

## Appendix E

If variation in *G. bulloides*  $\delta^{13}\text{C}$  records the variation in  $\delta^{13}\text{C}$  of seawater DIC experienced by coccolithophores during calcification, then the difference of  $\delta^{13}\text{C}$  of coccoliths and *G. bulloides* indicates the magnitude of isotopic fractionation in coccoliths ( $\varepsilon_{\text{coccolith}}$ ).  $\varepsilon_{\text{coccolith}}$  increases towards glacial periods and is particularly lower during the interglacial MIS 7 (Figure E1). The temporal variation is not an artifact of changing assemblage composition of the coccolith fraction as it is observed in a range of assemblage compositions. The magnitude of  $\varepsilon_{\text{coccolith}}$  may be sensitive to changes in dissolved  $\text{CO}_2$  concentrations, a function of water temperature and atmospheric  $\text{CO}_2$ , as well as algal growth rate. The difference in magnitude of vital effects in MIS 5 relative to MIS 7 may in part be due to the higher coccolithophore productivity recorded during early MIS 5, which may have offset the effect of increased interglacial  $\text{CO}_2$  availability during MIS 5.

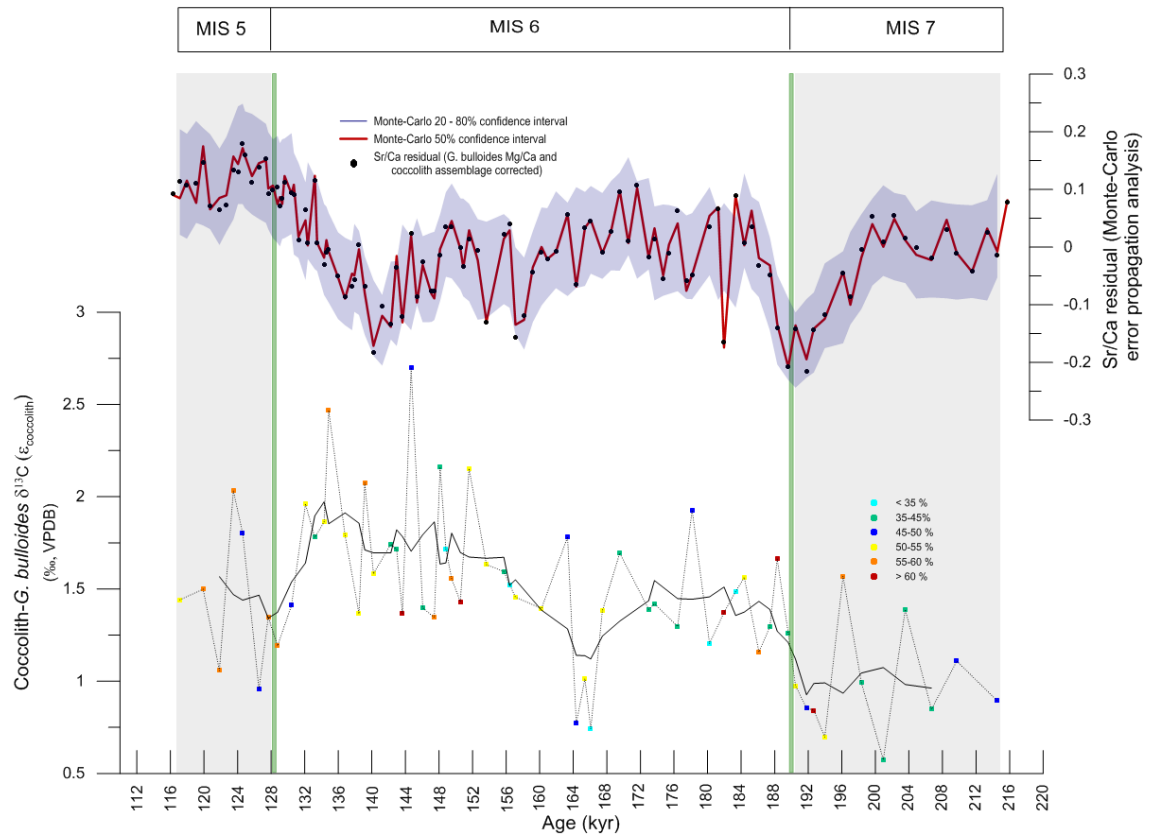


Figure E1. a) Coccolith Sr/Ca productivity record from core MD96-2080, as illustrated in Figure 3f. b) Difference of  $\delta^{13}\text{C}$  of the coccolith fraction and *G. bulloides* ( $\epsilon_{\text{coccolith}}$ ) in core MD96-2080 during the period of study showing increasing differences towards cold glacial periods and decreasing values during interglacials. 5 point running average is shown. Grey-shaded areas and green vertical lines as in Figure A1. Color coding of the sample points corresponds to the % carbonate contribution from larger coccoliths *C. leptoporus* and *H. carteri*.