
Organic carbon transfer and ecosystem functioning in the terminal lobes of the Congo deep-sea fan: outcomes of the Congolobe project

Rabouille C. ^{1,2}, Baudin F. ³, Dennielou Bernard ⁴, Olu Karine ⁵

¹ Univ Paris Saclay, UMR CEA CNRS UVSQ 8212, LSCE, Ave Terrasse, F-91198 Gif Sur Yvette, France.

² Univ Paris Saclay, IPSL, Ave Terrasse, F-91198 Gif Sur Yvette, France.

³ Sorbonne Univ, UPMC Univ Paris 06, CNRS, Inst Sci Terre IStEP, UMR 7193, 4 PI Jussieu, F-75005 Paris, France.

⁴ IFREMER, Ctr Brest, Dept REM, Unite Rech Geosci Marines, F-29280 Plouzane, France.

⁵ IFREMER, Ctr Brest, Dept REM, Unite EEP, Lab Environm Profond, F-29280 Plouzane, France.

1. Introduction: historical background, first observations and puzzling questions

The terminal lobe complex of the Congo deep-sea fan is a fascinating environment. It is located at 5000 meters depth and 750 km offshore Africa. It is currently connected with the Congo River by a canyon and deep-sea channel system starting in the estuary and still fed by turbidity currents (Babonneau et al., 2002; Savoye et al., 2000, 2009). It thus displays most of the characteristics of a subaquatic delta: large sedimentation rates, high concentration of organic carbon of terrestrial origin, active diagenesis. In addition, this peculiar zone of the deep ocean hosts unexpected biological assemblages resembling those of cold seeps: chemosynthetic vesicomyid bivalves, microbial mats and reduced sediments in metric size discrete habitats.

Previous observations in the Congo-Angola margin off the Congo River revealed that turbiditic activity is presently ongoing as shown by cable breaks (Heezen et al., 1964), direct observations of turbidity currents (Khripounoff et al., 2003; Vangriesheim et al., 2009) and sampling of turbiditic deposits from the last century (Savoye et al., 2009). Transport processes and recycling of biogenic particles (OC, BSi) across the margin are dominated by transfers through the canyon and deep-sea channel to the terminal lobe zone with very little lateral input mid-slope down to 4000 meters (Rabouille et al., 2009; Ragueneau et al., 2009). This material is essentially terrigenous which implies a direct origin from the river and low inputs of autochthonous marine production from the surface ocean (Baudin et al., 2010; Treignier et al., 2006).

The building of the present lobe complex started after the last avulsion of the turbiditic channel at an age estimated at about 4 ka BP (Picot, 2015). It covers an area of about 2500 km² with the last lobe showing an area of about 400 km² collecting the most recent deposits since the last 1000 y. However, the generally large sediment accumulation rates in the area (0.5 and 1.2 cm/y) suggest that the whole

41 lobe complex still collects significant amounts of sediments (Rabouille et al., 2017). The previous and
42 now abandoned lobe complex, which lies further north, was built between 6 and 4 ka (Picot 2015).

43 In the terminal lobe complex, two ROV video surveys conducted in 2000 (Savoie and Ondreas,
44 2000) detected the presence of patches of biological structures typically associated with cold seeps:
45 white patches of filamentous bacterial mats looking like those occurring in low-oxygen, high-sulfide
46 habitats (Nelson et al., 1989) and large bivalves visually attributed to the Vesicomidae family, known
47 from cold seeps to carry sulfide-oxidizing symbionts (Fiala-Médioni and Felbeck, 1990; Sibuet and Olu,
48 1998).

49 The existence of these dense ecosystems in the Congo lobe complex raised several questions
50 concerning their functioning: is it based, as in cold seeps, on chemosynthesis with methane and
51 sulphide oxidation? If so, is the source of reduced substances linked to the recycling of organic matter
52 in recent sediment layers, or is it related to deeper sources in fossil channels of the Congo deep-sea
53 fan? Are the inputs of terrestrial organic matter from the present Congo channel large enough and
54 easily metabolizable to start and sustain the development of these peculiar ecosystems? Do they
55 sustain high faunal densities? In the context of land-ocean carbon transfer, can we calculate mass
56 balances for the buried and mineralized carbon originating from the Congo River?

57 This special issue summarizes the first biological and biogeochemical survey performed in 2011-
58 2012 during the WACS (Olu, 2011) and CONGOLOBE (Rabouille, 2012) cruises on board the N.O.
59 Pourquoi Pas? within the framework of the Congolobe programme (www.congolobe.fr). It also
60 provides multidisciplinary answers to the above questions.

61

62 **2. Objectives and initial conceptual framework of the Congolobe project**

63 In this context, the Congolobe project aimed at establishing the functional relationship between
64 the organic matter input from the Congo canyon, its recycling in the sediment, and the structure and
65 functioning of the ecosystem based on chemosynthesis in the sedimentological context of the Congo
66 deep-sea fan.

67 The initial conceptual framework summarizing the hypotheses for the chemosynthesis-based
68 ecosystem functioning in the terminal lobe complex is represented in Figure 1.

69 It was hypothesized that large organic inputs from the Congo canyon are deposited in the terminal
70 lobes where they were mixed with pelagic inputs from the marine production. The proportions of the
71 two end-members were not known and represented one objective of the study in terms of quality and
72 biodegradability of the organic substrates. Indeed, depending on its reactivity, organic matter
73 degradation induces a series of microbial metabolisms which creates multiple geochemical and
74 microbiological gradients in the sediment: oxidants are generally consumed in the following order:
75 $O_2 \rightarrow NO_3^- \rightarrow Mn(IV) \rightarrow Fe(III) \rightarrow SO_4^{2-} \rightarrow CH_2O$, but alternative biogeochemical pathways can be present
76 in the lobe complex, as repartition and availability of oxidants in these turbidites are generally
77 unknown. The contribution of each pathway to early diagenesis was of prime interest in this study and
78 was extensively studied as some of these pathways lead to the production of sulfide which is the key
79 of symbiotic bivalves' colonization.

80 Indeed, sulfide plays a key role in the habitats where it fuels vesicomid symbionts and allows their
81 development (Fiala-Médioni and Felbeck, 1990). It can be produced by bacterial sulfate-reduction

82 which is a major pathway for organic matter mineralization in organic-rich sediments (Jorgensen, 1982)
83 and by Anaerobic Oxidation of Methane (AOM) performed by a microbiological consortium where
84 specific *Archaea* use methane as a reduced compound (Boetius et al., 2000). Constraining the balance
85 between these two processes and understanding the control on dissolved sulfide by iron were two
86 important objectives of this study. Furthermore, the understanding of the formation of methane in
87 deeper layers and its transport by diffusion or in microfaults throughout turbidites was another aim as
88 methane plays a key role in the production of sulfide by AOM.

89 The colonization and maintenance of discrete habitats by dense bivalve populations is another
90 striking feature of the lobe complex. The presence of chemosynthetic fauna was hypothesized to be
91 linked to the existence of “hot spots” of organic matter recycling and to the “shallow” early diagenesis
92 of the large organic inputs from the Congo River. Our project aimed at elucidating the functioning of
93 these ecosystems and calculating mass balances for the input of organic matter in the lobe complex.

94

95 The detailed objectives of the multi-disciplinary study of the Congolobe project were to:

96 - Determine quantitatively and qualitatively the origin and fate of biogenic particles (carbon and
97 silica) transferred to the terminal lobes. Using the data collected during the cruises, budgets of organic
98 carbon and biogenic silica were calculated especially recycling and burial. This study focused
99 particularly on the recycling of organic matter and the generation of reduced fluids which were linked
100 to chemosynthesis over different time scales, from years to thousand years.

101 - Study biological community structure from microbial communities to megafauna, and
102 ecosystem functioning in heterogeneous habitats of this lobe complex. The link with the direct input
103 of organic matter from the Congo channel and with the reduced fluids generated by the recycling of
104 formerly buried organic matter was investigated. The interactions between the living organisms’ size
105 classes from microbes to macrofauna were also considered. Functional biodiversity of the peculiar sites
106 observed in the lobe zone were compared to cold seep sites known from active “pockmarks” in this
107 region. Phylogenetic links among vesicomid bivalves from the lobes and from deep sea cold seeps
108 were investigated (Teixeira et al. 2013). Finally, the direct impact of turbiditic material deposits on the
109 benthic ecosystem outside of chemosynthesis-based habitats was investigated to assess the
110 consequence of this exceptionally high input of organic matter in deep sea sediment.

111 These two main objectives were divided in several tasks that served as building blocks of the
112 project:

113 - The understanding of sedimentological characteristics and history of the lobe complex

114 - The determination of terrestrial and marine organic matter inputs to the lobe zone using
115 biomarkers and isotopic tracers

116 - The estimation of recycling and burial of organic matter and biogenic silica in the lobe zone and
117 the establishment of mass balance for carbon and silica in the lobe area

118 - The assessment of the geographical distribution of megafauna and peculiar habitats, and their
119 functioning with respect to chemosynthesis *versus* detrital organic matter

120

121 **3. The multidisciplinary framework of Congolobe: articles’ description**

122 - *Geological context*

123 The Congo lobe complex is a large (2 525 km²) but thin (maximum 70 m thick) sedimentary body that
124 lies at an abyssal depth (5000 m) making its detailed structure and composition difficult to investigate
125 with ship-based acoustic tools. The acquisition of ROV-based high-resolution bathymetry and video
126 observations unveiled incredible and unsuspected morphologies that have provided new insights in
127 the build-up processes of channel-mouth lobes of silicoclastic mud-rich turbidite systems.

128 The general morphology and structure show that the lobe complex consists in five lobes that partly
129 overlap and have prograded onto the abyssal plain (Denniellou et al., 2017). The progradation,
130 aggradation and, probably, the size of a specific lobe are controlled by the local slope that allows the
131 turbidity currents to flow. When the top of a lobe becomes too flat, turbidity currents find their way
132 on steeper slopes on the side of the lobe and this process triggers the development of a new lobe
133 (Denniellou et al., 2017). In this respect, the most distal lobe receives more recent sediment than the
134 upstream lobes (Rabouille et al., 2017). However, sediment accumulation rates are very high over the
135 entire lobe complex, in the order of 0.3 to 12 cm.yr⁻¹ and show that turbidity currents flowing in the
136 feeding channel are largely unconfined (maximum channel depth is 40 m) and can spill over the whole
137 lobe complex (Rabouille et al., 2017).

138 The lobe complex is dominantly composed of turbiditic mud. A network of distributaries shows that
139 channelization and turbidity current spillover represent the sedimentary processes that first control
140 the development of the lobes. Sand is strictly confined in the feeding channel and distributaries and is
141 estimated to represent ca. 13% of the lobe complex (Denniellou et al., 2017). One striking feature is
142 the occurrence of mass wasting deposits nearly everywhere where cores were collected. Spectacular
143 pervasive mass wasting morphologies are also visible on the high-resolution bathymetry. Sliding and
144 blocks occur dominantly along the feeding channel and distributaries (Denniellou et al., 2017) but also
145 on the lobes rim (Croguennec et al., 2017) showing that mass wasting processes are a second order,
146 but intrinsic, control on the lobe build-up.

147 The extremely high accumulation rates also show that the lobe complex is young, probably about 4 ka
148 (Picot et al., 2016), and has significantly grown during the last century (Rabouille et al., 2017).
149 Moreover, transient state observed from pore water geochemistry (sulfate and methane) clearly
150 shows that sediment sliding has occurred during the last century (Croguennec et al., 2017). Thus, these
151 slidings were responsible for changes in the fluid flow within the sedimentary column, visible on the
152 pore water profiles.

153

154 - *The Congo terminal lobe complex: a large organic carbon burial center with hot spots of*
155 *anoxic recycling*

156 An intensive effort was carried out to quantify and characterize the organic matter using bulk
157 geochemical techniques (elemental and isotopic analyses, Rock-Eval pyrolysis), molecular
158 investigations (fatty acids, amino acids, hydrocarbon fractions, BIT index, molecular isotopy) as well as
159 optical observations of the organic particles. These studies allowed to assess the sources and the state
160 of preservation of the organic matter accumulated in the sediments of the lobe complex (background,
161 bacterial mats, vesicomid habitats) and to compare it with the bedload, the material recovered in
162 particle traps and suspended sediments from the Congo River watershed (Spencer et al., 2012; Talbot
163 et al., 2014).

164 The mud-rich sediments of the distal lobe contain high amounts of organic matter (~3.5 to 4.0 wt %
165 OC), the origin of which is a mixture of terrestrial higher-plant debris, soil organic matter and deeply
166 oxidized phytoplanktonic material (Baudin et al., 2017b; Stetten, 2015; Stetten et al., 2015). The
167 terrestrial fraction is dominant according to bulk and molecular organic geochemical signature. The
168 dominance of *n*-C₂₉ over aliphatic hydrocarbons, the occurrence of specific terpenes (gammacerene,
169 lupene, ursene or oleanene), the preeminence of very long chain fatty acids (≥ 24 atoms of C) and the
170 depleted isotopic ratios of most of the fatty acids (δ¹³C) all support the dominance of continental
171 derived C3 land plant organic matter in the sediment (Méganelle et al., 2017; Stetten, 2015; Stetten et
172 al., submitted). Optical analysis of the organic matter also supports this conclusion as delicate plant
173 structures, cuticle fragments and plant cellular material are well preserved in the distal lobe complex
174 (Schnyder et al., 2017).

175 The organic matter distribution seems homogeneous at different scales, from a single turbiditic event
176 to the entire lobe, and changes in accumulation rates have a limited effect on the quantity and quality
177 of the preserved organic matter, although the organic particles may be locally sorted by density or
178 buoyancy under the influence of turbidity currents (Schnyder et al., 2017). Nevertheless, particle
179 remobilization and oxygen diffusion in the first centimeter of the sediment could influence the
180 degradation of organic matter, but this process remains limited (Baudin et al., 2017a; Stetten, 2015).
181 The first degradation steps for terrestrial inputs in the suboxic sedimentary layers are reflected by high
182 phytadiene concentrations, while an even mode of C₁₂-C₁₈ saturated *n*-alkanes, presumably of bacterial
183 or fossil origin, is progressively enriched as the degradation proceeds further (Méganelle et al., 2017).
184 By contrast, organic matter recycling is particularly intense in vesicomid habitats and bacterial mats,
185 where biological action enriches the sediment in fatty acids and amino acids (Pruski et al., 2017). The
186 terrigenous organic matter, with an initial low reactivity, is bio-transformed and becomes more labile,
187 favoring its use by benthic macrofauna (Pruski et al., 2017).

188 This intense recycling of organic matter and the changes in the geochemical composition of pore
189 waters and sediments were investigated using a suite of *in situ* and *ex situ* techniques. Profilers and
190 benthic chambers were used *in situ* to determine diffusive and total recycling fluxes of oxygen
191 (Khrpounoff et al., 2017; Olu et al., 2017; Pastor et al., 2017; Pozzato et al., 2017), while porewater
192 composition was analysed using voltammetric microelectrodes and after extraction on board (Pastor
193 et al., 2017; Taillefert et al., 2017). The whole suite of electron acceptors and their by-products were
194 quantified both in background sediments, which constitutes most of the levees and channel sediments,
195 and within the different types of specific habitats such as microbial mats, vesicomid bivalves and
196 reduced sediments (Pastor et al., 2017; Pozzato et al., 2017; Taillefert et al., 2017). The biogeochemical
197 contrast is particularly striking between the background sediments and the habitats occurring in "hot
198 spots" patches. Background sediments are characterized by increased penetration of oxygen (≈ 1 cm),
199 the dominance of iron reduction (Beckler et al., 2016; Taillefert et al., 2017) and the absence of other
200 anoxic diagenetic pathways. Diffusive oxygen fluxes are around 2-5 mmol m⁻² d⁻¹ which is more active
201 than surrounding areas outside the Congo deep-sea fan indicating an increased mineralization activity
202 due to the inputs of terrigenous organic matter compared to nearby abyssal plains (Pozzato et al.,
203 2017; Wenzhoefer and Glud, 2002). Amorphous silica originating from the continent is recycled
204 together with the mineralization of terrigenous organic debris (Raimonet et al., 2015).

205 In contrast, pore water biogeochemistry and potential metabolic pathways are highly different in the
206 habitats. The oxygen penetration is limited to 1-2 mm, the oxygen flux rises to 8-40 mmol m⁻² d⁻¹ and,
207 in most instances, anoxic diagenesis involving complete sulfate reduction, methanogenesis and anoxic

208 oxidation of methane (AOM) is indicated by pore water composition (Pastor et al., 2017). The presence
209 of dissolved sulfide allows the colonization of habitats of several meters in size by compact vesicomyid
210 bivalves, often associated with white filamentous microbial mats looking like sulfur-oxidizing bacteria
211 and reduced black sediments. Evidence from microbiology indicates the simultaneous presence of
212 ANME (ANAerobic MEthanotrophs) and sulfate reducing bacteria suggesting that AOM is probably
213 active in these sediments although the sulfate-methane transition zone is not obvious (Pastor et al.,
214 2017). This observation could be linked to the deposition regime of the turbidites which involves high
215 accumulation rates ($0.5\text{-}20\text{ cm y}^{-1}$) and discontinuous deposition (1 turbidite every 6-17 years). The
216 edges of these habitats are well-defined and presumably linked to the precipitation of iron-sulfide
217 minerals (Taillefert et al., 2017), indicating that the coupling of iron and sulfur cycles plays a major role
218 in controlling the establishment and size of the habitats.

219 Mass balance indicates that burial largely outweighs remineralization processes. The averaged 3.5 to
220 4 wt% OC associated to high sedimentation rates in the Congo distal lobe complex implies a large burial
221 rate of organic carbon. Finally, the Congo deep-sea fan represents an enormous sink of terrestrial
222 organic matter when compared to other turbiditic systems over the world (Baudin et al., 2017a,
223 2017b). Recycling is limited by the short residence time in surface sediments and the refractory nature
224 of the substrate in most of the sediment except in the 'hot spots' colonized by the fauna which
225 reprocesses the initial organic matter deposition (Pruski et al., 2017).

226

227 *- Congo fan turbidites sustain high faunal densities and cold seep-like ecosystems*

228 Cold-seep like habitats have been observed, along about 70km in the five successive terminal lobes
229 from the entrance of the lobe complex to the main depositional area, but not further downstream.
230 Vesicomyid bivalve patches, microbial mats and reduced sediments were mapped using ROV still
231 imagery and mosaic processing along transects over several kilometers, and their distribution were
232 related to micro-topography (Sen et al., 2017). Vesicomyid clam patches were preferentially located
233 along the channel flanks affected by sliding, and on levees formed by channel overflow. They avoid the
234 channel center likely due to high current speed, but colonize it where it becomes wider and shallower
235 in the main depositional area, coping with exceptionally high sedimentation rates (Rabouille et al.,
236 2017). Locally, vesicomyids are closely related to slide scars or collapsed blocks of sediments, which
237 likely facilitate sulfide exhumation (Sen et al., 2017). Vesicomyid and microbial mats show a patchy
238 distribution, and a limited colonization of black sediment patches, reflecting a balance between
239 favorable habitat provision and disturbance by turbiditic currents.

240

241 Stable isotope and fatty acid analyses of faunal tissues evidenced that the fauna colonizing sulfide-rich
242 habitats mainly rely on microbial chemosynthesis, either through symbioses for vesicomyid clams, or
243 by the feeding on several microbial populations (aerobic or anaerobic methane oxidizers, sulfate-
244 reducers) by heterotrophic macrofauna (Pruski et al., 2017). The associated fauna shows a strong
245 similarity with cold seeps habitats, with characteristic patterns of high density and low diversity (Olu
246 et al., 2017). This macrofaunal community varies among habitats and sites and is dominated by several
247 sulfide-tolerant polychaete families which are hypothesized to change over time with habitat
248 geochemistry evolution and vesicomyid colonization, starting in microbial mats. Engineering effect on
249 geochemical gradients and infauna is hypothesized to differ between the two vesicomyid species,
250 *Christineconcha regab* and *Abyssogena southwardae*, related to their ability to burrow (Olu et al.,

251 2017). Indeed, these species differ in the properties of their respiratory pigments allowing them to
252 tolerate hypoxic or anoxic conditions (Decker et al., 2017).

253
254 Respiration rates were estimated in situ with ROV-manipulated benthic chambers above vesicomid
255 clusters, and on individual clams isolated from the sediment (Khripounoff et al., 2015; 2017). Both
256 techniques revealed high oxygen uptakes (from 6.2 to 22.9 $\mu\text{mol.gdw}^{-1}.\text{h}^{-1}$) and large fluxes of dissolved
257 inorganic carbon (DIC) and methane in the range of those measured for cold seep bivalves, showing
258 high levels of anaerobic metabolism.

259
260 The large terrestrial organic matter inputs also sustain exceptional macrofaunal densities in the whole
261 lobe sediments (on the levees, the channel flanks and in the depositional area, outside of
262 chemosynthesis-based habitats). In the sediments of the terminal lobe complex, these densities were
263 7 to 8 higher than in the abandoned lobe complex, where they seem also high compared to similar
264 depth abyssal sediment (Olu et al., 2017). Densities are consistent with total oxygen uptake, high and
265 quite homogeneous over the entire lobe complex, lower in the abandoned lobe and in abyssal
266 sediment at similar depth (Olu et al., 2017). Megafauna dominated by detritivores (large size
267 agglutinated foraminiferans, holothurians) and filter/suspension-feeders (sponges) rely on these
268 detritic, turbiditic, mainly terrestrial-originated inputs (Pruski et al., 2017; Sen et al., 2017;). This fauna
269 is supported by the exceptional amounts of organic carbon, which provide enough biopolymeric
270 carbon and proteinaceous nitrogen, despite poor digestibility of soil-derived organic matter (Pruski et
271 al., 2017). Nevertheless, macrofaunal communities, dominated in non-reduced sediments by
272 peracarid crustaceans but also by cossurid and spionid polychaetes were more closely related to those
273 colonizing low-flow cold seeps than those of typical deep-sea sediment, likely related to the high
274 organic content of the sediments (Olu et al., 2017).

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277 **4. The conceptual model revisited**

278 The new findings published in the 15 articles of this special issue provided a new conceptual framework
279 for the functioning of this complex area. The first striking feature is the existence of two different
280 environments at a distance of a few kilometers (and sometimes a few meters). This is what we named
281 “background” and “reduced habitats”. The reduced habitats are largely related to sedimentological
282 features such as large deposition of recent and relatively fresh terrestrial organic matter and to
283 sediment sliding on the flanks of the feeding channel (Croguennec et al., 2017; Dennielou et al., 2017).
284 These processes promote the production/exhumation of reduced substrates either from previously
285 buried sulfides or from intense diagenesis of recently deposited organic matter (Pastor et al., 2017;
286 Taillefert et al., 2017). They are colonized by dense populations of chemosynthetic vesicomid bivalves
287 and microbial mats (Olu et al., 2017; Sen et al., 2017) which feed on sulfide produced by sulfate
288 reduction or anaerobic oxidation of methane in proportion that remains unknown (Pastor et al., 2017).
289 In these reduced habitats, the penetration of oxygen is limited (Pozzato et al., 2017), the metabolism
290 is amplified as recorded by *in situ* incubation experiments (Khripounoff et al., 2017) and organic
291 substrates generation (Pruski et al., 2017). The macrofauna is dominated by species tolerant to or
292 taking advantage of sulfide with their symbionts and is very similar to cold seep assemblages (Decker
293 et al., 2017; Olu et al., 2017), though the source of methane might be located in shallow sediments
294 (Pastor et al., 2017).

295 On the contrary, most of the levees show very different biological assemblages dominated by large
296 sized foraminifera (*Bathysiphon* sp.; Sen et al. 2017). The biogeochemistry of this zone show large
297 content of terrigenous organic matter (Méjanelle et al., 2017; Schnyder et al., 2017) which is very
298 similar in concentration and type to the reduced habitats (Baudin et al., 2017b) but larger oxygen
299 penetration depth (Pozzato et al., 2017) and a noticeable absence of sulfide with large concentration
300 of dissolved iron in pore waters (Beckler et al., 2016). These features indicate a shift from sulfide-
301 dominated to iron-dominated diagenesis when moving from reduced habitats to the background
302 exemplified by the levees (Taillefert et al., 2017).

303 Overall, the terminal lobe complex is a hot spot of organic carbon burial compared to the hosting
304 abyssal plain (Baudin et al., 2017a) with specific burial of terrigenous organic carbon as large as 1000
305 g C m⁻² y⁻¹ in the most distal lobe and average values of around 100 g C m⁻² y⁻¹ over the 2525 km² of the
306 active lobes of the deep-sea fan. This mega-burial is accompanied by equivalent quantities of biogenic
307 silica originating from land (Raimonet et al., 2015). These values are 1000 to 10000 times larger than
308 average burial in the deep Atlantic Ocean (Rabouille et al., 2017; Stetten et al., 2015). Recycling
309 represents at most 25% of the deposited fluxes with lower values (5%) in the most distal lobes.

310

311 **5. An attempt to understand temporal evolution of the ecosystems**

312 The different types of reduced habitats (reduced sediment, microbial mats, vesicomid clusters) that
313 co-occur in the terminal lobe complex may correspond to a hypothetical sequence of successional
314 stages. Observations based on macrofaunal community patterns (composition, diversity, vertical
315 distribution), vesicomid population characteristics (size, species) and geochemical composition and
316 gradients can be interpreted as temporal evolution of these chemosynthesis-based habitats, with
317 interplay between the biotic and abiotic components involving biogeochemical as well as biological
318 controls (Olu et al., 2017; Taillefert et al., 2017). From these observations, it is hypothesized that
319 reduced sediments and microbial mats are the first to develop along the sulfide-rich patches generated
320 by exhumation or intensification of diagenetic processes. Vesicomid clams that colonize microbial
321 mats contribute to deepen the suboxic layer and displace the sulfide-rich layer deeper in the sediments
322 by bioturbation, bio-irrigation and sulfide uptake. Bio-irrigation also allows sulfate replenishment
323 leading to an increase of net sulfide production at depth (Pastor et al., 2017) and oxygen penetration
324 which contributes to favorable conditions for associated macrofauna. In the first steps of colonization,
325 the small size macrofauna may also modify the biotope, by grazing microbial mats, or by irrigation of
326 the first sediment layers. Indeed, the depth of the sulfide-rich layer has been observed to vary with
327 the vesicomid shell length and between the vesicomid species, assumed to differently bio-irrigate
328 the sediment due to their burrowing abilities (Decker et al., 2017; Olu et al., 2017). When moving 10
329 km away from the main active channel, the sulfide-enriched layer occurs deeper in the sediment
330 (Pastor et al., 2017; Taillefert et al., 2017), which is accompanied by the dominance of another
331 vesicomid bivalve of larger size, *A. southwardae*, assumed to burrow deeper and be adapted to
332 efficient O₂ storage and transport (Decker et al. 2017). The sulfidic layer eventually disappears in the
333 abandoned northern lobe (50km North) together with the cold-seep like habitats.

334 Although similarities with community succession at cold-seeps have been observed, the limited
335 resources, without methane stock at depth such as in cold seeps is assumed to limit the development
336 of chemosynthesis-based habitats to a very patchy distribution: only a limited number of black reduced
337 sediment patches observed in the youngest lobe are colonized by microbial mats and vesicomids,
338 (Sen et al., 2017). The high rate of disturbance by turbidite deposition also limits the development of

339 the habitats and is also hypothesized to increase the rate of faunal succession in very short-lived
340 habitats. Indeed, with a turbidite frequency estimated every 6 to 17 years (Dennielou et al., 2017), a
341 vesicomid cluster should develop within 5 to 10 years, which is consistent with the highest estimated
342 growth rates for cold seep vesicomids (see discussion in Sen et al., 2017). Long-term interdisciplinary
343 observation and sampling at inter-annual scale is clearly needed to test the proposed habitat
344 succession model, to assess vesicomid growth rates and species succession, to clarify the rate of
345 macrofaunal community dynamics in relation with biogeochemical and sedimentological changes as
346 they both are the final drivers of the ecosystem evolution.

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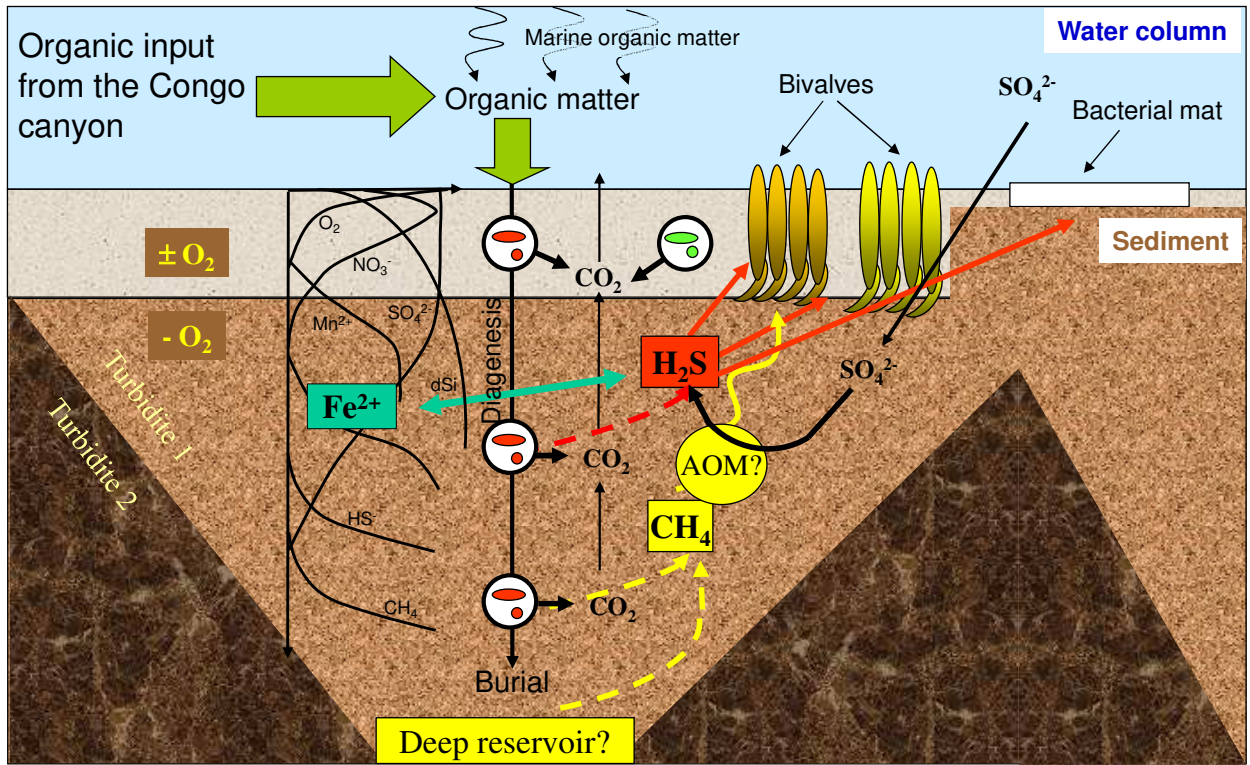
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498 Figure caption:

499 Figure 1: Initial conceptual framework for our study of the Congo lobe ecosystems.

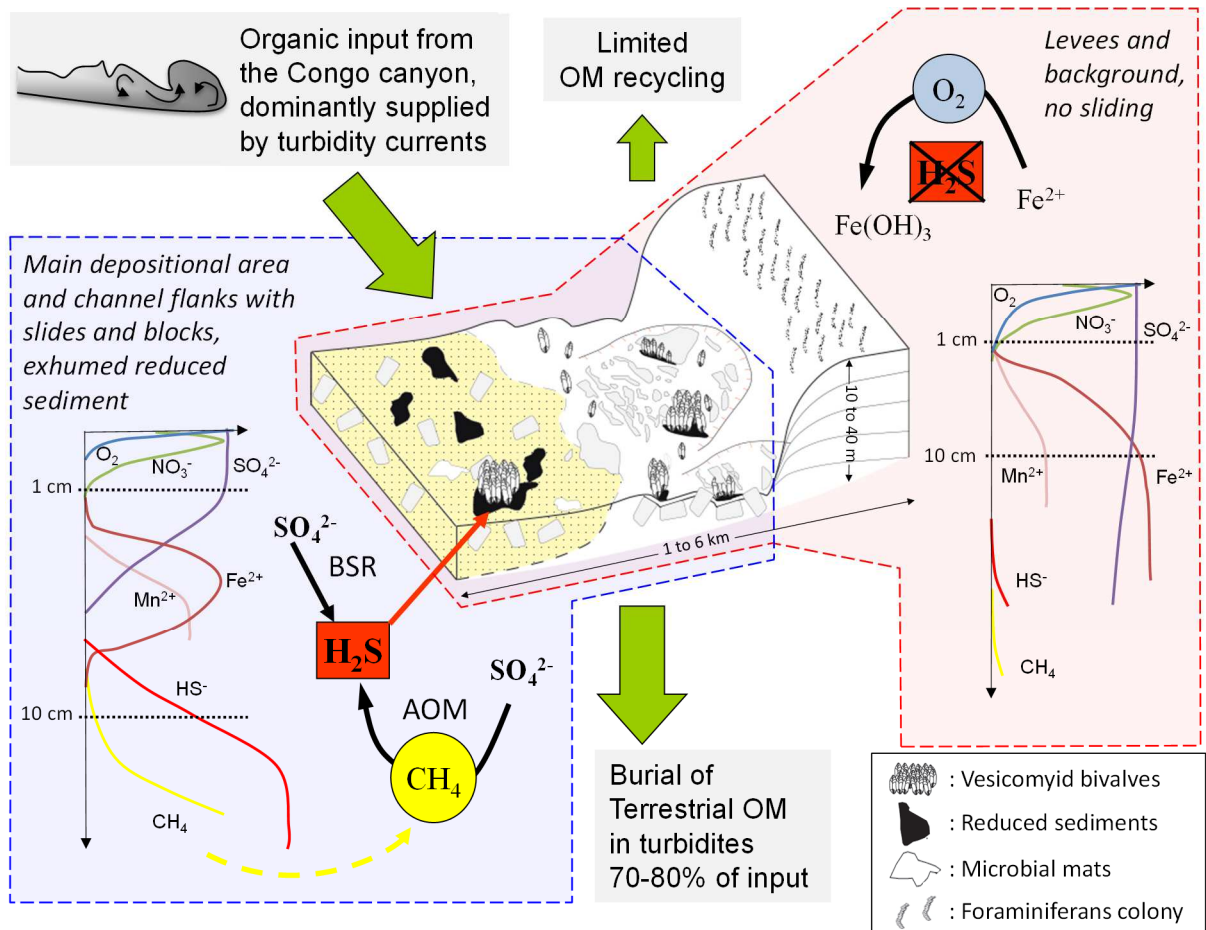
500 Figure 2: Final conceptual framework emphasizing the link between geological structures,
501 biogeochemical processing of organic matter and fauna. White parts of the schematic lobe complex
502 represent fine muds, while yellow dotted parts represent more sandy sediments.

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505 Figure 1

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508 Figure 2

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