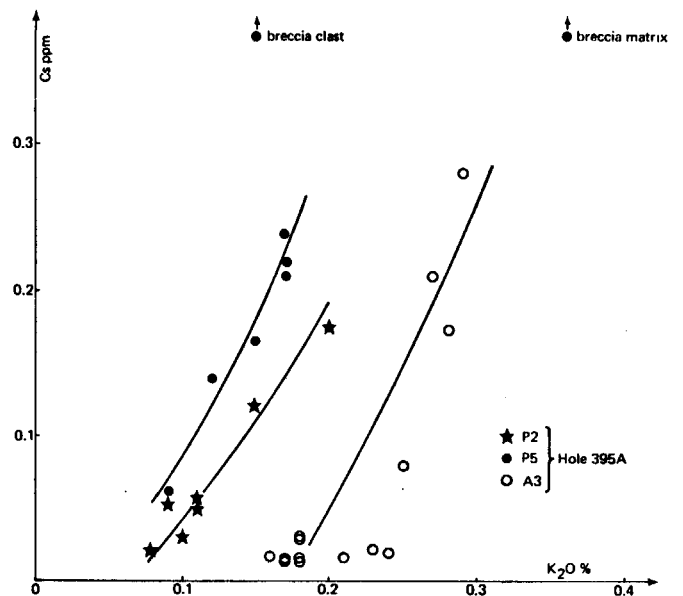


Figure 2. Cesium versus K₂O; Units P₂, P₅, and A₃ of Hole 395A.

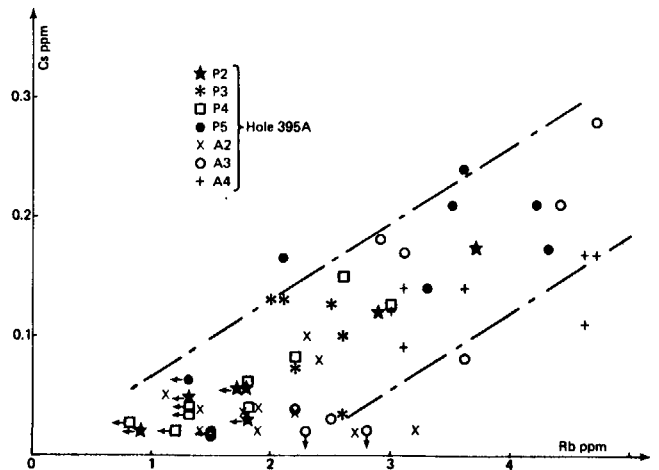


Figure 3. Cesium versus rubidium; all units of Hole 395A.

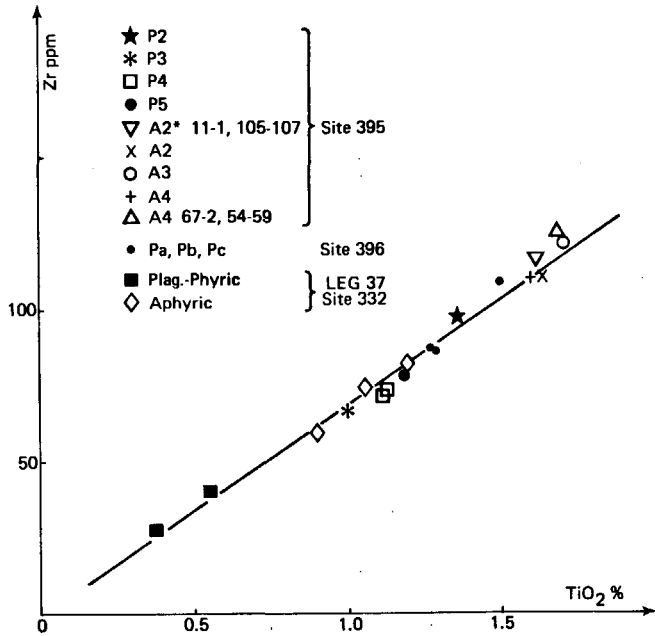


Figure 4. Zirconium versus titanium (TiO_2); comparisons between Leg 45 and Leg 37 results. Shifts along the line from Leg 37 to Leg 45 samples for phyric samples on one hand and for aphyric samples on the other.

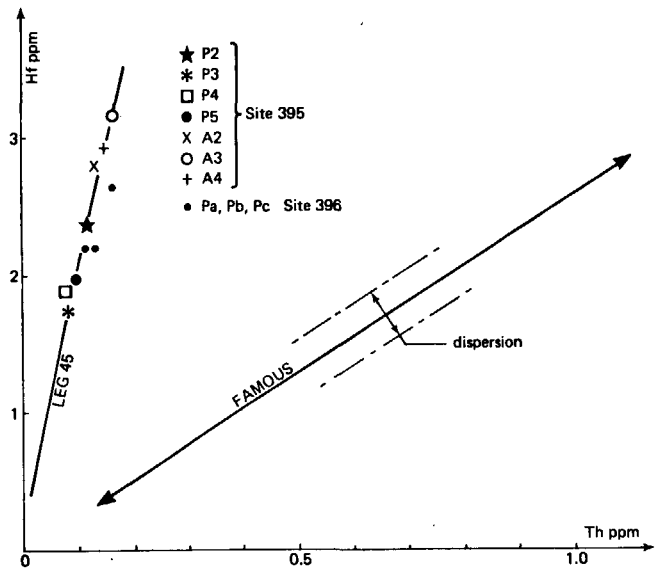


Figure 6. Hf versus thorium; an LPC element as a function of another very low LPC element, illustrating upper mantle heterogeneity plus variable partial melting (see text).

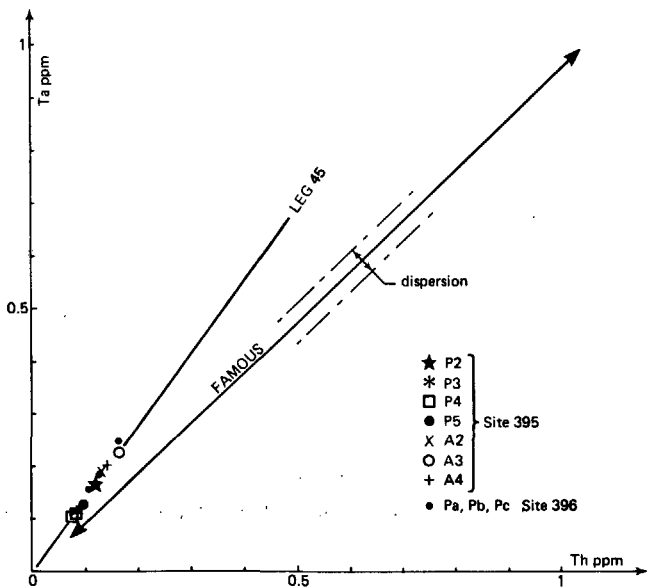


Figure 5. Tantalum versus thorium; a very low partition coefficient (LPC) hygromagmatophile element as a function of another very low LPC, hygromagmatophile element; upper mantle heterogeneity (see text).

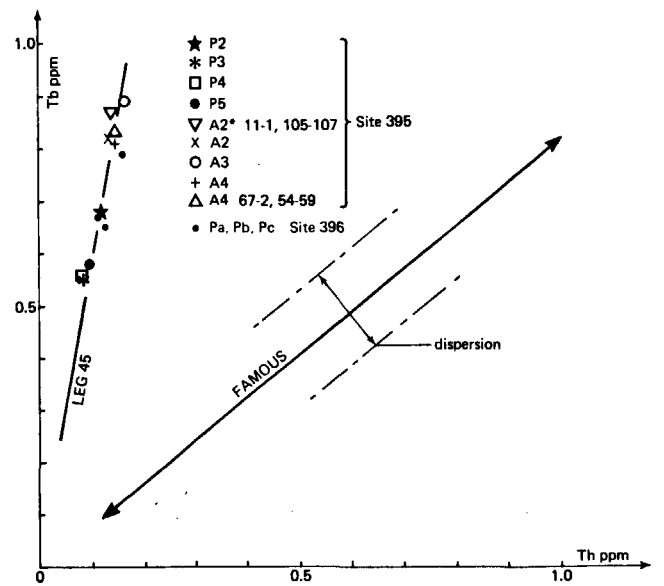


Figure 7. Terbium versus thorium; an LPC element as a function of another very low LPC element, illustrating mantle heterogeneity plus variable partial melting (see text).

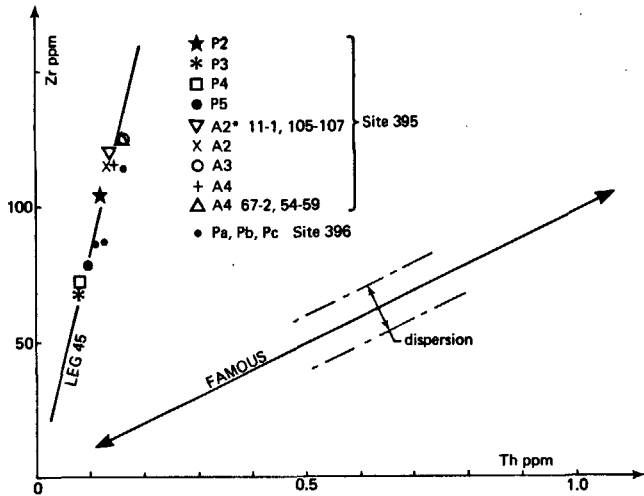


Figure 8. Zirconium versus thorium; an LPC element as a function of another very low LPC element, illustrating mantle heterogeneity plus variable partial melting (see text).

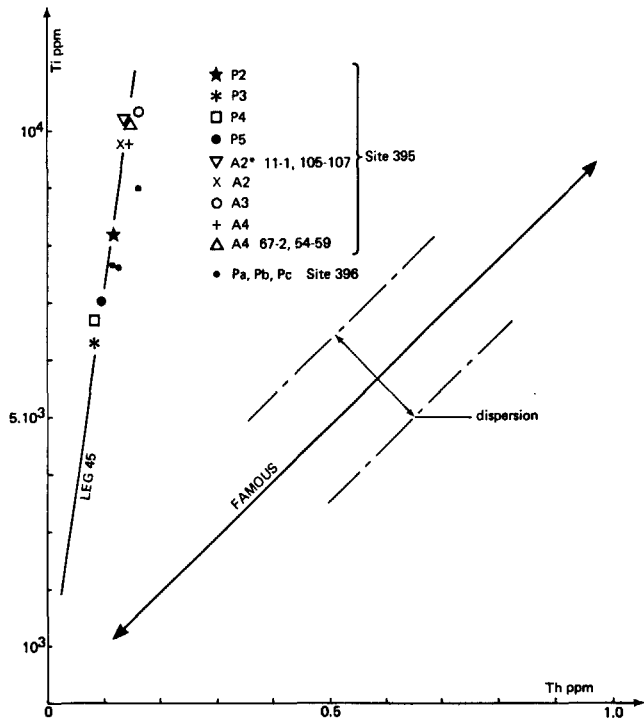


Figure 9. Titanium versus thorium; an LPC element as a function of another very low LPC element, illustrating mantle heterogeneity plus variable partial melting (see text).

TABLE 11A
Slope Values of Lines Corresponding to Element i Plotted Versus Element j

i		Th	Ta	Hf	Tb	Zr	Ti
Th	F	1	0.95	2.5	0.83	102	9500
	45	1	1.38	20	5.8	830	70000
Ta	F	1.05	1	2.63	0.87	107	10240
	45	0.725	1	14.5	4.2	602	50750
Hf	F	0.40	0.38	1	0.33	40.8	3900
	45	5 10 ⁻²	6.9 10 ⁻²	1	0.29	41.5	3500
Tb	F	1.20	1.14	3	1	122	11700
	45	0.17	0.23	3.4	1	141	11900
Zr	F	9.8 10 ⁻³	9.3 10 ⁻³	2.45 10 ⁻²	8.13 10 ⁻³	1	95.5
	45	1.2 10 ⁻³	1.65 10 ⁻³	2.4 10 ⁻²	6.96 10 ⁻³	1	84
Ti	F	1.02 10 ⁻⁴	9.7 10 ⁻⁵	2.55 10 ⁻⁴	8.46 10 ⁻⁵	1.04 10 ⁻²	1
	45	1.43 10 ⁻⁵	1.97 10 ⁻⁵	2.86 10 ⁻⁴	8.29 10 ⁻⁵	1.18 10 ⁻²	1

Note: The dashed line separates values involving either Th or Ta, or both, from values not involving these elements, but which were calculated relative to them. Leg 45 data indicated by 45, FAMOUS data by F.

TABLE 11B
Examples of Absolute Values of LPC Elements for Aphyric and Phyruc Samples^a

Unit A ₃		Th	Ta	Hf	Tb	Zr	Ti
ARP 7-7	45	0.164	0.228	3.16	0.89	125	10337
	F	0.86	0.85	2.37	0.71	110	8520
Unit P ₂		Th	Ta	Hf	Tb	Zr	Ti
ARP 31-36	45	0.120	0.161	2.37	0.68	97	8177
	F	0.32	0.28	0.94	0.29	40	3300

^aA comparison between Leg 45 (45) and FAMOUS area (F). Note: The dashed line separates representative aphyric or sparsely phyruc samples (A₃ and ARP 7-7) from phyruc samples (P₂ and ARP 31-36).

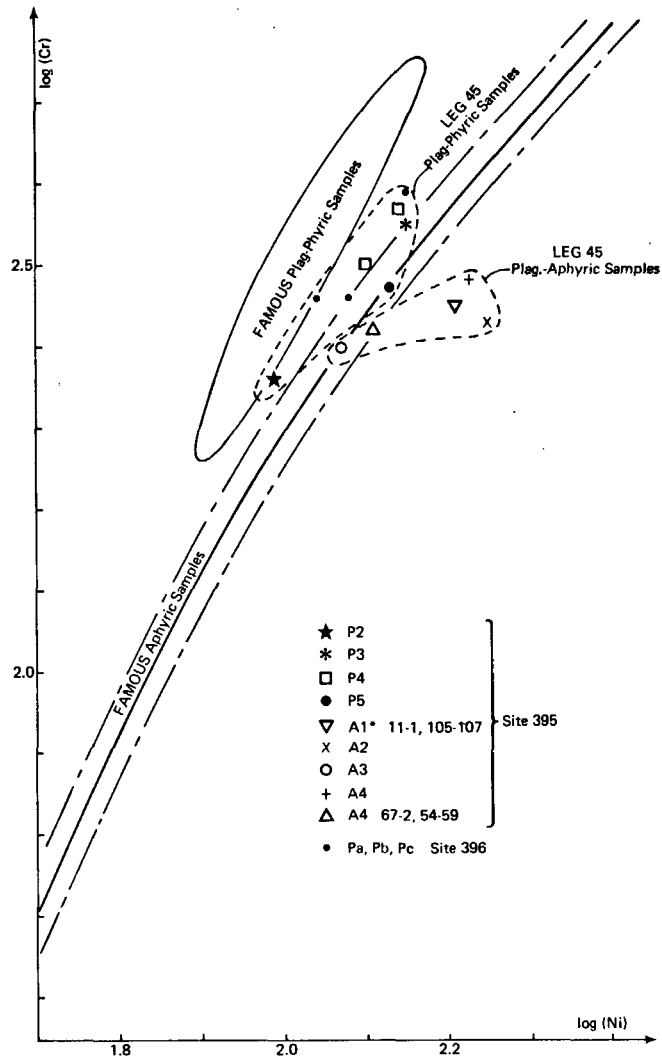


Figure 10. *Log (Cr) versus Log (Ni).*