

Supplementary information

Table S1. Cross-spectral analyses of *G. ruber* Mg/Ca-SST obtained for core MD05-2920 relative to benthic foraminiferal $\delta^{18}\text{O}$ records over different orbital bands

Orbital band	Coherency (lower/higher)	Phase ($^{\circ}$)	Phase (kyr)
<i>C. wuellerstorfi</i> $\delta^{18}\text{O}$			
100 kyr	0.99 (0.94/1.00)	12.7 \pm 10.3	3.5 \pm 2.8
41 kyr	0.95 (0.80/0.99)	2.3 \pm 5.1	0.3 \pm 0.6
23 kyr	0.98 (0.91/0.99)	2.2 \pm 0.9	0.1 \pm 0.1
<i>U. peregrina</i> $\delta^{18}\text{O}$			
100 kyr	0.99 (0.94/1.00)	12.6 \pm 10.1	3.5 \pm 2.8
41 kyr	0.95 (0.82/0.99)	3.4 \pm 4.7	0.4 \pm 0.5
23 kyr	0.97 (0.90/0.99)	2.1 \pm 1.3	0.1 \pm 0.1

Analyses were carried out using the Blackman-Tukey method with a bandwidth of 0.010 kyr⁻¹ and a confidence interval of 95%. No-zero coherence > 0.603.

Table S2. Cross-spectral analyses of bottom water temperature changes (ΔBWT) obtained for core MD05-2920 and for ODP1123 (Elderfield et al., 2012) relative to sea salt flux (Wolff et al., 2006), simulated sea-ice coverage in the Southern Ocean, and computed annual mean Southern Ocean net shortwave forcing over precession band (23 kyr)

Parameters	Coherency (lower/higher)	Phase ($^{\circ}$)	Phase (kyr)
$\Delta\text{BWT}_{\text{MD05-2920}}$			
Sea salt flux	0.81 (0.40/0.95)	11.1 \pm 0.4	0.7 \pm 0.0
Simulated sea-ice	0.86 (0.54/0.96)	59.3 \pm 50.6	3.8 \pm 3.2
Shortwave forcing	0.87 (0.57/0.97)	15.5 \pm 7.3	1.0 \pm 0.5
$\Delta\text{BWT}_{\text{ODP1123}}$			
Sea salt flux	0.85 (0.51/0.96)	6.8 \pm 2.3	0.4 \pm 0.1
Simulated sea-ice	0.91 (0.68/0.98)	55.8 \pm 49.2	3.6 \pm 3.1
Shortwave forcing	0.92 (0.71/0.98)	13.6 \pm 6.8	0.8 \pm 0.4

Analyses were carried out using the Blackman-Tukey method with a bandwidth of 0.010 kyr⁻¹ and a confidence interval of 95%. No-zero coherence > 0.603.

Supplementary figure caption:

Figure S1. Comparison of the past SST records based on foraminiferal Mg/Ca with zoom for the last four terminations ODP806B (0°19.1'N, 159°21.7'E, 2520 m; (Lea et al., 2000), MD97-2140 (2°02'N, 141°46'E, 2547 m; (de Garidel-Thoron et al., 2005), MD98-2162, (4°41.33'S, 117°54.17'E, 1855m; (Visser et al., 2003), MD01-2378 (13°4.95'S, 121°47.27'E, 1783 m; (Xu et al., 2008) and MD06-3067 (6°31'N, 126°30'E, 1575 m; (Bolliet et al., 2011).

Figure S2. Mg/Ca (mmol/mol) of *G. ruber* for the past 400 kyr together with individual foraminiferal test weight (250-355 μm) and loss of the tests during the cleaning. Bleu zones are glacial periods and numbers along upper x-axis indicate marine isotopic stages.

Figure S3. (a) Correlation between Nino 3.4 SSTA and tropical West Pacific SSTA field obtained from the Reynolds SST data set (Vazquez-Cuervo et al., 1998). (b) Present-day Monthly sea surface salinity in upper 50 m at 2.5S and 144.5E (Antonov et al., 2010). (c) Monthly climatological Chlorophyll-a concentration (1998/01 - 2007/12) at 2.9S and 144.4E. Data from <http://oceancolor.gsfc.nasa.gov/>.

Figure S4. Comparison between the simulated ice volume by gLOVE over the last 406 ka with reconstructed sea level evolution (Siddall et al., 2003; Waelbroeck et al., 2002).

Figure S5. Spectral analysis (multitaper method, MTM and Blackmann-Tuckey) of *G. ruber* Mg/Ca SST of core MD05-2920 performed with the Analyseries 2.0 software (Paillard et al., 1996). The power spectra and the significance value (>0.95) of the peaks clearly show that dominant frequencies correspond to the obliquity (41 kyr), the precession (23 kyr) and 100 kyr. Non-primary spectral peak represents the heterodyne of obliquity and the 23 kyr component of precession ($1/41+1/23 = 1/14.7$) also appears with semi-precession band (11 kyr) although these peaks are much smaller. BW=Band Width.

Figure S6. Comparison between residual Mg/Ca-SST variability that is unrelated to CO₂ forcing and seasonal SST bias related to *G. ruber* production over the past 400 kyr. The seasonal bias is estimated from chlorophyll-weighted simulated seasonal SSTs: $\sum_{i=1,4} \text{chl}(i) \text{SST}(i,t) / \sum_{i=1,4} \text{chl}(i)$, with $\text{chl}(i)$ representing the observed SeaWiFS seasonal mean chlorophyll concentration (Figure S2c) during December-January-February ($i=1$), March-April-May ($i=2$), June-July-August ($i=3$), September-October-November ($i=4$) and $\text{SST}(i,t)$ representing the SST in TR400 in season i and year t . The small seasonal bias and anti-relationship with the residual Mg/Ca-SST indicate that the seasonality is not the reason for residual Mg/Ca-SST variability.

Figure S7. Relationship between Antarctic air temperature anomaly relative to the present and WPWP SST anomaly for the last 400 kyr. Proxy relationship between Mg/Ca-SST of core MD05-2920 and the deuterium-based temperature reconstructions from Antarctica (Jouzel et al., 2007) is characterized by a slope of 2.9 ($R^2=0.7$). The simulated relationship based on TR400 shows a very similar slope of 2.8 ($R^2=0.5$)

Reference

Paillard, D., Labeyrie, L., and Yiou, P.: Macintosh program performs time-series analysis, EOS Trans. AGU, 77, 1996.

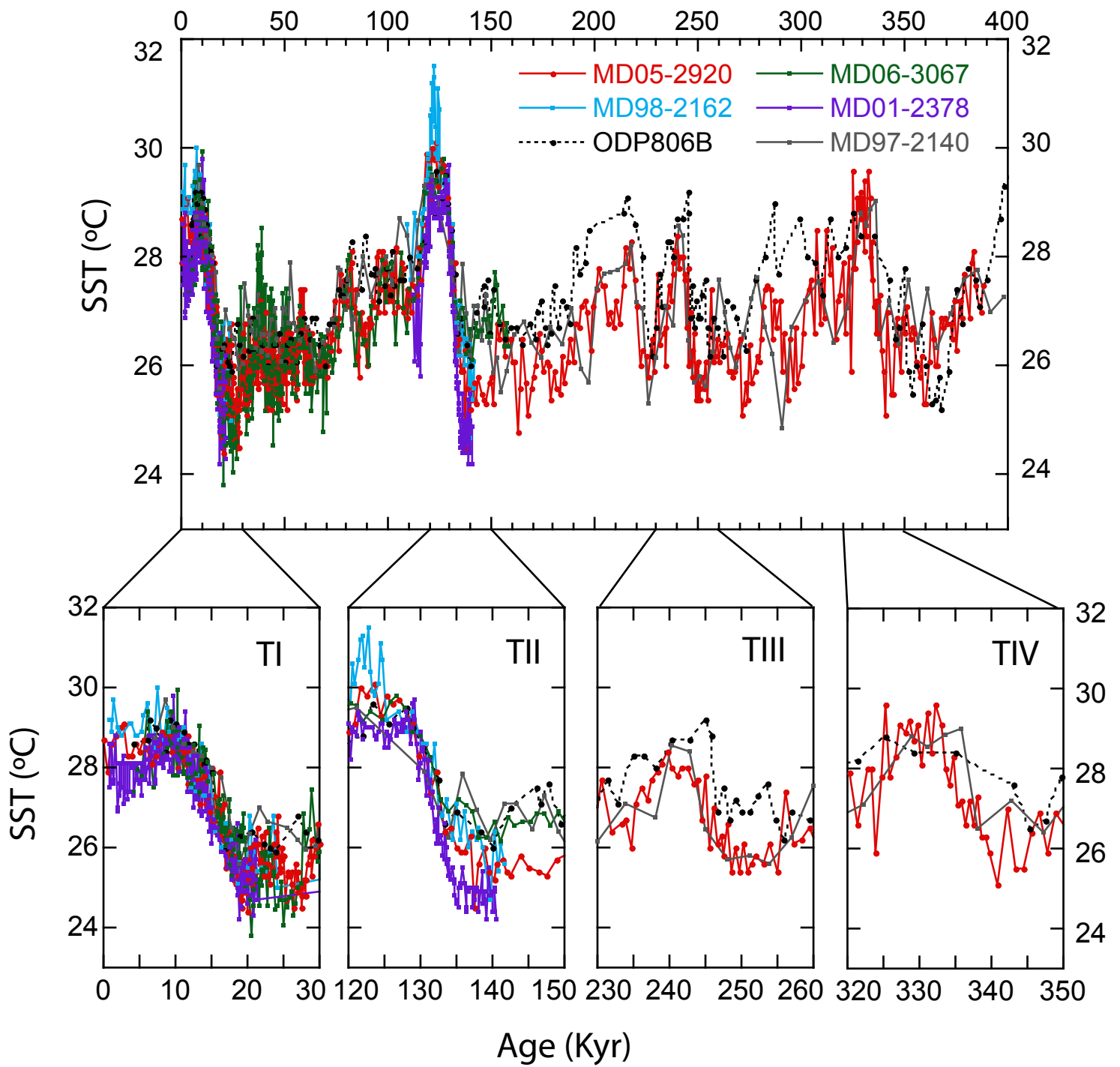


Figure S1

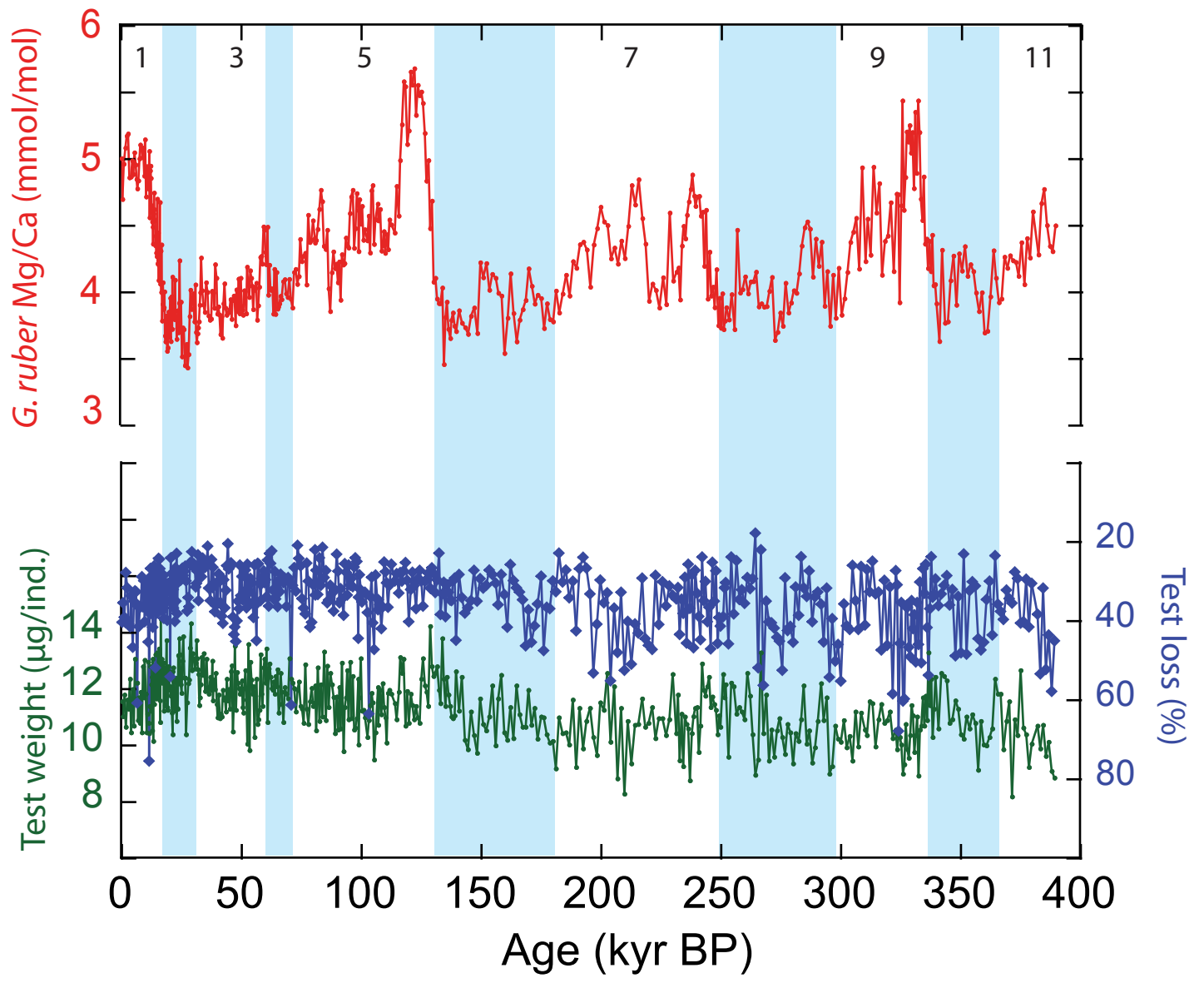


Figure S2

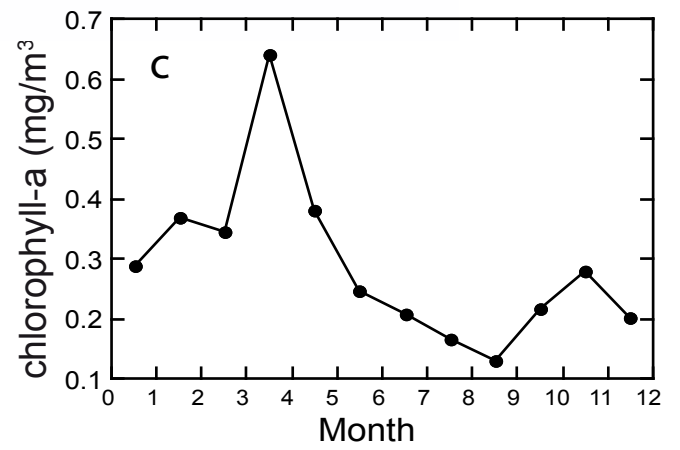
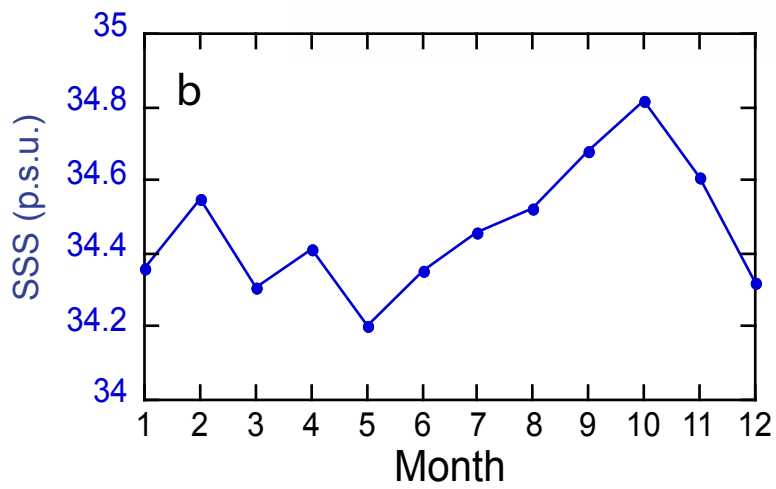
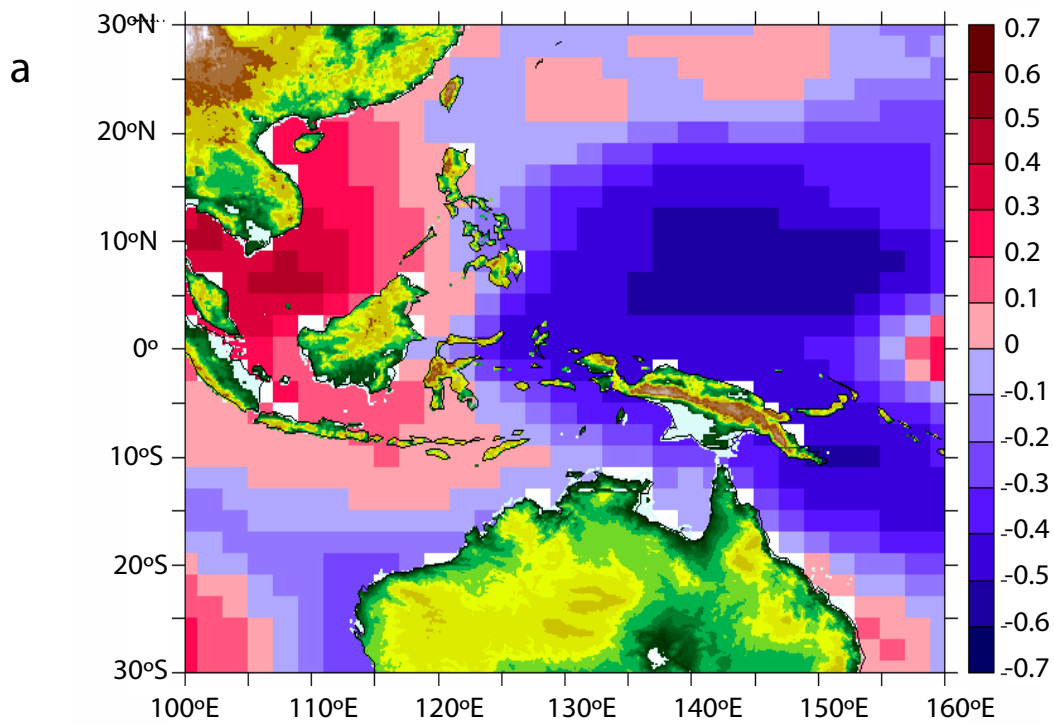


Figure S3

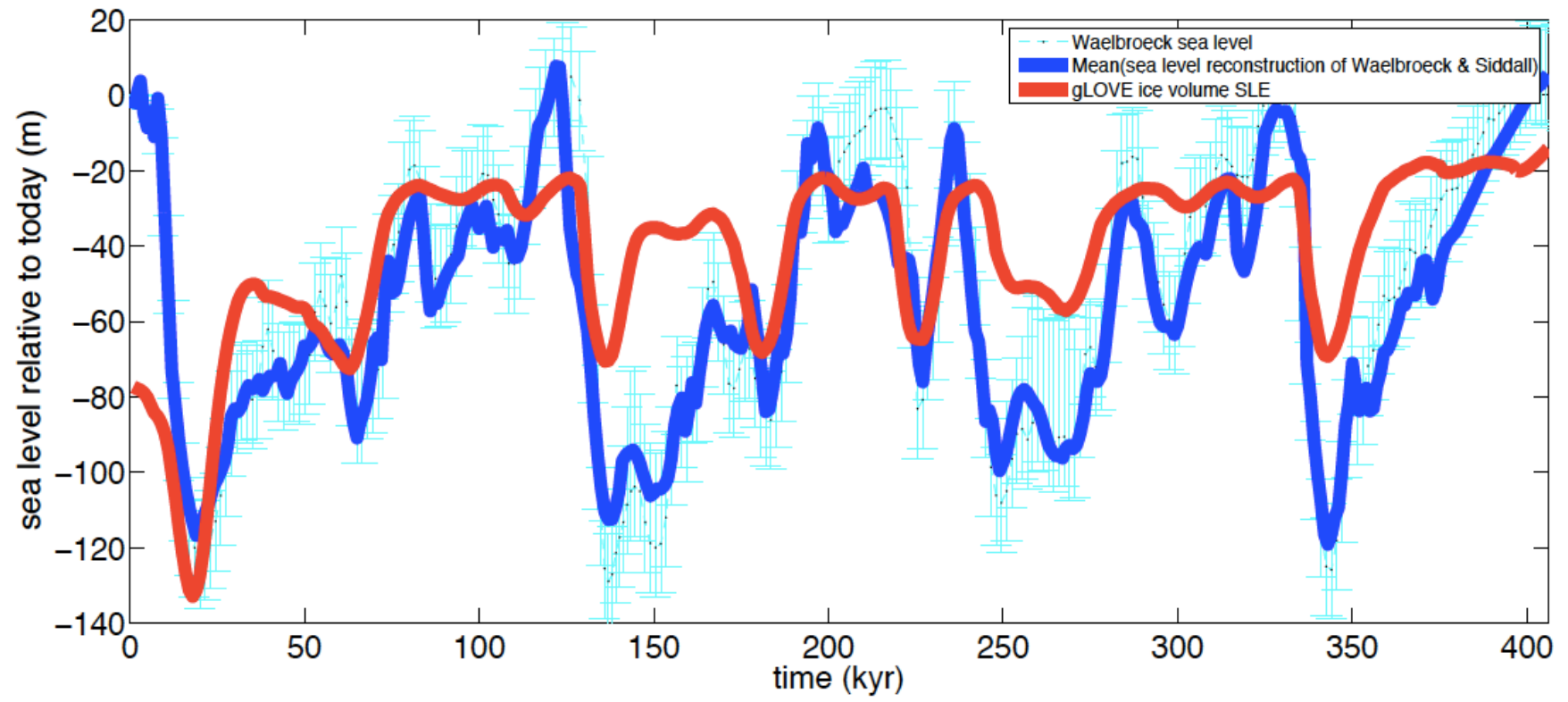


Figure S4

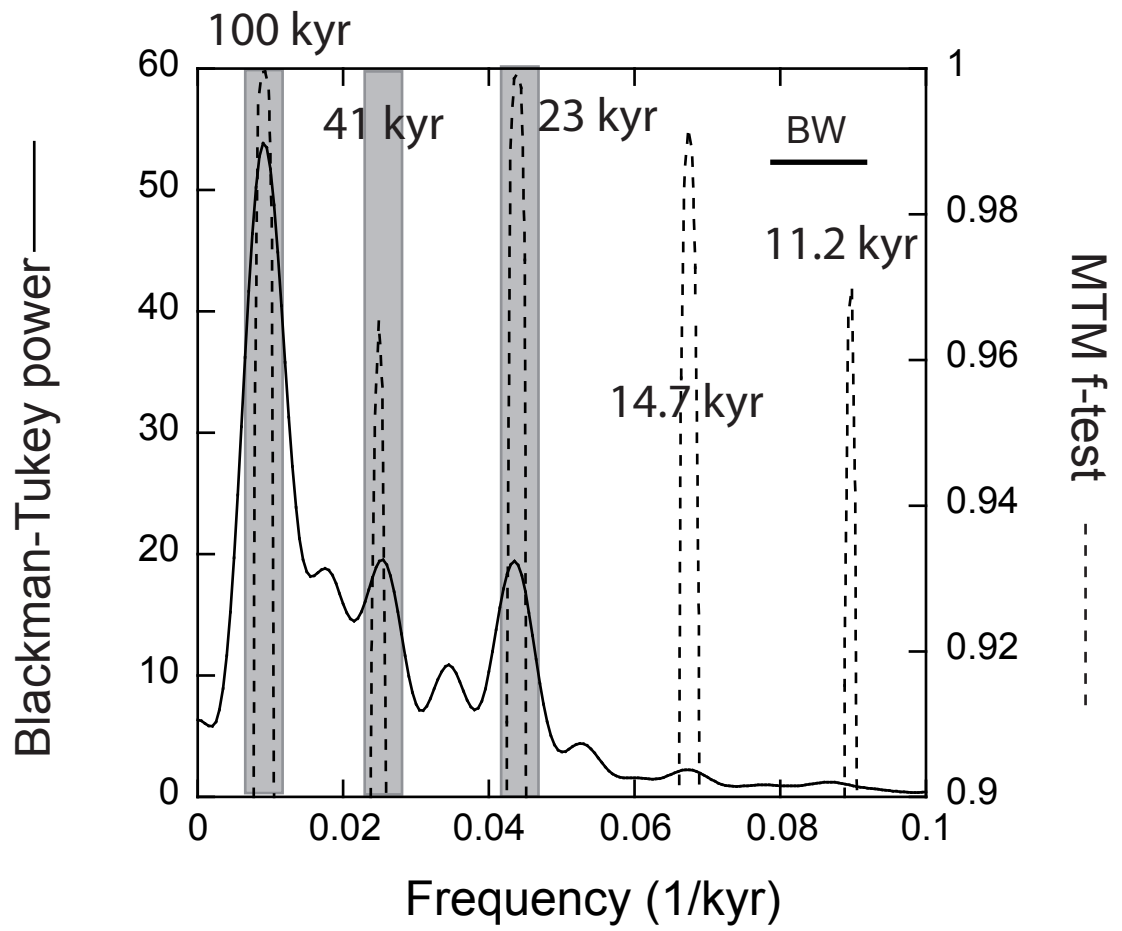


Figure S5

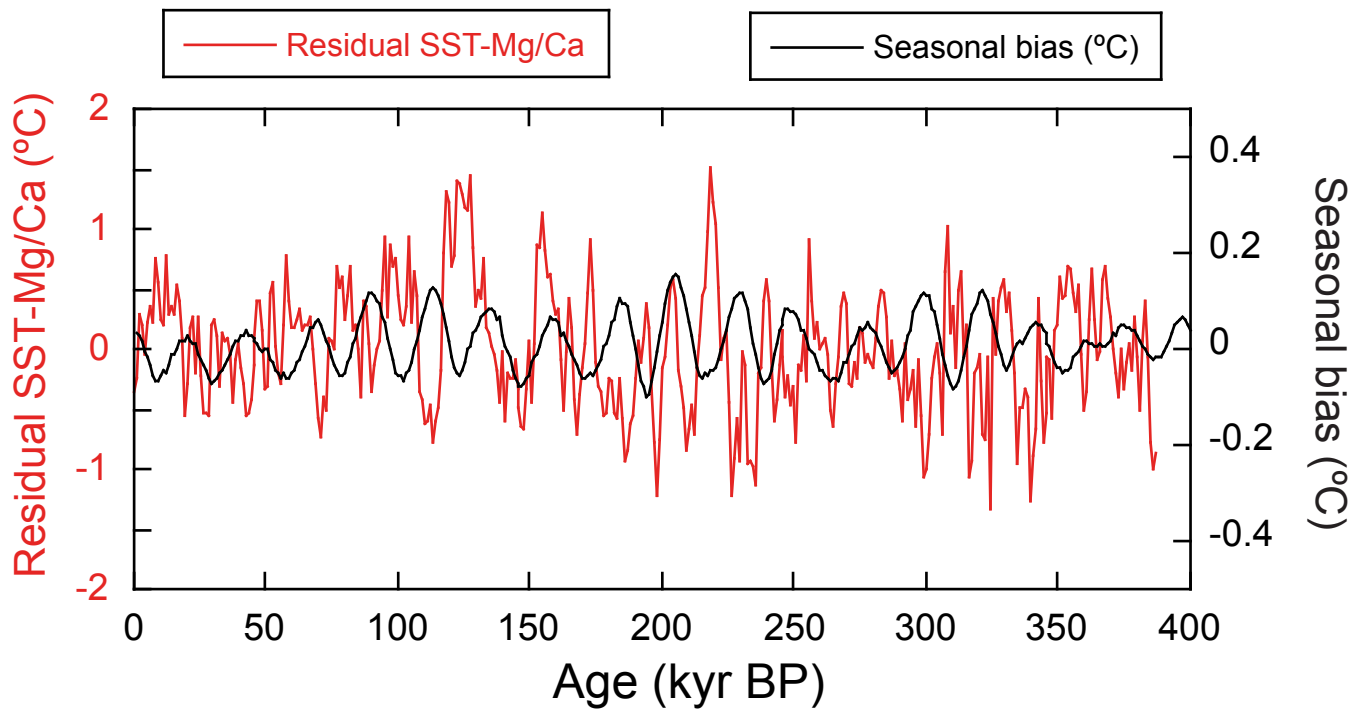


Figure S6

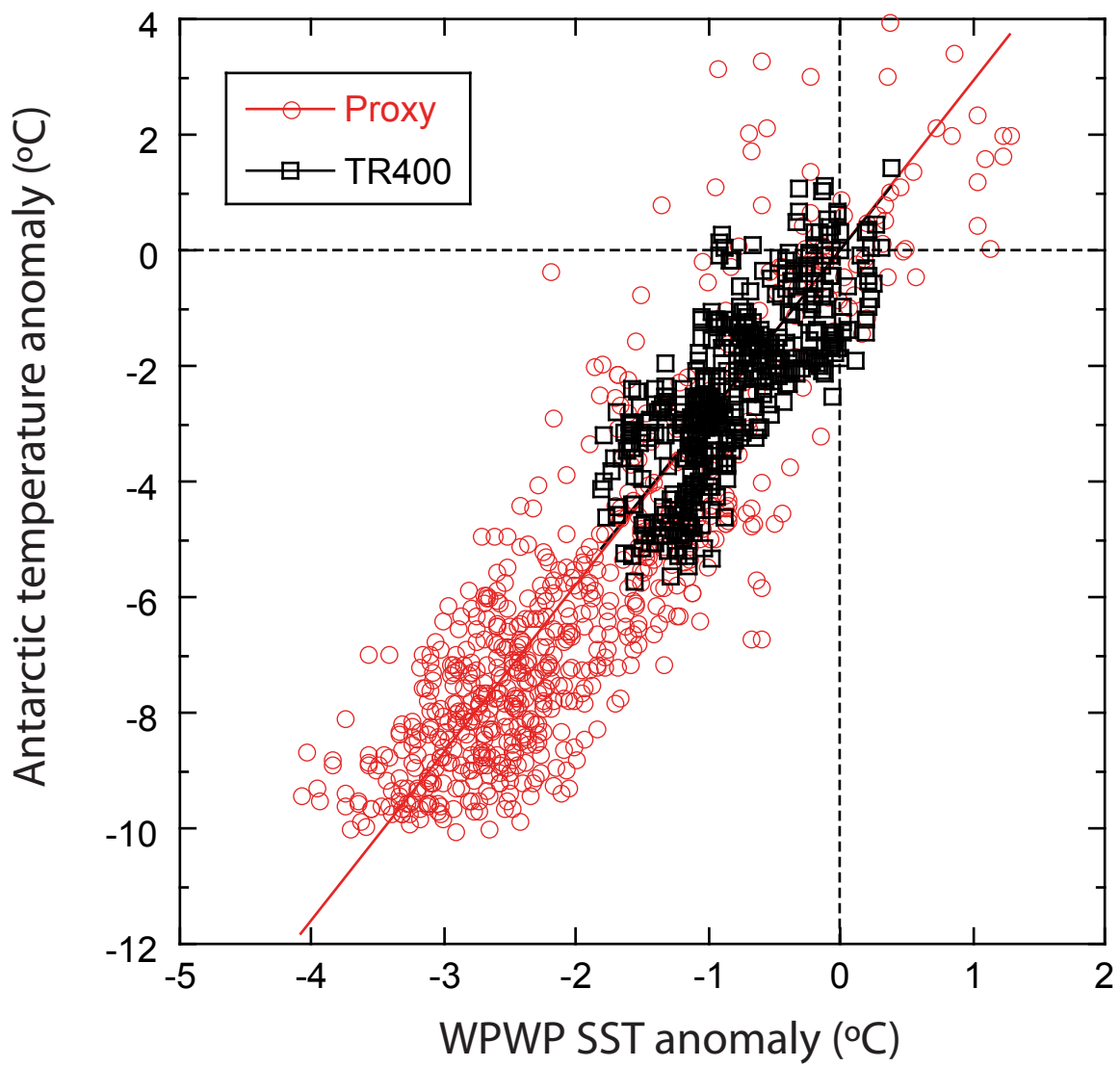


Figure S7