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OCEAN BIOGEOCHEMISTRY

Carbon at the coastal interface

The extent to which coastal-ocean regions act as a sink for carbon dioxide has been enigmatic. An estimate based on more than 3 million observations suggests a smaller sink than was thought, concentrated at high latitudes.

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he thin strip of coastal ocean at the interface between land and the vast areas of the open sea contains some of the most biologically productive areas of the world's oceans. These coastal regions support most of the marine resources harvested by humans¹, despite covering only a tiny fraction of the ocean's surface area. One would expect these regions to act as strong sinks for atmospheric carbon dioxide because of their high rates of photosynthetic carbon fixation. Indeed, an early estimate² from the late 1990s suggested a coastal-ocean sink strength of about 1 petagram of carbon per year (Pg Cyr⁻¹; 1 petagram is 10^{15} g) — a value that, when combined with the open-ocean sink strength³ of about 2.3 ± 0.7 Pg Cyr⁻¹, would make the ocean by far the largest sink for the CO₂ that humans emit every year into the atmosphere³. Reporting in Global Biogeochemical Cycles, Laruelle and coauthors⁴ now show that the global coastal ocean is a much smaller CO₂ sink than was thought, amounting to only about 0.2 Pg Cyr^{-1} .

The authors' result continues a trend toward progressively smaller estimates, with the most recent previous publication⁵ suggesting a value between 0.2 and 0.5 Pg C yr^{-1} . But previous global estimates were based on the extrapolation of a small number of observations (often not more than a few hundred), whereas Laruelle *et al.* constrained estimates of the coastal air–sea CO₂ fluxes nearly globally using more than 3 million observations of the surface-ocean concentration of CO₂ from a recently compiled database⁶.

The authors divided the global coastal ocean into 45 segments and then determined the air–sea CO_2 flux for each segment using an approach that depended on the data density. In well-sampled regions, each segment was sub-divided into a grid, data were 'binned' into the grid's cells, and flux was computed directly on the basis of the gridded data. In areas that were sampled less well, the data were grouped according to water depth and other parameters. If the coverage was extremely limited, the researchers simply averaged the few available observations. Through this adaptive extrapolation method, they were able to constrain the air–sea CO_2 fluxes for 96% of the global coastal



Figure 1 | **Global map of air-sea CO₂ fluxes.** Laruelle *et al.*⁴ report the average flux density of coastal ocean regions (strong colours) based on more than 3 million observations. Data for the open ocean¹² are shown in fainter colours, for comparison. Blue colours represent carbon sinks, other colours represent carbon sources. Units of flux are moles of carbon per square metre per year.

ocean. This great leap forward in data coverage and methodology compared with all previous studies makes this estimate by far the most solid and reliable, thus providing an important constraint for the global carbon cycle.

A striking observation arising from the coastal air–sea CO_2 fluxes is that not all regions are sinks — many are sources of CO_2 to the atmosphere (Fig. 1). The flux distribution follows largely a latitudinal trend, with coastal regions at low latitudes generally losing CO_2 to the atmosphere, whereas those at higher latitudes, and especially in the Arctic, take it up. To a zero-order approximation, this follows the distribution of CO_2 fluxes in the open ocean, where the main sources to the atmosphere are also found at low latitudes, and the largest sinks at mid- to high latitudes⁷.

At first thought, it is surprising that the distribution of CO_2 sources and sinks does not follow that of primary production — for example, there are modest differences in CO_2 flux between the coastal ocean and the adjacent open ocean, even though the magnitude of biological productivity in these two systems differs by more than fivefold. The explanation of this puzzling finding is that many processes ultimately determine whether a region is a source

or sink of atmospheric CO₂ (ref. 8). In general, the outcome depends on a complex interaction between the extent of warming or cooling (which alters the solubility of CO_2), the amount of biological activity that removes CO₂ from the system, and how much CO_2 is supplied by upwelling and mixing from deeper waters. In coastal systems, additional complexity arises from the supply of inorganic and organic carbon from rivers on the adjacent land. The general latitudinal trend observed by Laruelle and co-workers suggests that warming and cooling dominate the distribution of the air-sea CO₂ fluxes, but the many deviations from this trend remind us how crucial other processes are in modifying the coastal carbon cycle.

Despite the vast increase in the amount of data considered in the new study compared with previous reports, the spatial, and especially the temporal, coverage of the data are still quite poor. The full seasonal cycle is resolved in only a handful of regions, and spatial coverage is nearly always too coarse to properly take into account the high spatio-temporal variability that characterizes coastal systems. The uncertainty associated with the time-mean estimate provided by Laruelle *et al.* therefore remains rather large.

The data are also insufficient to assess whether there are any trends in coastal fluxes, which is a serious gap when considering that the influence of human activity on coastal systems is increasing rapidly⁹. Of particular interest is how changes in the lateral supply of carbon from land will alter the sink-source balance of coastal systems¹⁰, and how other perturbations, such as ocean acidification, deoxygenation and increasing nutrient load, will manifest themselves in the coastal carbon cycle and ultimately alter the coastal ocean's ability to take up atmospheric CO₂.

Finally, a global uptake flux of 0.2 Pg C yr^{-1} by the coastal ocean may seem modest, but it represents an invaluable ecosystem service.

Using the current CO_2 price of about $\in 5$ (US\$6) per tonne of CO_2 in the European Union emissions-trading system¹¹, this service may be worth about $\in 1$ billion per year.

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- Pauly, D. & Christensen, V. Nature 374, 255–257 (1995).
- 2. Tsunogai, S., Watanabe, S. & Sato, T. *Tellus B.* **51**, 701–712 (1999).
- Ciais, P. et al. in Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental

Panel on Climate Change (Stocker, T. F. et al.) Ch. 6, 465–570 (Cambridge Univ. Press, 2013).

- Laruelle, G. G., Lauerwald, R., Pfeil, B. & Regnier, P. *Glob. Biogeochem. Cycles* http://dx.doi. org/10.1002/2014GB004832 (2014).
- 5. Bauer, J. E. et al. Nature 504, 61–70 (2013).
- Bakker, D. C. E. et al. Earth Syst. Sci. Data 6, 69–90 (2014).
- 7. Gruber, N. et al. Glob. Biogeochem. Cycles 23, GB1005 (2009).
- 8. Takahashi, T. et al. Deep Sea Res. Part II Top. Stud. Oceanogr. **56**, 554–577 (2009).
- 9. Doney, S. C. Science 328, 1512-1516 (2010).
- 10. Regnier, P. et al. Nature Geosci. **6**, 597–607 (2013).
- 11.Koch, N. et al. Energy Policy **73**, 676–685 (2014).
- 12.Landschützer, P. et al. Glob. Biogeochem. Cycles 28, 927–949 (2014).

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