
Fuzzy Logic Applied to Track Generation Areas of Swell Systems Observed by SAR

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Abstract:

Recently, with the availability of a great number of synthetic aperture radar (SAR) wave mode spectra, it has been possible to derive a set of great circle lines of swell propagation whose intersection points indicate the position of the storm generating the observed swell field. However, due to the inherent limitations of SAR spectra, the locus of convergence of great circle of swell propagation can be sometimes diffuse or contain multiple convergence regions. In this letter, we adapted the fuzzy cluster logic method to identify the regions of convergence of SAR wave field rays. The analysis of the results of the fuzzy algorithm clearly indicates the ability of this statistical method to identify the cluster center region of swell fields observed in SAR wave mode images. The measure of success of the method was how well the generation center of the swell could be traced back to an existing strong storm system.

Keywords: Swell wave fields and fuzzy logic ; synthetic aperture radar (SAR)

1. Introduction

The wave field generated by severe storms are difficult to be accurately modeled [1]. From strong storms centers, swell systems propagate for long distances showing very little dissipation [2]. With the possibility of using sequences of directional wave spectra obtained from space-borne Synthetic Aperture Radar (SAR) Wave Mode images it is now possible to derive routinely a set of great circle lines of swell propagation. The intersection points of these lines can be used to indicate the possible region of the storm that generated the observed swell field [3], [4]. However, due to inherent uncertainties and limitations of a SAR spectra such as the non-linear mapping of sea surface wave spectrum into a SAR image spectrum [5], [6], the method used for determining SAR spectrum peak

24 period and wave propagation direction, among others, the locus of convergence of great circle of swell propagation
25 can sometimes be diffuse, or contain multiple convergence regions.

26 Considering the inherent inaccuracies of the data, the mathematical determination of the “best” convergence
27 region of a large number of great circle lines of swell propagation is not simple. The statistical approach of fuzzy
28 cluster logic is here proposed as a possible and novel solution for this problem. In this study, we adapted an
29 unsupervised Fuzzy Cluster Logic (FCL) method to identify the regions of convergence of the SAR wave field
30 rays. The FCL method is normally used in digital image pattern recognition, but it has also been applied for the
31 analysis of oceanic satellite data [7].

32 We present below an analysis of a series of ASAR/Envisat wave mode spectra for swell propagation in the
33 Atlantic ocean using the FCL method with the objective of showing its efficacy to determine the source regions of
34 observed swell wave fields.

35 We want to show that the proposed methodology is simple to be implemented, robust and can be a powerful tool
36 to identify the swell generation regions for large oceanic areas. The results obtained are illustrated for a few case
37 studies in which we compare the SAR located storm source regions with scatterometer wind (QuikScat) data.

38 II. DATA AND METHODS

39 A. ASAR Wave Mode data

40 The European Space Agency (ESA) ENVISAT ASAR derived spectra used in this paper are Level 2 products
41 provided by the Centre ERS d'Archivage et de Traitement (CERSAT) at Ifremer (France). The processing methodology
42 used to derive the wave spectra is described in detail by [8] [9].

43 The ASAR Wave Mode Level 2 directional spectra are given on a log-polar grid in wavenumber and direction
44 domains, $F(k, \phi)$. Nominal directional resolution is 10° and wavenumbers are given in logarithmic scale corre-
45 sponding to a nominal wavelength range between 30 and 800m [9]. The frequency $F(f)$ and directional spectra,
46 $\Psi(\phi)$ can then derived from the Level 2 spectra product. We used in this investigation a set of 10640 ASAR
47 directional spectra covering the period 2003 to 2009 and for the Atlantic sector between 10°S 60°N and 10°E
48 80°W . The spectra are provided in a grid of 100 km in the along track direction and 7° in the longitudinal direction.
49 The geophysical validation of the satellite wave fields shows good agreement in terms of statistical wave parameters
50 in comparison with WAM model [10].

51 B. Scatterometer data

52 The vector surface winds derived from the QuikScat satellite SeaWinds scatterometer used in this study were
53 also provided by CERSAT and are retrieved from scatterometer measurements and interpolated in space and time
54 as regular gridded fields at a spatial resolution of 25×25 km and daily mean. The QuikScat provides neutral
55 equivalent wind vectors (speed and direction) at a reference level of 10 m above sea surface [11]. All QuikScat
56 wind data used in this paper correspond to L2b products processed routinely by Jet Propulsion Laboratory (JPL).

57 The accuracy of the QuikScat winds has been investigated in a number of papers through comparisons against
 58 in-situ wind measurements (buoy and ships) [12], [13], [14], [15], [16]. The correlation coefficients between satellite
 59 and in-situ data for wind speeds and directions exceed 0.90 and 0.88 for wind speeds in the range of 0 to 30 ms⁻¹
 60 [17], respectively. The associated rms differences are about 1 ms⁻¹ and 18°.

61 C. Fuzzy C-Means Analysis (Fuzzy Logic Clustering)

62 In contrast to classical set theory where an element either belongs or not to a set, fuzzy set theory allows an
 63 element to have membership to one or more sets. [18] introduced fuzzy set theory in a mathematical context to
 64 handle imprecise information. In this context it is possible to extend the concepts of classical set theory allowing
 65 for partial or intermediate set membership values. It also allows for full (“crisp” or “hard”) memberships, being
 66 therefore a superset of classical set theory.

67 If we consider one universe that contains elements x that follow a rule given by $S=\{x\}$, under normal set theory
 68 the membership function $f_A(x)$, is either 1 or 0, if x belongs or not to a subset of A of S , Under fuzzy set theory,
 69 the membership function is altered to allow for graded memberships and is transformed into $0 \leq f_i(x) \leq 1$. Thus
 70 it is possible to obtain the partial set membership for each fuzzy set. In this sense $f_i(x)$ represents the probability
 71 that x belongs to the i th set. If there are c possible sets, then;

$$\sum_{i=1}^c f_i(x) = 1 \quad (1)$$

72 The Fuzzy C-Means (FCM) algorithm used here is a fuzzy technique developed by [19] and improved by [20].
 73 Frequently used in pattern recognition, the fuzzy technique is a method of clustering which allows one piece of
 74 data to belong to two or more clusters. In the most general case, clusters are defined as groups of points that are
 75 similar according to some measure of similarity. Usually, similarity is defined as proximity of the points according
 76 to a distance function.

77 Clusters representing one specific class in which each member has full membership, that is, they do not belong
 78 to any other class, are called discontinuous or discrete classes. On the other hand, classes in which each member
 79 may belong in some extent to every cluster or a partition are called continuous classes [21]. These classes are
 80 a generalization of the fuzzy set theory, and may have membership values ranging between 0 and 1, where 1
 81 represents the membership cluster center and 0 represents complete dissociation to the cluster.

82 The FCM algorithm attempts to minimize the objective function J_m ;

$$J_m = \sum_{j=1}^{N_s} \sum_{i=1}^{N_c} \mu_{ij}^m d^2(X_j V_i) \quad (2)$$

83 where: N_s is the total number of observations j in the data set; N_c is the number of separate clusters, i ; d is the
 84 Euclidean distance between an observation vector X_j and cluster center V_i ; μ_{ij} is the membership value of the
 85 j^{th} observation to the i^{th} cluster (a value between 0 and 1); and m is the weighting exponent (initially set to 2).
 86 High membership values indicates strong likelihood for the studied variable to be present at the cluster. The FCM

87 method allows a parameter to be part of more than one cluster. In our work we used only two data class clusters
88 (1 - a storm generation region and 2 - not a generation region).

89 *D. Swell Source Identification by Back Propagation*

90 From the ASAR wave mode product one can estimate the wave spectrum at a specific region, the peak period
91 and direction of propagation of the dominant wave field. From this information it is possible to use the physical
92 properties of the swell waves to propagate them backwards in time to estimate their generation region. [22] and [4]
93 showed that this methodology allows the tracking of swell fields in a range of 3 to 10 days. These studies have
94 demonstrated that the combination of SAR images at different times and places allows for the prediction of arrivals
95 of gravity waves in coastal regions, and to monitor the propagation patterns of wave fields.

96 Starting from a initial geographic position ϕ_0, λ_0 (longitude and latitude) and the direction of propagation θ_o ,
97 as obtained from a SAR wave directional spectrum, the swell field great circle trajectory in deep water can be
98 evaluated at a distance X and at a time t by $X = (t - t_o) \cdot C_g = (t - t_o) \cdot g / (4\pi f)$, where C_g is the deep water
99 group velocity (estimated from T_p), f is the spectra peak frequency and t_o is the time of SAR observation. This
100 corresponds to a spherical distance $\alpha = X/R$ along the great circle, where R is the radius of the Earth. Figure 1
101 indicates the complexity one is faced when trying to identify the generation zone of a particular swell field using
102 the backward propagation from a series of observed wave mode image spectra.

103 Here we use sequences of SAR Wave Mode spectra solely to identify the formation region of observed swell
104 fields. For this, we generate from each Wave Mode spectra the backward trajectory using the great circle approach.
105 The probable location of swell generation is defined as the convergence zone of the different trajectories for similar
106 times.

107 We initiate the processing using all satellite tracks passing through the study region within an established time
108 interval; tests indicated that 3 days was the best choice. Thus, all ASAR Wave Mode spectra present in the CERSAT
109 database for the study region and for the chosen three day period were included for each specific case study. In each
110 case analyzed, between 250-400 ASAR wave spectra were used. The initial pre-processing included the extraction
111 from each spectra its geographical location, the peak period and propagation direction. In the sequence the back
112 propagation of each swell field was done using the above described methodology. In this study we choose a
113 backward propagation time step of 2 hours. Thus for each spectra, a set of propagation points is generated. All
114 trajectories coincident points (intersection points) were captured along all the back propagation fields. Normally,
115 this resulted in a set of about 3×10^3 intersection points per case study. A small fraction of the derived trajectories
116 do not converge, however, to the same regions. These isolated intersection points, normally associated with higher
117 frequency wind-sea waves, were considered spurious for the swell storm generation analysis and were discarded.

118 The collected coincident points, X_j in equation (2), were then used in the fuzzy C-means algorithm to determine
119 the spatial clusters. We used in this study three variables: latitude, longitude and time, of each back propagating
120 swell great circle points. After testing different cluster numbers (N_c in equation (2)) we observed that best results,
121 i.e., better defined clusters, were obtained using two clusters. The cluster that shows the highest convergence of

122 points in space is assumed corresponding to the swell generation region. The second cluster (normally with a much
123 more diffuse spreading of points) is discarded in the analysis.

124 As explained above, the membership values are allowed to vary from 0 to 1 for each cluster; the maximum
125 value corresponding to the center of the cluster. We take the region next to the center of the cluster as the swell
126 generation region. We have set this region being limited by membership values between 0.95 and 1. All other points
127 corresponding to smaller memberships were filtered out.

128 To verify the consistency of the fuzzy model we plotted the estimated swell generation region ($0.95 \leq f_i(x) \leq 1$)
129 over a map of corresponding scatterometer wind data. The presence of strong winds or storm center coincident
130 with fuzzy derived cluster analysis was used as a validation of the method.

131 III. RESULTS AND DISCUSSION

132 From the Level 2 ASAR directional spectra available at CERSAT for the years 2003-2009, for the study region,
133 we were able to test the FCM method for almost 40 cases. The main motivation was to verify whether the fuzzy
134 clusters determined in an unsupervised algorithm had a consistency when compared against strong wind generation
135 areas, possible candidates responsible for the observed swell fields. In the following we present and discuss four
136 of these case studies.

137 Figure 2 shows the swell back propagation intersection points and their membership values relative to cluster 1
138 (in color) obtained from ASAR spectra from days 25, 26 and 27 January, 2003. The intersection points corresponds
139 approximately to a decimation of 50% of total number of back propagation points analyzed. Note that a well defined
140 region, with membership values near 1, is identified by the fuzzy algorithm. We claim that this high membership
141 zone can be used as the generation region of the observed swell. The high membership region corresponds to
142 January 24, 2003.

143 Figure 3 shows the results obtained by the method for January, 2003 (described above), February 2004, March
144 2005 and January 2008, where regions of high membership values (above 0.95) are superimposed over QuikScat
145 wind fields. A clearer delimitation of the swell source region is obtained by filtering out all points with membership
146 values less than 0.95. Observe that the fuzzy centers are located at oceanic regions showing very high winds, ranging
147 from 22 m s^{-1} for case B to 26 m s^{-1} for case C. For cases A and D, the fuzzy centers were located almost on
148 top of a closed storm center. For cases B and C, although not coincident with storm center, the fuzzy centers were
149 located near maximum winds. Table 1 shows the distances of the fuzzy centers (membership = 1) and the storm
150 centers given by scatterometer wind data. Note that the greatest distance observed between the databases is 3.5° .
151 We noticed in analyzing all 40 cases that fuzzy cluster maximum is able to precisely pinpoint storm center when it
152 is sharply defined. When storm center is more diffuse the fuzzy algorithm is still, in the majority of cases capable
153 of capturing its neighboring area. Another possible cause of fuzzy center being displaced to observed QuikScat
154 storm center is the rapid displacement of some of these storms.

155 Long period and high significant wave height swell systems have been observed in the Norwegian Sea shelf
156 region having their origin in strong storms present in the northwestern Atlantic [23]. Normally these swell events

157 were observed being generated by intense extra tropical cyclonic storms that move rapidly towards northeast. The
158 cases presented before fit to this scenario, that is, the SAR observed swell systems were generated in the NW
159 Atlantic by strong low pressure systems moving from SW to NE. [24] shows strong wave fields associated to the
160 climatology of local NH winter storms. High winds associated with the low pressure systems passing through the
161 mid-latitude storm tracks generate large wind waves with high peak phase speeds. Normally, the result of this is a
162 southward propagation of swell systems out of the storm tracks during the NH winter [24]. Throughout the year
163 the high winds along the North Atlantic storm tracks are more intense in January, February and March. Therefore,
164 it is expected that swell originates predominantly from these regions during this period [25].

165 IV. CONCLUSIONS

166 The analysis of the results of the fuzzy algorithm applied to our database of SAR wave spectra clearly indicates
167 the ability of the unsupervised fuzzy statistical method to identify the cluster center regions of swell fields observed
168 in SAR wave mode images. The measure of success of the method was how well the generation center of the
169 swell could be traced back to an existing strong storm system. Out of a total of about 40 case studies for the years
170 2003 and 2009, the fuzzy algorithm showed good results for about 90% of the tests cases. The unsuccessful cases
171 seemed to be related to high frequency wind-sea dominant cases and very little swell trajectory convergence.

172 It should be noted, however, that for cases where only weak weather systems are present, it can happen that the
173 intersection points of great circle propagation are rather diffuse and the algorithm is not capable of converging to
174 a tight generation region.

175 Since coincident points used for the definition of the fuzzy center include measurements at different times for
176 the same day, the results obtained are suitable only for a daily average estimate, i.e., the storm center obtained
177 represents an average position of the storm during the day.

178 The accuracy of the method is expected to be much better for swell waves than for wind sea systems since SAR
179 wave mode spectra is much more reliable under the linear imaging mechanism in the long wave spectral domain
180 [26], [8]. However, considering that swell systems are much more predominant in the global ocean [27], with the
181 few exceptions of enclosed or semi-enclosed seas [24], the proposed method can be a useful tool to study wave
182 propagation systems.

183 Accurate analysis of swell propagation is an important factor in evaluating potentially hazardous conditions for
184 ocean and coastal zones. To our knowledge this is the first application of fuzzy methodology to identifying swell
185 storm generation regions. The method here proposed can be implemented to run in a completely autonomous mode
186 in a normal desktop computer in less than one hour per case. This method can be used to check the origin of strong
187 long period waves and additionally as a tool to validate weather forecast and wave models accuracies. If sufficiently
188 long series of SAR wave mode spectra are available, the application of the method can be used to improve the
189 climatology of swell generating regions in the global ocean.

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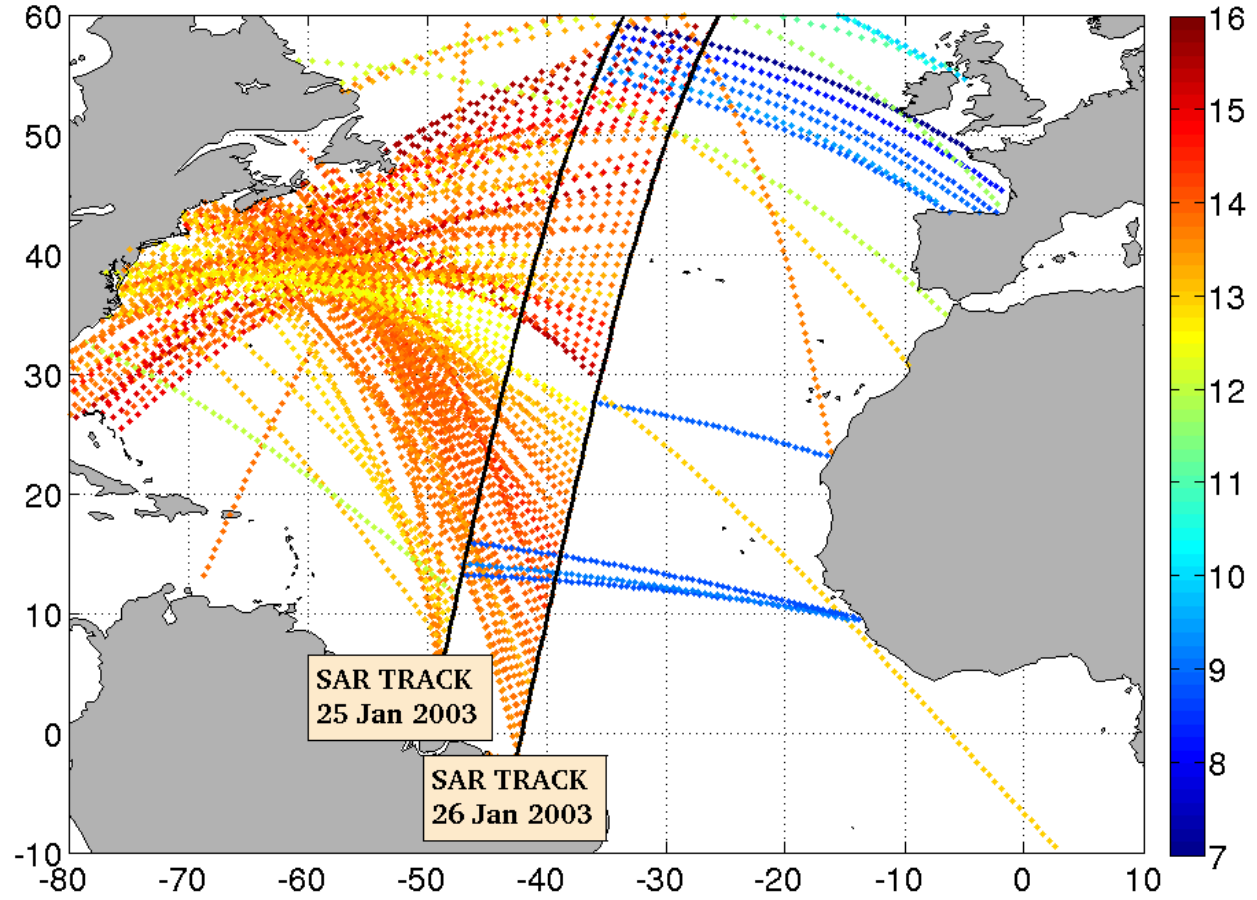


Fig. 1. Map showing the back propagating points and respective periods using two SAR tracks for our study region. The period of the wave field (in seconds) is indicated by the color bar. The lines in black represent the ASAR tracks for Jan. 25 and 26, 2003

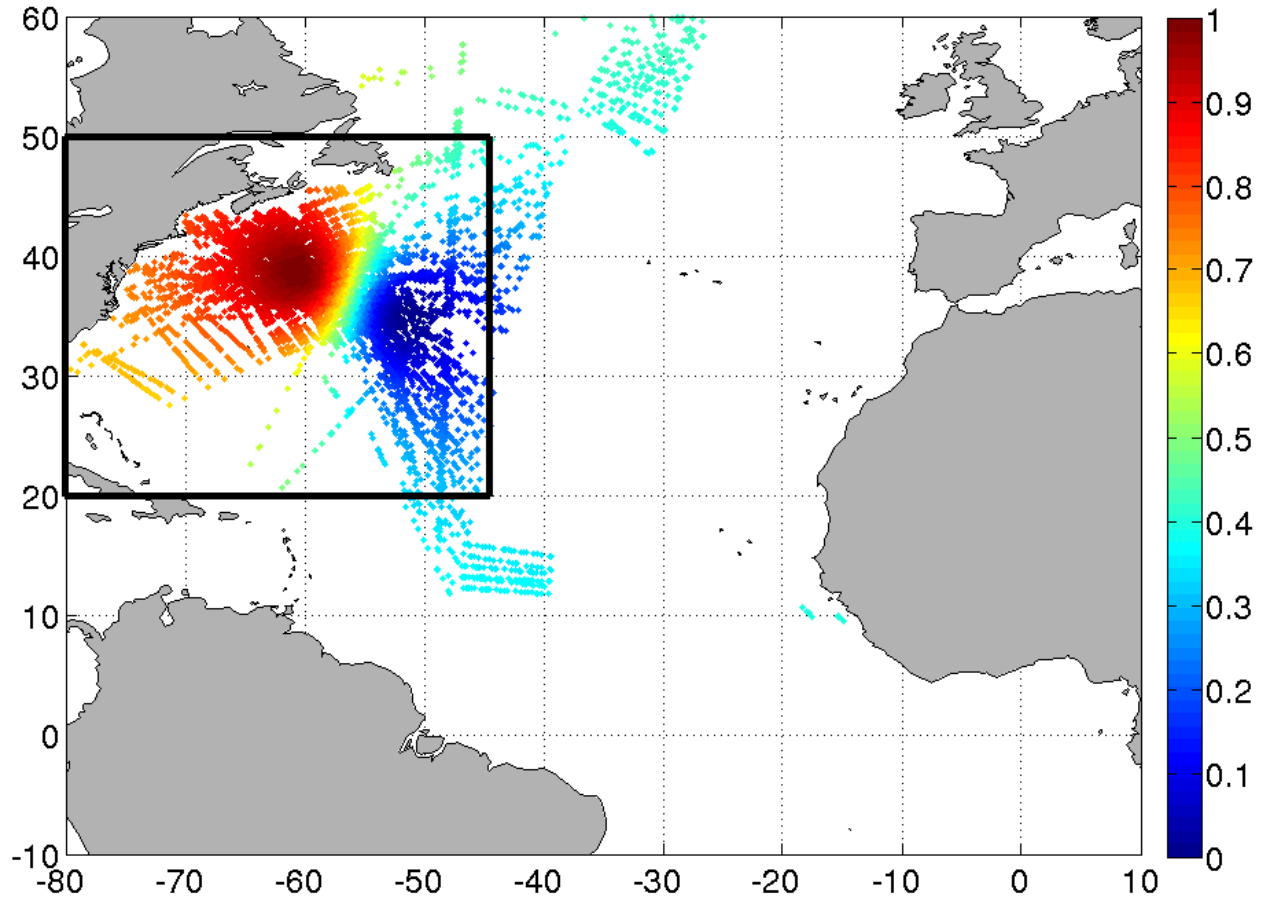


Fig. 2. Map showing the membership values of the swell back propagation intersection points, for the ASAR wave spectra from Jan. 25-27, 2003. The membership values are indicated by the color bar at right. The black rectangle represents the sub-region (45° to 80° W and 20° to 50° N) of swell convergence rays.

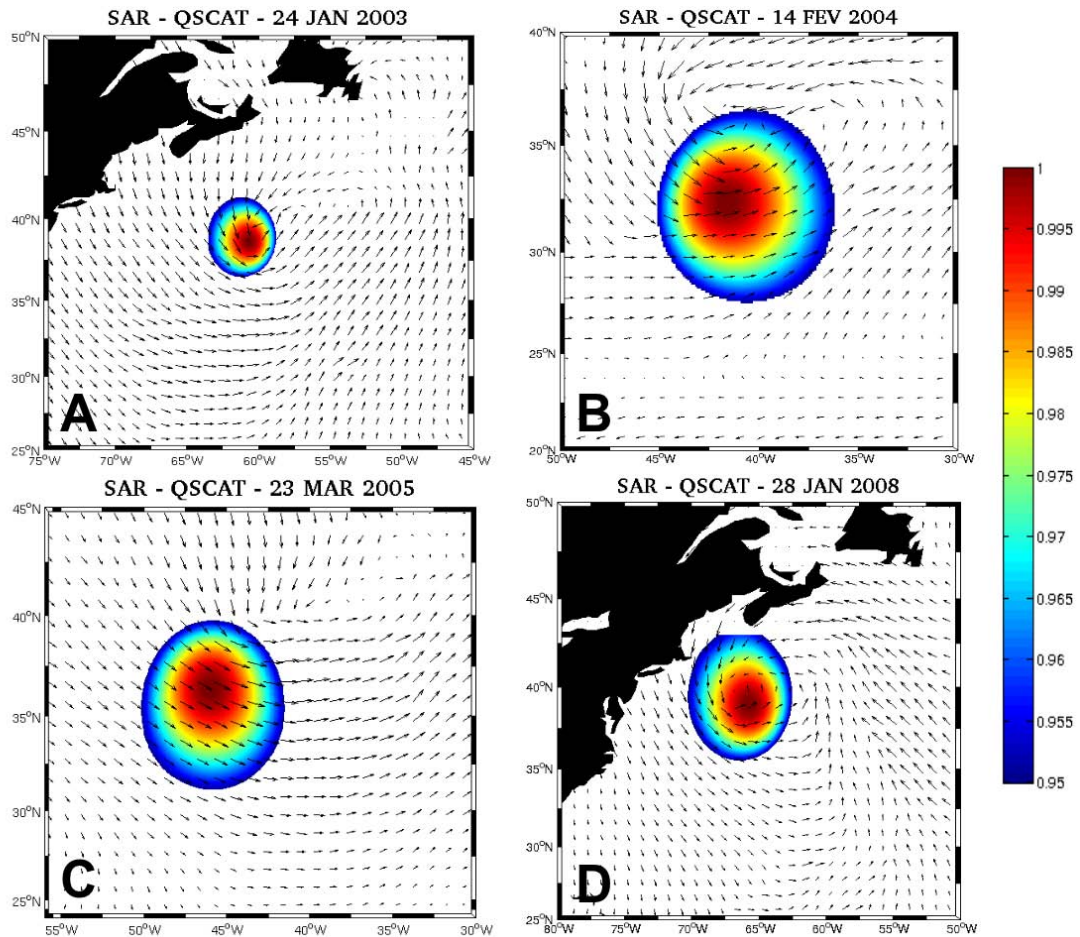


Fig. 3. Maps of QuikScat wind fields and fuzzy clusters: (A) January 24, 2003; (B) February 14, 2004; (C) March 23, 2005, and (D) January 28, 2008. Maximum QuikScat wind intensity for each case was: (A) 23 m s^{-1} ; (B) 22 m s^{-1} ; (C) 26 m s^{-1} , and (D) 25 m s^{-1} . The membership values are indicated by the color bar at right.

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I DETECTION PERFORMANCE 13

TABLE I
DETECTION PERFORMANCE

Storm ocurrence	Distance between fuzzy center and the low pressure center
January 24, 2003	0°
February 14, 2004	3°
March 23, 2005	3.5°
January 28, 2008	0.7°