
Compact symmetric Poisson equation discretization for non-hydrostatic sigma coordinates ocean model

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Abstract :

In anticipation of relaxing the hydrostatic assumption in a sigma coordinates primitive equations ocean model, we show how a projection method can be designed with the use of a compact symmetric 15-point stencil for the Poisson equation. This is achieved by recognizing that, owing to the non-orthogonality of the grid, the velocity has a contravariant and covariant set of components. The two sets play a different role in the primitive equations: the contravariant components enter the definition of the model fluxes, whereas the covariant components experience the forces and in particular the pressure gradient. By treating these two sets separately, the discretized gradient and divergence operators are simple finite differences. The two sets of components are related via a linear transformation, the metric tensor, which is entirely determined by the kinetic energy. We show how the spatial discretization of the kinetic energy fully controls the Poisson equation discretization, including its boundary conditions. The discretization of the Poisson equation is shown to converge at second order and to behave as well or better than alternative methods. This approach is a prerequisite to implement later an efficient Poisson solver, such as a multigrid algorithm.

Highlights

► A new formulation of the non-hydrostatic primitive equations in sigma coordinates is proposed. ► Momentum and flux are two different objects related via a linear transformation. ► The Poisson equation for pressure is discretized with a compact symmetric stencil. ► The approach is a prerequisite to implement a fast Poisson solver for a forthcoming non-hydrostatic version of ROMS.

Keywords : Non-hydrostatic, Sigma coordinates, Poisson equation

1 **1. Introduction**

2 With the growing interest on processes at submesoscale (Nikurashin et al.,
3 2013; Molemaker et al., 2015; Callies et al., 2015; Gula et al., 2015) the ocean
4 modelling community needs non-hydrostatic (NH) ocean circulation models.
5 This will allow to investigate in realistic contexts the full range of internal
6 waves, symmetric instability, convection processes, etc. and to bridge the gap
7 with LES (Sullivan et al., 1994). A few circulation models do already handle
8 the non-hydrostatic physics properly: MITgcm (Marshall et al., 1997), POM
9 (Kanarska and Maderich, 2003), ROMS (Kanarska et al., 2007), Symphonie
10 (Auclair et al., 2011), GETM (Klingbeil and Burchard, 2013). However, NH
11 simulations on large grids, in realistic regional configurations are still awaited.
12 It is timely to resolve this. The reason for the NH circulation models rarity
13 is that relaxing the hydrostatic assumption, while keeping the Boussinesq
14 assumption, is not straightforward. At least two levels of difficulties must
15 be faced: i) the requirement to solve a 3D Poisson equation for the pressure
16 at each time step, made more complex in the case of sigma coordinates by
17 ii) the non-orthogonality of the grid. This second difficulty is absent in z -
18 coordinates models. Incidentally, for models with a time-splitting on the
19 free surface, a third difficulty is to properly handle the coupling between the

20 non-hydrostatic pressure and surface waves (Auclair et al., 2011).

21 The first problem amounts to solving a linear system of equations with
22 as many equations as number of grid points. For the intended purpose
23 of high-resolution turbulent simulations we have in mind grids as large as
24 $2000 \times 2000 \times 1000$ grid points, corresponding to $N \sim 10^9$ coupled equations.
25 It is a classical well known problem of High Performance Computing (HPC).
26 A possible way to circumvent the problem consists in relaxing the incom-
27 pressibility constraint, which forces to cope with sound waves. The pressure
28 can then be computed by integrating in time the compressible physics using a
29 time-splitting technique. In the atmospheric community, this is the approach
30 implemented in WRF (Skamarock and Klemp, 2008). In the ocean the ratio
31 between advective speed and the phase speed of sound waves is much more
32 unfavorable, leading to a required time split that is two orders of magnitude
33 larger. An approach that uses artificial incompressibility to slow down the phase
34 speed of sound waves and hence reduce the required ratio in split time steps
35 is currently under development as part of the French regional ocean modeling
36 project CROCO, but not available in a published manuscript (Francis Au-
37 clair, private communication). Klingbeil and Burchard (2013) follow a similar
38 approach where a time-splitting on the baroclinic time step is introduced to
39 achieve the incompressibility constraint. These methods trade the price of
40 solving the Poisson problem with accepting a non-strict divergence-freeness
41 of the flow.

42 Otherwise, the numerical implementation necessarily requires the use of
43 an iterative solver to solve the Poisson equation. Many different numerical
44 methods exist but only a few of them succeed at maintaining a uniform

45 convergence rate for large grids. This is the case of the multigrid technique
46 whose computational cost scales as $N \log N$, which is considered to be the
47 optimal scaling. It has been shown recently to perform very well with large
48 grids on massively parallel cluster, up to 10^{10} degrees of freedom and 65536
49 cores, in the context of atmospheric modelling (Müller and Scheichl, 2014).
50 To get the best performances, the stencil for the Poisson equation needs
51 to be symmetric and compact. These two properties reduce the number of
52 floating point operations, the amount of data storage and the data movement
53 between the cpu and the memory. In modern computer architecture, data
54 movement is the bottleneck for high performances (Williams et al., 2009).
55 Any numerical method that reduce data transfer is worth it. In the present
56 study, the compactness means a 15-point stencil, compared to the usual 25-
57 point one (Auclair et al., 2011). The symmetry allows to almost halve the
58 number of coefficients for the matrix since only 8 coefficients are needed for
59 a 15-point stencil (one main and 7 lower diagonals). Overall the present
60 matrix offers a reduction by a factor 3 on the number of matrix coefficients
61 compared to existing methods. In anticipation of implementing such solution
62 into a sigma-coordinates model we present here a way to achieve compactness
63 and symmetry for the Poisson equation.

64 The second problem is due to the non-orthogonality of the grid. One
65 consequence is that the horizontal pressure gradient can not be computed as
66 a simple finite difference between two adjacent horizontal grid cells. Chain
67 rule and vertical interpolation should be used (Shchepetkin and McWilliams,
68 2003) resulting in wide stencils for the gradient operator. The divergence of
69 vertical flux behaves similarly because the vertical flux involves horizontal

70 velocities. Since the Poisson equation arises as the successive action of the
71 gradient and the divergence, the resulting stencil for the Poisson equation
72 is tall in the vertical, involving overall 25 points (Auclair et al., 2011). The
73 Poisson equation couples points over 5 levels in the vertical and 3 in both
74 horizontal directions. The stencil can be made symmetric by defining the gra-
75 dient as minus the adjoint of the divergence. This is the so-called compatible
76 discretization (Taylor and Fournier, 2010). It is also possible to discretize
77 the Poisson matrix directly and independently of how the gradient and the
78 divergence are discretized (Kanarska et al., 2007). In that case, the stencil
79 is compact with 15 points. The major drawback of such a method is that it
80 fails at maintaining the compatibility between operators which impacts the
81 energy conservation.

82 In this paper we show how the two problems can be solved jointly. The
83 first step is to recognize that because of the non-orthogonality of the coordi-
84 nates system, the velocity has two sets of components: a contravariant one
85 and a covariant one. The two sets are related via a linear transformation,
86 the metric tensor, which is invertible. The metric tensor is completely deter-
87 mined by the kinetic energy. The two sets of components play a different role
88 and must be treated distinguishly. The contravariant components appear in
89 the definition of fluxes (volume, tracer, momentum), whereas the covariant
90 components appear in the force budget. Using the framework of the discrete
91 differential geometry (Desbrun et al., 2008) we show that the contravariant
92 components must be discretized at the cell faces of the primal grid whereas
93 the covariant components should be discretized at the edges of the dual grid
94 (the lines joining the cell centers). Such discretization is in line with the

95 C-grid staggering, and the finite volume discretization for tracers. It also
96 completely adheres to Thuburn and Cotter (2012) way to discretize shallow
97 water equations on non-orthogonal grids. The divergence and the gradient
98 operators have then trivial discretization: they are simple finite differences.
99 In the spirit of Molemaker et al. (2005) and Dubos et al. (2015) we then
100 deduce the discrete Poisson equation directly from the discretized kinetic en-
101 ergy and the discretized divergence operator. The price of this approach is
102 to prognose either the covariant components or the contravariant ones. For
103 sake of completeness we present the momentum equations written for the
104 covariant components in both vector-invariant form and in flux-form. The
105 vector-invariant form is similar to a standard Cartesian formulation whereas
106 the flux-form involves additional pseudo-force terms. Up to now, sigma-
107 coordinates models prognose Cartesian components of the velocity, namely
108 horizontal and vertical, which are a blend of contravariant and covariant
109 components. This causes unnecessary wide stencil for the divergence and
110 the gradient and finally a too wide stencil for the Poisson equation. By ex-
111 pliciting and clarifying the subtleties induced by non-orthogonal coordinates,
112 this paper offers a roadmap for an efficient implementation of a projection
113 method.

114 The paper is organized as follows. We define in section 2 the contra-
115 and covariant sets of components for the velocity and present the model
116 continuous equations. In section 3, we present the spatial discretization of
117 the variables, the divergence and the discretized kinetic energy and derive all
118 the other discretizations from them. In section 4 we discuss the properties
119 of the discretization on a test-case. A conclusion is given in section 5.

120 **2. Formulation**

121 As stated in the introduction, the objective of this study is to provide a
 122 discrete Poisson operator that can be solved efficiently by iterative methods
 123 on very large grids. In this section, we will introduce a dual representation
 124 of the velocities on the grid, which will prove to be advantageous in arriving
 125 at a compact, symmetric, discrete operator. Subsequently, we will present
 126 the governing equations of motion for these prognostic variables.

127 *2.1. Flux and momentum*

128 In orthogonal curvilinear coordinates (ξ, η, z) , the velocity of a fluid par-
 129 cel is naturally expressed as $\mathbf{u} = h_\xi \dot{\xi} \mathbf{i} + h_\eta \dot{\eta} \mathbf{j} + \dot{z} \mathbf{k}$, where the dot denotes
 130 the Lagrangian time derivative, $\mathbf{i}, \mathbf{j}, \mathbf{k}$ are the unit vectors in horizontal and
 131 vertical directions, and (h_ξ, h_η) are the Lamé coefficients in the horizontal di-
 132 rections. We define $(u_c, v_c, w_c) = (h_\xi \dot{\xi}, h_\eta \dot{\eta}, \dot{z})$ and denote them, with a slight
 133 abuse of usage, the Cartesian velocity components. The terrain following
 134 sigma coordinates are (ξ, η, σ) , for which the vertical position is $z(\xi, \eta, \sigma, t)$.
 135 The time dependency in z is driven by the sea surface height variations and
 136 yields a breathing grid. Using the chain rule, this causes

$$\mathbf{u} = u_c \mathbf{i} + v_c \mathbf{j} + (u_c s_\xi + v_c s_\eta + h_\sigma \dot{\sigma} + \partial_t z) \mathbf{k}, \quad (1)$$

137 where

$$s_\xi = \frac{1}{h_\xi} \frac{\partial z}{\partial \xi}, \quad s_\eta = \frac{1}{h_\eta} \frac{\partial z}{\partial \eta} \quad (2)$$

138 are the slopes of σ -surfaces, h_σ is the Lamé coefficient in the vertical direction
 139 and $\partial_t z = \partial z / \partial t$ accounts for the breathing of grid. We denote $(U, V, W) =$

140 $(h_\xi \dot{\xi}, h_\eta \dot{\eta}, h_\sigma \dot{\sigma})$ the sigma coordinate velocity components. They are related
 141 to the Cartesian velocity components by

$$\begin{aligned} u_c &= U \\ v_c &= V \\ w_c &= U s_\xi + V s_\eta + W + \partial_t z. \end{aligned} \quad (3)$$

142 Hereafter we distinguish

$$W^a = W + \partial_t z = w_c - U s_\xi - V s_\eta, \quad (4)$$

143 the vertical velocity relative to a fixed grid, from W , the vertical velocity
 144 through the breathing grid. In the sequel, we refer to (U, V, W^a) as the con-
 145 travariant components of velocity. Notice that it is a slight abuse of language,
 146 since strictly speaking the contravariant components of velocity are just the
 147 material derivatives of the coordinates, i.e. $(\dot{\xi}, \dot{\eta}, \dot{\sigma})$, while (U, V, W^a) incor-
 148 porate the Lamé coefficients $(h_\xi, h_\eta, h_\sigma)$. With this definition, (U, V, W^a) are
 149 close to the Cartesian components typically manipulated by existing hydro-
 150 static models, which simplifies the implementation of our method in such
 151 models. A similar caveat applies to the covariant velocity components de-
 152 fined below.

153 The velocity can be written in vectorial form as

$$\mathbf{u} = U (\mathbf{i} + s_\xi \mathbf{k}) + V (\mathbf{j} + s_\eta \mathbf{k}) + W^a \mathbf{k}, \quad (5)$$

154 for which it is clear that $(\mathbf{i} + s_\xi \mathbf{k}, \mathbf{j} + s_\eta \mathbf{k}, \mathbf{k})$ are the basis vectors aligned with
 155 the coordinates (ξ, η, σ) . Interestingly there are two possible interpretations
 156 for the contravariant components: i) according to (3) and (4), U and V are

157 horizontal components and W^a is the vertical component normal to sigma iso-
 158 surfaces ; ii) according to (5), U and V are along coordinates lines and W^a is
 159 the vertical component. This dual interpretation is a source of confusion. The
 160 first interpretation is motivated by discrete differential geometry (Desbrun
 161 et al., 2008), the second interpretation is more in line with the physicist way
 162 of seeing the contravariant components. In this paper we will adopt the first
 163 interpretation and motivate it more thoroughly in section 3.

164 Now that we have the contravariant components of the velocity we can
 165 define the covariant components¹. For that we use the kinetic energy K as
 166 a natural inner product, allowing to define the metric tensor. In terms of
 167 (U, V, W^a) K reads

$$K = \frac{1}{2} [U^2 + V^2 + (Us_\xi + Vs_\eta + W^a)^2] , \quad (6)$$

168 where we have used W^a , not W , because the grid motion should not be
 169 accounted for in the kinetic energy. The covariant set of components (u, v, w)
 170 is defined such that

$$K = \frac{1}{2}(uU + vV + wW^a). \quad (7)$$

¹As for the contravariant velocity components, our definition of covariant components is slightly abusive. The standard covariant components are obtained by multiplying ours by the corresponding Lamé coefficients. Again, our definition results in (u,v,w) being closer to the Cartesian components, simplifying the implementation in existing hydrostatic models while ensuring key relationships (7,8).

171 We can find these components by differentiating K with respect to U, V, W^a

$$\begin{aligned} u &= \partial K / \partial U = U(1 + s_\xi^2) + V s_\xi s_\eta + W^a s_\xi \\ v &= \partial K / \partial V = U s_\xi s_\eta + V(1 + s_\eta^2) + W^a s_\eta \\ w &= \partial K / \partial W^a = U s_\xi + V s_\eta + W^a. \end{aligned} \quad (8)$$

172 The covariant components can be related to the Cartesian ones by combining
173 equations (3) and (8)

$$\begin{aligned} u &= u_c + w_c s_\xi \\ v &= v_c + w_c s_\eta \\ w &= w_c. \end{aligned} \quad (9)$$

174 We may use these components to rewrite the velocity in vectorial form

$$\mathbf{u} = u \mathbf{i} + v \mathbf{j} + w (\mathbf{k} - s_\xi \mathbf{i} - s_\eta \mathbf{j}). \quad (10)$$

175 Again we have a double interpretation for the covariant components: i) ac-
176 cording to (9) (u, v, w) are aligned with the coordinates lines and in particular
177 w is the vertical component ; ii) according to (10), (u, v, w) are orthogonal
178 to coordinates lines. To be consistent with our previous choice, we need to
179 chose the first interpretation.

180 The system (8) defines the metric tensor relating contravariant to co-
181 variant components. It is a 3×3 full matrix. Interestingly the inverse
182 transformation has less entries

$$\begin{aligned} U &= u - w s_\xi \\ V &= v - w s_\eta \\ W^a &= -u s_\xi - v s_\eta + w(1 + s_\xi^2 + s_\eta^2). \end{aligned} \quad (11)$$

183 We define

$$\mathbf{T} = \begin{pmatrix} 1 & 0 & -s_\xi \\ 0 & 1 & -s_\eta \\ -s_\xi & -s_\eta & 1 + s_\xi^2 + s_\eta^2 \end{pmatrix} \quad (12)$$

184 the matrix transforming (u, v, w) into (U, V, W^a) . It can be used to express
185 the kinetic energy

$$K = \frac{1}{2} (u, v, w) \mathbf{T} (u, v, w)^T, \quad (13)$$

186 seen as an inner product between covariant components. The \mathbf{T} matrix is
187 the inverse of the metric tensor, it plays a central role in the discretization
188 we propose.

189 At this stage we have set the framework. We have three different sets of
190 velocity components available: Cartesian, contravariant and covariant. The
191 Cartesian one is appealing but it does not correspond to either the coordi-
192 nates lines or to the lines orthogonal to the coordinates hyperplanes. By
193 contrast the contravariant and the covariant sets are natural. None of them
194 is superior and they both have their interest. The covariant components
195 are the natural prognostic variables of the momentum equations, while the
196 contravariant components appear in all the fluxes and in particular into the
197 incompressibility equation. For these reasons, (u, v, w) will be called momen-
198 tum and (U, V, W^a) flux. In this approach flux and momentum are clearly two
199 different objects. They handle their “horizontal” and “vertical” components
200 very differently.

201 Finally because the two objects are related by the kinetic energy, their
202 relation depends on whether the hydrostatic assumption is made or not.

203 Within the hydrostatic assumption, K simplifies into

$$K_{hydro} = \frac{1}{2} [U^2 + V^2] . \quad (14)$$

204 This implies $w = 0$, which expresses the fact that the vertical momentum
 205 equation is a diagnostic equation. It also implies $(u, v) = (U, V) = (u_c, v_c)$,
 206 which means that the distinction between contravariant and covariant van-
 207 ishes. Therefore everything can be cast in terms of Cartesian components
 208 and these subtleties can be ignored. The case of W^a is special because it
 209 keeps its definition (4), albeit with $w = 0$. W , the flux through the mov-
 210 ing grid, is called the omega velocity. The distinction between Cartesian,
 211 contravariant and covariant components is thus essentially a non-hydrostatic
 212 issue.

213 2.2. Equations in momentum variables

214 We now derive the momentum equations for the newly defined momentum
 215 components (u, v, w) , used as prognostic variables. As starting point, we
 216 will repeat equations (3.8 – 3.10) together with equations (3.13) and (3.16)
 217 from Kanarska et al. (2007) that define the momentum equations in terms of
 218 substantial derivatives of Cartesian velocity components, i.e. in flux form,

$$\begin{aligned} \frac{\mathcal{D}u_c}{\mathcal{D}t} &= -\mathcal{A}_\xi \left. \frac{\partial(p^h + q)}{\partial\xi} \right|_z + \mathcal{V}(C_{u_c} + G_{u_c} + F_{u_c}) \\ \frac{\mathcal{D}v_c}{\mathcal{D}t} &= -\mathcal{A}_\eta \left. \frac{\partial(p^h + q)}{\partial\eta} \right|_z + \mathcal{V}(C_{v_c} + G_{v_c} + F_{v_c}) \\ \frac{\mathcal{D}w_c}{\mathcal{D}t} &= -\mathcal{A}_\sigma \frac{\partial q}{\partial\sigma} + \mathcal{V}(C_{w_c} + F_{w_c}) \end{aligned} \quad (15)$$

219 where $(\mathcal{A}_\xi, \mathcal{A}_\eta, \mathcal{A}_\sigma) = (h_\eta h_\sigma, h_\xi h_\sigma, h_\xi h_\eta)$ are elementary surfaces and $\mathcal{V} =$
 220 $h_\xi h_\eta h_\sigma$ is the elementary volume. The gravitational term and the vertical

221 hydrostatic pressure gradient are absent in (15) because they balance. The
 222 two contributions p^h and q to the total pressure thus play a completely dif-
 223 ferent role: p^h accounts for the gravitational effects whereas q enforces the
 224 incompressibility. The additional subscripts c emphasize that the momenta
 225 are the Cartesian velocity components. C_{u_c, v_c, w_c} represent the Coriolis force,
 226 G_{u_c, v_c} represent the curvilinear terms arising from the curvature of the grid in
 227 the horizontal plane, and F_{u_c, v_c, w_c} represent viscous dissipation, mixing and
 228 forcing terms. Note that there is no curvature term in the vertical direction
 229 and that $C_{w_c} = 0$ if the traditional assumption is made. The substantial
 230 derivative

$$\frac{\mathcal{D}^*}{\mathcal{D}t} = \frac{\partial}{\partial t} (\mathcal{V}^*) + \frac{\partial}{\partial \xi} (\mathcal{A}_\xi U^*) + \frac{\partial}{\partial \eta} (\mathcal{A}_\eta V^*) + \frac{\partial}{\partial \sigma} (\mathcal{A}_\sigma W^*) \quad (16)$$

231 uses the flux through the breathing mesh. Note that in the present formalism
 232 (16) is not impacted by the new definition of momentum because the flux
 233 defined in (3) is identical to the one in Kanarska et al. (2007), up to the area
 234 metric factors. Consequently the continuity equation remains identical

$$\frac{\partial}{\partial t} \mathcal{V} + \frac{\partial}{\partial \xi} (\mathcal{A}_\xi U) + \frac{\partial}{\partial \eta} (\mathcal{A}_\eta V) + \frac{\partial}{\partial \sigma} (\mathcal{A}_\sigma W) = 0. \quad (17)$$

235 We can now recast the equations in terms of momentum. First we apply
 236 the total derivative (16) to (9)

$$\begin{aligned} \frac{\mathcal{D}u}{\mathcal{D}t} &= \frac{\mathcal{D}u_c}{\mathcal{D}t} + s_\xi \frac{\mathcal{D}w_c}{\mathcal{D}t} + w_c \frac{\mathcal{D}s_\xi}{\mathcal{D}t} \\ \frac{\mathcal{D}v}{\mathcal{D}t} &= \frac{\mathcal{D}v_c}{\mathcal{D}t} + s_\eta \frac{\mathcal{D}w_c}{\mathcal{D}t} + w_c \frac{\mathcal{D}s_\eta}{\mathcal{D}t} \\ \frac{\mathcal{D}w}{\mathcal{D}t} &= \frac{\mathcal{D}w_c}{\mathcal{D}t}, \end{aligned} \quad (18)$$

237 then substitute the material derivatives of (u_c, v_c, w_c) with (15) to get

$$\begin{aligned}\frac{\mathcal{D}u}{\mathcal{D}t} &= -\mathcal{A}_\xi \left. \frac{\partial(p^h + q)}{\partial\xi} \right|_z - \mathcal{A}_\sigma s_\xi \frac{\partial q}{\partial\sigma} + w \frac{\mathcal{D}s_\xi}{\mathcal{D}t} + \mathcal{V}(C_u + G_u + F_u) \\ \frac{\mathcal{D}v}{\mathcal{D}t} &= -\mathcal{A}_\eta \left. \frac{\partial(p^h + q)}{\partial\eta} \right|_z - \mathcal{A}_\sigma s_\eta \frac{\partial q}{\partial\sigma} + w \frac{\mathcal{D}s_\eta}{\mathcal{D}t} + \mathcal{V}(C_v + G_v + F_v) \\ \frac{\mathcal{D}w}{\mathcal{D}t} &= -\mathcal{A}_\sigma \frac{\partial q}{\partial\sigma} + \mathcal{V}(C_w + F_w),\end{aligned}\quad (19)$$

238 where $(F_u, F_v, F_w) = (F_{u_c} - s_\xi F_{w_c}, F_{v_c} - s_\eta F_{w_c}, F_{w_c})$ and $(G_u, G_v) = (G_{u_c}, G_{v_c})$.

239 The Coriolis terms $C_{u,v,w}$ are detailed in Appendix A. Using the chain rule

$$\mathcal{A}_\xi \left. \frac{\partial*}{\partial\xi} \right|_z = \mathcal{A}_\xi \frac{\partial*}{\partial\xi} - \mathcal{A}_\sigma s_\xi \frac{\partial*}{\partial\sigma} \quad (20)$$

240 simplifies the expression for the non-hydrostatic pressure gradient

$$\begin{aligned}\frac{\mathcal{D}u}{\mathcal{D}t} &= -\mathcal{A}_\xi \left. \frac{\partial p^h}{\partial\xi} \right|_z - \mathcal{A}_\xi \frac{\partial q}{\partial\xi} + w \frac{\mathcal{D}s_\xi}{\mathcal{D}t} + \mathcal{V}(C_u + G_u + F_u) \\ \frac{\mathcal{D}v}{\mathcal{D}t} &= -\mathcal{A}_\eta \left. \frac{\partial p^h}{\partial\eta} \right|_z - \mathcal{A}_\eta \frac{\partial q}{\partial\eta} + w \frac{\mathcal{D}s_\eta}{\mathcal{D}t} + \mathcal{V}(C_v + G_v + F_v) \\ \frac{\mathcal{D}w}{\mathcal{D}t} &= -\mathcal{A}_\sigma \frac{\partial q}{\partial\sigma} + \mathcal{V}(C_w + F_w).\end{aligned}\quad (21)$$

241 (21) is the new system of momentum equations we propose to use to integrate

242 the primitive equations in sigma-coordinates model when the hydrostatic

243 assumption is relaxed. The terms in (21) involving the substantial derivatives

244 of s_ξ and s_η can also be written as

$$\begin{aligned}w \frac{\mathcal{D}s_\xi}{\mathcal{D}t} &= w\mathcal{V} \frac{\partial s_\xi}{\partial t} + wU \mathcal{A}_\xi \frac{\partial s_\xi}{\partial\xi} + wV \mathcal{A}_\eta \frac{\partial s_\xi}{\partial\eta} + wW \mathcal{A}_\sigma \frac{\partial s_\xi}{\partial\sigma} \\ w \frac{\mathcal{D}s_\eta}{\mathcal{D}t} &= w\mathcal{V} \frac{\partial s_\eta}{\partial t} + wU \mathcal{A}_\xi \frac{\partial s_\eta}{\partial\xi} + wV \mathcal{A}_\eta \frac{\partial s_\eta}{\partial\eta} + wW \mathcal{A}_\sigma \frac{\partial s_\eta}{\partial\sigma}.\end{aligned}\quad (22)$$

245 This derivation makes (22) inherently new terms compared to Kanarska et al.

246 (2007). The system (21) can be derived independently by application of

247 Hamilton's principle (Appendix C). It yields the vector-invariant form.
 248 After a bit of algebra the flux form can be recovered along with the expression
 249 for the curvilinear terms. This alternative derivation blends the contributions
 250 of (22) and $\mathcal{V}G_{u,v}$ into a same term. Finally note that the chain rule (20)
 251 can also be applied to the hydrostatic pressure gradient term in (21). This
 252 introduces a buoyancy term in the horizontal components of the momentum
 253 equations, like in the variational approach (Appendix C). In practice the
 254 hydrostatic pressure gradient and this buoyancy term are computed together
 255 to limit the pressure gradient error (Shchepetkin and McWilliams, 2003).

256 Finally we want to emphasize that there is no necessity to prognose the
 257 covariant components. A perfectly valid choice could be to prognose the
 258 contravariant ones (U, V, W^a). In fact, it seems that the good choice of prog-
 259 nostic variables depend on the formulation of the momentum equations: in
 260 vector-invariant, the covariant components appear to be natural (see Eq C.4),
 261 in flux-form, the contravariant components are more natural.

262 3. Discrete operators for the non-hydrostatic pressure

263 In this section we first show how the Poisson equation for pressure can be
 264 derived from the definitions of the kinetic energy and the divergence operator.
 265 We then discretize the kinetic energy and deduce the discretizations for the
 266 \mathbf{T} matrix and the Poisson equation.

267 3.1. Pressure projection in matrix formulation

268 Following Kanarska et al. (2007), we will use the language of pressure
 269 projection methods (Chorin, 1968; Bell et al., 1989). These methods refer to
 270 a class of fractional step, implicit algorithms for unsteady flows that satisfy a

271 non-divergence constraint. In the fractional step method a provisional veloc-
 272 ity field is constructed first by advancing the momentum equations in time
 273 without enforcing the no-divergence constraint. Subsequently, the divergent
 274 part of velocity field is removed by a projection to the no-divergence sub-
 275 space. Here, we will present a derivation following Molemaker et al. (2005)
 276 to arrive, first at an expression for the operators in matrix form and second,
 277 specifically in terms of discrete coefficients.

278 We define the diagonal matrices \mathbf{L} , \mathbf{A} , \mathbf{V} corresponding to grid lengths,
 279 cell face areas and cell volumes working on velocity points. The elements
 280 of \mathbf{L} are the Lamé coefficients. The matrices are related by $\mathbf{V} = \mathbf{A}\mathbf{L}$. We
 281 define \mathbf{d} the differentiation matrix acting on cell centers (rho-points) and
 282 returning gradient components at face centers (velocity points). \mathbf{d} carries no
 283 dimension. The divergence is $\mathbf{D} = -\mathbf{d}^T\mathbf{A}$. Note the minus and the transpose.
 284 The divergence works on fluxes and returns finite volume scalars (scalar times
 285 the elementary volume).

286 Using (12), we can write the relation between fluxes and momenta com-
 287 pactly in matrix form as:

$$X^a = \mathbf{T} x \quad (23)$$

288 where $x = (u, v, w)^T$ is the solution vector in terms of momenta and $X^a =$
 289 $(U, V, W^a)^T$ the equivalent in terms of fluxes. We use the domain integrated
 290 kinetic energy $\mathcal{K} = \frac{1}{2}x^T\mathbf{V}X^a$ to define the inner product of two momenta x
 291 and y

$$\langle x, y \rangle = x^T\mathbf{V}\mathbf{T}y. \quad (24)$$

292 To be an inner product $\mathbf{V}\mathbf{T}$ needs to be symmetric. This is achieved by

293 discretizing the kinetic energy in such way that

$$\mathcal{K} = \frac{1}{2}(\mathbf{L}x)^T \tilde{\mathbf{T}}(\mathbf{L}x), \quad (25)$$

294 where $\tilde{\mathbf{T}}$ is a symmetric matrix. The discretization of \mathbf{T} is then deduced

$$\mathbf{T} = \mathbf{A}^{-1} \tilde{\mathbf{T}} \mathbf{L}. \quad (26)$$

295 The non-divergence condition is a linear constraint of the flow that, using
296 (17), can now be written in terms of the momenta as:

$$\mathbf{D}\mathbf{T} x_D = 0 \quad (27)$$

297 where the subscript D indicates a solution vector that satisfies the no-divergence
298 constraint (27).

299 We are now sufficiently equipped to project from a provisional velocity
300 field x , to a non-divergent field x_D . If ϵ is the difference vector between
301 the provisional solution and its projection onto the no-divergence divergence
302 subspace:

$$\epsilon = x - x_D, \quad (28)$$

303 we demand that this correction vector is orthogonal to the subspace of no-
304 divergence solutions, or:

$$\langle x_D, \epsilon \rangle = 0. \quad (29)$$

305 The correction is of the form

$$\epsilon = \mathbf{V}^{-1} \mathbf{D}^T p, \quad (30)$$

306 with p a scalar field, because

$$\langle x_D, \mathbf{V}^{-1} \mathbf{D}^T p \rangle = x_D^T \mathbf{T}^T \mathbf{D}^T p = (\mathbf{D}\mathbf{T}x_D)^T p = 0. \quad (31)$$

p acts as a Lagrange multiplier. Its function is to help satisfy the linear constraint of non-divergence. Physically speaking, the pressure functions as such for an incompressible flow, and we see that the correction vector ϵ is, in fact, the pressure gradient. We now have the following system:

$$\mathbf{V}\mathbf{T}x - \mathbf{V}\mathbf{T}x_D = (\mathbf{D}\mathbf{T})^T p \quad (32)$$

$$\mathbf{D}\mathbf{T} x_D = 0. \quad (33)$$

307 From which we can easily derive an equation for p

$$(\mathbf{D}\mathbf{V}^{-1}\mathbf{T}^T\mathbf{D}^T) p = \mathbf{D}\mathbf{T} x. \quad (34)$$

308 (34) is the Poisson equation for pressure and

$$\mathbf{P} = \mathbf{D}\mathbf{V}^{-1}\mathbf{T}^T\mathbf{D}^T \quad (35)$$

309 is the Poisson matrix. The matrix can be simplified further into

$$\mathbf{P} = \mathbf{d}^T \tilde{\mathbf{T}} \mathbf{d}, \quad (36)$$

310 which is symmetric if $\tilde{\mathbf{T}}$ is. We have done all this effort to arrive at the Pois-
 311 son equation for the pressure (34). We see that once the kinetic energy is
 312 discretized in a symmetric form using the $\tilde{\mathbf{T}}$ matrix, the divergence operator
 313 \mathbf{D} , and the volume matrix \mathbf{V} are fixed, the definition of the other operators
 314 follow. The flux-momentum approach yields a compact 15-point stencil be-
 315 cause \mathbf{P} can be cast into the simplified form (36) for which the differential
 316 operator \mathbf{d} has a minimal form.

317 3.2. Variable staggering

318 We describe here how the momentum and flux can be discretized in such
 319 way that the operators grad and div operating on them acquire the simplest

320 possible discretization, namely basic finite differences. For that we use the
321 framework of discrete differential geometry (DDG) (Desbrun et al., 2008;
322 Grady and Polimeni, 2010), which extends the common concepts of finite
323 volume and C-grid staggering to a higher degree of generality. The frame-
324 work also allows to rationalize why the contravariant components should be
325 interpreted as normal to coordinate hypersurfaces and not as aligned with
326 the coordinate lines. The interpretation we present is also completely aligned
327 with Thuburn and Cotter (2012) for non-orthogonal grids in the context of
328 shallow water equations.

329 The goal of this section is to introduce all the elements present in figure 1,
330 adapted from Desbrun et al. (2008), which summarizes the various aspects of
331 the spatial discretization. The (ξ, η, σ) coordinates are discretized in logically
332 rectangular cells, with indices (i, j, k) pointing to cell vertices. The cell edges
333 are along coordinates lines, cell faces are along coordinates hyperplanes and
334 finally each cell form a volume. This defines the primal grid. The dual grid
335 is defined by setting vertices at the primal cell centers, then dual grid edges
336 connect these dual grid vertices through primal faces. Dual faces connect
337 dual edges. Finally the dual volume cell is defined as the interior of dual
338 faces, it surrounds each primal vertex.

339 The deep breakthrough of DDG is to see the usual operators (gradient,
340 curl and div) as transformations between objects: the gradient transforms
341 vertices to edges, the curl edges to faces and the divergence faces to volume.
342 Each operator is then simply a finite difference between quantities. We detail
343 this below.

344 Scalar quantities are naturally discretized at either cell centers or cell

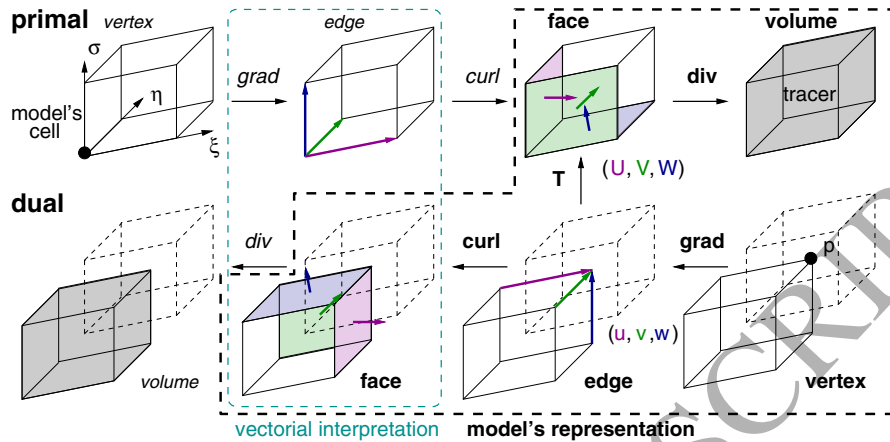


Figure 1: Staggering of variables on the primal and dual grids according to their nature: velocity, pressure or tracer. Contravariant components are defined on the primal, covariant components on the dual. They are related by the \mathbf{T} transformation. By interpreting the contravariant as flux, living on faces, and the covariant as gradient terms, living on the edges, this makes the pressure gradient operation and the divergence of the flux very simple. This interpretation is consistent with a discretization of the pressure at the primal cell centers (dual vertices) and a finite volume interpretation of tracers (primal volume). Each object is connected to its right-neighbour (primal) and left-neighbour (dual) by a differentiation operator which, depending on the object, is the gradient, the curl or the divergence. The vectorial interpretation is illustrated by the cyan dashed box. The model's objects and operators (bold) are delimited by the black dashed box.

345 volumes depending on whether the quantity is a true scalar, like pressure, or
 346 an integrated quantity, like tracer content. Vectors are naturally discretized
 347 at either cell edges or cell faces. Contravariant components are defined on
 348 the primal whereas the covariant components are on the dual. The two in-
 349 terpretations for the contravariant components can now be enriched by this
 350 geometric framework: the contravariant components are either i) on the faces
 351 of the primal or ii) on the edges of the primal, which is the vectorial inter-

352 pretation. Interpretation i) is the one we made, it makes the flux naturally
 353 transformed into a volume via the divergence operator. With interpretation
 354 ii), the divergence of the contravariant components is not a trivial operation.
 355 Similarly, the momentum should behave like a gradient, because it is covari-
 356 ant. It can be discretized either i) on the edges of the dual or on ii) the
 357 faces. To be consistent, it has to be one edges. It makes the computation of
 358 the pressure gradient a simple finite difference. By using discretization ii),
 359 the expression of the pressure gradient acting on the momentum would be
 360 complicated.

361 The consistency of i) with the C-grid discretization practice is even deeper:
 362 the finite volume discretization of tracers implies that the flux sits at the pri-
 363 mal faces, and the discretization of pressure at cell centers implies that the
 364 covariant velocity sits at the dual edges. Finally, having the contravariant
 365 components on the primal faces makes the operation $\text{div}\mathbf{U}$ covariant, i.e.
 366 coordinates independent.

367 The \mathbf{T} matrix transforms objects from the dual edges to the primal faces.
 368 Lastly, from figure 1 it is clear that the vorticity, defined as the curl of
 369 the momentum lives at the dual faces. It inherits covariant components.
 370 The vorticity appears explicitly in the vector-invariant form. It is only a
 371 diagnostic quantity in flux-form.

372 From now on, each variable is indexed with a triplet of indices (i, j, k) .
 373 Faces and edges with index (i, j, k) are located to the west, south and bottom
 374 of the cell (i, j, k) . To lighten the notation the indices will be indicated only
 375 when they differ from i, j or k , e.g. ϕ_{i-1} will stand for $\phi_{i-1,j,k}$. There is
 376 n_z cells in the vertical and $n_z + 1$ w -levels. We define the right difference

377 operator $\delta_{i,r}$ acting on the field $\phi_{i,j,k}$ in the direction ξ as $\delta_{i,r}[\phi] = \phi_{i+1} - \phi$ and
 378 the left difference operator $\delta_{i,l}[\phi] = \phi - \phi_{i-1}$. We define the right averaging
 379 operator $\bar{\phi}^{i,r} = (\phi_{i+1} + \phi)/2$ and the left averaging one $\bar{\phi}^{i,l} = (\phi + \phi_{i-1})/2$.
 380 Likewise we define the same operators acting in direction η with index j , and
 381 in direction s with index k .

382 Grid spacing $(h_\xi, h_\eta, h_\sigma)$, and local slopes (s_ξ, s_η) are also defined at cell
 383 centers with

$$s_\xi = \left[\frac{\delta_{i,l}[z]}{h_\xi^u} \right]^{i,r} \quad \text{and} \quad s_\eta = \left[\frac{\delta_{j,l}[z]}{h_\eta^v} \right]^{j,r}. \quad (37)$$

384 A superscript of u, v, w denotes a quantity that has been averaged to the
 385 position of its respective velocity, e.g. $h_\xi^u = \delta_{i,r}[h_\xi]$. In particular h_σ^w is only
 386 half of its value at the bottom ($k = 1$) and top level ($k = n_z + 1$). Faces areas
 387 are defined at velocity points using product of two lengths, e.g. $\mathcal{A}_\xi^u = h_\eta^u h_\sigma^u$.
 388 Volumes are defined at velocity points using the product of the lengths in
 389 each direction, e.g. $\mathcal{V}^u = h_\xi^u h_\eta^u h_\sigma^u$.

390 3.3. Discretization of the gradient and divergence

391 As announced, the discretization of the gradient and divergence boil down
 392 to simple finite differences. If we denote $X = (U, V, W)^T$ the flux vector and
 393 $\delta = \mathbf{D}X$ the result of the action of \mathbf{D} , then the discretization of the divergence
 394 reads

$$\delta_{i,r}[\mathcal{A}_\xi^u U] + \delta_{j,r}[\mathcal{A}_\eta^v V] + \delta_{k,r}[\mathcal{A}_\sigma^w W] = \delta. \quad (38)$$

395 The definition is valid at all grid points, including at the boundaries. The
 396 bottom boundary condition is handled outside of the divergence by simply
 397 setting $W = 0$. The gradient is defined as the skew-adjoint of \mathbf{D} . For interior

398 points it is easy to show that

$$h_i^u u_i = \delta_{i,l}[q], \quad h_i^v v_i = \delta_{j,l}[q], \quad h_i^w w_i = \delta_{k,l}[q], \quad (39)$$

399 where q is a scalar defined on the dual vertices and (u, v, w) are the (covariant)
400 gradient components. The expression for the curl is given in Appendix C.
401 It is basically computed by estimating the circulation around each face.

402 3.4. Discretization of the kinetic energy

403 As said above, the discretization of the \mathbf{T} matrix derives from \mathcal{K} written
404 in the form (25). To discretize \mathcal{K} we start from the continuous expression

$$\mathcal{K} = \iiint \left[\frac{1}{2} (u^2 + v^2 + \gamma w^2) - (s_\xi u w + s_\eta v w) \right] \mathcal{V} d\xi d\eta d\sigma, \quad (40)$$

405 with $\gamma = 1 + s_\xi^2 + s_\eta^2$, and split it into several contributions

$$\mathcal{K} = \mathcal{K}_{int} + \mathcal{K}_{cross} + \mathcal{K}_{bottom}, \quad (41)$$

406 where \mathcal{K}_{bottom} accounts for all terms involving the bottom vertical momentum
407 w_b (w at $k = 1$), \mathcal{K}_{cross} accounts for the cross terms uw and vw , and \mathcal{K}_{int} is
408 the remainder. The interior kinetic energy

$$\mathcal{K}_{int} = \frac{1}{2} \sum_{i,j} \left\{ \sum_{k=1}^{n_z} u^2 \mathcal{V}^u + \sum_{k=1}^{n_z} v^2 \mathcal{V}^v + \sum_{k=2}^{n_z+1} \gamma^w w^2 \mathcal{V}^w \right\} \quad (42)$$

is the simplest term. Each component is weighted with its volume defined at the component location. This part of the kinetic energy encodes the diagonal terms of the $\tilde{\mathbf{T}}$ matrix, which are $(A^u/h_\xi^u, A^v/h_\eta^v, c^w A^w/h_\sigma^w)$. The cross terms

contribution reads

$$\begin{aligned}
\mathcal{K}_{cross} = & -\frac{1}{2} \sum_{i,j} \left[s_\xi h_\eta \overline{[h_\xi^u u]}^{i,r} + s_\eta h_\xi \overline{[h_\eta^v v]}^{j,r} \right]_{k=1} [h_\sigma^w w]_{k=2} \\
& - \sum_{i,j} \sum_{k=2}^{n_z-1} \left[s_\xi h_\eta \overline{[h_\xi^u u]}^{i,r} + s_\eta h_\xi \overline{[h_\eta^v v]}^{j,r} \right] \overline{[h_\sigma^w w]}^{w,r} \\
& - \frac{1}{2} \sum_{i,j} \left[s_\xi h_\eta \overline{[h_\xi^u u]}^{i,r} + s_\eta h_\xi \overline{[h_\eta^v v]}^{j,r} \right]_{k=n_z} [h_\sigma^w w]_{k=n_z} \\
& - \sum_{i,j} \left[s_\xi h_\eta \overline{[h_\xi^u u]}^{i,r} + s_\eta h_\xi \overline{[h_\eta^v v]}^{j,r} \right]_{k=n_z} [h_\sigma^w w]_{k=n_z+1}. \quad (43)
\end{aligned}$$

409 All terms are defined at cell centers (rho points). The volume elements have
410 been carefully split into three length elements to form elementary circulations
411 to ensure (25). The choice to average the terms at cell centers allows for the
412 maximum of terms cancellation in the resulting pressure matrix. The sum
413 has been split in horizontal slabs: in the bottom level ($k = 1$) we have
414 removed the part involving w_b , in the top level ($k = n_z$) we have introduced
415 an extra factor 2 to compensate for h_σ^w being halved.

The bottom contribution accounts for the kinetic energy density $K = \frac{1}{2} \gamma w^2 - (s_\xi u w + s_\eta v w)$ integrated over the bottom cells lower half volume. The u^2 and v^2 are already accounted in \mathcal{K}_{int} . Because w at the bottom is not a prognostic variable, it should be substituted with its diagnostic relation $\gamma w = s_\xi u + s_\eta v$, obtained from (11), the definition of W^a , and the fact that the breathing vanishes at the bottom boundary ($\partial_t z = 0$). After substitution, the kinetic energy density reads $K = -\frac{1}{2\gamma} [(s_\xi u)^2 + (s_\eta v)^2 + 2s_\xi s_\eta uv]$, which is the continuous density we use for the discretization at the bottom. The

discretized volume-integrated bottom kinetic energy then reads

$$\mathcal{K}_{bottom} = -\frac{1}{4} \sum_{i,j,k=1} \left\{ \overline{\left[\frac{s_\xi^2}{\gamma} \right]}^{i,l} u^2 \mathcal{V}^u + \overline{\left[\frac{s_\eta^2}{\gamma} \right]}^{j,l} v^2 \mathcal{V}^v + 2 \frac{s_\xi s_\eta h_\sigma}{\gamma} \overline{[h_\xi^u u]}^{i,r} \overline{[h_\eta^v v]}^{j,r} \right\}, \quad (44)$$

416 where the u^2 and v^2 have been discretized at their velocity point and the
 417 uv cross term has been averaged at cell center. The whole term is indeed
 418 weighted with a half volume cell because of the extra factor $1/2$ that com-
 419 pensates for \mathcal{V}^u , \mathcal{V}^v and h_σ being full cell metrics.

420 3.5. Discretization of the fluxes

421 The flux discretization, that is the action of \mathbf{T} on the momentum, can now
 422 be deduced either upon identification of the $\tilde{\mathbf{T}}$ with (25) and then application
 423 of (26) or by differentiating \mathcal{K} with respect to the momentum, e.g.

$$U_{i,j,k} = \frac{1}{\mathcal{A}_\xi^u} \frac{\partial \mathcal{K}}{\partial (h_\xi^u u_{i,j,k})}. \quad (45)$$

424 We have for the horizontal flux U

- 425 • In the interior ($1 < k < n_z$)

$$U = u - \frac{1}{\mathcal{A}_\xi^u} \overline{s_\xi h_\eta [h_\sigma^w w]}^{w,r,i,l} \quad (46)$$

- 426 • At the bottom $k = 1$

$$U = \left(1 - \overline{\left[\frac{s_\xi^2}{2\gamma} \right]}^{i,l} \right) u - \frac{1}{2\mathcal{A}_\xi^u} \left(\overline{\frac{s_\xi s_\eta h_\sigma}{\gamma} [h_\eta^v v]}^{j,r,i,l} + \overline{\frac{s_\xi h_\eta}{\sigma} [h_\sigma^w w]}_{k=2}^{i,l} \right) \quad (47)$$

- 427 • At the top $k = n_z$

$$U = u - \frac{1}{2\mathcal{A}_\xi^u} \left(\overline{s_\xi h_\eta [h_\sigma^w w]_{k=n_z}} + 2\overline{s_\xi h_\eta [h_\sigma^w w]_{k=n_z+1}} \right)^{i,l} \quad (48)$$

428 The discretization for V is similar. It is obtained by applying the permutation

429 $(\xi, u, i) \leftrightarrow (\eta, v, j)$ on the above expressions.

430 The discretization of the vertical flux W^a reads

- 431 • at the bottom $k = 1$

$$W^a = 0, \quad (49)$$

- 432 • in the interior $1 < k < n_z + 1$

$$W^a = \gamma^w w - \frac{1}{\mathcal{A}_\sigma^w} \left(\overline{s_\xi h_\eta [h_\xi^u u]^{i,r,k,l}} + \overline{s_\eta h_\xi [h_\eta^v v]^{j,r,k,l}} \right), \quad (50)$$

- 433 • and at the surface $k = n_z + 1$

$$W^a = \gamma^w w - \frac{1}{\mathcal{A}_\sigma^w} \left[\overline{s_\xi h_\eta [h_\xi^u u]^{i,r}} + \overline{s_\eta h_\xi [h_\eta^v v]^{j,r}} \right]_{k=n_z}. \quad (51)$$

434 These expressions complete the definition for the discrete \mathbf{T} matrix, that
435 transforms momentum into flux.

436 3.6. Discretization of the Poisson operator

437 The Poisson matrix is obtained from (35) with the use of the discretization
438 of the fluxes and the divergence. Its discretization is given in appendix Ap-
439 pendix B. Incidentally the compactness of the stencil does not change the
440 matrix condition number, compared to any other discretization. Indeed, the
441 condition number is controlled by the grid characteristics: grid steps and
442 number of grid points.

443 The combination of the discrete divergence and the discrete kinetic en-
 444 ergy fully determine the boundary conditions on the pressure equation. At
 445 the surface the boundary condition is of Dirichlet type ($p = 0$) whereas at
 446 the bottom and laterally it is of Neumann type. To exhibit them we use
 447 the property that the flux is divergent-free $\langle p, \mathbf{D}\mathbf{T}x_D \rangle = 0$, then make use
 448 of the adjoint to find $\langle \mathbf{D}^T p, \mathbf{T}x_D \rangle = 0$ which expresses that the pressure
 449 force does not work. By expanding this inner product as a sum on velocity
 450 points we recover the discrete kinetic energy for which one of the momentum
 451 component is replaced with the pressure gradient. In particular the w^2 term
 452 at $(i, j, n_z + 1)$ in \mathcal{K}_{int} is now $-\gamma^w \mathcal{V}^w w p_{i,j,n} / h_{i,j,n_z+1}^w$, from which we deduce
 453 the vertical pressure gradient at the surface is $\partial_z p = -p_{i,j,n} / h_{i,j,n_z+1}^w$. This
 454 discretization is indeed a vertical gradient if $p_{i,j,n_z+1/2} = 0$, thus proving that
 455 the surface boundary condition is of Dirichlet type. The surface boundary
 456 condition is realized in the Poisson stencil by the A_σ^w / h_σ^w term at $k = n_z + 1$ in
 457 the diagonal coefficient P_1 (see (B.2) in Appendix B). The bottom boundary
 458 condition is handled by \mathcal{K}_{bot} and behaves such that $W = 0$ at the bottom.
 459 The pressure gradient at the bottom is not defined but its associated verti-
 460 cal flux ($\mathbf{T}\mathbf{D}^T p$) is zero there, which corresponds to a Neumann boundary
 461 condition. The lateral boundary conditions is treated by assuming the co-
 462 variant velocity, either u or v , vanishes. This removes in the Poisson matrix
 463 all connections between interior and outer pressure points.

464 4. Discussion

465 We now assess the properties of the discretization on a test-case. The
 466 goal is test the discretized projection in a configuration simple enough, yet

467 ensuring that none of the 15 coefficients of the Poisson matrix vanishes. The
 468 comparison with an analytical solution would have been ideal but we have
 469 found none with a genuinely three-dimensional topography.

470 The test-case, expressed in Cartesian coordinates, consists in projecting a
 471 divergent velocity field $(u, v, w) = (0, 0, \cos(\pi x / L))$ onto the divergent-free
 472 subspace in the closed domain $(x, y, z) \in [0, L] \times [0, L] \times [-H + h(x, y), 0]$,
 473 where $h(x, y) = h_0 \exp(-\alpha (x - L/2)^2 - \alpha (y - L/2)^2)$ is a sea-mount topogra-
 474 phy. The numerical values are $L = 10$ km, $H = 4$ km, $h_0 = 2$ km, $\alpha = 25/L^2$.
 475 The test-case is particularly severe for the bottom boundary condition be-
 476 cause the initial flow field is already divergence-free but it is far from being
 477 the solution because it does not satisfy the no-flow bottom boundary condi-
 478 tion. The right hand side for the Poisson equation is thus zero everywhere
 479 except along the topography.

480 Even though the test-case is relatively simple, it has no analytical solu-
 481 tion. It is therefore difficult to prove that the numerical solution obtained
 482 with the 15-point stencil is correct. To circumvent this issue we have com-
 483 puted the solution using two other methods: the tall 25-point stencil obtained
 484 with Cartesian components and the 7 points stencil derived using a z-level dis-
 485 cretization. The 25-point stencil matrix can be written as $\mathbf{P}_w = \mathbf{D}_w \mathbf{V}^{-1} \mathbf{D}_w^T$,
 486 where the divergence operator \mathbf{D}_w has a wider stencil than \mathbf{D} because of
 487 the extra derivatives coming from the chain rule. The approach closely fol-
 488 lows Auclair et al. (2011) but it differs by explicitly obtaining a symmetric
 489 matrix. In the z-levels discretization, the topography is handled via a mask
 490 array that flags cells outside the domain. The resulting matrix is similar
 491 to what is implemented in MITgcm (Marshall et al., 1997). In both cases,

492 the bottom boundary condition is trivially satisfied by resorting a divergence
 493 matrix that does not uses the w of the bottom level to compute the diver-
 494 gence in the first level. In addition to these three matrices we also test the
 495 impact of the extra-care needed to handle to bottom boundary condition in
 496 the 15-point stencil case by implementing a degraded version of it in which
 497 we simply remove in the \mathbf{T} matrix all the contributions from the \mathcal{K}_{bot} term.
 498 This matrix is obtained by setting the flag $\mu = 0$ in the Poisson stencil given
 499 in appendix.

500 We solve the problem on a series of grids $(n_x, n_y, n_z) = 2^n \times (25, 25, 10)$
 501 with doubling resolution. The matrices are implemented in Matlab. Because
 502 the grids need to fit in RAM of a single chipset, we were able to test the
 503 range $n = 0, 1, 2, 3$ for the four matrices. The solution is a non trivial
 504 flow. Figure 2 shows the Cartesian u_c and w_c components of the flow in
 505 the mid-plane $y = L/2$ for $n = 3$. The flow has on one side a down-welling
 506 branch emanating from the surface and an upwelling branch on the other. In-
 507 between, the flow is passing around (not shown) and over the sea-mount. The
 508 four implementations give an identical flow. At this resolution the differences
 509 are several orders of magnitude smaller than the solution. To quantitatively
 510 measure the convergence we compute the kinetic energy \mathcal{K} of each solution.
 511 Results are shown in Figure 3. They clearly converge to the same value \mathcal{K}_∞ ,
 512 indicated by the tick black line, albeit with a different convergence rate. The
 513 15- and the 25-point stencils converge more rapidly to \mathcal{K}_∞ . Specifically their
 514 increment of accuracy $\delta\mathcal{K} = \mathcal{K}(n) - \mathcal{K}(n - 1)$ decreases by a factor of 4 when
 515 the resolution is doubled, implying that \mathcal{K} is of the form $\mathcal{K} = \mathcal{K}_\infty + c4^{-n}$,
 516 where c is a constant. We can then use the Richardson extrapolation on the

517 15-point results to estimate this limit value as $\mathcal{K}_\infty \approx \mathcal{K}(n) + \delta\mathcal{K}(n)/3$.

518 This allows to monitor the absolute error $|\mathcal{K} - \mathcal{K}_\infty|$ (inset of Fig. 3). This
 519 figure confirms that the 15- and 25-point stencil converge to the solution at
 520 second order, as indicated by the -2 slope. Interestingly, the 15-point sten-
 521 cil is slightly more accurate than the 25-point stencil. This can be traced
 522 back to slightly smoother vertical shears near the bottom. The z-levels only
 523 converges at first order, which is consistent with the discretization of the to-
 524 pography that is only first order. For that case the solution has the footprint
 525 of steps near the topography (not shown). The degraded 15-point stencil
 526 also has a global order 1 convergence. Comparisons with the full 15-point
 527 stencil reveal that the horizontal flux vertical profile is discontinuous at the
 528 bottom level, generating a maximum of error there. Because the problem
 529 is essentially global this impacts the solution everywhere. The contribution
 530 of the horizontal cross terms in the bottom level ensures a uniform second
 531 order estimate of the velocities everywhere.

532 5. Conclusion

533 We have presented an alternative system of equations for the primitive
 534 equations written in sigma coordinates with a breathing mesh. It is based on
 535 the recognition of two sets of components for the velocity: a contravariant
 536 set, involved in the flux definitions, and a covariant one, involved in the
 537 momentum definition. The two sets are related by a linear transformation,
 538 the metric tensor, fully determined by the kinetic energy. We provide the
 539 momentum equations written for the covariant components in both vector-
 540 invariant and flux-form. We show that within the hydrostatic assumption,

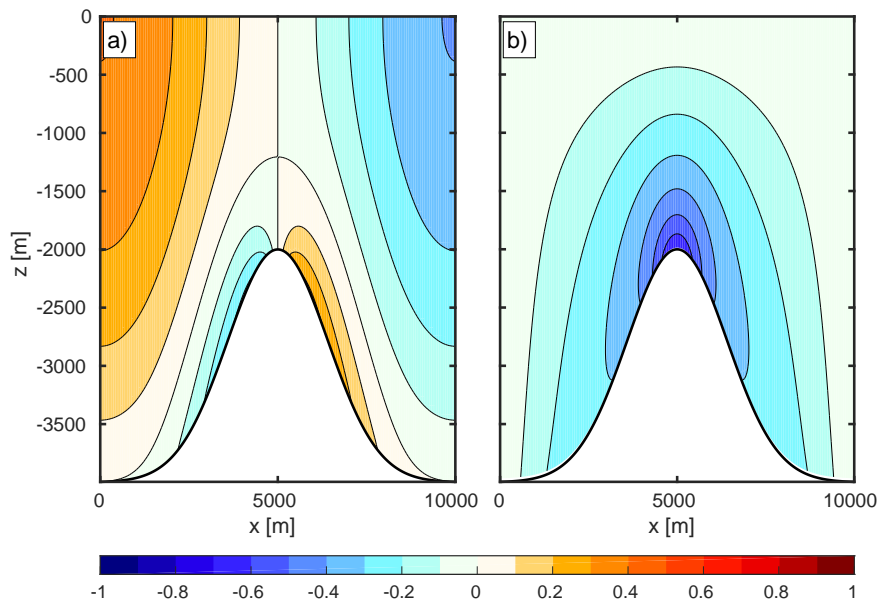


Figure 2: Vertical section of the velocity Cartesian components w_c (a) and u_c (b) at $y = L/2$, for the full solution obtained with the 15-point stencil matrix.

541 the distinction between contra- and covariant vanishes, which explains why
 542 the distinction has been largely ignored so far and why it is genuinely a
 543 non-hydrostatic issue. In flux-form the new system of momentum equation
 544 possesses extra pseudo-force-like terms in the horizontal. They involve the
 545 partial derivatives of the sigma slopes. The chief gain of this system, and
 546 the reason we present it, is to remove chain rule operations in the gradient
 547 and divergence operators. In the NH case this allows, in turn, to express the
 548 Poisson equation for the pressure very compactly. We show how this property
 549 can be implemented numerically yielding to a compact symmetric stencil for
 550 the Poisson matrix with only 15 points. The discretization is done at the level

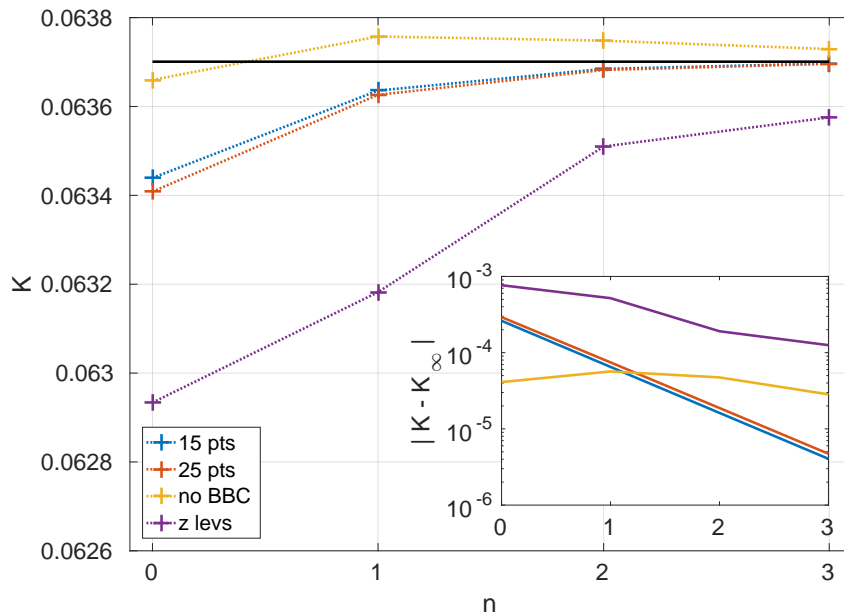


Figure 3: Kinetic energy as a function of resolution n for four different Poisson matrices: the 15-point stencil \mathbf{P} matrix (blue), the 25 points Cartesian matrix \mathbf{P}_w (red), the 15-point matrix with the bottom boundary condition removed (yellow), the 7 points z-levels matrix (purple). The thick line is the limit value \mathcal{K}_∞ obtained by Richardson extrapolation. The inset shows the absolute error $|\mathcal{K} - \mathcal{K}_\infty|$ as a function of resolution.

551 of the kinetic energy, which ensures that the pressure work vanishes globally.
 552 It also ensures the symmetry of the stencil. We show how the no-flow bottom
 553 boundary condition can be discretized with a second order accuracy. This
 554 discretization is also done directly at the level of the kinetic energy.

555 This approach opens a promising way of solving NH physics in a sigma
 556 coordinates model. The full implementation of this system of equations in
 557 an ocean model will be reported in a forthcoming paper.

558 **Appendix A. Coriolis terms**

559 In the non-traditional approximation the Coriolis terms involve the three
560 components of the Earth angular speed Ω , namely

$$\begin{aligned} f_\xi &= 2\Omega \cos \phi \sin \alpha \\ f_\eta &= 2\Omega \cos \phi \cos \alpha \\ f_z &= 2\Omega \sin \phi, \end{aligned} \quad (\text{A.1})$$

561 where α is the angle between the ξ and the zonal direction and ϕ is the
562 latitude. From Appendix B of Kanarska et al. (2007) the Coriolis terms for
563 Cartesian velocity components are

$$\begin{aligned} C_{u_c} &= -v_c f_z + w_c f_\eta \\ C_{v_c} &= +u_c f_z - w_c f_\xi \\ C_{w_c} &= -u_c f_\eta + v_c f_\xi. \end{aligned} \quad (\text{A.2})$$

564 Using the relations (9) yields the momentum based expression

$$\begin{aligned} C_u &= C_{u_c} + s_\xi C_{w_c} \\ C_v &= C_{v_c} + s_\eta C_{w_c} \\ C_w &= C_{w_c}. \end{aligned} \quad (\text{A.3})$$

565 Rearranging using (3) and (9) again, then defining

$$f_\sigma = f_z - s_\xi f_\xi - s_\eta f_\eta, \quad (\text{A.4})$$

566 yields

$$\begin{aligned} C_u &= -V f_\sigma + W^a f_\eta \\ C_v &= +U f_\sigma - W^a f_\xi \\ C_w &= -U f_\eta + V f_\xi. \end{aligned} \quad (\text{A.5})$$

567 Note that the Coriolis force involves the vertical flux W^a relative to a fix grid,
 568 not W . These terms are obviously energy neutral because the energy equa-
 569 tion is obtained by forming the inner product of the momentum equations
 570 with the flux (U, V, W^a) . The traditional approximation case is obtained by
 571 setting $f_\xi = f_\eta = 0$.

572 Appendix B. Poisson matrix

The discretization of the Poisson matrix follows from (35). Its action on a discrete scalar field $q_{i,j,k}$ can be written as

$$\begin{aligned}
 P_1 q + P_2 q_{k-1} + P_{2,k+1} q_{k+1} + P_3 q_{j-1,k+1} + P_{3,j+1,k-1} q_{j+1,k-1} \\
 + P_4 q_{j-1} + P_{4,j+1} q_{j+1} + P_5 q_{j-1,k-1} + P_{5,j+1,k+1} q_{j+1,k+1} \\
 + P_6 q_{i-1,k+1} + P_{6,i+1,k-1} q_{i+1,k-1} + P_7 q_{i-1} + P_{7,i+1} q_{i+1} \\
 + P_8 q_{i-1,k-1} + P_{8,i+1,k+1} q_{i+1,k+1} + P_9 q_{i-1,j-1} + P_{9,i+1,j+1} q_{i+1,j+1} \\
 + P_{10} q_{i-1,j+1} + P_{10,i+1,j-1} q_{i+1,j-1} = r_{i,j,k} \quad (\text{B.1})
 \end{aligned}$$

573 where the coefficients P_l , $l = 1, 2, \dots, 10$, define the stencil of the matrix and
 574 $r_{i,j,k}$ is the result. Coefficients pointing to indices outside of the index range
 575 are set to 0. To lighten the notation the indices are indicated only when they
 576 differ from i, j or k , e.g. P_2 stands for $P_{2,i,j,k}$. Each coefficient has a general
 577 expression for interior levels $1 < k < n_z$ but differ at the top and bottom
 578 levels. We define the indicatrix function $\Phi_{l,i,j,k}$ that is 0 for $k \neq l$ and 1
 579 for $k = l$. We also define a flag $\mu = 1$ for the full 15-point Poisson matrix
 580 and $\mu = 0$ for the degraded 15-point matrix in which the bottom boundary

581 condition is not explicitly handled. The diagonal coefficient reads

$$P_1 = \frac{A_\sigma^w}{h_\sigma^w}(1 - \Phi_1) + \frac{A_\sigma^w}{h_\sigma^w} \Big|_{k+1} + \frac{A_\eta^v}{h_\eta^v} + \frac{A_\eta^v}{h_\eta^v} \Big|_{j+1} + \frac{A_\xi^u}{h_\xi^u} + \frac{A_\xi^u}{h_\xi^u} \Big|_{i+1}. \quad (\text{B.2})$$

Recall that A_σ^w/h_σ^w is well defined at $k = n_z + 1$. The other coefficients read

$$P_2 = \frac{A_\sigma^w}{h_\sigma^w} \gamma^w (1 - \Phi_1) \quad (\text{B.3})$$

$$P_3 = \frac{1}{4} (s_\eta h_\xi + s_{\eta,j-1} h_{\xi,j-1}) (1 - \Phi_{n_z}) \quad (\text{B.4})$$

$$P_4 = \frac{A_\eta^v}{h_\eta^v} \left(1 - \frac{1}{4} \left[\frac{s_\eta^2}{\gamma} + \frac{s_{\eta,j-1}^2}{\gamma_{j-1}} \right] \Phi_1 \mu \right) + \frac{1}{4} (s_\eta h_\xi - s_{\eta,j-1} h_{\xi,j-1}) (\Phi_{n_z} + \Phi_1) \quad (\text{B.5})$$

$$P_5 = -\frac{1}{4} (s_\eta h_\xi + s_{\eta,j-1} h_{\xi,j-1}) (1 - \Phi_1) \quad (\text{B.6})$$

$$P_6 = \frac{1}{4} (s_\xi h_\eta + s_{\xi,i-1} h_{\eta,i-1}) (1 - \Phi_{n_z}) \quad (\text{B.7})$$

$$P_7 = \frac{A_\xi^u}{h_\xi^u} \left(1 - \frac{1}{4} \left[\frac{s_\xi^2}{\gamma} + \frac{s_{\xi,i-1}^2}{\gamma_{i-1}} \right] \Phi_1 \mu \right) + \frac{1}{4} (s_\xi h_\eta - s_{\xi,i-1} h_{\eta,i-1}) (\Phi_{n_z} + \Phi_1) \quad (\text{B.8})$$

$$P_8 = -\frac{1}{4} (s_\xi h_\eta + s_{\xi,i-1} h_{\eta,i-1}) (1 - \Phi_1) \quad (\text{B.9})$$

$$P_9 = -\frac{1}{8} \left(\frac{s_{\xi,i-1} s_{\eta,i-1} h_{\sigma,i-1}}{\gamma_{i-1}} + \frac{s_{\xi,j-1} s_{\eta,j-1} h_{\sigma,j-1}}{\gamma_{j-1}} \right) \Phi_1 \mu \quad (\text{B.10})$$

$$P_{10} = \frac{1}{8} \left(\frac{s_{\xi,i-1} s_{\eta,i-1} h_{\sigma,i-1}}{\gamma_{i-1}} + \frac{s_{\xi,j+1} s_{\eta,j+1} h_{\sigma,j+1}}{\gamma_{j+1}} \right) \Phi_1 \mu \quad (\text{B.11})$$

582 As it is seen, only 8 coefficients are needed away from the top and bottom
 583 boundaries. This is a manifestation of the both the compactness and the
 584 symmetry of the stencil.

585 **Appendix C. Variational derivation**

586 In this section we recover the momentum equations (21) with a varia-
 587 tional derivation using Hamilton's principle. However we were able to do
 588 it only for the case of a non-breathing grid, so that all terms involving $\partial_t z$,
 589 $\partial_t s_\xi$ and $\partial_t s_\eta$ vanish. We also remove the full Coriolis force because the vari-
 590 ational approach requires to have horizontal Lamé coefficients (h_ξ, h_η) that
 591 vary in the vertical, which corresponds to the deep atmosphere assumption
 592 in atmospheric models (Tort and Dubos, 2014).

593 To derive the momentum equations we follow Holm et al. (1998) vari-
 594 ational derivation and adapt it to the case of non-orthogonal curvilinear
 595 coordinates. We start from the Lagrangian

$$\mathcal{L}(U_i, \rho, b, p) = \iiint \left[\rho \left(\frac{1}{2} u_i U_i + bz \right) - p \left(\frac{\rho}{\rho_0} - 1 \right) \right] \mathcal{V} d\xi d\eta d\sigma, \quad (\text{C.1})$$

596 where the index notation $i = 1, 2, 3$ stands for (ξ, η, σ) , U_i 's are the compo-
 597 nents of the flux \mathbf{U} , u_i 's are the component of the momentum \mathbf{u} and b is the
 598 buoyancy relative to the Boussinesq reference density ρ_0 . The variations of
 599 the Lagrangian are

$$\begin{aligned} \frac{1}{\rho} \frac{\delta \mathcal{L}}{\delta U_i} &= u_i \\ \frac{1}{\rho} \frac{\delta \mathcal{L}}{\delta b} &= z \\ \frac{\delta \mathcal{L}}{\delta \rho} &= \frac{1}{2} u_i U_i + bz - \frac{p}{\rho_0} \\ \frac{\delta \mathcal{L}}{\delta p} &= \frac{\rho}{\rho_0} - 1. \end{aligned} \quad (\text{C.2})$$

600 The total pressure $p = p^h + q$ arises as a Lagrange multiplier enforcing
 601 the incompressibility constraint $\rho = \rho_0$. The Euler Poincaré equation, or

602 momentum equation, reads (Holm et al., 1998)

$$\partial_t \mathbf{u} + \text{curl } \mathbf{u} \times \mathbf{U} + \nabla (\mathbf{u} \cdot \mathbf{U}) - \nabla \frac{\delta \mathcal{L}}{\delta \rho} + \frac{1}{\rho} \frac{\delta \mathcal{L}}{\delta b} \nabla b = 0, \quad (\text{C.3})$$

603 namely

$$\partial_t \mathbf{u} + \text{curl } \mathbf{u} \times \mathbf{U} = -\nabla \left(\frac{1}{2} \mathbf{U} \cdot \mathbf{u} + \frac{p}{\rho_0} \right) + b \nabla z. \quad (\text{C.4})$$

604 (C.4) is the momentum equation in vector-invariant form. It is identical to
 605 Tort and Dubos (2014). The formalism comes with the tracer conservation
 606 equation $\partial_t \phi + \text{div}(\mathbf{U} \phi) = 0$, that holds for any tracer. In particular it
 607 translates into the continuity equation (17) for $\phi = \rho$. We now show how
 608 (C.4) can be transformed into the flux form (21) plus a pseudo force $\tilde{\mathbf{G}}$. The
 609 idea is to prove that

$$\left[\text{curl } \mathbf{u} \times \mathbf{U} + \nabla \left(\frac{1}{2} \mathbf{U} \cdot \mathbf{u} \right) \right] \mathcal{V} = \partial_j (\mathcal{A}_j U_j \mathbf{u}) - \tilde{\mathbf{G}} \mathcal{V}. \quad (\text{C.5})$$

610 In components the curl term reads

$$\begin{aligned} \mathcal{V} (\text{curl } \mathbf{u} \times \mathbf{U})|_i &= \epsilon_{ijk} \epsilon_{jlm} \partial_l (u_m h_m) U_k \frac{\mathcal{A}_k}{h_i} \\ &= (\delta_{kl} \delta_{im} - \delta_{km} \delta_{il}) \partial_l (u_m h_m) U_k \frac{\mathcal{A}_k}{h_i} \\ &= \partial_k (U_k \mathcal{A}_k u_i) - U_k \mathcal{A}_i \partial_i u_k + U_k \frac{\mathcal{A}_k}{h_i} (u_i \partial_k h_i - u_k \partial_i h_k), \end{aligned} \quad (\text{C.6})$$

611 where we have used $\partial_k (U_k \mathcal{A}_k) = 0$. The next step is to rearrange

$$\begin{aligned} \mathcal{V} \nabla \left(\frac{1}{2} \mathbf{U} \cdot \mathbf{u} \right)|_i &= \frac{1}{2} \mathcal{A}_i \partial_i (U_k u_k) \\ &= \mathcal{A}_i U_k \partial_i u_k + \frac{1}{2} \mathcal{A}_i (u_k \partial_i U_k - U_k \partial_i u_k) \\ &= \mathcal{A}_i U_k \partial_i u_k + \frac{1}{2} \mathcal{A}_i u_k u_j \partial_i T_{jk}. \end{aligned} \quad (\text{C.7})$$

612 Summing (C.6) and (C.7) gives (C.5) with the general expression for the
613 pseudo-force

$$\tilde{\mathbf{G}}_i = \frac{U_k}{h_i h_k} (u_k \partial_i h_k - u_i \partial_k h_i) - \frac{1}{2h_i} u_k u_j \partial_i T_{jk}. \quad (\text{C.8})$$

614 Using the expression for T_{ij} allows to explicit each component

$$\begin{aligned} \tilde{G}_u &= +(v \partial_\xi h_\eta - u \partial_\eta h_\xi) \frac{V}{h_\xi h_\eta} + \frac{wW}{h_\xi h_\sigma} \partial_\xi h_\sigma + \frac{wU}{h_\xi} \partial_\xi s_\xi + \frac{wV}{h_\xi} \partial_\xi s_\eta \\ \tilde{G}_v &= -(v \partial_\xi h_\eta - u \partial_\eta h_\xi) \frac{U}{h_\xi h_\eta} + \frac{wW}{h_\eta h_\sigma} \partial_\eta h_\sigma + \frac{wU}{h_\eta} \partial_\eta s_\xi + \frac{wV}{h_\eta} \partial_\eta s_\eta \\ \tilde{G}_w &= 0, \end{aligned} \quad (\text{C.9})$$

615 where we have used the shallow fluid approximation $\partial_\sigma h_\xi = \partial_\sigma h_\eta = 0$ and
616 $\partial_\xi h_\sigma = h_\xi \partial_\sigma s_\xi$, $\partial_\eta h_\sigma = h_\eta \partial_\sigma s_\eta$. This pseudo-force can be transformed further
617 into the form of (21–22), namely

$$\begin{aligned} \tilde{G}_u &= G_u + \frac{wU}{h_\xi} \partial_\xi s_\xi + \frac{wV}{h_\eta} \partial_\eta s_\xi + \frac{wW}{h_\sigma} \partial_\sigma s_\xi \\ \tilde{G}_v &= G_v + \frac{wU}{h_\xi} \partial_\xi s_\eta + \frac{wV}{h_\eta} \partial_\eta s_\eta + \frac{wW}{h_\sigma} \partial_\sigma s_\eta, \end{aligned} \quad (\text{C.10})$$

618 with

$$\begin{aligned} G_u &= +(V \partial_\xi h_\eta - U \partial_\eta h_\xi) \frac{V}{h_\xi h_\eta} \\ G_v &= -(V \partial_\xi h_\eta - U \partial_\eta h_\xi) \frac{U}{h_\xi h_\eta}, \end{aligned} \quad (\text{C.11})$$

619 where we have used the property

$$\frac{1}{h_\xi} \partial_\xi s_\eta - \frac{1}{h_\eta} \partial_\eta s_\xi = \frac{1}{h_\xi h_\eta} (s_\xi \partial_\eta h_\xi - s_\eta \partial_\xi h_\eta). \quad (\text{C.12})$$

620 The pseudo-force differs from the hydrostatic case by the presence of terms
621 proportional to w , the vertical momentum. They vanish in the hydrostatic
622 case because the vertical momentum is identically zero.

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626 **References**

627 M. Nikurashin, G. K. Vallis, A. Adcroft, Routes to energy dissipation for
628 geostrophic flows in the Southern Ocean, *Nature Geoscience* 6 (1) (2013)
629 48–51.

630 M. J. Molemaker, J. C. McWilliams, W. K. Dewar, Submesoscale instability
631 and generation of mesoscale anticyclones near a separation of the California
632 Undercurrent, *Journal of Physical Oceanography* 45 (3) (2015) 613–629.

633 J. Callies, R. Ferrari, J. M. Klymak, J. Gula, Seasonality in submesoscale
634 turbulence, *Nature communications* 6.

635 J. Gula, M. J. Molemaker, J. C. McWilliams, Gulf Stream dynamics along
636 the southeastern US seaboard, *Journal of Physical Oceanography* 45 (3)
637 (2015) 690–715.

638 P. P. Sullivan, J. C. McWilliams, C.-H. Moeng, A subgrid-scale model for
639 large-eddy simulation of planetary boundary-layer flows, *Boundary-Layer*
640 *Meteorology* 71 (3) (1994) 247–276.

641 J. Marshall, C. Hill, L. Perelman, A. Adcroft, Hydrostatic, quasi-hydrostatic,
642 and nonhydrostatic ocean modeling, *J. Geophys. Res.* 102 (1997) 5733–
643 5752.

- 644 Y. Kanarska, V. Maderich, A non-hydrostatic numerical model for calculat-
645 ing free-surface stratified flows, *Ocean Dynamics* 53 (3) (2003) 176–185.
- 646 Y. Kanarska, A. Shchepetkin, J. McWilliams, Algorithm for non-hydrostatic
647 dynamics in the regional oceanic modeling system, *Ocean Modelling* 18 (3)
648 (2007) 143–174.
- 649 F. Auclair, C. Estournel, J. W. Floor, M. Herrmann, C. Nguyen,
650 P. Marsaleix, A non-hydrostatic algorithm for free-surface ocean modelling,
651 *Ocean modelling* 36 (1) (2011) 49–70.
- 652 K. Klingbeil, H. Burchard, Implementation of a direct nonhydrostatic pres-
653 sure gradient discretisation into a layered ocean model, *Ocean Modelling*
654 65 (2013) 64–77.
- 655 W. C. Skamarock, J. B. Klemp, A time-split nonhydrostatic atmospheric
656 model for weather research and forecasting applications, *Journal of Com-
657 putational Physics* 227 (7) (2008) 3465–3485.
- 658 E. H. Müller, R. Scheichl, Massively parallel solvers for elliptic partial dif-
659 ferential equations in numerical weather and climate prediction, *Quarterly
660 Journal of the Royal Meteorological Society* 140 (685) (2014) 2608–2624.
- 661 S. Williams, A. Waterman, D. Patterson, Roofline: an insightful visual per-
662 formance model for multicore architectures, *Communications of the ACM*
663 52 (4) (2009) 65–76.
- 664 A. F. Shchepetkin, J. C. McWilliams, A method for computing horizontal
665 pressure-gradient force in an oceanic model with a nonaligned vertical co-
666 ordinate, *Journal of Geophysical Research: Oceans* 108 (C3).

- 667 M. A. Taylor, A. Fournier, A compatible and conservative spectral element
668 method on unstructured grids, *Journal of Computational Physics* 229 (17)
669 (2010) 5879–5895.
- 670 M. Desbrun, E. Kanso, Y. Tong, Discrete differential forms for computational
671 modeling, in: *Discrete differential geometry*, Springer, 287–324, 2008.
- 672 J. Thuburn, C. J. Cotter, A framework for mimetic discretization of the ro-
673 tating shallow-water equations on arbitrary polygonal grids, *SIAM Journal*
674 *on Scientific Computing* 34 (3) (2012) B203–B225.
- 675 M. J. Molemaker, J. C. McWilliams, I. Yavneh, Baroclinic instability and loss
676 of balance, *Journal of Physical Oceanography* 35 (9) (2005) 1505–1517.
- 677 T. Dubos, S. Dubey, M. Tort, R. Mittal, Y. Meurdesoif, F. Hourdin, DY-
678 NAMICO, an icosahedral hydrostatic dynamical core designed for consis-
679 tency and versatility, *Geoscientific Model Development Discussions* 8 (2)
680 (2015) 1749–1800.
- 681 A. J. Chorin, Numerical solution of the Navier-Stokes equations, *Mathemat-*
682 *ics of computation* 22 (104) (1968) 745–762.
- 683 J. B. Bell, P. Colella, H. M. Glaz, A second-order projection method for the
684 incompressible Navier-Stokes equations, *Journal of Computational Physics*
685 85 (2) (1989) 257–283.
- 686 L. J. Grady, J. Polimeni, *Discrete calculus: Applied analysis on graphs for*
687 *computational science*, Springer Science & Business Media, 2010.

688 M. Tort, T. Dubos, Dynamically consistent shallow-atmosphere equations
689 with a complete Coriolis force, Quarterly Journal of the Royal Meteorological
690 Society 140 (684) (2014) 2388–2392.

691 D. D. Holm, J. E. Marsden, T. S. Ratiu, The Euler-Poincaré equations
692 and semidirect products with applications to continuum theories, Adv.
693 in Math. 137 (1998) 1–81.

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