

Planktic foraminifera Water masses Frontal zones Diversity NE Atlantic

Planktic foraminifera as North Atlantic water mass indicators

Foraminifères planctoniques Masses d'eau Région frontale Diversité Atlantique NE

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ABSTRACT

Distribution patterns of planktic foraminifera, collected from Eastern North Atlantic surface waters during August and September 1986, are related to measured hydrographical parameters. Three main water masses, Subpolar Water (SW), the North Atlantic Current (NAC), and the Azores Current (AC) are all characterized by distinct faunal assemblages. A fourth, distinct water mass is here referred to as North Atlantic Transitional Water (NATW). The distinction, based on species clusters, between the NATW and the southern AC, is less evident than between the NATW and the northern SW and the NAC. This is the combined result of low absolute frequencies, a high number of species and a superimposed complex hydrography in the southern water masses.

The oceanic frontal zones between the NAC, NATW and the AC coincide with faunal boundaries. Along the boundaries, broad mixing zones can be recognized by their highly diverse foraminiferal assemblages of relatively low equitability.

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RÉSUMÉ

Les foraminifères planctoniques comme indices des masses d'eau de l'Atlantique Nord

Des modèles de répartition des foraminifères planctoniques prélevés dans les eaux superficielles à l'est de l'Atlantique Nord en août et septembre 1986, sont comparés aux paramètres hydrologiques mesurés à la même époque. Trois masses d'eau importantes, l'eau subpolaire (SW), le courant nord-atlantique (NAC), et le courant des Açores (AC) sont caractérisés par des associations de foraminifères distinctes les unes des autres. Une quatrième masse d'eau distincte peut être mentionnée : l'eau intermédiaire nord-atlantique (NATW). La distinction entre la NATW et le AC du Sud, d'après les groupes d'espèces, est moins évidente que celles entre la NATW, la SW du Nord et le NAC. Ceci est dû aux faibles fréquences absolues, au nombre élevé d'espèces et à l'hydrologie complexe dans les masses d'eau australes.

Les zones océaniques frontales entre le NAC, la NATW et le AC coïncident avec les limites faunistiques, au long desquelles des zones de mélange étendues peuvent être identifiées par des associations de foraminifères très diverses et d'une équitabilité relativement basse.

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INTRODUCTION

Paleoceanographic and paleoclimatic reconstructions based on planktic foraminifera (e.g.: CLIMAP, 1976; 1981; Vincent and Berger, 1981) make use of the traditional distinction between "arctic", "subarctic", "transitional", "subtropical" and "tropical" faunal provinces (Bé and Tolderlund, 1971). However these provinces do not necessarily match the planktic foraminiferal patterns in surface waters. Faunal frequency patterns as well as the peak frequencies of different species in fact undergo substantial changes throughout the year, induced by seasonally changing hydrographic conditions (e. g.: Cifelli and Smith, 1974; Cifelli and Benier, 1976; Pujol et al., 1976; Brummer and Kroon, 1988). A detailed knowledge of hydrography can clarify the controls on the differences between surface water and bottom assemblages. For example, water circulation can create and maintain water bodies with particular species groups. Information concerning the seasonal effect in sediment assemblages can be obtained from water samples collected at identical stations during different seasons. By quantitatively determining specific living seasonal assemblages and their horizontal variation, it can be established to what extent a particular season is represented in the annual shell accumulations on the sea floor and we can consequently determine average paleotemperatures. For example, if a certain species is solely present during summer in a certain area, an average summer temperature for the sedimentary assemblage in that area can be reconstructed from the oxygen isotope ratio in the foraminiferal shell of this particular species (Ganssen, 1983).

The objective of this paper is to characterize water masses in terms of planktic foraminifera (frequency distributions, cluster analyses and information indices) in conjunction with recorded hydrographical parameters such as temperature, salinity and density. The study covers the Subpolar Water, the North Atlantic Transitional Water and the North Atlantic and Azores Currents. The author calculated information indices for use as a research tool on sediment assemblages comprising extinct species, as they do not depend on the assumption of a temporally invariant species ecology or of a particular species composition. The described material was collected during the R/V Tyro APNAP-I cruise (August and September 1986) in the Central North Atlantic (Ganssen et al., 1986). APNAP (Actuomicropaleontology Paleoceanography North Atlantic Project) investigates an area east of the Mid-Atlantic Ridge from 60° N to 27° N (Fig. 1 a, b) during several seasons. In the course of this cruise, planktic foraminifera were collected from surface waters, the watercolumn and the seafloor.

The present paper concerns the study of 93 surface water samples with the aim of applying derived foraminiferal water mass characteristics to bottom sediments.

METHODS

Whith the research vessel under way, surface water (0-5 m) was continuously filtered through 75 µm mesh nylon plankton nets by means of a deckwash pump with an attached flow meter. A collecting interval of between three and six hours was used, depending on productivity. Upon recovery, part of the samples were examined under a light microscope to verify that the amount of filtered water was sufficient to obtain a valuable data set. The temperature and salinity of the surface waters were logged at the time of recovery using continuous CTD recording, when available, or were measured using hand methods (Tab. 1). Additionally, 14 CTD profiles were recorded. Only patterns are used for the reason that the CTD data underwent no calibration. The plankton samples were stored in an alcohol compound (95 % ethanol, 5 % methanol). In the laboratory the samples were washed over a 63 µm sieve to remove the alcohol. The residue was dried at 50° C and then combusted in a Low Temperature Asher, which oxidizes organic matter, concentrating the non-organic contents (e.g. calcareous shells). Species were counted seprately for the 125-250 µm and 250-500 µm fractions. Although the coarse fraction (250-500 µm) adds only less than one per cent to the total sample (>125 μ m; Tab. 2 a and b) and its signals will therefore be obscured, we did separate it, because it permits rapid determination and is commonly used for isotopic measurements. Therefore its very variable and low absolute frequencies are not considered in quantitative data analysis such as cluster analysis on samples and . equitability computations. For each fraction, large samples were split into suitable aliquots of at least 200 specimens; sample fractions containing less than 200 specimens were counted completely. Matrices of planktic foraminiferal relative frequencies were clustered both for samples and for species. A Principal Component Analysis (PCA) did not result in a significant variation reduction and is considered no further in this paper. In addition, information indices of Shannon-Wiener for diversity and Buzas-Gibson for equitability (Lipps et al., 1979) were computed for this dataset. Since equitability values depend in great measure on the number of specimens, especially when the species count is low (Sheldon, 1969), the coarse fraction with its low and variable frequencies is not considered in the equitability computations. The resulting values of the fine fraction were been contoured on the track chart. Computer contouring of these and of surface temperature and salinity data used the weighted mean method and an adjustable digital Butterworth filter (Slootweg, 1978). The material has been filed in the collection of the Geomarine Center, Amsterdam.

Hydrography

GENERAL

Surface water temperatures of the Eastern North Atlantic are closely related to latitude and currents, and to a with





Figure 1

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a) R/V Tyro track of cruise APNAP I (August and September, 1986) showing CTD stations and dates. b) Idem showing numbers and intervals of surface water sampled. Dots at the end of each sampling interval also indicate samples used for temperature and salinity measurements.

DATE	NUMBER	TIME START	TIME END	LAT START	LAT END	LON START	LON END	P	TEMP	SALI
170886	T86-F1	00.00	07.00	50,57.7	51,18.6	11,15.2	13,06.6	y ı	15.05	35.12
170886	T86-F2	07.30	13.00	51,20.2	51,33.8	13,15.4	14,46.3	Y 1	14.75	35.15
170886	186-F4	16.00	19.00	51,53.8	51,41.1	14,46.3	15,32.1	y ı	115.10	35.05
170886	T86-F5	19.00	22.30	51,49.0	51,58.8	16,21.4	17,21.5	Y Y	1 15.25	35.30
180886	T86-F6	01.00	07.00	52,04.6	52,25.6	18,01.3	19,55.1	Ŷ	14.80	35.00
180886	T86-F7 T86-F8	07.00	12.30	52,25.6	25,35.8	19,55.1	21,13.2	YI	14.70	35.10
180886	T86-F9	16.00	19.20	52,43.7	52.51.8	22,12.1	22, 12.1 23.11.0	V I	14.50	35.17
180886	T86-F10	19.20	22.00	52,51.8	52,59.6	23,11.0	23,38.5	y i	13.90	35.15
190886	T86-F11	22.00	01.00	52,59.6	53,06.4	23,38.5	24,31.5	УЗ	14.15	35.00
190886	T86-F13	01.00	04.00	53,00,4	53,13.5	24,31.5	25,23.0	Y I	14.35	34.94
190886	T86-F14	07.00	10.00	53,23.3	53,29.9	26,14.0	27,04.6	y i	14.50	34.78
190886	T86-F15	10.00	16.30	53,29.9	53,29.9	27,04.6	27,08.1	У	13.90	34.75
190886	T86-F16	16.30	23.00	53,29.2	53,24.6	27.08.1	27,34.9	Y	13.95	34.65
200886	T86-F18	03.00	06.00	52.54.4	52.16.7	27.58.8	28.11.8	Y Y	14.40	34.80
210886	T86-F19	08.45	16.20	51,44.1	50,19.3	27,26.3	27,03.6	Y I	15.00	35.15
220886	T86-F20	16.20	00.00	50,19.3	50,09.8	27,03.6	26,59.4	УJ	1 15.45	34.58
220886	T86-F21	00.00	03.15	50,09.8	50,09.7	26,59.4	26,57.4	УI	14.70	35.10
220886	T86-F23	09.30	14.30	49,19.2	48.39.8	26,27.8	25,50.9	y i v i	115.90	34.97
220886	T86-F24	14.30	17.00	48,39.8	48,43.9	25,50.9	25,49.5	YI	1	33.00
220886	T86-F25	18.00	23.30	48,40.5	48,42.4	25,50.1	26,02.8	Y I	16.20	35,10
230886	180-F26 T86-F27	23.30	11.30	48,42,4	47,44.1	26,02.8	25,41.6	ny	16.00	34.90
230886	T86-F28	11.30	17.45	46,52.9	46,54.2	25,22.2	25,25.3	יין עריק	1 10.13	55.00
240886	T86-F29	03.45	10.30	45,38.0	44,59.0	26,05.0	26,25.4	Y y	19.00	34.70
240886	T86-F30	10.30	19.00	44,59.0	45,06.6	26,25.4	26,25.2	уг	18.70	34.92
250886	T86-F32	01.00	01.00	44,29.0	43,57 2	25,34 0	25,34.0	y I v ·	118.95	35.30
250886	T86-F33	07.15	14.15	43,57.3	43,56.0	25,01.0	24,57.6	y '	20.20	34.90
250886	T86-F34	14.15	23.00	43,56.0	43,51.7	24,57.6	25,00.0	y i	20.45	34.75
260886	T86-F35	23.00	04.00	43,51.7	43,02.5	25,00.0	25,22.4	ΥI	20.15	35.38
260886	T86-F37	09.30	18.00	42.17.8	42,17.6	25,22,4	25,38.6	y i v i	120.08	35.42
260886	T86-F38	18.00	23.15	42,13.6	42,07.9	25,39.2	25,44.2	y i	21.30	35.30
270886	T86-F39	23.15	04.30	42,07.9	41,26.9	25,44.2	26,36.7	X J	21.60	35.50
270886	T86-F40	104.30	10.30	41,26.9	40,47.8	26,36.7	27,27.8	УY	21.40	35.64
280886	T86-F42	21.30	04.30	40.39.5	39,40.0	27.32.0	28.02.0	v r	121.55	35.85
280886	T86-F43	04.30	10.15	39,40.0	38,46.0	28,02.0	28,29.0	y i	24.35	35.80
290886	T86-F44	01.00	05.30	38,28.4	37,46.9	28,39.1	29,21.8	УI	23.20	
290886	T86-F45	05.30	11.45	37,46.9	37,06.8	29,21.8	30,03.1	У	1 23.80	35.95
300886	T86-F47	22.30	05.00	37.06.1	36.26.0	30,03.1	31,22 2	Y I	24.30	35.50
300886	T86-F48	05.00	12.00	36,26.0	35,46.0	31,22.2	32,31.4	ý i	25.78	36.44
300886	T86-F49	12.00	23.30	35,46.0	35,41.9	32,31.4	32,31.9	Уз	1 25.79	36.41
310885	T86-F50	23.30	05.30	35,41.9	34,59.1	32,31.9	33.29.3	УI	125.90	36.35
310886	T86-F52	11.30	23.00	34,19.0	34,11.3	34,20.0	34,20.0	Y Y	24.25	36.52
010986	T86-F53	23.00	10.30	34,11.3	32,46.3	34,25.3	34,43.2	Y	26.30	36.80
020986	T86-F54	13.45	01.30	32,45.6	32,45.8	34,43.0	34,41.6	נ ע	26.69	36.78
020986	T86-F56	06.00	14.10	32,45.9	31,26.6	34,41.6	35,44.4	Y Y	26.88	36.90
030986	T86-F57	14.10	01.45	31,26.6	31,25.7	36,15.0	36,18.5	Y	27.18	36.56
030986	T86-F58	01.45	10.45	31,25.7	30,32.1	36,18.5	36,53.8	YI	1 27.20	36.36
040986	186-F59	10.45	05 00	30,32.1	30,24.1	36,53.8	37,03.5	Y I	27.35	36.53
040986	T86-F61	05.00	11.15	29,45.0	29,05.6	38,01.0	38,51.3	l, ,	27.82	37,06
050986	T86-F62	11.15	02.45	29,05.6	29,06.4	38,51.3	38,50.3	y i	27.76	37.07
050986	T86-F63	06.15	11.10	29,07.0	28,35.4	38,49.7	38,44.7	Y :	27.91	37.21
060986	T86-F65	21.15	02.00	28,33.1	29,07.0	38,44.7	38,41.7	Y Y	1 28.22	37.18
060986	T86-F66	02.00	06.30	29,07.0	29,40.8	37,55.0	37,14.4	y 1	27.66	37.00
060986	T86-F67	06.30	13.00	29,40.8	30,30.1	37,14.4	36,10.0	א ו	27.84	36.83
070986	180-108 186-F60	13.00	20.40	30,30.1	31,22.3	36,10.0	34,54.0	Y 1	27.21	36.70
070986	T86-F70	01.00	08.30	31,55.4	32,49.9	31,07.2	32,55.1	י ז'ן י ז'ן	26.71	36,73
070986	T86-F71	08.30	13.20	32,49.9	33,28.4	32,55.1	32,00,1	y i	28.86	36.78
070986	T86-F72	13.20	19.30	33,28.4	34,11.7	32,00.1	30,54.9	y ı	100 01	1
080986	186-F74	19.30	01.15	34,55 5	35,54 6	29.54 2	29,54.2	ניצו	125.84	36.64
080986	T86-F75	09.00	13.30	34,54.6	36,29.0	28,35.2	27,44.8	ני ען	24.37	36.02
080986	T86-F76	13.30	19.15	36,29.0	37,07.2	27,44.8	26,51.9	y i	23.05	35.97
090986	186-F77	19.15	01.00	37,07.2	37,53.9	26,51.9	25,53.2	Y 1	22.57	35.79
090986	T86-F79	08.00	13.45	38.48.4	39,18.1	25,53.2	23,23 2	v i	22.53	35.74
090986	T86-F80	13.45	19.15	39.18.1	39,53.3	23,23.2	22,10.6	Y I	21.71	35.75
100986	T86-F81	19.15	02.00	39,53.3	40,26.6	22,10.6	21,04.7	Y	22.18	35.93
100986	T86-F82	02.00	09.00	40,26.6	41,13.5	21,04.7	19,46.4	YI	120.55	35.61
100986	T86-F84	13.30	19.00	41,29.0	42,00.0	19,17.0	18,17.0	V Y	20.45	35.60
110986	T86-F85	19.00	01.15	42,00.0	42,36.4	18,17.0	17,03.6	ý v	19.91	35.58
110986	T86-F86	01.15	09.00	42,36.4	43,30.1	17,03.6	15,23.0	уī	19.10	35.48
110986	186-F87	09.00	13.30	43.30.1	43,54.2	15,23.0	14,30.1	Y	19.67	35.50
120986	T86-F89	19.45	01.30	44.41 4	45,12 4	13,14 /	12,08 1	Y I	1 18.98	35.31
120986	T86-F90	01.30	10.00	45,12.4	45, 57.3	12,08.1	10,38.7	y i	18.14	35.42
120986	T86-F91	10.00	14.00	45,57.3	46,27.3	10,38.7	09,41.6	УЗ	18.00	35.32
120986	T86-F92	14.00	19.15	46,27.3	46,59.9	09,41.6	08,30.5	Y	18.48	35.16
130300	100-193	13.12	01.30	+0, 59.9	*1,22.3	00,30.5	01,40.1	X Z	11.37	35.13

Table 1

Timing and location of sampled intervals and temperature and salinity measurements at interval end. $P = plankton \ sample \ (n = no \ recovery) \ and \ I = water \ sample \ for \ isotope \ analysis \ (n = not \ collected).$

Strong seasonal variations in temperature are encountered especially at the higher latitudes, while salinity remains relatively constant throughout the year. North of the subtropical convergence at about 33° N, with a seasonal shift of 3°(Käse and Siedler, 1982), salinities and temperatures primarily depend on currents of Gulf Stream origin (Tchernia, 1980). In the north, the Subarctic Front (SF), associated with strong horizontal temperature gradients, separates cold Subpolar Water (SW) fromwarmer Gulf Stream waters (Krauss and Käse, 1984). The Gulf Stream branches in the vicinity of 40° N-40° W into a strong North Atlantic Current (NAC) and the Azores Current (AC; Käse *et al.*, 1986). The northern boundary of the NAC is also formed by the so-called Subarctic Front, although temperature gradients are less distinct in the eastern basin (Krauss, 1986). The NAC dies out southward, yet strong current regions (NAC s.s.) are separated by frontal jets from regions with indistinct currents (Krauss, 1986), referred to as North Atlantic Transitional Water (NATW) in this paper. At about 34° N the perpetual, basically west-east meandering Azores Frontal Zones (AFZ) is associated



Figure 2

Surface water masses as described in the literature for the North East Atlantic Basin.

		1	-1			_	-	··											
sn	sac rub	aeq ter	bul	rubs	glu	inf	dut	inc	qui	ru	sci	fal	Cra	con	pac	res	hs	SW	BG
1	4.1 0.	3 0 0	64 1	٥	0.5	1 0	0	20.5	7.8		0	0	0 ·	-	0	0	1490	1 1	0 375
2	1.1 1.	1 0 0	40.2	Ō	1.1	3.9	ŏ	26.8	25.7	ŏ	ō	õ	ŏ	ŏ	ŏ	õ	1432	1.35	0.549
3	3.5 0.	6 0 0	54.3	0	0.6	1.7	0	27.7	10.4	0	0	0	0	0	0	0	2753	1.23	0.38
5	0.4 0		48.6	0	0	2.9	0	8.6	62.9	0	0	0	0	0	0	0	35	1.12	0.611
6	1.7 0	0 0	49.2	Ő	0	1	ŏ	29.3	18.9	ŏ	0	ŏ	ŏ	õ i	0	õ	2376	1.14	0.624
7	0 0	0 0	29.4	0	0.5	4.4	0	47.5	17.6	0	0	0	0	0	0	0	3264	1.21	0.559
9	0 0		23	0	0	0.9	0	23.8	58.8	0	0	0 0	0	0	0	0.1	3904	1.15	0.792
10	0 0	1.8 0	18.7	Ō	0.4	0.4	ō	9.4	69.1	ō	Ō	ō	Ō	0	ō	0.4	2224	0.92	0.36
11		2 0 0	16	0	0	1.2	0	15.3	67.1	0	0	0	0	0	0	0	1126	0.92	0.42
13	0.2 0	0.5 1	38.5	ŏ	ŏ	0.4	0.1	7.2	52.1	Ö	0.2	ŏ	õ	ŏ	0	0	6468	1.01	0.344
14	0 0	0 0	29.	0	0	0.5	0	4.7	65.5	0	Ō	Ō	0	0	Ó	0	6176	0.81	0.56
15		0 0	43.2	0		1	0	6 8 5	49.8	0	0	0	0		0	0	2520	0.92	0.629
17	o o	0.5 0	48.7	ŏ	0	0.7	0	4.7	45.3	õ	õ	0	õ	ŏ	ō	0	4400	0.92	0.501
18	0 0	2.9 0	44.2	0	1.5	4.1	0	38.6	7.6	0	0	1.2	0	0	0	0	2736	1	0.676
20	0 0	1.7 0	50.1	. 0	0.1	3.1	0	35.4	14.7	0	0		0	0	0	0	41024	1.34	0.635
21	0 0	3.2 0	45.5	0	0.5	3.6	Ō	34.8	11.4	0	Ō	ō	Ō	Õ	Ō	0.1	28256	1.28	0.4
22		1 0 0	5 42.7	0	0	'6.7	0	39.8	10.2	0	0	0.8	0	0	0	0.1	49025	1.18	0.542
24	0 0.	4 0.2 0	24.1	0.1	0.1	8.2	0	57.7	8.6	ŏ	0	ŏ	0	ŏ	0	ŏ	1654	1.13	0.387
25	0 0	0 0	46.2	0	0.6	4.6	0	44.8	3.8	0	0	0	0	0	0	0	1384	1.01	0.55
20		0 0	65.4	0	0	8.4	-	22.8	3.4	0	ñ	-	0	ō	0	ō	- 610	- 0 94	- 639
28	0 0	1.5 0	42.1	Ō	Ō	7.7	Ő	47.3	1.2	ŏ	õ	ŏ	D	õ	õ	0.4	1616	0.97	0.294
29	1.3 0.	5 1 0	67.1	. 0	0	4.4	0.2	21.2	4.4	0	0	0	0	0	0	0	19745	1.01	0.343
31	2.6 0.	6 1.8 0	73.1	Ö	lő	4.8	õ	11.8	5.2	0	0	0	õ	ŏ	0	0	8032	0.99	0.336
32	6.2 3.	5 4 0	60.5	0	0.5	5.9	0	12.6	6.9	0	0	0	0	0	0	0	3444	1.36	0.487
33	25.6 10.	8 5.2 0	34.1	0.5	4.8	3.8	0.5	2.9	1.7	0	0	3.5	0	0	0	0.3	4752 2940	1.8	0.464
35	41.3 11.	5 1.7 0	4 27.7	1.3	2.4	0.4	0	11.5	1.3	õ	Õ	0.4	0	ŏ	ŏ	0.4	940	1.57	0.402
36	29.2 12.	3 8.1 0	26.3	0.8	1.7	1.3	0	9.7	1.7	0	0	0	0	0	0	0	1889	1,85	0.635
38	38.4 13	2.5 2	5 37.2	0.3	0.3	0.8	0	2.8	2	0	0	0	0	ŏ	õ	0.6	2832	1.45	0.355
39	53.2 16	4.5 5	1 14.7	0.6	0	0.6	Ö	5.1	0	0	0	0	0	0	0	0	2497	1.42	0.517
40	52.7 4. 48.3 10.	6 0 2 1 13.4 4	2 7.8	0.8	9.7	1.6	0.8	16.7	1.2	0	0	0	0	0	0	0	3976	1.56	0.432
42	46.4 11.	3 6.7 5	9 18.7	0.8	1.3	0.5	1.8	3.8	1.8	0.3	0.3	õ	õ	ŏ	õ	ŏ	1561	1.72	0.37
43	40.3 10.	3 0.8 8	$\frac{3}{2}$ 7.1	5.1	10.3	2.4	0	8.3	4.3	0	0	0 [0	0 (0	0.7	1004	1.99	0.523
45	27.1 19.	4 0.7 7	4 18.7	3.7	11	0.7	0	5	2.3	2.7	0.3	ŏ	0	ŏ	0	0.3	1214	1.31	0.491
46	33.5 17.	1 2.2 6	3 13.8	0.7	9.7	0	0.4	4.8	7.1	0.4	2.2	0	0	0	1.1	0	2152	2.01	0.533
47	15 7 21	4 2.5 7	1 15.4	2.4	112	0.4	0	6.2	2.1	0.4	0.4		0		0	0	1102	1.99	0.524
49	12.8 22.	6 0.9 13	7 1.3	2.2	17.7	0.9	0.4	2.7	4	õ	8	ŏ	2.2	ŏ	5.3	4	908	2.34	0.519
50	16.1 31	4.2 9	9	2.5	11.6	0	0	3.2	0	0.3	5.5	0	1.3	0	1	2.9	620	2.21	0.48
52	12.8 19.	7 0 33	5 4.8	3.2	12.2	0.5	0	1.6	2.3	0.5	3.2	ö	0	4.3	2.5	1.5	753	2.03	0.447
53	6.5 52.	306	2 8.5	0.8	17.3	0.4	0	3.5	1.9	0.8	1.5	0	0	0	0	0.4	1040	1.58	0.374
54	4.1 21	$\begin{bmatrix} 1 & 8 \\ 1 & 6 & 3 \end{bmatrix}$	$\frac{2}{2}$ $\frac{1}{7}$	3.6	39.5	0.5	0.5	8.2	9.2	0	2.6	0	0	0	0	0	788	1.59	0.328
56	4.1 36.	3 4.4 4	4 3.2	0.9	23.3	1.9	1.3	5.4	6.6	0	1.9	0.3	0.9	0	1.6	1.6	317	2.1	0.431
57	1.2 41.	5 2.1 0	8 6.2	2.5	13.7	2.1	0	2.1	12.4	5	8.3	0.4	0.4	0	0	1.2	482	1.98	0.402
59	1.3 49.	4 0.8 4	9 3.8	0	12.9	4.6	2.3	0	9.7	0.4	4.9	0.9	0.3	0.3	0	0.9	1575	1.98	0.374
60	3.4 58.	2 0.8 4	9 3.8	0	12.9	4.6	2.3	0	4.9	0	2.7	0.4	0	0	0	1.2	526	1.58	0.345
62	1.7 70.	5 0.8 0	4 1.3	1.3	12.2	0	0	2.5	3	5.5	0	0.6	0.3	0	0	1.2	309	1.08	0.226
63	1.6 62.	2 0.5 0	5 7.3	0.3	10.2	0.5	0.3	5.4	0.8	1,1	1	0	0.3	ō	0	0.3	372	1.29	0.402
65	2.2 52	2 1.5 2	2 11.9	2	13.4	6.8 1 S	0	11 9	2	0	1.4		0	0	0	2.1	296	1.31	0.336
66	0 57	0 0	8 6.	1.6	24.2	õ	ō	3.3	ō	5.5	ŏ	õ	õ	ō	ŏ	2.3	128	1.34	0.238
67 68	0.5 53	0 1	8 2.3	0.9	17.1	0.9	0	5.5	1.4	13.8	0.5	0	0	0	0	2.3	217	1.51	0.379
69	0.5 71.	1 0.5 0	5.8	0	15.8	ŏ	0.4	1.6	0.5	3.7	0	ŏ	õ	0.5	0	0.7	760	1.0	0.301
70	3.2 40.	8 0.8 23	2 0	0	25.6	0	4.8	0	0	0	0.8	0	0	0	Ó	0.8	508	1.48	0.49
71	7.2 20	1.1 54	5 0.7	1.5	48.6	1	0	2	0	0.5	1		0	0	0	2	204	1.42	0.343
73	13.9 46.	9 4.8 9	1 2.4	1	15.3	0	0	1.4	0	5.3	õ	ŏ	0	0	Õ	0	418	1.64	0.513
74	10.2 15.		8 0	0.3	6.8	0	0.3	0	0	0	0.2	0.2	0	0.3	0	0	5192	1.11	0.304
76	50.6 22.	4 2.7 8	6 0.4	2	12.9	õ	0	ŏ	0	ŏ	Ō	0.5	0	ŏ	0	0.4	2049	1.35	0.471
17	59.5 15	7.2 5	9 3.6	0	8.2	0	0	0.3	0	0	0	0	0	0.3	0	0	1224	2.04	0.479
79	63.1 13.	5 3.1 7	0 2.2 1 1	3.4	6.1	U 0.3	0	0		0	U 07	0.2	0 2	0	0	0.6 1	2592	1.18	0.362
80	80 6.	1 3.3 0	6 1.	0	5	0.6	õ	ĭ.1	0.6	0.6	0	õ l	0	ŏ	õ	ō	720	1.88	0.218
81	84.2 2.	7 2.1 2	1 2.1	0	6.1	0,	0	0.3	0.3	0	0	0	0	0	0	0	2632	0.69	0.25
83	63 8.	1 1.7 4	2 18.2	ō	3.4	0.6	0	0.6	0.3	0.4	0	0	0	0	0	0	1428	1.2	0.288
84	47.6 8.	5 0.5 1	1 41.	0.2	0.5	0	0	0	0.2	0	0	0	0	0	0	0	1254	1.05	0.357
85	51.9 7	$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	2 38.8 1 32 1	1.2	0.4	1.2	0	0	0	0	U D	0	0	0	0	0	245	1.17	0.403
87	13 3.	5 0.9 0	46.1	. 0	Ó	3.5	0.9	22.6	7	õ	õ	ō.9	õ	õ	ŏ	ĭ.7	114	1.57	0.482
88	13.4 5.	$\frac{9}{8}$ $\frac{1.10}{2}$ $\frac{1}{1}$	72.6	0.5	1.1	1.6	0	2.2	1.6	0	0	0	0	0	0	0	186	1.01	0.305
90	10.4 6.	2 0 0	81.1	0.7	Ô	0.2	ō	ō	0.2	ŏ	0	ŏ	1	0	ō	0	408	0.69	0.285
91	7.1 5.	7 0 0.	82.2	0.4	0.7	0.7	0	0.7	1.8	0	0	0	0	0	0	0	281	0.74	0.234
92	10.4 4.	3 0.9 0	67.9	0.6	1.7	U.8 1.9	0	4.3	6	0	0	0	1.6	0	0	0.2	503 471	1.05	0.475
E	1	9	1														1		

Table 2 a

Relative frequencies of the planktic foraminiferal species (125-250 μ m) of the counted samples. Abbreviations: sn = sample number; sac = G. sacculifer; rub = G.ruber white; aeq = G. aequilateralis; ten = G. tenellus; bul = G. bulloides; rubs = G.rubescens; glu = G. glutinata; inf = Gt. inflata dut = N. dutertrei; inc = N. incompta; qui = G. quinqueloba; ru = G. ruber pink; sci = Gt. scitula; fal = G. falconensis; cra = Gt. crassaformis; con = G. conglobatus; pac = N. pachyderma; res = indet. specimens; hs = total number of specimens/sample; SW = Shannon-Wiener index; and BG = Buzas-Gibson index.

sn	sac	rub	aeq	ten	bu1	rus	glı	i inf	dut	inc	qui	rup	uni	men	sci	fal	trs	trđ	con	hír	pul	res	hs	SW	BG
$\begin{array}{c}1\\2\\3\\4\\5\\6\\7\\8\\9\\1\\1\\1\\2\\2\\2\\3\\4\\5\\6\\7\\8\\9\\0\\1\\1\\2\\2\\2\\3\\3\\4\\5\\5\\6\\7\\7\\8\\9\\0\\1\\2\\2\\2\\2\\3\\3\\4\\5\\5\\5\\6\\6\\6\\6\\6\\6\\6\\6\\6\\6\\6\\6\\6\\6\\6\\6$	$\begin{array}{c} 0\\ 0\\ 12.5\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\$	$\begin{array}{c} 0 \\ 0 \\ 16.7 \\ 60 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ $	$\begin{array}{c} 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 $				$\begin{array}{c} 9.1.\\ 9.1.\\ 9.1.\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\$	$ \begin{array}{c} 1,11\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\$	$\begin{array}{c} 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 $	$\begin{array}{c} 0.5\\ 25\\ 0\\ 20\\ 7.7\\ 19.7\\ 9.9\\ 0\\ 5.6\\ 0\\ 3.1\\ -\\ 0\\ 9.9\\ 39.1\\ 31.9\\ 32.4.6\\ 32.4\\ 3\\ 3.3\\ -\\ 40\\ 5.4\\ 2.8\\ 1.8\\ 3.3\\ 3\\ -\\ 40\\ 5.4\\ 4.9\\ 1.3\\ 3.3\\ 0\\ 11.6\\ 0\\ 5.4\\ 5.4\\ 1.6\\ 0\\ 3.7\\ -\\ 0\\ 1.5\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\$	$\begin{array}{c} \mathbf{q}_{11} \\ 0$	0 0 0 0 0 0 0 0 0 0 0 0 0 0	$\begin{array}{c} 0\\ 0\\ 13.3\\ 0\\ 7.7\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\$	$ \begin{array}{c} 1.8 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ $				$ \begin{bmatrix} 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0$	$\begin{array}{c} 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ 0\\ $	0 0 0	$\begin{array}{c} 0 \\ $	$\begin{array}{c} 1 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\$	$\begin{array}{c} 1.3\\ 40\\ 8\\ 8\\ 15\\ 5\\ 5\\ 6\\ 42\\ 174\\ 151\\ 151\\ 151\\ 151\\ 152\\ 2\\ 362\\ 2\\ 37\\ 142\\ 2\\ 2352\\ 2352\\ 1984\\ 490\\ 179\\ 34\\ 425\\ 1984\\ 190\\ 179\\ 33\\ 22\\ 37\\ 534\\ 4800\\ 179\\ 37\\ 534\\ 445\\ 150\\ 0\\ 0\\ 118\\ 168\\ 150\\ 0\\ 10\\ 118\\ 148\\ 150\\ 0\\ 10\\ 17\\ 118\\ 67\\ 7\\ 118\\ 144\\ 122\\ 445\\ 150\\ 10\\ 10\\ 22\\ 3\\ 3\\ 1\\ 10\\ 10\\ 22\\ 12\\ 12\\ 12\\ 12\\ 12\\ 12\\ 12\\ 12\\ 12$	$\begin{array}{c} - & & - & \\ & - & & & \\ & - & & & \\ & & & &$	- 0.806 0.862 0.467 0.727 0.583 0.472 0.553 0.472 0.556 0 0.427 0.556 0 0.427 0.556 0 0.427 0.556 0 0.427 0.437 0.778 0.437 0.778 0.437 0.778 0.437 0.437 0.434 0.947 0.556 0.326 0.361 0.652 0.361 0.556 0.361 0.556 0.366 0.361 0.552 0.361 0.556 0.366 0.366 0.566 0.366 0.662 0.544 0.556 0.662 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.566 0.681 0.749 0.415 0.566 0.881 0.566 0.866 0.866 0.866 0.866 0.866 0.945 1 0.943 1 0.943 1 0.943 1 0.943 1 0.943 1 0.943 1 0.943 1 0.943 1

Table 2 b

 $(250-500 \ \mu n)$ further idem a, except for the following abbreviations: rup = G.ruber pink; uni = o. universa; men = Gt. menardii; trs = Gt. truncatulinoides sinistral; trd = Gt. truncatulinoides dextral; hir = Gt. hirsuita; pul = P. obliquiloculata. Note the differences in headsums (hs).

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with the eastward AC (Siedler *et al.*, 1985). The AFZ separates southern regions with a permanent thermohaline mixed layer and a subtropical thermocline from northern regions with a seasonal thermocline formation, comprising water masses of the NAC, which show winter convective and advective mixing (Käse and Siedler, 1982). The AC branch of the Gulf Stream (Käse *et al.*, 1986; Käse and Zenk, 1987), also forms the northern boundary of the subtropical gyre, which encompasses the Subtropical Mode Water (McCartney and Talley, 1982; Käse *et al.*, 1985).

APNAP I HYDROGRAPHY

To determine the representativity of the APNAP I temperature data, we compared our data with temperature values derived from long-term mean August and September isotherms, averaged over a period of 25 years by Robinson *et al.*, 1979 (r = 0.98, *see* also Fig. 3 *a*, *c* and *d*).

APNAP I salinity data are compared with salinity values derived from interannual mean isohalines (Robinson et al., 1979 give no record of monthly mean isohalines), giving a lower correlation (r = 0.93) than temperature. Our August salinities are generally lower than the annual average, but still fall within the normal annual variation (see also Fig. 3b and e). These excellent correlations confirm that we sampled during a representative summer. Comparison of our patterns of surface water isotherms and isohalines (Fig. 3 a and 3 b) with the patterns of the mean August and September isotherms over the same period (Fig. 3 c and d) and the annual mean isohalines (Fig. 3 e; Robinson et al., 1979) reveals the following differences. From 47° N to 44° N our surface water isotherms show a sudden increase in density of the isolines absent from the Robinson et al., 1979 map. A north-south meridional isotherm cross-section along about $30^{\circ}W$ (Fig. 4 *a*) shows that this column down to 800 m, forming an uplift of deeper isolines. The





Figure 3

Isoline maps: a isotherms based on surface water temperature data of APNAP I; b isohalines based on surface water salinity data of APNAP I; c mean isotherms of August; d idem for September; e annual mean isohalines; c, d and e all after Robinson et al. (1979). Dots refer to APNAP measurements.





Isoline sections of: a temperature in °C; b salinity in practical units %; c density in t (after Ganssen et al., 1986). For station locations see Figure 1 a. isohalines at the surface do not give such a strong signal, but the isohaline meridional section (Fig. 4 b) does show a clear signal in the shape of a relatively fresh water lens. In the isopycnal section (Fig. 4 c), this feature is less evident, but still apparent. Because of its presence in the surface waters and the water column, we conclude this feature to be associated with a frontal zone, rather than a local eddy. At 34° N the isotherms show a second increase in density as well as a northwest deflection. The surface isohalines reflect the same feature, in a somewhat weaker form. Examination of the CTD N-S sections of temperature, salinity and density (Fig. 4) reveals a sharp downward bend in the isolines, indicating a frontal zone. This frontal zone coincides with the Subtropical Front, also referred to as the AFZ (Siedler *et al.*, 1985), associated with the AC. The observed deflection at the surface may indicate meandering or may be an effect of the 8-day time lag between our measurements. At about 51°N the isotherms have a low gradient in map view but only the CTD sections show a sudden dip of the isolines, which is interpreted as the boundary between the SW and the NAC. The observed features are in general less evident in the upper 25 m of water, because of superimposed seasonal surface warming or mixing. Our observations are summarized in a hydrographic interpretation chart for this area, showing the position of the most evident water masses during APNAP I (Fig. 5).



Figure 5

Hydrographic interpretation of the studied area for August and September 1986, based on both surface water and CTD data. Approximate current velocities are for NATW after Krauss (1986) and for the NAC and AC after Krauss (1986) and Käse and Siedler (1982).

Planktic foraminiferal assemblages

Weighted pair group cluster analyses on samples (fraction 125-250 μ m, Fig. 6 *a* and *b*) and curves of relative species frequencies (Fig. 7 *a* and *b*) are intended to describe the possible relationship between distributional patterns of foraminifera and water masses. We excluded species with a lower frequency than 2 % from all analyses. Water masses recorded in the studied area are not restricted to a certain latitude, but also

extend longitudinally. Therefore analyses were executed on two sets of data. Samples collected along the meridional section will give sharper boundaries (Fig. 1 b) and an improved correlation at subcluster level (Fig. 6 b). This is because the meridional section crosses most of the water masses obliquely and water mass boundaries tend to fade out towards the east (Krauss and Käse, 1984). The total number of samples shows possible longitudinal shifts of the faunal boundaries towards the east. The four conspicuous sample groupings, shown by



Figure 6

Weighted pair group cluster analysis on counted planktic foraminiferal samples, based on a correlation matrix of relative frequencies, excluding species with a relative abundance lower than 2%.

Both sample sets show at least four clusters; in cluster II both sample sets give a coincident subdivision.



Figure 7

Relative frequencies of planktic foraminifera from surface water along the sample track (see also Fig. 1 b). a) 125-250 μ m sieve fraction; b) 250-500 μ m sieve fraction. G. ruber=G. ruber white. AFZ = Azores Frontal Zone. For further explanations of the water mass abbreviations see Figure 5. N* and S* indicate respectively the northernmost and southernmost samples. Full, half and new moon are indicated at the top of the figures.

cluster analysis of the two data sets (assemblage I-IV; Fig. 6 a and b) almost coincide, when plotted on the map, although a shift of two samples is present at some boundaries (Fig. 8). Cluster analysis on species with a higher relative frequency than 5 % (Fig. 9) was carried out on both size fractions. The coarse fraction gives lower correlations than the fine fraction due to the occasionally low frequencies in the former one and the use of relative frequencies. The fine fraction (Fig. 9 a) clusters in four assemblages, which correspond to the four groupings in the dendrogram of the coarse fraction, but contain different species (Fig. 9 b). Some artificial subclusters also occur, such as a subcluster comprising Globigerina bulloides d'Orbigny and G. quinqueloba Natland together with a subcluster of Globigerinoides ruber (d'Orbigny) pink variety and Globorotalia truncatulinoides (d'Orbigny) sinistral. If the curves of the relative frequencies are examined, it becomes apparent that the species confining these subclusters are, from an ecological point of view, in no way closely related (Bé, 1977). The only two samples containing the latter two have very low absolute frequencies causing artificial relative frequencies and subsequently subclustering. To relate species clusters to sample clusters, curves of the relative frequencies along the sample track (Fig. 7), are examined. This results in specific planktic foraminiferal assemblages (Tab. 3).

Watermasses

SPECIES DISTRIBUTION AND HYDROGRAPHY

The four major faunal assemblages correspond to, but do not always completely coincide with, the water masses defined hydrographically in terms of temperature and



Faunal boundaries, based on cluster analyses of the samples. The two cluster analyses (see also Fig. 6) do not always give coincident sample assemblages; at the boundary a shift of two samples is sometimes present. Solid lines indicate boundaries based on all samples; dotted lines indicate boundaries on the meridional sections; dashed line is based on a subcluster present in both sample sets. For further explanation of assemblages 1-4 see Table 3.

		species	clusters	sample clusters	
(640N)	HYDROGRAPHIC UNITS used in this paper	125-250 um SIEVE FRACTION	250-500 um SIEVE FRACTION	ASSEMBLAGE	
	SUBPOLAR	G. quinqueloba	G. bulloides G. quinqueloba	I	- (640N)
520N	NORTH ATLANTIC CURRENT	G. bulloides Gt. inflata N. incompta	N. incompta Gt. inflata	11	
340N	NORTH ATLANTIC TRANSITIONAL WATER	G. sacculifer G. aequilateralis	G. sacculifer O. universa	111	
(15ºN)	AZORES CURRENT (subtropical)	G. ruber G. glutinata G. tenellus Gt. scitula G. rubescens N. dutertrei	G. ruber N. dutertrei Gt. menardii G. aequilateralis G. conglobatus P. obliquelocelata	IV	(15°N)

Table 3

Relation of specific planktic foraminiferal assemblages I-IV, derived from sample clusters, with species clusters and described water masses for August and September 1986. No polar water is encountered during APNAP I cruise (latitudes at boundaries are only valid along the eastern track).



Figure 9

Weighted pair group cluster analysis on species, based on a correlation matrix of relative frequencies. Species with a lower relative abundance then 5 % are excluded.

a) dendrogram of the 125-250 µm fraction; b) dendrogram of the 250-500 µm fraction. Both sieve fractions show four clusters, although the coarse fraction shows artificial subclustering (see text).

salinity (Wüst, 1935; Williams et al., 1968 b; compare Fig. 5 and 8).

For example, the sharp faunal boundary between assemblage I and II at 52° N corresponds largely to the turnover of Subpolar Water and the North Atlantic Current (NAC) at about 51° N. North of this boundary the foraminifers Globigerina auinqueloba and only large G. bulloides are the frequent species; south of this the species Neogloboquadrina incompta (Cifelli), Globorotalia inflata (d'Orbigny) and G. bulloides are most abundant (Tab. 3). At 47° N a weak boundary can also be seen, considering the clustered samples, although this boundary has no clearly defined hydrographical counterpart. This feature, which cannot be traced in clusters based on species, possibly corresponds to the northern part of the frontal zone observed at the same latitude. A more evident faunal boundary is shown at 44°N along with the transition between the NAC and the North Atlantic Transitional Water (NATW, at about 45°N), forming also the southern part of the frontal zone. The NATW is characterized by an assemblage of the species Globigerinoides sacculifer (Brady), Globigerinella aequilateralis (Brady) and to a lesser extent Orbulina universa (d'Orbigny). Around 34°N a distinct faunal boundary is present, coinciding with a strong frontal zone (34-36°N), which confines the Azores Current (AC) to the North. The AC is characterized by a high number of species, with Globigerinoides ruber, Globigerinita glutinata (Egger), Globigerinoides tenellus Parker, Globigerina rubescens Hofker, and Globorotalia scitula (Brady) as dominant species in the fine fraction .



In the coarse fraction the AC contains the species G. ruber, Neogloboquadrina dutertrei (d'Orbigny) and to a lesser extent G. aequilateralis, Globorotalia menardii (Parker, Jones and Brady), Globigerinoides conglobatus (Brady) and Pulleniatina obliquiloculata (Parker and Jones).

Information indices and hydrography

Information indices add valuable water mass characteristics to those already defined on foraminiferal assemblages and physico-chemical environmental properties. As expected (Shih, 1979; Pianka, 1966; Pielou, 1975; 1979; van Soest, 1979) the contoured Shannon-Wiener index (Fig. 10 a and b) shows that in general the diversity increases southward and westward (see also the diversity curves of Fig. 12). The coarse fraction (250-500 µm, Fig. 10 b) has more variation in diversity than the fine fraction (125-250 µm, Fig 10 a), because the Shannon-Wiener index takes both the number of species and the relative abundance of each species into consideration. Since the coarse fraction has very variable frequencies, although the number of species remains equal, this results in artificially variable diversities. The Subpolar Water contains a low diverse planktic fauna. In the NAC the diversity increases only slightly southward especially in the coarser fraction. This low diversity can be related to seasonal variability of the environment and periodic high nutrient supplies (see McArthur, 1965; Hallock, 1987). The NATW shows a



August-September 1986 planktic foraminiferal diversities (Shannon-Wiener index) contoured at a 0.2 interval. a) 125-250 µm sieve fraction ; b) 250-500 µm sieve fraction.

Also shown are sample track and water mass boundaries; main boundaries are indicated by solid arrows, open arrows indicate boundaries only present in the meridional samples (for legend see Fig. 5). Note the maximum diversity at the FZ north of the AC.

Figure 10

relatively high diversity in the west, decreasing rapidly eastward and being highest towards its boundaries, pointing to broad mixing zones between the water masses. Diversity in the AC is very high although it decreases south of 31°N and strongly towards the east. Equitability contours after the Buzas-Gibson index (Fig. 11) have a more complex pattern (see also the equitablity values plotted along the sample track, Fig. 12). Equitability in the Subpolar Water varies from intermediate to high towards the boundaries. No dominant species are present. The NAC has an intermediate to high equitability, especially along the boundaries, decreasing eastward. The NATW has an intermediate to low equitability. The AC shows an intermediate equitability strongly decreasing southward. When the diversity and equitability curves are compared (Fig. 12), all show in their eastern part (sample numbers 65-93) scattered values. This can be attributed to the divergence of current systems towards the eastern boundary of the Northeast Atlantic basin (Krauss, 1986) with mixed faunas of the different water masses as a result.

DISCUSSION

It is shown that water masses can be characterized on the basis of the relative frequencies of planktic foraminifera established by cluster analyses and information indices related to measured hydrographical parameters. Differences within faunal water masses still appear to be considerable in comparison with the differences between the water masses. Samples are relatively easy to cluster; however the species do not cluster mutually to any great extent. This is especially accentuated in the southern watermasses with their low absolute frequencies and their high number of species, wich may be attributed in part to the result of dense sampling in which small scale temporal variations will also be recorded.

In addition, the southern currents are more irregular than their northern counterparts, causing less distinct water masses (Weyl, 1978) and are therefore more difficult to characterize. Vertical migration will further alter the



Figure 11

August-September 1986 planktic foraminiferal equitabilities (Buzas-Gibson index) contoured at a 0.1 interval for the 125-250 µm sieve fraction only. Also shown are sample track and water mass boundaries (see also Fig. 12).



Figure 12

Diversity (S/W:125-250 μ m and 250-500 μ m) and equitability (B/G:125-250 μ m) plotted along the sample track (see also Fig. 1 b). Dots (diversity) and squares (equitability) refer to actual measurements; the solid line, a running average, shows the general trend. Breaks in the direction of the sample track are indicated by NE, N, and SW at the basis of the figure (for the abbreviations of water masses see Fig. 2). Note the high diversity at the NATW-AC boundary and its relatively low equitability.

planktic patterns, because we sampled only the upper 5 m of the surface water at different times.

Broad mixing zones along water masses and especially along frontal zones are reflected in the diversity gradients. These frontal zones are hydrographically distinct water mass boundaries, where species of the adjacent water masses are mixed by converging surface waters, giving rise to a high diversity. But, since mixing also leads to "artificially" strong dominance of those species that occur in both water masses, other than normal (Lloyd and Ghelardi, 1964) this high diversity is associated with a low to intermediate equitability. Mixing also enhances primary productivity, increased turnover rates (Fournier, 1978; Tranter et al., 1983; Angel, 1989) and mortality within displaced foraminiferal populations (Berger, 1971) resulting in a high sediment output, which can characterize frontal zones in the sediments. The high diversity of NATW and the AC probably has its origins in the Northern Sargasso Water (Cifelli and Smith, 1974) and is carried eastward in the AC. When environmental stress becomes too high, part of the coexisting species will disappear. The Sargasso Sea has stable, oligotrophic waters, which are more transparent than eutrophic waters. This results in a broad photic zone with,



Table 4

Climap assemblages related to APNAP I hydrographic units and assemblages. See also Table 3.

seasonally, almost uniform conditions with permit a higher degree of specialization (Hallock, 1987) and therefore a higher diversity. The very low diversity in the northeast might be the result of high nutrient fluxes from the nearby European shelf waters.

The planktic foraminiferal assemblages presented in this paper, characterizing hydrographically defined water masses, suggest that an adjustment of the biogeographic assemblages for zooplankton of CLIMAP, 1981 (Tab. 4), which are supposed to "closely follow the current systems and surface water masses", is called for. CLIMAP's subpolar and transitional/subtropical assemblages coincide with our SW, NAC and NATW. Their eastern tropical assemblage is partially represented in our AC, although we consider these waters to be of subtropical origin. Equitability, being part of the general diversity index, has been related to resource limitations (Pulliam et al., 1967) and tentatively to water mass properties and seasonal movements of foraminiferal populations (Balsam et al., 1980). Further, mixing of populations can be traced by a lower equitability than one would expect, given the high diversity of the sample and following examination of the surrounding environments (Gibson and Buzas, 1973). Our observations confirm this. The NATW is such an example of a water mass comprising mixed populations (i.e. of NAC and AC), shown by its low equitability coupled to a high diversity. Frontal zones can also be recognized by mixed populations. Equitability should, however, not be interpreted without diversity and vice versa. The discussed information indices also show potential for interpreting sediments, such as the recognition of paleo-frontal zones. Such interpretations should also take account of the fact that taphonomic processes and time averaging will alter the original information of living populations.

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