



Quantifying Biogenic Versus Detrital Carbonates on Marine Shelf: An Isotopic Approach

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The terrigenous sedimentary budget of passive margins, records variations in past sedimentary fluxes, and thus can be used to infer past variations of Earth surface deformation processes or climate change. Accurate estimates of sediment fluxes over various times and spatial scales are therefore crucial. Traditionally, offshore sediment volume determination only considers siliciclastic accumulation, the carbonate fraction (i.e., CaCO₃) being considered only as *in situ* production. Here we propose a new geochemical methodology to decipher and quantify the number of detrital carbonates in comparison to *in situ* produced biogenic carbonates. This isotopic approach enables considering the export of detrital carbonates and investigating its effect on sediment budgets. This study, located in the Gulf of Lion, is based on a 300 m long sediment borehole located near the shelf break and covering the last 500 000 years (i.e., five glacial-interglacial periods). Strontium isotope (⁸⁷Sr/⁸⁶Sr) of carbonate fractions (0.70809 to 0.70858) are significantly less radiogenic than modern seawater (i.e., 0.7092) and show fluctuations in agreement with stratigraphic and climatic variations. These results suggest an unsuspected high export of detrital carbonates from the catchment area during both glacial (between 55 and 85% of the sedimentary carbonate fraction) and interglacial (between 30 and 50%) conditions. Thus, not only do detrital carbonate fluxes need to be factored into sediment flux calculations, but these results also suggest that detrital carbonate components could potentially have a strong influence on bulk carbonate ⁸⁷Sr/⁸⁶Sr ratios when not obtained from micro drilled biogenic carbonates, such as the entirety of the Precambrian Sr chemostratigraphic record.

Keywords: ⁸⁷Sr/⁸⁶Sr carbonate, detrital carbonate export, source-to-sink, glacial – interglacial, strontium isotope stratigraphy

INTRODUCTION

An extensive dataset has been collected over the last decade on marine carbonates and fossils, to document past variations in the strontium isotopic composition record (⁸⁷Sr/⁸⁶Sr), as a tool to reconstruct changes in the seawater composition through Earth's history. These past variations are of interest for two reasons: first, ⁸⁷Sr/⁸⁶Sr ratios measured on marine carbonate offer a widely

used chronostratigraphic tool (SIS – Strontium Isotope Stratigraphy; for reviews of SIS the readers should refer to Elderfield, 1986; McArthur, 1994; Veizer et al., 1997, 1999; McArthur et al., 2012); second, secular changes in Sr isotope composition provide information about the geochemical cycling of strontium in the ocean, weathering processes, hydrothermal circulation, and carbonate dissolution at the sea-floor (Burke et al., 1982; De Paolo and Ingram, 1984; Veizer, 1989; Prokoph et al., 2008; Allègre et al., 2010 among others).

Since the rapid increase of calcifying organisms (i.e., Mesozoic), the SIS method relies on two assumptions: (i) strontium isotopes composition measurements from well preserved, non-altered, shell material (i.e., bivalves, rudist, belemnites, planktonic foraminifera) are assumed to reflect the seawater Sr isotope composition from which they precipitated (Veizer, 1989; McArthur et al., 2012); (ii) the world's ocean is homogenous with respect to $^{87}\text{Sr}/^{86}\text{Sr}$, and always has been. Such uniformity is expected because Sr residence time in the ocean (4 Myr; Broecker, 1963; Goldberg, 1963; Hodell et al., 1990) is far longer than their mixing time ($\sim 10^3$ years). Since the 1980s, the SIS tool relies on the extensive compilation of data, cross calibration between carbonate producers (ammonites, foraminifera, calcareous nannofossils) and/or other geochemical proxies ($\delta^{13}\text{C}$, $\delta^{18}\text{O}$). Extensive compilations of measured $^{87}\text{Sr}/^{86}\text{Sr}$ in marine material yield to the development of numerical age determination tools (McArthur et al., 2012 for a thorough review). The degree to which SIS numerical dating is correct mostly depends on the slope and the accuracy of the age model during the given time interval, as well as the quality and state of preservation of the studied material. Given the difficulty of dating sedimentary material, especially during the Paleozoic and Precambrian era, it could be tempting to use a “bulk” SIS approach by measuring the $^{87}\text{Sr}/^{86}\text{Sr}$ ratio of the entire carbonate fraction preserved in the sediment. However, such a carbonate component may not necessarily entirely be produced *in situ* and could include a significant detrital proportion. If so, not only would the SIS-derived ages be biased, but the inferred flux of detrital material to total sediment would also be underestimated. This is of great importance when calculating sedimentation rates in stratigraphic simulations studies (Allen, 1974; Castellort and Van Den Driessche, 2003; Allen, 2008; Armitage et al., 2011, 2013; Simpson and Castellort, 2012; Romans et al., 2016); but also when reconstructing temporal variations in the marine $^{87}\text{Sr}/^{86}\text{Sr}$ or to understand the relative Sr fluxes related to continental weathering versus hydrothermal inputs over time (e.g., Burke et al., 1982; Veizer, 1989; Veizer et al., 1999; Halverson et al., 2007; Prokoph et al., 2008).

In this study, we examined the Sr isotope compositions preserved in the sediment carbonate fractions from the Gulf of Lion which were deposited during one glacial advance and the following retreat between 160 to 120 ka, i.e., over the MIS 6 to MIS 5 transition. The goals of this exploratory project were first to test whether detrital carbonates could be preserved in shelf accumulation and developed an isotopic approach that could help in their quantification.

We secondly explore their potential effect on sedimentary fluxes calculations.

GEOLOGICAL BACKGROUND AND ANALYTICAL RESULTS

The Gulf of Lion (GoL), located in the North-Western Mediterranean, is characterized by a wide continental shelf (70 km) that was sub-aerially exposed during glacial periods over the Late Quaternary period (Rabineau et al., 2005). The sedimentation is mainly dominated by the Rhône River inputs which currently provide about 80% of the total sediment flux (Aloisi et al., 1977; de Madron et al., 2000; Molliex et al., 2016). Rivers from the Pyrenees and Languedoc (Hérault, Orb, Aude, Agly, Tech and Têt; **Figure 1**) supply the remaining fraction. The GoL catchment is composed of (i) crystalline rocks, located in mountainous areas (Inner Alps, Massif Central, Pyrenees), and (ii) a large part of carbonated rocks, mostly marl and limestone from the Mesozoic era (Jurassic and Cretaceous) located in Alpine foreland. Some Cenozoic carbonates (Eocene and Miocene), mostly bioclastic and continental sandy limestone, are also present in the downstream part of the catchment and in the foreland of the Alps and Pyrenees and (iii) Pliocene-Quaternary formations which consist of fluvial deposits (Molliex et al., 2016). This study is based on sediment samples collected from borehole PRGL1-4 (**Figure 1**), drilled in the framework of the EU PROMESS project¹, which sampled a 300 m long continuous record spanning the last five glacial-interglacial cycles. The sedimentary succession consists of five progradational units related to the 100-kyr glacio-eustatic cyclicity (Rabineau et al., 2005, 2006). In addition to the moving (regression-transgression) of the shoreline and associated sedimentary environments, the sea-level strongly controls the connection of the riverine inputs with the upper slope setting. As a consequence, the sediment column at PRGL1-4 is essentially composed of fine siliciclastic grains detrital sediment with several interbedded cm-thick sandy-size layers, made mostly of foraminifera shells accumulation, marking the periods of shelf maximum flooding during interglacial periods (**Figure 1**; Sierro et al., 2009; Frigola et al., 2012). The stacking of 100-kyr sequences is favored by a high subsidence rate in that area (Rabineau et al., 2014). Sediment provenance studies indicate a predominance of Rhône river sediment (Revillon et al., 2011). These studies however, only considered siliciclastic material and, no reliable information on carbonate export is available, although carbonated rocks constitute more than 50% of the Gulf of Lion catchment area, and 40% of the eroded volumes (Molliex et al., 2016).

A total of 12 sediment samples from PRGL1-4 were selected along the glacial retreat from MIS 6 (i.e., 160 ka) toward the MIS 5 climatic optimum (i.e., 120 ka), together with 10 samples from modern Rhône river tributaries riverbeds (see **Figure 1** for location). In the present study we analyzed the

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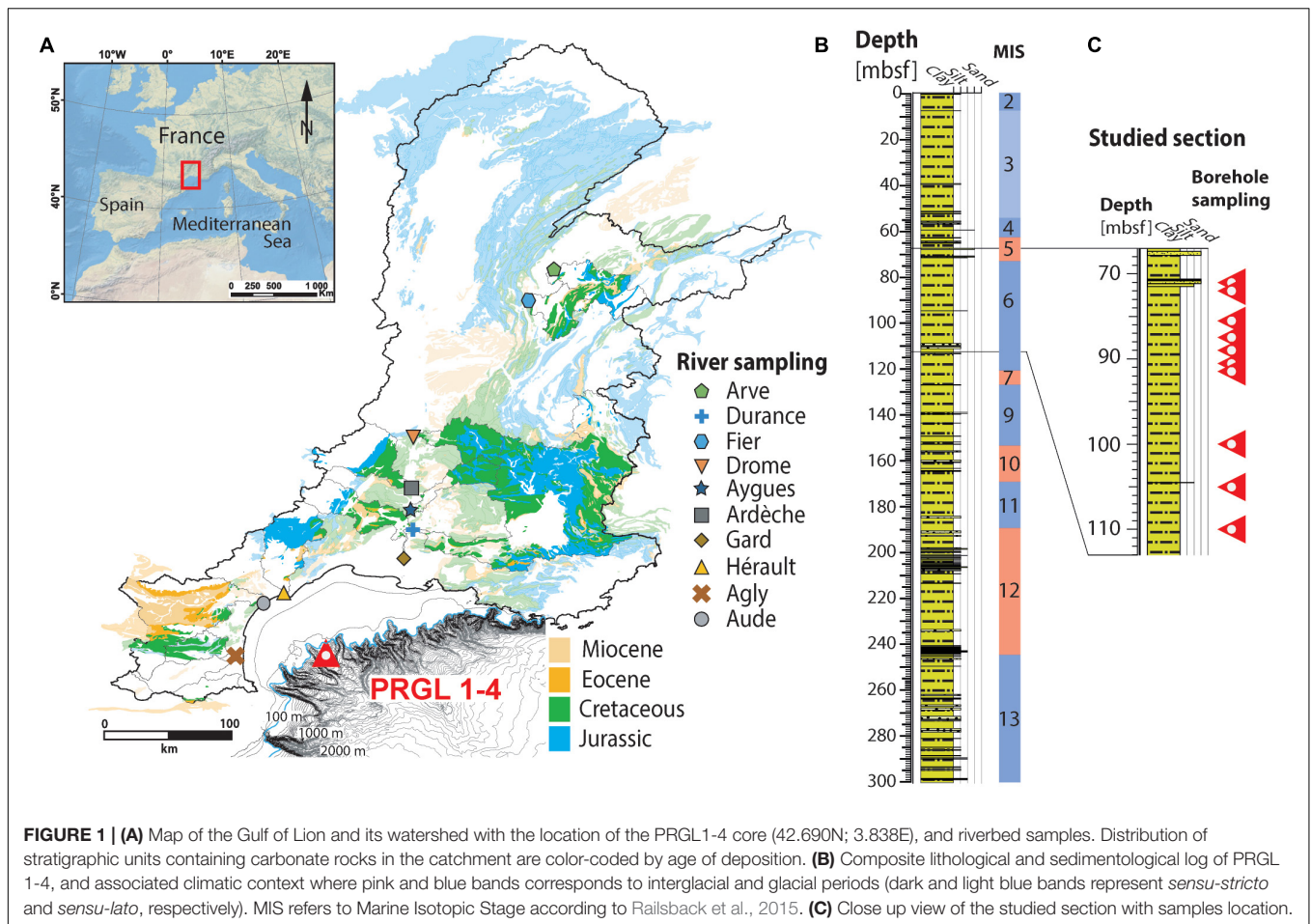


FIGURE 1 | (A) Map of the Gulf of Lion and its watershed with the location of the PRGL1-4 core (42.690N; 3.838E), and riverbed samples. Distribution of stratigraphic units containing carbonate rocks in the catchment are color-coded by age of deposition. **(B)** Composite lithological and sedimentological log of PRGL 1-4, and associated climatic context where pink and blue bands corresponds to interglacial and glacial periods (dark and light blue bands represent *sensu-stricto* and *sensu-lato*, respectively). MIS refers to Marine Isotopic Stage according to Railsback et al., 2015. **(C)** Close up view of the studied section with samples location.

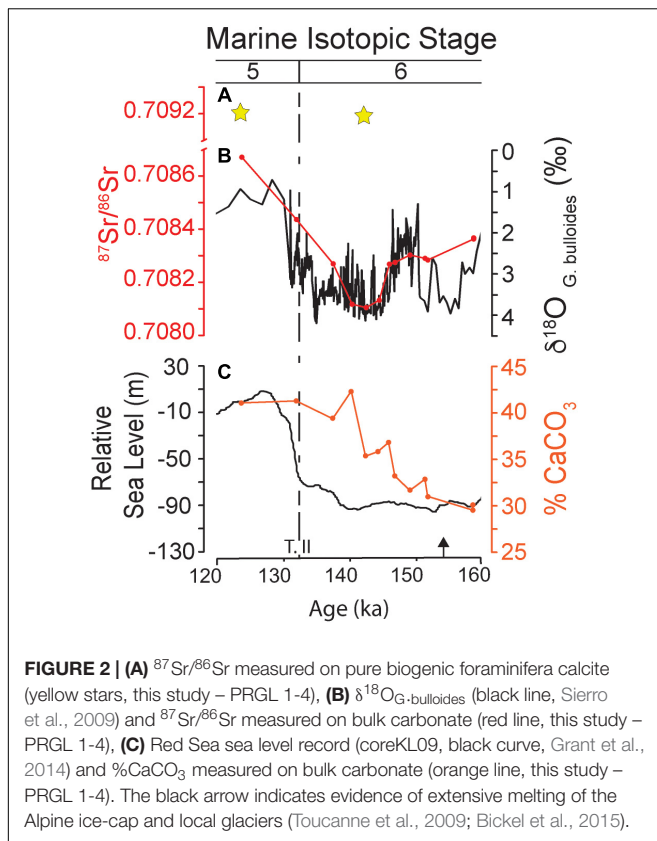
Sr isotope compositions, trace elements, and CaCO_3 contents of bulk carbonate samples extracted (i.e., leached) from both marine (i.e., PRGL1-4) and river sediments (i.e., from very coarse sand to clay; **Supplementary Table S1**), using 5% acid acetic digestion. Strontium was isolated from the matrix by column chromatography using a Sr-Spec resin (Eichrom®) prior to analysis by TIMS (Thermo Fisher Scientific TRITON) at the Pôle Spectrométrie Océan (Brest, France). Purified Sr were loaded on single W and measured on static mode. All measured Sr ratios were normalized to $^{86}\text{Sr} / ^{88}\text{Sr} = 0.1194$. During the course of analysis, Sr isotope compositions of standard solution NBS987 gave $^{87}\text{Sr}/^{86}\text{Sr} = 0.710259 \pm 7$ (2σ , $n = 9$, recommended value 0.710250). Total procedural blanks were <200 pg of Sr and therefore negligible in all cases (see **Supplementary Material** for a complete description of the methodology).

Throughout the studied interval, the Sr isotopic compositions of the carbonate fractions show significant variations from 0.70809 to 0.70858, whereas carbonate contents vary from 29.5 to 42.5 wt.%, (**Figure 2**). Within river samples, $^{87}\text{Sr}/^{86}\text{Sr}$ varies between 0.70741 and 0.70855 and the carbonate content ranges between 2.0 and 40.3 wt.% (**Supplementary Figure S1**). No clear trend was observed, neither in marine nor in river sediments, between $^{87}\text{Sr}/^{86}\text{Sr}$ ratio and CaCO_3 content (% CaCO_3), nor with

[Sr], (**Supplementary Figure S2**). All data reported in this study are available within the **Supplementary Files**, and online at <http://www.pangaea.de/Projects/PROMESS1/>.

EVIDENCE FOR A DETRITAL CARBONATE INPUT

Seawater $^{87}\text{Sr}/^{86}\text{Sr}$ variations are often used to infer changes in the global strontium geochemical cycles, long-term variations of carbonate rocks erosion, or variations in the marine strontium reservoir through time. Most recently, $^{87}\text{Sr}/^{86}\text{Sr}$ variations in carbonates have been also used to establish continuous high-resolution seawater curves for the last 500 Myr (Howarth and McArthur, 1997; McArthur et al., 2001 among others). In this case-study, sediment was deposited offshore between 120 and 160 ka, a time-span shorter than the residence time of strontium in the ocean, i.e., 4 Myr (Broecker, 1963; Goldberg, 1963; Hodell et al., 1990) and during which the Mediterranean basin remained well connected to the open world ocean via the Strait of Gibraltar (Hernández-Molina et al., 2014; Rohling et al., 2014). Therefore, variations in the carbonate $^{87}\text{Sr}/^{86}\text{Sr}$ ratios analyzed here cannot reflect changes in seawater Sr isotope composition and the observed deviation between our $^{87}\text{Sr}/^{86}\text{Sr}$ results from



the present-day seawater Sr isotope composition has to be related to another process. In other words, if the carbonate components preserved in bulk marine samples were purely biogenic then their isotopic composition should be similar to the seawater: $0.709170 (\pm 4.10^{-6})$, Mokedem et al., 2015; Meknassi et al., 2018 and references therein). In order to further confirm the invariance of the sea-water isotopic composition we performed laser-ablation – MC-ICPMS Sr isotope composition analyses on pure foraminifera and bivalves calcites on the most radiogenic samples (i.e., 120 ka, S.90-21/22) and the least radiogenic samples (i.e., 142 ka, S.110-59/60). Both samples are statistically indistinguishable from modern sea water composition with isotopic values of $0.70919 \pm 0.0001 (2\sigma)$ and $0.70921 \pm 0.0002 (2\sigma)$ respectively. This clearly demonstrates the continuous connectivity with the global ocean and the fact that short timescale variability recorded in our bulk $^{87}\text{Sr}/^{86}\text{Sr}$ cannot be explained by changes in the isotopic composition of the parent fluid (i.e., seawater). How then can such variability be explained?

In the Gulf of Lion context, the most likely mechanism to explain the observed data is through the export of detrital carbonates from the catchment area. Indeed, carbonates represent about 50% of the drainage area (relics of the Tethys Ocean, see Figure 1). Recent studies show that long-term denudation of carbonate rocks within the GoL catchment is significant and slope-dependent (Godard et al., 2016; Thomas et al., 2017), enhancing the transfer of carbonates through rivers' suspended material (i.e., 30 to 60% of the coarse

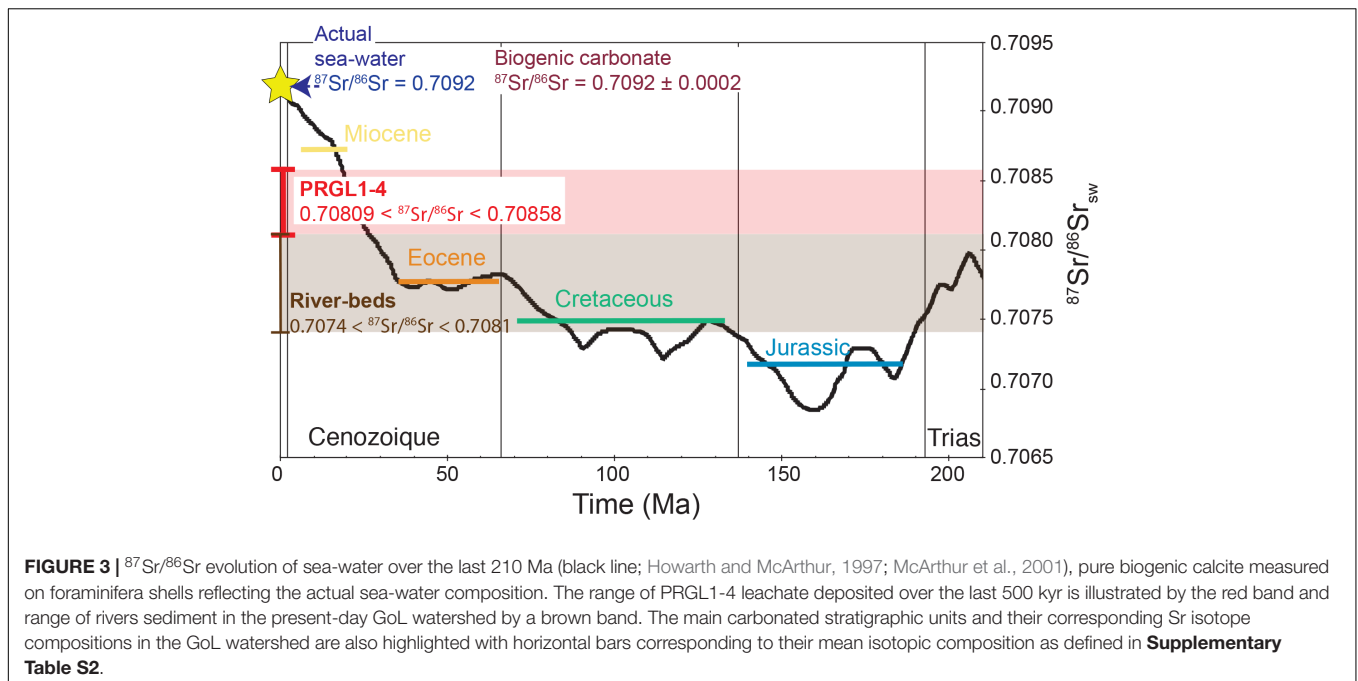
fraction; Pont et al., 2002). Our river samples contain between 2.0 and 40.3 wt.% of carbonates (Supplementary Table S1 and Supplementary Figure S1) which strongly support these conclusions. Moreover, the riverine bedloads are characterized by Sr isotope compositions ranging between 0.70741 and 0.70855 (Supplementary Table S1 and Supplementary Figure S1) which is close to the $^{87}\text{Sr}/^{86}\text{Sr}$ ratios of the different carbonates (i.e., Miocene, Eocene, Cretaceous, Jurassic) present in the catchment area (ranging from 0.70685 to 0.70895; Figure 3 and Supplementary Table S2). Thus, our river ratios likely represent a mixture of the different carbonates end-members (of different age) exposed in the catchment areas (Figure 3). The observed inter-river isotope variability may result from different contributions of mechanical and chemical weathering processes in each watershed and/or reflect the relative proportion of each end-member carbonate unit. In this context, the export of detrital carbonates into marine sediment appears as the best mechanism to explain the observed dataset.

VARIATION THROUGH TIME: COMPOSITION OR PROPORTION?

A clear distinction in $^{87}\text{Sr}/^{86}\text{Sr}$ of leached carbonate fractions is observed between glacial and interglacial intervals (Figure 2) as deduced from the oxygen isotopic curve obtained in planktonic foraminifera (*Globigerina bulloides*) and associated age-model from the same core (Siirro et al., 2009; Pasquier et al., 2017). The two samples from interglacial MIS 5 are characterized by $^{87}\text{Sr}/^{86}\text{Sr}$ ratios of 0.70838 and 0.70858 ($n = 2$) in contrast with less radiogenic composition in glacial sediment (MIS 6) which range from 0.70809 to 0.70832 ($n = 11$, Figure 2). The lowest $^{87}\text{Sr}/^{86}\text{Sr}$ ratios (i.e., 0.70810 ± 0.00001 , $n = 3$) correspond to maximal ice extension and sea-level lowstand.

Contrary to $^{87}\text{Sr}/^{86}\text{Sr}$ ratios, the $\% \text{CaCO}_3$ does not track climatic conditions, and is not modulated by depositional conditions across the termination II (T.II ~130 ka). Instead, $\% \text{CaCO}_3$ shows a gradual rise from approx. 30 wt.% at the onset of MIS 6 (i.e., 160 ka) to 42 wt.% at the end of the penultimate glacial maxima (i.e., 140 ka); then it slightly decreases down to 39 wt.% during the beginning of T.II, before finally rising and reaching a constant value of ~41 wt.% during the MIS 5 (Figure 2). During glacial times, we observe a greater CaCO_3 proportion of what is unexpected considering the low carbonate productivity observed in the western Mediterranean Sea (Hoogakker et al., 2004; Toucanne et al., 2015). Interestingly, this increase in carbonate content is concomitant with a higher detrital flux (Cortina et al., 2013, 2016; Pasquier et al., 2017, 2018); as the rise around 155 ka is synchronous with an extensive melting episode of Alpine ice-caps and local glaciers (Toucanne et al., 2009; Bickel et al., 2015 and reference therein), (Figure 2). This suggests a significant increase of detrital carbonates exported during glacial conditions. This increase might be due to enhanced mechanical glacial and peri-glacial processes such as frost cracking or ablation by glaciers.

The temporal variations in the $^{87}\text{Sr}/^{86}\text{Sr}$ ratios of the carbonate fraction can be explained either by changes in the relative



proportions of biogenic versus detrital carbonates, changes in the isotope composition of detrital carbonates exported into GoL sediment or a combination of both processes.

In the former, the resulting isotope composition of biogenic plus the detrital carbonate mixing processes is typically related to their respective end-member proportion and compositions. Changes in the $^{87}\text{Sr}/^{86}\text{Sr}$ of the carbonate assemblage could be the natural result of an increase or decrease of biogenic carbonate production related to modification of *in situ* primary production through time. In shelf environments, the biogenic carbonate content is generally controlled by the surface water carbonate productivity, the total amount of carbonate being further controlled by terrigenous sediment dilution effects (Cremer et al., 1992; Hoogakker et al., 2004; Toucanne et al., 2015), with low biogenic CaCO_3 production during glacial periods and high biogenic CaCO_3 content during interglacial ones. These changes in carbonate production should affect the *in situ*-biogenic versus detrital carbonate proportions, leading to more (less) radiogenic $^{87}\text{Sr}/^{86}\text{Sr}$ during interglacial (glacial) times, as observed in our dataset. However, downcore fluctuations of CaCO_3 do not follow the $\delta^{18}\text{O}$ records. Instead, we observed the most important increase in $\%\text{CaCO}_3$ during the glacial *sensu stricto* period (Figure 2). This reveals that, at our site, the $\%\text{CaCO}_3$ is not strictly controlled by primary production. Thus, changes in the relative proportion of biogenic carbonates (within the carbonate mixing) cannot on their own explain the observed variation in $^{87}\text{Sr}/^{86}\text{Sr}$.

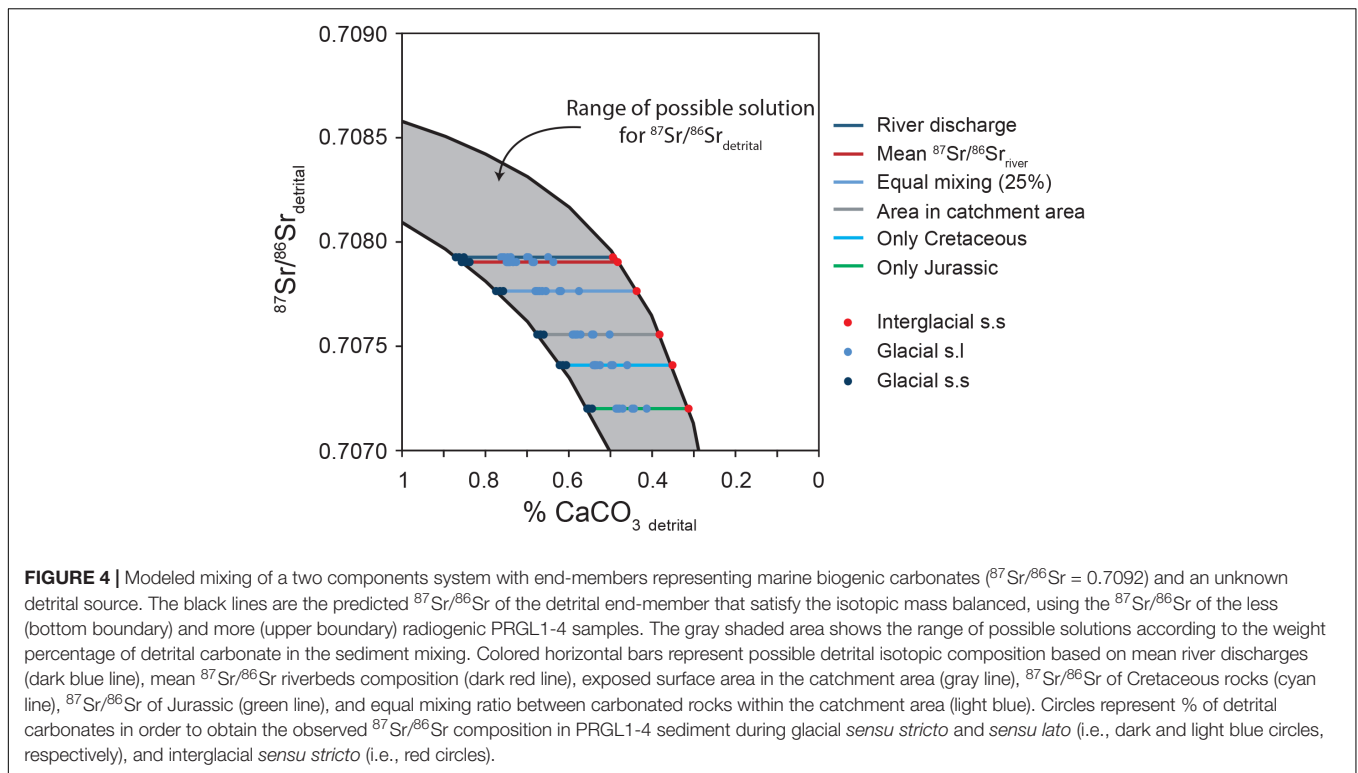
Interestingly, this increase in $\%\text{CaCO}_3$ happened during the period of lower sea level (i.e., glacial maxima), and the less radiogenic values are observed during the lowest sea level conditions (i.e., glacial *sensu stricto*). At that time, due to sea level fall, PRGL1-4 borehole site is closer to shore (~10 km) and is therefore more prone to receive and preserve detrital materials.

During interglacial times, the sea-level rise and the biogenic carbonate production increase leading to a relative decrease in the proportion of detrital carbonates preserved in PRGL1-4 sediment. This is also observed at finer timescales during the entire penultimate glaciation where variability of $^{87}\text{Sr}/^{86}\text{Sr}$ closely mimics planktonic oxygen isotopes and sea-level reconstruction (Figure 2 and Supplementary Figure S3). Therefore, $^{87}\text{Sr}/^{86}\text{Sr}$ fluctuations may be related to variation in the relative proportion of exported carbonates. Mixing calculations between biogenic and detrital carbonates are used to test this hypothesis.

$$(^{87}\text{Sr}/^{86}\text{Sr})_m = [(^{87}\text{Sr}/^{86}\text{Sr})_{sw} \cdot X] - -[(^{87}\text{Sr}/^{86}\text{Sr})_d \cdot (1 - X)]$$

Where $(^{87}\text{Sr}/^{86}\text{Sr})_m$ represent the Sr isotopic composition measured in PRGL1-4 carbonate fractions; $(^{87}\text{Sr}/^{86}\text{Sr})_{sw}$ corresponds to seawater composition and $(^{87}\text{Sr}/^{86}\text{Sr})_d$ refers to the isotopic composition of the detrital carbonates. In this equation X represents the percentage of *in situ* biogenic carbonates, and $(1-X)$ the required proportion of detrital carbonates in order to satisfy the isotopic mass balance.

Calculation results are shown in Figure 4 where the shaded area illustrates the range of possible $^{87}\text{Sr}/^{86}\text{Sr}$ ratios for the detrital assemblage when considering mixing processes between biogenic carbonates (i.e., $^{87}\text{Sr}/^{86}\text{Sr} = 0.70917$) and the minimum and maximum $^{87}\text{Sr}/^{86}\text{Sr}$ ratios recorded in the PRGL1-4 sediment carbonate fraction (bottom and top black line, respectively). In this space, the percentage of detrital carbonate can be predicted in order to satisfy the mass balance for a given detrital Sr isotope composition. We also investigated the impact of detrital assemblage isotope composition on the percentage of detrital carbonate required to fulfill the mixing mass balance equation (Figure 4). This detrital assemblage



isotope composition was calculated by averaging the mean river $^{87}\text{Sr}/^{86}\text{Sr}$ ratios (red line on **Figure 4**), using either a single carbonate value [as highlighted in **Figure 3** (i.e., using the average of the LOWESS numerical fitted curve of McArthur et al. (2012) giving Cretaceous $^{87}\text{Sr}/^{86}\text{Sr} = 0.707406$, Jurassic $^{87}\text{Sr}/^{86}\text{Sr} = 0.707199$, respectively cyan and green line on **Figure 4**)], or resulting from mixing in rivers. In the latter, we used several multi-component mixing, using individual river $^{87}\text{Sr}/^{86}\text{Sr}$ composition and assuming different relative proportions as a function of: (1) river discharges (i.e., $^{87}\text{Sr}/^{86}\text{Sr} = 0.70793$; dark blue), (2) percentage of the exposed carbonated surface in the catchment area (i.e., $^{87}\text{Sr}/^{86}\text{Sr} = 0.70756$; gray line), and (3) assuming an equal mixing between all carbonated rocks in the catchment (i.e., 25% of each carbonate unit exposed; $^{87}\text{Sr}/^{86}\text{Sr} = 0.7077$; light blue), see **Figure 4** and **Supplementary Table S2**. Our results show variation of the relative contribution of detrital carbonates (i.e., % CaCO_3 detrital) over glacial-interglacial cycles (**Figure 4**). Using the different scenarios presented above, we also suggest that a higher proportion of detrital carbonates are preserved during glacial *sensu stricto* (55 and 87%) compared to interglacial (31 to 49%) samples, respectively dark blue and red points in **Figure 4**. We also note that this observation is independent of the isotopic composition of the detrital exported material because whatever the Sr isotope composition of the detrital carbonate fraction is, the required amount of detrital carbonate is always higher during glacial periods. Therefore, $^{87}\text{Sr}/^{86}\text{Sr}$ fluctuation may result from sea-level variability, with sea-level lowstand allowing a better deposition/preservation of detrital carbonates on the GoL's upper slope.

Alternatively, if we assume that detrital carbonates are exported at a constant rate, variation in their strontium isotopic composition could also explain our data. As we showed previously, the GoL catchment area is mainly composed of Mesozoic and Cenozoic marine limestone that are characterized by a large range of $^{87}\text{Sr}/^{86}\text{Sr}$ ratios, and where Miocene carbonate $^{87}\text{Sr}/^{86}\text{Sr} >$ Eocene carbonate $^{87}\text{Sr}/^{86}\text{Sr} >$ Cretaceous carbonate $^{87}\text{Sr}/^{86}\text{Sr} >$ Jurassic carbonate $^{87}\text{Sr}/^{86}\text{Sr}$ (**Figure 3**). Changes in the Sr isotope composition of the PRGL1-4 carbonate fraction can result from variations in the Sr isotope composition of the detrital end-member. Thus, our data could indicate sedimentary mixing of biogenic carbonate with less radiogenic detrital carbonates during glacial times, possibly enhanced by greater inputs of Jurassic or Cretaceous units. Indeed, at the time of minimum $^{87}\text{Sr}/^{86}\text{Sr}$ ratio (glacial) our mixing model predicts less radiogenic detrital carbonates, as indicated by the lower boundary of the shaded area (**Figure 4**). By contrast, more radiogenic detrital carbonates are required during interglacial stages, as indicated by the upper boundary (**Figure 4**). For instance, if we consider that 90% of the carbonates result from detrital export, then the isotopic composition of the detrital source should have evolved from 0.7080 during the glacial *sensu stricto* to 0.7850 during the interglacial *sensu stricto*. Such variability in the detrital $^{87}\text{Sr}/^{86}\text{Sr}$ ratio should be interpreted as the natural response to various mechanical and chemical weathering processes and/or spatial changes in sediment provenance in the catchment.

The directionality of the relation between sea-level change and $^{87}\text{Sr}/^{86}\text{Sr}_{\text{carb}}$ observed here, and the timing between the increase in % CaCO_3 and the Alpine ice cap collapse around 150

ka, argue for the export and preservation of detrital carbonates from the catchment to the GoL shelf. This provides a powerful way to reconstruct past detrital carbonate exports in Source-to-Sink systems.

IMPLICATION FOR FLUXES RECONSTRUCTIONS

Stratigraphic knowledge of the area relies on previous studies based on seismic and PROMESS drilling data in the Gulf of Lion (Rabineau, 2001; Rabineau et al., 2005, 2006; Bassetti et al., 2008; Leroux et al., 2016 among others; see detailed methodology in the Section “Materials and Methods”).

We focused our sediment budget within the so-called Sequence 3, in which enough high-resolution seismic data allowed identification of two sedimentary units, U75 and U80, that, respectively correspond to an interglacial and glacial period (respectively Marine Isotopic Stages 9 and 8), (**Supplementary Figure S4**).

Previously reported sediment budgets obtained for these units were first corrected from porosity to calculate “deposited” terrigenous solid volumes. Considering that %CaCO₃ reflects *in situ* biogenic primary production, “deposited” sediment volumes are corrected to get the “detrital” sediment budget.

In order to investigate the impact of the detrital/biogenic ratio on sediment flux calculation, we then applied the detrital carbonate estimations established for MIS 5 (i.e., 30–50%) and 6 (i.e., 55–85%), assuming it is suitable for MIS 8 and MIS 9, in order to calculate “true” detrital sediment budgets (**Supplementary Figure S5** and **Supplementary Table S3**). Considering the only uncertainties about detrital carbonate content, we observe that “true” detrital sediment estimates within a single unit yield a systematic increase whatever the climatic conditions. Additional detrital sediment represents an addition of 78 to 130 km³.Myr and 128 to 199 km³.Myr, respectively for MIS 9 and MIS 8, corresponding to +20 to +33% and +23 to +36% detrital fluxes, respectively. Our conclusions imply an overall under-estimate of (detrital) sediment supply in Source-to-Sink studies when carbonate content is not examined. This observation is valid for both glacial and interglacial periods, with variable magnitude but is high enough to be considered. Consequently, corrections from *in situ* carbonate production appears critical for future quantitative studies in Source-to-Sink routing systems.

As the terrigenous sediment budget of passive margin basins records variations in the continental relief, triggered by either deformation or climate, it becomes a major challenge to determine sediment accumulation histories in a large number of basins found in various geodynamic contexts (Guillocheau et al., 2012). Usually, Source-to-Sink studies try to relate significant changes in sediment flux to significant changes in terms of climate or deformation, through the geological history of the studied area. But the concept of “significance” is largely dependent on the uncertainties and the time-resolution of the study. And, as previous authors underlined, “the assessment of the associated uncertainties are as important as the accumulation

values themselves.” Indeed, if uncertainties are underestimated, some “apparent” changes in sediment fluxes may not be significant in terms of climate or tectonic changes and may lead to misinterpretations.

However, assessing these uncertainties still remains tricky. To do so, we need to consider and quantify to what extent each factor (i.e., autogenic and allogenic) impacts sediment budget measurement, in addition to the uncertainties strictly related to the method itself (e.g., such as the seismic resolution and borehole age uncertainties). Further work is therefore needed to refine our method, particularly in order to better define the characteristics of the detrital end-member and to therefore more accurately estimate its magnitude.

Moreover, Source-to-Sink sedimentary systems are important settings of carbon cycling, serving as sites of carbon transfer between terrestrial and marine reservoirs, and as the primary locations for organic carbon burial on Earth (Leithold et al., 2016). Whereas the order of magnitude for *in situ* carbonate correction that is measured in this study appears lower than many other uncertainties in some cases, it should be nevertheless interesting to analyze Sr isotopes in Source-to-Sink studies where *in situ* carbonate production (autogenic factor) can “introduce noise, lags and/or completely mask signals of external forcings” (Romans et al., 2016). It may be particularly true for basins with low sediment accumulation rates where small variations of detrital sediment fluxes, including detrital carbonate, can be significant in term of climate or tectonic change. As well as in Source-to-Sink systems (i) exhibiting small sediment supplies, (ii) including few carbonated rocks in the catchment area (i.e., detrital carbonate flux toward the sink), or (iii) bad preservation of detrital sediment (sediment by-pass). Conversely, Sr isotope analyses may also be relevant for areas where biogenic production (and carbonate preservation) is particularly high or where productivity (and associated export in the surrounding basin) is very unstable throughout time, such as a carbonated/mixed platform.

We can now wonder (i) how far this under-estimate can change from one basin to another, and (ii) how far this error could have led to potential previous misinterpretations in term of “true” changes on sediment history and processes at their origin, especially for studies exploring glacial and interglacial cycles.

Overall, as already postulated by Helland-Hansen et al. (2016), this study highlights the usefulness of developing isotopic tools and using combined approaches to better refine our understanding of Source-to-Sink systems and to unravel past source terrains.

IMPLICATIONS FOR DEEP-TIME SR ISOTOPE CHEMOSTRATIGRAPHY

Temporal variations in the marine ⁸⁷Sr/⁸⁶Sr record globally reflect variations in the relative age-weighted fluxes of continental weathering relative to hydrothermal inputs over time, although other sources such as oceanic islands have to be considered (Revillon et al., 2007). Such variations are linked to supercontinent breakup and assembly and sea-level changes

(e.g., Burke et al., 1982; De Paolo and Ingram, 1984; Haq et al., 1988; Veizer, 1989; Veizer et al., 1999; Prokoph et al., 2008) and are used extensively in the Precambrian era. Moreover, in the absence of robust biostratigraphic records and absolute chronological constraints, $^{87}\text{Sr}/^{86}\text{Sr}$ records are a critical tool for inter- and intra-basinal correlations in the Precambrian era. However, the absence of calcifying organisms means that all Precambrian (and many early Paleozoic) $^{87}\text{Sr}/^{86}\text{Sr}$ values come from microdrilled bulk carbonate samples (e.g., Halverson et al., 2007).

Most of the available geologic records of marine sedimentary rocks predominantly preserve sediment deposited on the continental paleo-shelf (Peters and Husson, 2017), where carbonate facies are abundant (Grotzinger, 1990; Grotzinger and Knoll, 1999; Higgins et al., 2009). Consequently, bulk samples can include detrital carbonates.

Still, if detrital carbonates had the ability to reach marine sediment since the emergence of the continent, we can wonder how much carbonated outcrops were likely to be weathered (and preserved over paleo-shelf). Currently, there is no reconstruction of such flux, but a simple geochemical argument can be made: CO_2 has been emitted into the ocean-atmosphere as long as volcanic activity existed, i.e., since the very beginning (Tailor and McLennan, 1985; Jacobsen, 1988; Ying et al., 2011). As CO_2 cannot indefinitely accumulate in the atmosphere, it sank as carbonate mineral and organic matter, at least since the last 3.8 billion years (Schidlowski, 2001). Carbon isotope compilation over the geological record (e.g., Schidlowski, 2001) suggests that approximately 80% of the CO_2 source has been removed as carbonates over that entire time. This is a first-order approximation, and there is many details and caveats, but it reveals that carbonates have been formed over much of Earth's history. Today most carbonates form biologically, while in the Precambrian and early Proterozoic era they formed abiotically (Higgins et al., 2009). In addition, and in contrast with the Phanerozoic era, past ocean chemistry (i.e., high silica content) promoted silicification of most Precambrian and early Proterozoic depositional environments (Siever, 1992; Treguer et al., 1995; Hofmann and Wilson, 2007 among others). Meaning that deep-time carbonates were probably more competent upon uplift and exposure, and therefore were even more likely to be transported to the ocean and “contaminate” the bulk Sr samples.

The above considerations suggest that variations in the resulting $^{87}\text{Sr}/^{86}\text{Sr}$ values within and between basins could potentially also reflect differential detrital carbonate components, impacting both stratigraphic correlations as well as our understanding of tectonic evolution over time. Therefore, it seems critical to consider the potential impact that detrital carbonate may have on these records.

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DATA AVAILABILITY

Publicly available datasets were analyzed in this study. This data can be found here: <https://www.pangaea.de/?q=PROMESS1>.

AUTHOR CONTRIBUTIONS

VP, SR, and MR conceived the work. VP, SR, LM, and SM organized the sampling. VP and SR carried out the geochemical analyses. VP, SR, EL, SM, and MR wrote the manuscript and the **Supplementary Information**. All authors discussed the interpretation of the results and contributed to the manuscript.

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/feart.2019.00164/full#supplementary-material>

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Conflict of Interest Statement: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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