

---

## On Dual Co-Polarized SAR Measurements of the Ocean Surface

V.N. Kudryavtsev<sup>a</sup>, B. Chapron<sup>b</sup>, A.G. Myasoedov<sup>a</sup>, F. Collard<sup>c</sup> and J.A. Johannessen<sup>d</sup>

<sup>a</sup> Satellite Oceanogr. Lab. (SOLab), Russian State Hydrometeorological Univ., St. Petersburg, Russia

<sup>b</sup> Institut Français de Recherche pour l'Exploitation de la Mer, 29280 Plouzané, France.

<sup>c</sup> Division of Radar Applications, CLS, 29280 Plouzané, France.

<sup>d</sup> Nansen Environmental and Remote Sensing Center, 5006 Bergen, Norway.

\*: Corresponding author : V.N. Kudryavtsev, email address : e-mail: [kudr@rshu.ru](mailto:kudr@rshu.ru)

---

### Abstract:

An effective methodology using satellite high-resolution polarized information to interpret and quantitatively assess various surface ocean phenomena is suggested. Using a sample RADARSAT-2 quad-polarization ocean synthetic aperture radar (SAR) scene, the dual co-polarization (VV and HH) radar data are combined into polarization difference, polarization ratio, and nonpolarized components. As demonstrated, these field quantities provide means to distinguish Bragg scattering mechanism and radar returns from breaking waves. As shown, quantitative characteristics of the surface manifestation of ocean currents, slicks, and wind field features in these dual co-polarization properties are very different and may be effectively used in the development of new SAR detection and discrimination algorithms.

**Keywords:** Radar cross sections ; radar signal analysis ; synthetic aperture radar (SAR) imaging

### I. Introduction

---

The goal of this study is to promote a very effective methodology using satellite high resolution dual copolarized information to interpret and quantitatively assess various surface ocean phenomena. Spaceborne synthetic aperture radar (SAR) has already proven to be a very useful tool to assess and reveal various ocean-atmosphere processes (see, e.g., [1]). Accordingly, methods have been demonstrated to advance quantitative interpretations of the SAR image intensity contrasts associated with local ocean surface roughness variations linked to changes in the near-surface winds, waves, and currents, as well as the presence of surface contaminants.

With the advent of new SAR instrument systems, dual- and quad-polarization SAR data are now made available, to possibly yield more useful information than conventional single polarization SAR observations. This can help to possibly go beyond the present geophysical retrieval algorithms. For instance, Zhang *et al.* [2] demonstrated such a new potential to map oil slick using polarimetric SAR decomposition parameters obtained from a RADARSAT-2 quad-polarization image. Indeed, as previously suggested by Schuler and Lee, ambient clean sea and slick areas can be well separated

using the polarimetric matrix decomposition, particularly suggesting that non-Bragg scattering becomes dominant over oil slicks.

However, without taking into account the full amplitude and phase information of each resolved pixel of quad-polarization observations, dual co-polarized radar instrument could be already ideally suited. As analyzed by Mouche *et al.*, copolarized VV and HH data can be combined to infer both local wind characteristics and then minimize the wind-induced surface motion to estimate surface current. According to different asymptotic electromagnetic models, the polarization sensitivity for both Doppler shifts and radar intensity signals seems readily exploitable. The combined use of dual co-polarized measurements can then become a very efficient tool to better understand, discriminate, and quantify the different scattering mechanisms responsible for the manifestation of surface currents, slicks, and wind field features in SAR images.

In this letter, the proposed methodology is tested and assessed with a RADARSAT-2 fine-mode quad-polarization SAR image, but we solely focus on the dual co-polarized VV and HH products. Hereafter, we take advantage of this dual information to effectively and quickly identify different scene areas where wind direction changes occur or where surface wave–current interactions strongly enhance surface breaking waves. The proposed methodology also helps to identify and interpret oil spill manifestations, to possibly complement analysis using the full-polarimetric SAR information.

## II. Observations and interpretation

---

### A. SAR Data

This study is based on a quad-polarized RADARSAT-2 SAR image acquired over the Mediterranean Sea, i.e., the coastal area near the town of Begur (Spain), which is acquired at 17:40 UTC on December 18, 2010 (see Fig. 1). The SAR image roughly covers a 40 km × 50 km area and has pixel spacing of about 5 and 10 m in slant and azimuth directions, respectively. In this image, the incidence angle varies from 32.7° to 35.7°. The RADARSAT-2 Wide Fine Quad-Pol imaging mode provides single-look complex data in HH, HV, VH, and VV polarizations. VV- and HH-polarized images are shown in Fig. 1(a) and (b). The corresponding image of the polarization ratio (HH over VV in linear units, hereinafter PR) and the polarization

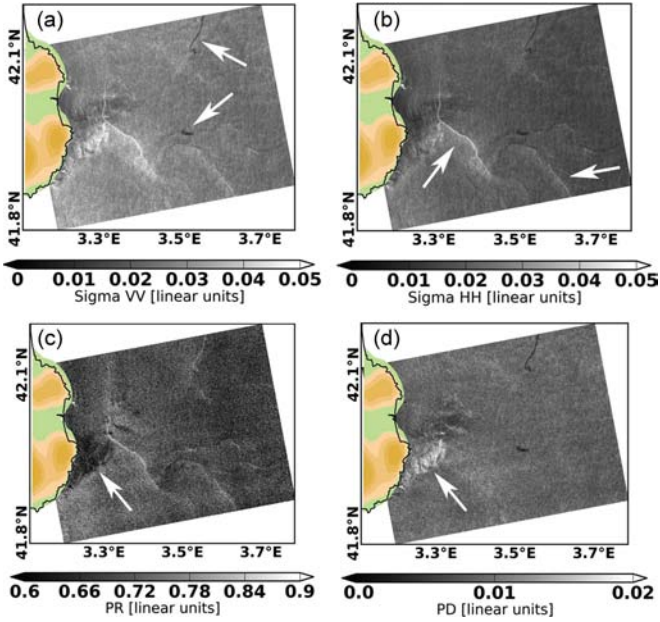


Fig. 1. RADARSAT-2 SAR image of a coastal area in the Mediterranean Sea acquired at 17:40 UTC on December 18, 2010 in terms of (a) VV-polarization, i.e.,  $\sigma_0^{VV}$  (in linear units); (b) HH-polarization, i.e.,  $\sigma_0^{HH}$ ; (c) PR  $\sigma_0^{HH}/\sigma_0^{VV}$  (in linear units); and (d) PD  $\Delta\sigma = \sigma_0^{VV} - \sigma_0^{HH}$  (in linear units). The arrows indicate manifestations of (a) surface slicks, (b) surface current signatures, and (c) and (d) wind field features.

81 difference (VV minus HH in linear units, hereinafter PD) are  
82 shown in Fig. 1(c) and (d), respectively. It was checked that  
83 signal intensity is everywhere ten times above the thermal noise  
84 level, even for HH signals over the reduced backscatter slick  
85 areas. As such, it was not necessary to account for the thermal  
86 noise in the normalized radar cross section (NRCS) calibration  
87 and subsequent PR and PD calculations.

88 Fig. 2(a) provides Meteo-France AROME model information  
89 about the surface wind field conditions. Mostly blowing from  
90 the south, an offshore wind has developed near the coast, with  
91 chances of convective clouds and associated rain cells at the  
92 atmospheric front, as seen on the SAR scenes. The Doppler  
93 frequency shift in Fig. 2(b) confirms the near-coastal wind  
94 direction change predicted by the atmospheric model.

95 In Fig. 1(a) and (b), both VV and HH SAR images exhibit  
96 distinct linear bright/dark signatures, which can be interpreted  
97 as manifestations of surface currents, and two dark areas pre-  
98 sumably caused by the surface slicks. The distinct current-  
99 induced features are well expressed in the PR field but almost  
100 entirely removed in the PD field (see Figs. 1 and 4). The clear  
101 drop of the Doppler velocity in Fig. 2(b) across the bright SAR  
102 signatures confirms that this signature traces the surface current  
103 shear. Unlike the current-induced features, the surface slicks are  
104 detected in both PR and PD fields. However, within the slicks,  
105 the PR values are larger over the ambient area, whereas PD  
106 values are lower to reflect the surface roughness suppression  
107 in the slicks (see Figs. 1 and 3).

## 108 B. Model Approach

109 To interpret these observed SAR features, the HH and NRCS  
110 components are represented as a sum of polarized scattering,

which is associated with the conventional two-scale Bragg scat- 111  
tering  $\sigma_{0B}^{pp}$  and the nonpolarized (NP) scattering from breaking 112  
waves  $\sigma_{wb}$  [5], [6], i.e., 113

$$\sigma_0^{pp} = \sigma_{0B}^{pp} + \sigma_{wb}. \quad (1)$$

The NP scattering originates from radar returns from very 114  
“rough” and “steep” surface patches, and it is the same at both 115  
polarizations (see, e.g., [5], [7], and corresponding references). 116  
Using the NRCS model (1), the PR reads 117

$$P \equiv \frac{\sigma_0^{hh}}{\sigma_0^{vv}} = \frac{\sigma_{0B}^{hh} + \sigma_{wb}}{\sigma_{0B}^{vv} + \sigma_{wb}}. \quad (2)$$

PR values have already received attentions reported in the liter- 118  
ature, mostly to discuss departure from tilted Bragg scattering 119  
mechanism due to the impact of breaking waves (see, e.g., [7] 120  
and [8]). In particular, Mouche *et al.* reported comprehensive 121  
analysis of dual-polarized C-band airborne radar measurements 122  
and found a remarkable deviation of the observed PR values 123  
from the Bragg scattering predictions, showing, in turn, signifi- 124  
cant contribution of the NP scattering to the sea surface NRCS. 125

The wave breaking contribution can be thus removed using 126  
the PD between VV and HH NRCSs, i.e., 127

$$\Delta\sigma_0 \equiv \sigma_0^{vv} - \sigma_0^{hh} = \sigma_{0B}^{vv} - \sigma_{0B}^{hh} \quad (3)$$

which is controlled by the surface roughness produced by wave 128  
components close to the Bragg wave number. Since C-band 129  
Bragg waves, which are around 5-cm wavelength, have “quick 130  
response” to wind forcing, with the spatial relaxation scale of 131  
an order of 10 m, we can anticipate that PD  $\Delta\sigma_0$  should closely 132  
reflect the near-surface wind variability and presence of the 133  
slicks. 134

From (1) and (3), the NP component, i.e.,  $\sigma_{wb}$ , can be 135  
assessed as 136

$$\sigma_{wb} = \sigma_0^{vv} - \Delta\sigma_0 / (1 - p_B) \quad (4)$$

where  $p_B$  is the PR for the two-scale Bragg scattering, i.e., 137

$$p_B = \sigma_{0B}^{hh} / \sigma_{0B}^{vv}. \quad (5)$$

For the two-scale Bragg scattering,  $p_B$  is mostly governed 138  
by the local geometry and tilting effects. At moderate incidence 139  
angles, i.e.,  $\theta > 25^\circ$ , relation for  $p_B$  can be simplified to (see, 140  
e.g., [7, eq. (31)]) 141

$$p_B = \frac{|G_h|^2}{|G_v|^2} \left[ 1 + (g_h - g_v) s_i^2 \right] \quad (6)$$

where  $|G_p|$  is the scattering coefficient (depending on the 142  
incidence angle),  $g_p = 1/2 \cdot \partial^2 |G_p| / \partial^2 \theta$ , and  $s_i^2$  is the mean 143  
square slope of tilting waves (waves with wavelength longer 144  
than few times the Bragg wavelength) in the incidence plane 145  
direction. The second term in the square bracket describes the 146  
impact of tilting waves to  $p_B$ . In the range of incidence angles 147  
from  $25^\circ$  to  $45^\circ$ , it varies from 0 (calm) to about 0.5 at moderate 148  
winds. Thus,  $p_B$  may be assumed to be weakly dependent 149  
on wind speed and almost independent on azimuth (due to 150

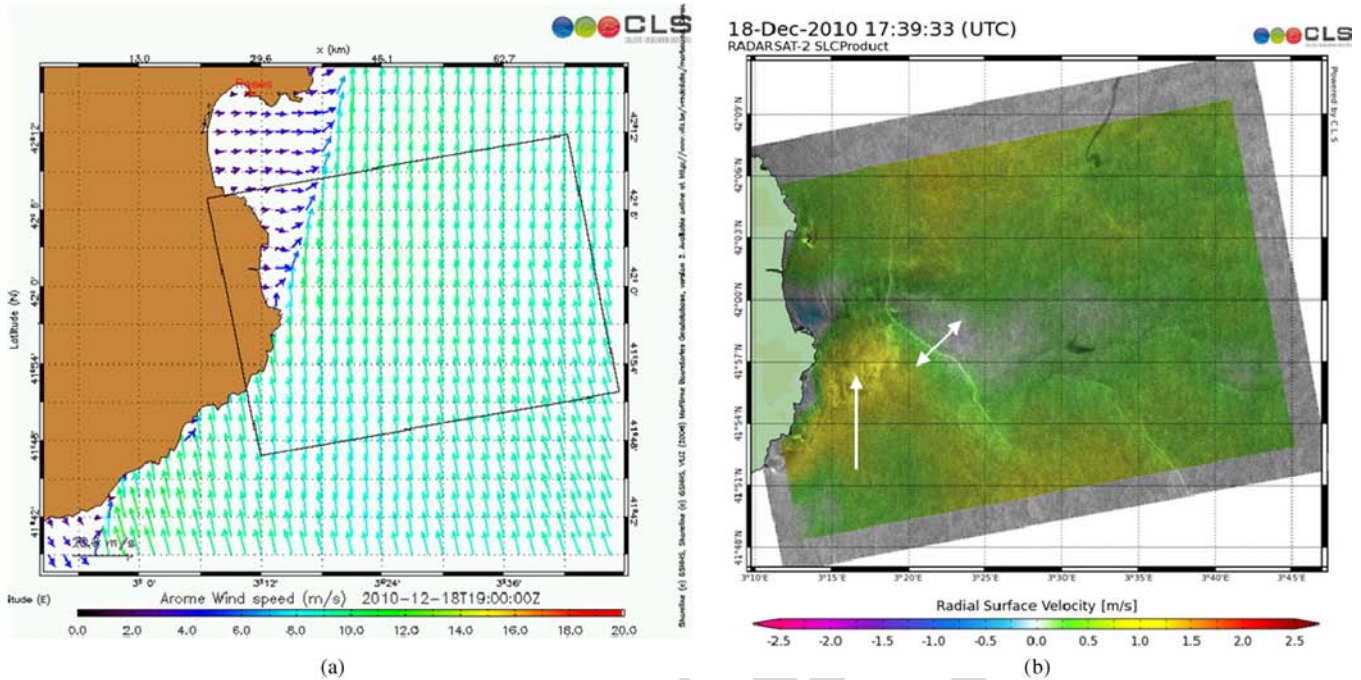


Fig. 2. (a) Contemporaneous Meteo-France AROME model surface wind field. (b) Doppler velocity field overlapped the VV NRCS image. The box in (a) indicates the position of the SAR image. The positive Doppler velocity is directed toward the SAR look direction. The arrow in (b) indicates the area with “large” positive values of Doppler velocity spatially coincident with the manifestation of the wind field feature marked in Fig. 1(c) and (d). The double-sided arrow marks the Doppler velocity drop over the bright NRCS signature. Near the coastal area, the Doppler changes sign to confirm the wind direction change becoming downwind, where the Doppler is negative. Over the main area, the Doppler values are small, confirming crosswind radar conditions.

151 approximately azimuthal isotropy of tilting waves slopes; see  
152 also [8, Fig.16]). For the considered RADARSAT-2 image, two-  
153 scale model simulations give  $p_B$  to vary from 0.56 to 0.45 in  
154 linear units for incidence angles ranging from  $32.7^\circ$  to  $35.7^\circ$ .

155 Based on model (1), the original co-polarized VV and HH  
156 images can be thus transformed to new PD and NP images de-  
157 fined by (3) and (4), respectively, which possess information on  
158 very different radar scattering mechanisms, i.e., the polarized  
159 Bragg scattering provided by short fast-response wind waves  
160 and NP radar returns from breaking waves in a wide spectral  
161 range. Interplay between these mechanisms is included in the  
162 PR image defined by (2). Due to different sensitivity of short  
163 wind waves and wave breaking to various ocean phenomena,  
164 this set of new images can then serve as an effective tool for  
165 SAR data interpretation.

### 166 C. Scattering Mechanism Analysis

167 In the present case, the mean PR is about 0.7 in linear units  
168 [about  $-1.5$  dB; see Fig. 1(c)], except within the “dark” area  
169 near the coast [see the arrow in Fig. 1(c)], where PR drops  
170 to 0.55 (or about  $-2.5$  dB). In this area, the wind direction  
171 is offshore, close to the downwind radar look direction. In  
172 the downwind direction, the model suggests that the Bragg  
173 component dominates over the wave breaking contribution that  
174 results in a lower PR (see also , Fig. 16[8]). In addition to the  
175 atmospheric model [see Fig. 2(a)], the offshore wind direction  
176 in this area is also confirmed by the positive Doppler velocity  
177 [see Fig. 2(b)]. In this localized coastal area, the PR values are  
178 in closer agreement with the two-scale Bragg scattering model.  
179 Outside this coastal area, the PR values are larger, about 0.7,

confirming the significant contribution of wave breaking to the  
180 total NRCS (about 40%).

181  
182 The bright/dark quasi-linear current-induced signature in the  
183 PR field (see Figs. 1(c) and 4) can be further interpreted as en-  
184 hancement/suppression of wave breaking due to wave–current  
185 interactions. As found, the PR values related to this bright  
186 quasi-linear feature are close to 1 in linear units (0 dB), in-  
187 dicated that NRCS is mostly dominated by rough patches of  
188 breaking waves, and the effect of Bragg wave modulation by  
189 current is negligible.

190 Over the slick area, PR values are also larger than the  
191 background values (see Fig. 3). Surface films (either biogenic  
192 or oil) suppress short wind waves and thus reduce resonant  
193 Bragg scattering (see, e.g., [9]). However, wind waves with  
194 wavelength on the order of a decimeter and longer are not  
195 damped by the surface films. Hence, the contribution of these  
196 longer waves within slicks is probably the same as outside,  
197 leading to increased PR values (see Fig. 3).

### 198 D. Polarization Decomposition

199 As anticipated, a PD image should closely reflect the near-  
200 surface wind variability and slicks. Near the coast, where off-  
201 shore winds are expected due to breezes, the NRCS difference  
202 is clearly enhanced [see Fig. 1(d)], confirming a downwind  
203 radar look direction, for which Bragg scattering is maximal.

204 Fig. 1 suggests that the bright/dark quasi-linear feature (as-  
205 sociated with the surface currents manifestation) well visible  
206 in VV, HH, and PR scenes is not distinguishable in the PD  
207 scene (see also Fig. 4). This indicates that SAR signatures over  
208 oceanic fronts mostly result from enhanced/suppressed wave

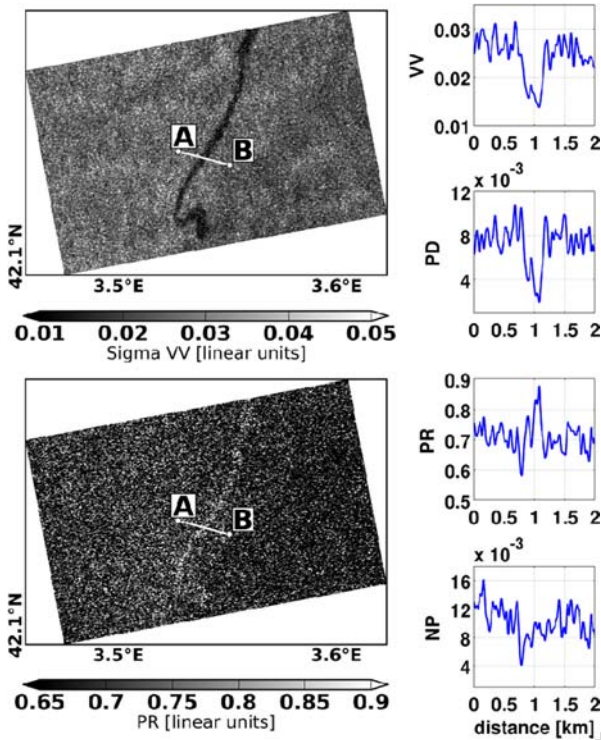


Fig. 3. (Upper left) Enlarged fragment of the slick area in Fig. 1(a). (Lower left) Corresponding enlargement of the PR in Fig. 1(c), showing the same slicks, which manifests as a bright area. (Right column) Cross section across the slick area along line A-B for (top to bottom) VV NRCS  $\sigma_0^{VV}$ , PR  $\sigma_0^{hh}/\sigma_0^{VV}$ , PD  $\Delta\sigma = \sigma_0^{VV} - \sigma_0^{hh}$ , and NP component  $\sigma_{wb}$  of the NRCS defined by (4). All quantities are represented in linear units.

209 breaking, which provide NP radar returns. On the other hand,  
 210 the slick areas correspond to dark areas due to strong damping  
 211 of the Bragg waves (see Figs. 1(d) and 3). Notice that the  
 212 background-to-slick ratio for the PD is about 3 (or 4.8 dB),  
 213 whereas that for the VV NRCS is only about 1.8 (or 2.6 dB).  
 214 This shows that PD  $\Delta\sigma_0$  is more sensitive to the presence of  
 215 the surface slick, in comparison with  $\sigma_0^{VV}$ , as removal of  $\sigma_{wb}$  is  
 216 less affected by slicks. The PD thus contains information about  
 217 “fast-response” spatial changes of short-scale Bragg waves,  
 218 which are mainly caused by variable wind field and/or the  
 219 presence of surface slicks.

#### 220 E. Wave Breaking Contribution

221 NP component of the NRCS defined by (4) is shown in  
 222 Figs. 4 and 5, using  $p_B = 0.5$ . First, the surface currents’ fea-  
 223 tures are remarkably emphasized. The current-over-background  
 224 ratio for  $\sigma_{wb}$  is about 4 (6 dB), which is significantly larger  
 225 than the similar ratio for VV, which is about 1.6 (or 2 dB).  
 226 Second, surface slicks are almost not detected in the NP field  
 227 (see Fig. 3). This indicates that, unlike Bragg waves, longer  
 228 scale roughness is less affected by slicks, although some small  
 229 suppression of  $\sigma_{wb}$  in the slick area can be revealed.

230 Near the coast where the wind is blowing offshore, the  
 231 impact of wave breaking to the NRCS is also noticeably  
 232 weaker. Such lower values of  $\sigma_{wb}$  in this area may partly result  
 233 from azimuthal dependence of  $\sigma_{wb}$ . Following [7], the radar  
 234 return from breaking waves is maximal in the upwind direction

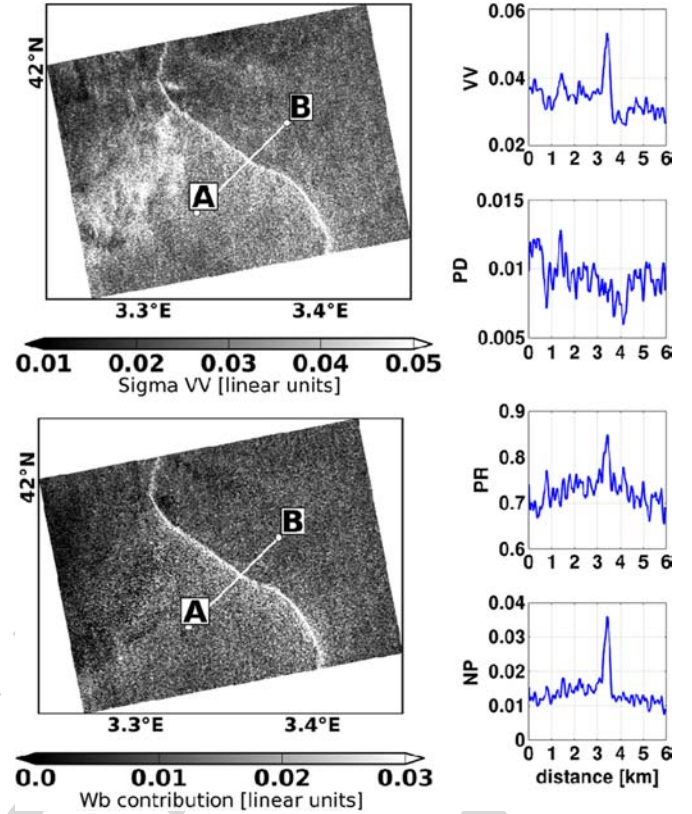


Fig. 4. (Upper left) VV SAR image enlargement of the current feature in Fig. 1(a). (Lower left) NP part of the NRCS of the same signature. (Right column) Cross section across the current feature along line A-B for (top to bottom) VV NRCS  $\sigma_0^{VV}$ , PR  $\sigma_0^{hh}/\sigma_0^{VV}$ , PD  $\Delta\sigma = \sigma_0^{VV} - \sigma_0^{hh}$ , and NP part  $\sigma_{wb}$  of the NRCS defined by (4). All quantities are represented in linear units.

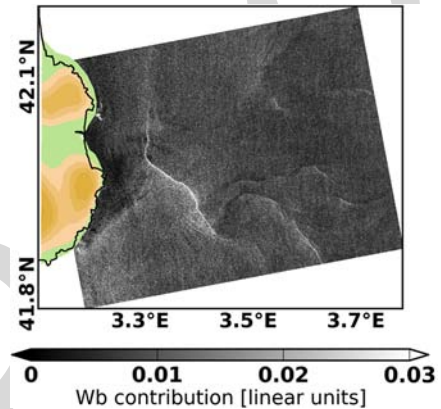


Fig. 5. Wave breaking contribution (NP component of the NRCS) estimated from (4) via  $\sigma_0^{VV}$  and  $\Delta\sigma_0$  fields shown in Fig. 1.

and minimal in the downwind direction. Analysis of NSCAT 235 AQ7  
 Ku-band dual-polarized satellite data also confirms this as- 236  
 237 sumption [6], [10]. Lower values of both  $\sigma_{wb}$  and PR near the  
 238 coast represent a consistent pattern caused by local offshore  
 239 wind condition. Notice that shorter fetch in the offshore wind  
 240 reduces the NRCS and, more specifically, its NP part. 241

### 242 III. CONCLUSION

Dual co-polarization SAR data can thus provide a tool 243  
 to distinguish between different mechanisms affecting ocean 244

245 radar backscatter. NRCS difference between VV and HH co-  
246 polarized data (PD image) enhances variability produced by  
247 the resonant scattering mechanism. Since short Bragg waves  
248 are “fast-response” waves, a PD image carries information on  
249 the spatial variability of surface wind velocity field (including  
250 gustiness) and surface slicks.

251 As demonstrated, combination of VV and PD images can  
252 be then used to derive the NP component of the NRCS (NP  
253 image), essentially dominated by breaking waves. Since wave  
254 breakings are very sensitive to the nonuniform surface current,  
255 an NP image shall reflect surface manifestations of sub- and  
256 mesoscale ocean currents.

257 The use of the PD and NP images per se or in their  
258 combination with original VV- and HH-images or/and a PR  
259 image can thus become a very powerful tool for detection and  
260 discrimination of various ocean surface phenomena such as  
261 surface currents, slicks, and wind field features.

262 Over the main area of the radar scene considered in this letter,  
263 the radar look direction corresponds to the crosswind direction.  
264 As have been shown by Kudryavtsev *et al.*, such observation  
265 geometry is optimal for SAR detection of surface current fea-  
266 tures, which are distinctly visible in both VV and HH images as  
267 linear bright/dark features linked to the surface current velocity  
268 gradients identified via Doppler velocity measurements. As  
269 even better detected in the NP image, the origin of such SAR  
270 signatures is associated with the enhancement/suppression of  
271 wave breaking by the surface current shear. Remarkably, the  
272 surface currents are then not visible in the PD image, confirm-  
273 ing that the resonant Bragg waves have very short relaxation  
274 time and, thus, are not modulated by the surface currents.

275 The surface slicks are manifested as dark areas in both  
276 VV and HH SAR images due to the damping of the Bragg  
277 waves. However, slicks appear as “bright” areas in PR images.  
278 Though short waves are damped, the radar return supported  
279 by longer wave breaking is almost unaffected in slicks. As a  
280 result, slicks are well visible in the PD image and almost not  
281 visible in NP images. Moreover, the apparent difference of slick  
282 signatures in VV image (HH image) and in the PR image opens  
283 a promising opportunity to discriminate slicks from a look-alike  
284 feature associated with low-wind conditions and surface current  
285 effects. This result is in line with the previously reported full-  
286 polarimetric analysis [2], [3].

287 Wind speed variability is well traced in both the PD image  
288 and from the Doppler frequency. Wind variability can be thus  
289 more effectively “filtered out” from dual co-polarized SAR  
290 images to help remove the wind-wave-induced Doppler contri-  
291 bution in [12] to better quantify the surface current signatures in  
292 the NP image. In addition, strong azimuthal wind dependence  
293 of PR and NP signals may serve as an indicator of wind  
294 direction changes. In this study, this feature helped to reveal a  
295 localized zone with offshore winds. Future investigations shall  
296 further dwell on the present analysis of dual co-polarized SAR

data to confirm the radar backscatter polarization sensitivity and  
to better assess the Doppler frequency polarization sensitivity.  
As foreseen, combined dual co-polarized SAR data can then  
help and refine the interpretation of the various detected fea-  
tures (wind field changes, surface currents, and oil slicks versus  
their look-alikes).

## ACKNOWLEDGMENT

The authors would like to thank the support from the SeaU  
EU FP7 Project 263246 and ANR-BLAN-0365 REDHOTS. RADARSAT  
RADARSAT is an official mark from the Canadian Space  
Agency (CSA). RADARSAT-2 SAR (CSA) data were acquired  
within the framework of the VIGISAT project, in the framework  
of a FEDER Grant Présage No. 32635, and a partnership with  
CLS for providing the data, under Contracts 08 GET 13M and  
09 GET 11M.

## REFERENCES

- [1] B. Holt, “SAR imaging of the ocean surface,” in *Synthetic Aperture Radar Marine User’s Manual*, C. R. Jackson and J. R. Apel, Eds. Silver Spring, MD: NOAA, 2004, pp. 263–275.
- [2] B. Zhang, W. Perrie, X. Li, and W. G. Pichel, “Mapping sea surface oil slicks using RADARSAT-2 quad-polarization SAR image,” *Geophys. Res. Lett.*, vol. 38, no. 10, p. L10 602, May 2011.
- [3] D. L. Schuler and J. S. Lee, “Mapping ocean surface features using biogenic slick-fields and SAR polarimetric decomposition technique,” *Proc. Inst. Elect. Eng.—Radar Sonar Navig.*, vol. 153, no. 3, pp. 260–270, Jun. 2006.
- [4] A. A. Mouche, B. Chapron, N. Reul, and F. Collard, “Predicted Doppler shifts induced by ocean surface wave displacements using asymptotic electromagnetic wave scattering theories,” *Waves Random Complex Media*, vol. 18, no. 1, pp. 185–196, Feb. 2008.
- [5] O. M. Phillips, “Radar returns from the sea surface—Bragg scattering and breaking waves,” *J. Phys. Oceanogr.*, vol. 18, no. 8, pp. 1065–1074, Aug. 1988.
- [6] Y. Quilfen, B. Chapron, A. Bentamy, J. Gourrion, T. El Fouhaily, and D. Vandemark, “Global ERS 1 and 2 and NSCAT observations: Upwind/crosswind and upwind/downwind measurements,” *J. Geophys. Res.*, vol. 104, no. C5, pp. 11 459–11 469, May 1999.
- [7] V. N. Kudryavtsev, D. Hauser, G. Caudal, and B. Chapron, “A semiempirical model of the normalized radar cross-section of the sea surface 1. Background model,” *J. Geophys. Res.*, vol. 108, no. C3, pp. 8054–8077, Jan. 2003.
- [8] A. A. Mouche, D. Hauser, and V. Kudryavtsev, “Radar scattering of the ocean surface and sea-roughness properties: A combined analysis from dual-polarizations airborne radar observations and models in C band,” *J. Geophys. Res.*, vol. 111, no. C9, p. C09 004, Sep. 2006.
- [9] M. Gade, W. Alpers, H. Huehnerfuss, H. Masuko, and T. Kobayashi, “Imaging of biogenic and anthropogenic ocean surface films by the multifrequency/multipolarization SIR-C/X-SAR,” *J. Geophys. Res.*, vol. 103, no. C9, pp. 18 851–18 866, Aug. 1998.
- [10] B. Chapron, V. Kerbaol, and D. Vandemark, “A note on relationships between sea-surface roughness and microwave polarimetric backscatter measurements: Results from POLRAD’96,” presented at the Proc. POLRAD Int. ESA Workshop, ESTEC, Noordwijk, The Netherlands, Apr. 29, 1997, ESA WPP-135.
- [11] V. N. Kudryavtsev, D. Akimov, J. A. Johannessen, and B. Chapron, “On radar imaging of current features: 1. Model and comparison with observations,” *J. Geophys. Res.*, vol. 110, no. C7, p. C07 016, Jul. 2005.
- [12] B. Chapron, F. Collard, and F. Ardhum, “Direct measurements of ocean surface velocity from space: Interpretation and validation,” *J. Geophys. Res.*, vol. 110, no. C7, p. C07 008, Jul. 2005.