

Western Mediterranean Sea Paleothermometry Over the Last Glacial Cycle Based on the Novel RI-OH Index

Nina Davtian^{1*}, Guillemette Ménot², Yoann Fagault¹, and Edouard Bard¹

¹CEREGE, Aix Marseille Univ, CNRS, IRD, INRA, Collège de France, Technopôle de l'Arbois, BP 80, 13545 Aix-en-Provence, France

²Univ Lyon, ENS de Lyon, Université Lyon 1, CNRS, UMR 5276 LGL-TPE, F-69364, Lyon, France

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Introduction

The Supporting Information contains two Figures and one Table. Figure S1 illustrates how we identified and quantified OH-GDGTs. Figure S2 illustrates our assessment of biases for TEX₈₆. Table S1 summarizes our sampling strategy.

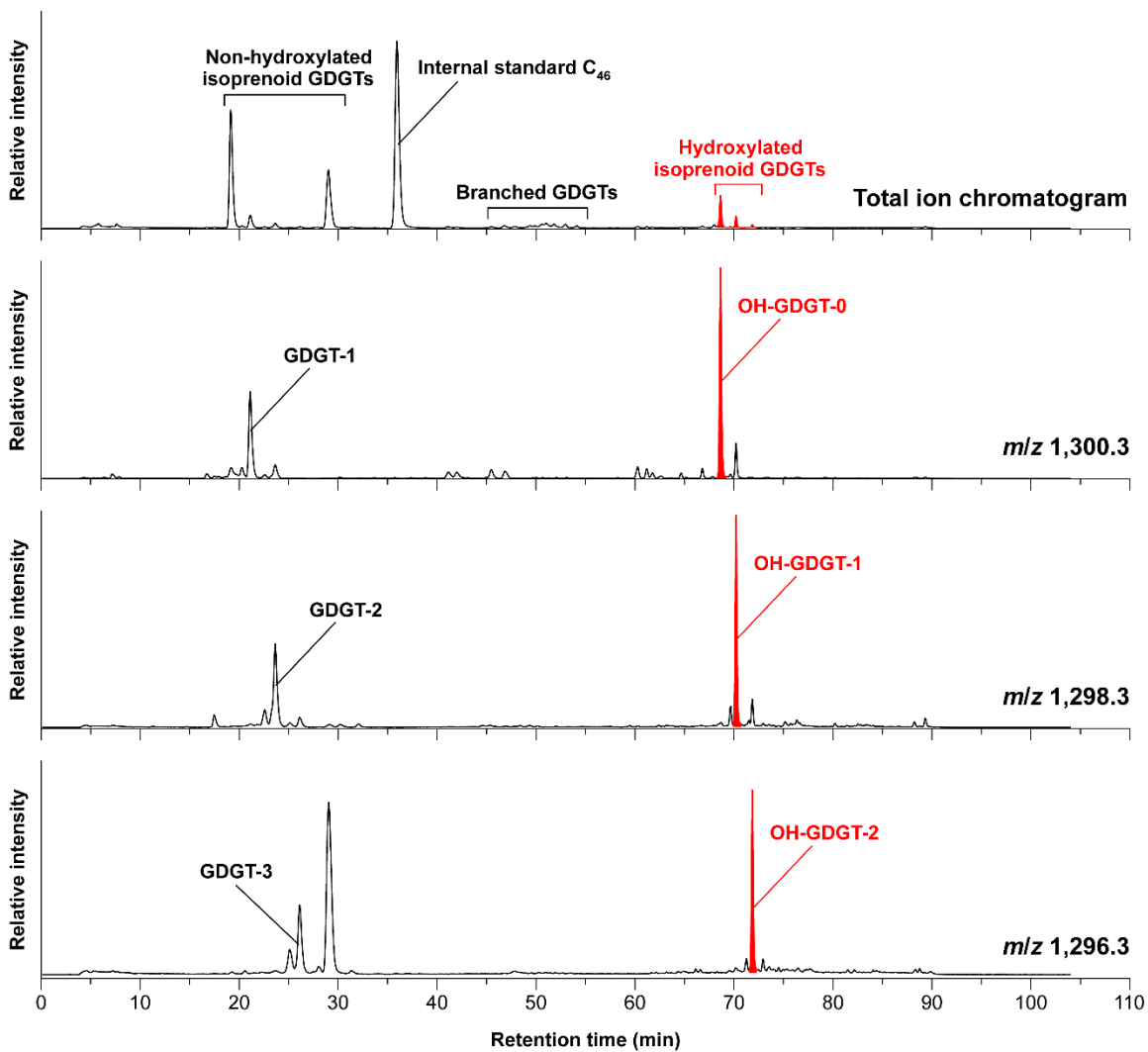


Figure S1. Total ion chromatogram and m/z 1,300.3, 1,298.3, and 1,296.3 single ion monitoring scans for the sediment sample MD99-2348_S12_50-52cm. OH-GDGT peaks are highlighted in red. Other families of GDGTs, the internal standard C₄₆ (Huguet et al., 2006), and some iGDGTs are identified as well.

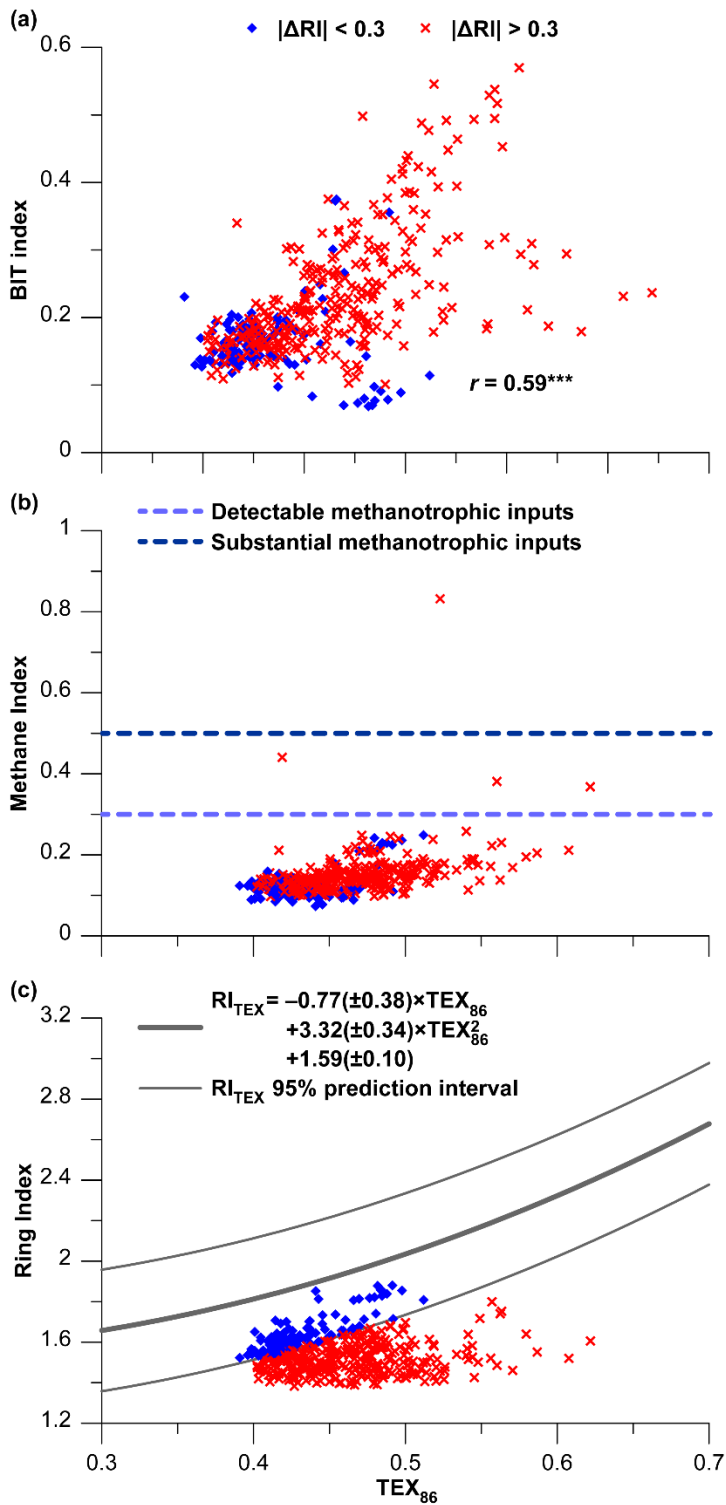


Figure S2. Assessment of biases for TEX_{86} (Schouten et al., 2002) in MD99-2348/PRGL1-4 sediments. (a) BIT index (Hopmans et al., 2004) vs TEX_{86} for terrestrial inputs. (b) Methane Index (Zhang et al., 2011) vs TEX_{86} for methanotrophic inputs. (c) Ring Index (Zhang et al.,

2016) vs TEX_{86} for all biases. Sediment samples with $|\Delta RI|$ values below and above 0.3 are distinguished.

Core/Borehole	<i>n</i>	Depth range (mbsf)	MIS	Sampling interval (cm)	Support
MD99-2348	11	0.01–0.21	1 and 2	2	U-Channels
MD99-2348	124	0.31–22.42	2	18 on average	U-Channels
PRGL1-4	86	19.20–47.75	2 and 3	34 on average	U-Channels
PRGL1-4	48	47.95–62.90	3 and 4	32 on average	Cubes
PRGL1-4	74	63.15–72.50	5	13 on average	Cubes
PRGL1-4	82	73.11–109.54	6	45 on average	Cubes

Table S1. Sampling strategy of core MD99-2348 and borehole PRGL1-4 for GDGT analysis. Sediment samples taken from U-Channels are 2-cm-thick slices. Cubes have a volume of 8 cm³ each and were taken following as closely as possible the sampling strategy of Cortina et al. (2015, 2016). MIS = Marine Isotope Stage.

Data Set S1. Metadata (core/borehole, section, and sample name) and original data (depth, age, and tetraether-based data). The 1σ analytical uncertainties are 0.009 for RI-OH, 0.003 for TEX_{86} , BIT, and MI, and 0.019 for RI. These correspond to ranges of 0.5 and 0.2 °C in RI-OH-temperatures and TEX_{86} -temperatures, respectively. The age model is based on Sierra et al. (2009) and Frigola et al. (2012). The calibration used for RI-OH-based sea surface temperature (SST) reconstructions is $SST = (RI-OH - 1.11)/0.018$ (Lü et al., 2015).

Data Set S2. Annual sea surface temperature (SST) anomalies for the latest two glacial-interglacial transitions per site and per proxy. The anomalies are defined as the difference between the interglacial optimum and the preceding glacial maximum, rounded to the closest unit – $SST \text{ anomaly} = SST \text{ interglacial optimum} - SST \text{ glacial maximum}$. Anomalies MIS 1 – 2 and MIS 5e – 6 refer to Holocene optimum – Last Glacial Maximum anomalies and Eemian optimum – Marine Isotope Stage 6 anomalies, respectively.

Data Set S3. Sea surface temperature (SST) anomalies between the mid-Holocene (6 ka BP) and the Last Glacial Maximum (LGM, 21 ka BP) based on all Paleoclimate Modeling Intercomparison Project Phase 3 (PMIP3) models with the corresponding runs – IPSL-CM5A-LR (Dufresne et al., 2013), FGOALS-g2 (Li et al., 2013), CNRM-CM5 (Voldoire et al., 2013), MIROC-ESM (Watanabe et al., 2011), MRI-CGCM3 (Yukimoto et al., 2012), CCSM4 (Gent et al., 2011), MPI-ESM-P (Giorgetta et al., 2013), and GISS-E2-R (Shindell et al., 2012). The anomalies are defined as the difference between the mid-Holocene and the LGM, rounded to the closest unit – $\text{Modeled SST anomaly} = \text{Modeled SST mid-Holocene} - \text{Modeled SST LGM}$. Model data refer to the sites with corresponding proxy data (see Data Set S2).