

# Supporting Information for "Large scale forces under surface gravity waves at a wavy bottom: a mechanism for the generation of primary microseisms"

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## Introduction

### Text S1. Modulation factor for the wavenumber over a cosine bottom

For small bottom amplitudes, the non-dimensional coefficient  $\alpha$ , is the sum of the two modulation effects listed in the paper: the variation of the local wavenumber, and the

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variation of the bottom pressure amplitude, which itself can be split into a modulation of the surface to bottom transfer function, and a modulation of the surface elevation amplitude, conserving the energy flux. This gives  $\alpha = \alpha_1 + (\alpha_2 + \alpha_3)$ .

We first estimate the wavenumber, given by Taylor expansion around the depth  $D_0$ ,

$$k(x) = k_0 - d_0 \cos(k_b x) \frac{\partial k_0}{\partial D_0} + O\left(d_0^2 \frac{\partial^2 k}{\partial D^2}\right). \quad (1)$$

Differentiating the dispersion relation

$$\sigma^2 = gk(x) \tanh[k(x)D(x)]. \quad (2)$$

gives for a constant  $\sigma$  [Ardhuin and Herbers, 2002, appendix D, eq. (D.10)],

$$2\sigma \frac{\partial \sigma}{\partial D} = g \tanh(kD) \frac{\partial k}{\partial D} + \frac{gk}{\cosh^2(kD)} \left(k + D \frac{\partial k}{\partial D}\right) = 0 \quad (3)$$

which for  $D = D_0$  gives,

$$\frac{\partial k_0}{\partial D_0} = \frac{-2k_0^2}{2k_0 D_0 + \sinh(2k_0 D_0)}. \quad (4)$$

Second, we estimate the spectrum of a constant amplitude signal that is modulated in wavenumber. For this we write the phase  $S(x)$  using  $k(x)$  given by eq. (1),

$$S(x) = \int k(x) dx \simeq \int \left(k_0 - d_0 \cos(k_b x) \frac{\partial k_0}{\partial D_0}\right) dx = k_0 x + \epsilon \sin(k_b x) + O\left(d_0^2 \frac{\partial^2 k_0}{\partial D_0^2}\right), \quad (5)$$

where we have used the non-dimensional modulation index  $\epsilon$  given by

$$\epsilon = -\frac{d_0}{k_b} \frac{\partial k_0}{\partial D_0}. \quad (6)$$

Following Cuyt et al. [2008], the Fourier decomposition of a signal with such a modulated phase is given by the Laurent expansion of  $\exp[iz \sin(k_b x)]$ ,

$$\cos[k_0 x + \epsilon \sin(k_b x) - \sigma t] = \sum_{n=-\infty}^{\infty} J_n(\epsilon) \cos[(k_0 + nk_b)x - \sigma t], \quad (7)$$

where  $J_n(\epsilon)$  is the  $n$ -th order Bessel function of the first kind. When  $\epsilon \ll 1$ , we may replace the Bessel function by its asymptote  $J_{-1}(\epsilon) \simeq -\epsilon/2$ . We define

$$\alpha_1 = -\frac{\epsilon}{k_0 d_0}. \quad (8)$$

Considering only the lowest harmonics in (7) and in particular the long wavelength component,

$$\cos[S(x) - \sigma t] \simeq \cos(k_0 x - \sigma t) - \alpha_1 k_0 d_0 \frac{1}{2} \cos(k_0 x - k_b x - \sigma t) + \dots \quad (9)$$

Using  $k_b \simeq k_0$ , the value of  $\alpha_1$  simplifies to

$$\alpha_1 = -\frac{d_0}{k_0^2 d_0} \frac{-2k_0^2}{2k_0 D_0 + \sinh(2k_0 D_0)} = \frac{2}{2k_0 D_0 + \sinh(2k_0 D_0)}. \quad (10)$$

The neglected terms in eq. (5) can also give long wavelength components but only at third order with an amplitude  $O(d_0^3 \partial^3 k / \partial D^3 / k_b)$ . Indeed, the second order term is proportional to  $\cos^2(k_b x)$  and thus gives Fourier components  $\cos[(k_0 + 2nk_b)x - \sigma t]$  that cannot produce long wavelengths for  $k_b \simeq k_0$  but would contribute for  $k_b \simeq k_0/2$  when a wide bottom spectrum is considered.

## **Text S2. Modulation factors for the bottom pressure amplitude over a cosine bottom**

Linear wave theory gives

$$p(z = -D) = \rho_w g \frac{a(x) \cos[S(x) - \sigma t]}{\cosh(k(x)D(x))}. \quad (11)$$

With  $y = k_0 D_0$ , this expands as

$$p(z = -D) = \rho_w g \frac{a(x) \cos[S(x) - \sigma t]}{\cosh y} \frac{\cosh y}{\cosh[k(x)D(x)]} \quad (12)$$

$$\simeq \rho_w g \frac{a(x) \cos[S(x) - \sigma t]}{\cosh y} [1 - \alpha_2 k_0 d(x)], \quad (13)$$

with

$$\alpha_2 = - \left( 1 + \frac{D_0}{k_0} \frac{\partial k_0}{\partial D_0} \right) \tanh(k_0 D_0), \quad (14)$$

$$= -(\sinh^2(y)) / [y + \sinh(y) \cosh(y)] \quad (15)$$

Similarly, the variation of  $a(x)$  given by  $C_g(x)a^2(x) = C_{g0}a_0^2$  can be expanded as

$$a(x) \simeq a_0 [1 - \alpha_3 k_0 d(x)] \quad (16)$$

$$\text{with } \alpha_3 = -\frac{0.5}{C_{g0}k_0} \frac{\partial C_{g0}}{\partial D_0}. \quad (17)$$

The linear group velocity  $C_g = \sigma/k(0.5 + kD/\sinh(2kD))$  gives

$$\frac{\partial C_g}{\partial D} = \frac{\sigma}{k_0} \left[ \frac{1}{\sinh(2kD)} \frac{\partial(kD)}{\partial D} - 2kD \frac{\partial(kD)}{\partial D} \frac{\cosh(2kD)}{\sinh(2kD)^2} \right] - \frac{\partial k}{\partial D} \frac{C_g}{k} \quad (18)$$

$$= \frac{\sigma}{k_0} \frac{\partial(k_0 D_0)}{\partial D_0} \left[ \frac{1}{\sinh(2y)} - 2y \frac{\cosh(2y)}{\sinh^2(2y)} \right] - 0.5 \frac{\sigma}{k_0^2} \frac{\partial k_0}{\partial D_0} \left[ 1 + \frac{2y}{\sinh(2y)} \right]. \quad (19)$$

### Text S3. Combined modulation factor

We now combine  $\cos[S(x) - \sigma t]$  using eq. (9) with the results on the bottom pressure amplitude, giving

$$p(z = -D) \simeq \rho_w g a_0 \cos(k_0 x - \sigma t) / \cosh(k_0 D_0) + p_{ls} + \dots \quad (20)$$

with the large scale components containing a phase  $[(k_0 - k_b)x - \sigma t]$  given by

$$p_{ls} = -(\alpha_1 + \alpha_2 + \alpha_3) \frac{k_0 d_0}{2} \cos(k_0 x - k_b x - \sigma t) \frac{\rho_w g a_0}{\cosh(k_0 D_0)}. \quad (21)$$

For  $\epsilon \ll 1$ ,  $\alpha$  is a function of  $y = kD$  that is the sum of the three parts

$$\alpha_1 = \frac{2}{2y + \sinh(2y)} = \frac{1}{y + \sinh(y) \cosh(y)}, \quad (22)$$

$$\alpha_2 = -\frac{\sinh^2 y}{y + \sinh y \cosh y}, \quad (23)$$

$$\alpha_3 = -\frac{2 \sinh(2y) - 4y \sinh^2 y}{(2y + \sinh(2y))^2} = -\frac{\sinh y \cosh y - y \sinh^2 y}{(y + \sinh y \cosh y)^2}. \quad (24)$$

Putting all three terms with a common denominator gives,

$$\alpha(y) = \frac{y + \sinh(y) \cosh(y) - [\sinh(y) \cosh(y) + \cosh(y) \sinh^3(y)]}{[y + \sinh(y) \cosh(y)]^2} = \frac{y - \cosh(y) \sinh^3(y)}{[y + \sinh(y) \cosh(y)]^2}. \quad (25)$$

**Figure S1.** (a) example of the surface elevation and bottom pressure field for  $D_0 = 35$  m and  $f = 0.085$  Hz. (b) same surface and elevation with an exponential attenuation of the energy flux  $\exp(nx)$  with  $n = 10^{-5}\text{m}^{-1}$ . (c) and (d) corresponding spectra of the bottom pressure estimated using the numerical data, or from the theoretical large-scale only given by eq. (13) of the paper.

**Figure S2.** Matlab code used to check the analytical results

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