



The CORA 5.2 dataset: global in-situ Temperature and Salinity measurements dataset. Data description and validation.

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1 **Abstract:** We present the Copernicus in-situ ocean dataset of temperature and salinity (version V5.2). The
2 ocean subsurface sampling varied widely from 1950 to 2017, as a result of changes in the instrument
3 technology and development of in-situ observational networks (in particular, tropical moorings, ARGO
4 program). The global ocean temperature data coverage on an annual basis grows thus from 10% in 1950
5 (30% for the North Atlantic basin) to 25% in 2000 (60% for the North Atlantic basin) and reaches a plateau
6 exceeding 80% (95% for the North Atlantic Ocean) after the deployment of the ARGO program. The
7 average depth reached by the profiles also increases from 1950 to 2017. The validation framework is
8 presented, and an objective analysis-based method is developed to assess the quality of the dataset
9 validation process. Analyses of the ocean variability are calculated without taking into account the data
10 quality flags (raw dataset OA), with the near real time quality flags (NRT dataset OA) and with the delayed
11 time mode quality flags (CORA dataset OA). The comparison of the objective analysis variability shows
12 that the near real time dataset managed to detect and to flag most of the large measurement errors, reducing
13 the analysis error bar compared to the raw dataset error bar. It also shows that the ocean variability of the
14 delayed time mode validated dataset is almost exempt from the random error induced variability.

Keywords: Global dataset, In-situ, Temperature and salinity profiles



1. Introduction

5 Estimating the temperature and salinity ocean state is critical for documenting the evolution of the ocean and its role in the present climate. To do so, the scientific community relies on *in situ* measurements at a global scale, gathered by the scientific community and clustered into global datasets.

Among the global datasets widely used in the scientific community, one can cite the world ocean database (Boyer et al, 2013, hereafter WOD) and the EN4 database (Good et al. 2013, www.metoffice.org) distributed by the UK Meteorological Office. Here, we present CORA, a dataset distributed by Copernicus Marine Service (hereafter
10 CMEMS) and produced by Coriolis. CORA differs from these earlier datasets by choices in the construction and the production of the dataset. Indeed, WOD is validated with the highest quality control methods at 102 vertical levels, whereas the EN4 profiles are limited to a maximum of 400 vertical levels and is automatically validated (Ingleby and Huddleston, 2007), whereas CORA retains data at the highest vertical resolution. The choice of
15 reducing the number of levels in the data validation and in the dataset, construction helps to quickly cluster new measurements to the dataset and provides easy to handle datasets. On the other hand, these methodologies results in a loss of measurements potentially available for the scientific community, through the vertical sampling of the profiles or in the data validation. In the construction of CORA, all the measurements available are kept, then an automatic validation is first performed followed by a manual/individual check (Gaillard et al. 2009, Cabanes et al, 2013). This validation framework requires the production of two datasets, a near real time validated dataset,
20 distributing the profiles within days after collection, and a delayed time mode validated dataset, covering in year *n* the historical period up to year *n*-1. This choice, made in the early versions of CORA, has been retained in the latest one that we describe here.

The global ocean heat content increase has been observed on decadal scale timeseries, whether it is in the upper layers of the ocean (Domingues et al, 2008, Ishii and Kimoto, 2009, Levitus et al, 2009), below the thermocline
25 (Von Schuckmann and Le Traon, 2011) or in the abyss (Purkey and Johnson, 2010). Beside the influence of the mapping method and the Baseline climatology (Abraham et al, 2013, Cheng et al, 2015, Boyer et al. 2016), the data validation performed on *in-situ* measurements has a direct influence on the estimation of global ocean indicators such as global heat content, global freshwater content and sea level height (Abraham et al. 2013). As an example, differences in the ocean heat content estimation in the Johnson et al, 2010 analysis compared to the
30 Lyman et al. (2010) analysis has been shown to be a consequence of quality control issues. The particular case of XBT measurements (Levitus et al, 2009, Cheng et al. 2009) influence on the global heat content estimation is well documented. Systematic errors in other instruments types may also introduce systematic biases leading to biases in the global heat content estimation (Lyman et al, 2006, Willis et al, 2011). The validation of a quality control method is thus a critical task to ensure that the dataset flags are accurate enough to flag erroneous measurements
35 without biasing the dataset. The uncertainty surrounding the quality assessment of large oceanographic dataset being a critical topic in the ocean climate studies, we propose here a method of global dataset quality assessment and we apply it to the near real time validated and delayed time mode validated datasets.



We will first list the data sources of the CORA measurements in section 2. A description of the CORA data space and time repartition will be reported on section 3. Then, the quality control procedure will be described in section 4. Lastly, gridded temperature and salinity fields are calculated using an objective mapping that is presented in section 5. The results of the dataset validation and quality assessment are finally discussed on section 6.

2. Data providers

The CORA 5.2 dataset is an incremental version of the previous CORA datasets, covering the period 1950 to now and distributed by CMEMS. Most of the CORA profiles are first collected by the Coriolis data center and validated in near real time mode. Coriolis is a Global Data Assembly Centre (DAC) for the ARGO program (Roemmich et al. 2009). It collects ARGO profiles from the regional Data Assembly Centers (DACs) and distributes them to the community. Coriolis also collects XBTs, CTDs and XCTDs measurements from French and European research programs as well as from the Global Communication System (GTS), Voluntary Ship System (VOS), subtropical mooring networks(TAO/TRITON/RAMA/PIRATA programs from PMEL). A major effort has also been made to include smaller datasets to the Coriolis dataset that are available in delayed time mode, such as the ITP and CTD profiles from the ICES program, Sea Mammals measurements from MEOP (<http://www.meop.net>) and validated surface drifter data. Delayed time mode measurements have also been loaded from the World Ocean Database (WOD13) and the French Service Hydrographique de la Marine (SHOM). It should be noted that in the case of a profile distributed by Coriolis in real time mode and by one of these datasets in delayed time mode, the delayed time mode validated profile replaces the real time mode profile in the CORA database.

Last, recent comparison of the CORA profile positions with the EN4 dataset (metoffice.gov.uk) have shown that some of the profiles distributed in EN4 were not in CORA previous versions. A partnership with the EN4 teams allowed us to detect and to import most of those profiles. 5069864 profiles have been imported in this way, covering the period 1950-2015. However, contrary to the other measurements, the profiles from the EN4 database are not reported with a pressure measurement, but instead with depth and with a maximum number of reported levels in an individual profile set to 400. The issue of the inhomogeneity in the dataset with respect to the vertical sampling, will be discussed.

3. Dataset description

The CORA dataset aims to provide a comprehensive dataset of in-situ temperature and salinity measurements from 1950 to 2017. The oceanic temperature and salinity measuring instruments have however radically changed during the last 70 years.

As a result, the origin and characteristics of data distributed in CORA dataset widely varied in time (Fig : 1) Most of the profiles collected prior to 1965 are mechanical bathythermographs (MBT) measurements or Nansen casts. From 1970 to 1990, the most common profile are from the expendable bathythermographs (XBT), developed



75 during the 1960s and widely used by navies. Most of the XBT profiles collected during this period are T4 type sensor, measuring temperature above -460 meter depth.

The development of the Sippican T-7 instrument with a maximum depth of 1000m slowly increases the number of measurements between 460m and 1000m during the 1980s (see Fig : 2 for the dataset measurements distribution with depth). An instrument capable of measuring conductivity, temperature and pressure (CTD) was developed in the 1960s, allowing an accurate estimation of sea salinity and temperature. The yearly amount of CTD profiles in the CORA dataset then slightly increased reaching a plateau of about 20000 profiles in the early 1990s.

80 During this period, the largest density of profiles is found in the North Atlantic Ocean, with a coverage ratio, calculated on a 3° per 3° grid with a one year time step, increasing from 30% in 1950 to a plateau of 60-70% in the 1970s (Fig: 3). The North Pacific mean sampling rate is lower than 10% before 1965, with the largest portion of the collected profiles located close to the Japanese and North American coasts and along a transect connecting the USA West coast to the Hawaiian archipelago (not presented). It quickly increases from 1965 to 1970 to reach about 50% in the early 1980s with a more homogeneous spatial resolution. Before 1974 in the other ocean basins, most of the collected profiles are found in the coastal zone and along a few ship tracks. The coverage then slightly increases in the western part of the Indian Ocean and in the eastern part of the South Pacific Ocean, increasing the associated basin sampling rate from 10 % in 1965 to 20-25% in 1990. The Austral Ocean sampling rate remains however around 5% during the whole period.

90 During the 1990 decade, the yearly number of XBT profiles strongly decreases while the number of bottles and CTD profiles slightly increases leading to a more homogeneous but barely smaller dataset. The measurements provided are however deeper than in the previous decade, leading to a better coverage below 500m depth (Figure 2). The profile number then exponentially increases since the development of the TAO/RAMA/PIRATA equatorial mooring program throughout the 1990s. During this time, the North Atlantic and the North Pacific Ocean spatial sampling rates decreases, and the global ocean sampling rates barely maintain a plateau at 20%. The ocean sampling rate rapidly increases in the early 2000s thanks to the development of autonomous profilers and the worldwide ARGO program.

95 The global ocean sampling rate reaches 70% before the mid-2000s with a maximum of 85% in the northern Atlantic Ocean. Notice the simultaneous growth of the autonomous profiler measurements on figure 1, and the increasing amount of measurements below 1000m depth on figure 2. In the Austral Ocean, the sampling is sharply increasef from 8 to 40% in 2005-2006, and then grows slowly up to 50% in 2017. This increase in the Austral ocean coverage is a combined consequence of ARGO deployments mostly north of 55° S and the collection from CTD casts mounted on sea mammals, in particular between Kerguelen Island and the Antarctic continent (Turpin et al. 2011).

100 It must be emphasized that a fraction of the profile number sharp increase of the early 2000s results from the data acquisition from high frequency measurement devices such as the ocean drifters, the thermosalinographs (TSGs), both near the ocean surface, or undulating CTDs either towed or untowed (scanfish, seasoar, gliders,...). Indeed, each undulating CTD profile and each independent TSG or drifter measurement is treated as an independent profile in the CORA dataset, which leads to a more homogeneous dataset file structure. This dataset structure is also adopted for the mooring measurements which in some cases are also collecting data at high frequency. This large



number of mooring data induces a large increase of measurements such as 250m and 500m depths, whereas at the surface, the large increase is due to data from TSGs and drifting buoys.

115 4. Data quality control

The measurements collected by the Coriolis data center are distributed to the scientific community with a near real time quality control flag within days from the data reception and with a delayed time mode validation quality control within a year. The Coriolis datacenter validation workflow scheme is given on figure 4.

120 Numerous measurements distributed by Coriolis have however been validated by scientific teams prior to the integration in the Coriolis streamflow. The most important of these datasets are the delayed time mode validated ARGO profiles, the tropical mooring dataset, distributed by PMEL, the sea mammal measurements validated by the MEOP project and the TSG measurements validated by the GO-SUD project. In such cases, the current practice at Coriolis is to retain the flags from the imported database and to run the delayed time mode tests afterwards.

125 4.1 Near real time validation

The near real time dataset validation tests are mostly taken from the ARGO real time quality control tests (Wong et al. 2009). The goal is to distinguish the spurious measurements from the good measurements and to flag them quickly. The test checks are designed to detect well known types of errors. A global range test and a regional range test are performed to detect obvious errors with respect to known ocean variability. The bounds of those two tests are very large with respect to the known ocean variability to ensure that no bad flag would be incorrectly attributed. A spike test and a gradient test are performed to detect measurement spikes in the temperature and salinity fields. The test thresholds are set large enough to lower the number of incorrect spike detections corresponding to a sharp, yet correct, thermocline or halocline. The stuck value test aims to detect temperature or salinity profiles with a constant value within the vertical reported inaccurately.

130 A second step in the near real time quality control is performed daily on the ARGO profilers distributed by Coriolis using an objective mapping detection method (Gaillard et al. 2009). Following the framework developed by Bretherton et al. (1976), the residual of the objective analysis depends on the covariance from data point to data point. This second check step aims thus to detect measurements departing from other data in its vicinity. The correlation scale in the objective analysis varies with depth and latitude. Spurious detections can however occur when profiles located on both sides of a frontal zone are within a correlation radius. Therefore, detected profiles are visually checked by a PI to distinguish erroneous measurements from possible measurements.

135 Lastly, a quality control based on altimetry comparisons is also performed on a quarterly basis to improve the real time validated dataset (Guinehut et al. 2009). A PI investigation is also performed on profiles flagged as suspicious by comparison with altimetric sea level.

140 4.2 Delayed time mode validation tests

The delayed time mode validation is performed on a yearly basis. This validation framework is based on tests more stringent than the near real time validation process, which requires a systematic visual control by an oceanographer. The controlled profiles are those who have not been controlled in the previous version of CORA. Therefore, most



150 of the controlled profiles for a given version of CORA are the profiles measured during the previous year, but not controlled for the earlier version. The profiles for which the measurements have been updated or adjusted since the latest version are however controlled. Last, some datasets covering the historical period may have been incorporated in the Coriolis dataset, which are then controlled in delayed time mode in CORA.

The delayed time mode validation process is schematized on Figure 4. The profiles to be validated are first checked by the CORA tests. The checks raise an alert flag on suspicious profiles, which are then visually checked. For 155 CORA, the validation checks are applied until all the tests are successful. If a single check test fails, the profile is put aside for visual check and the following tests are not applied. The profiles undetected by the CORA tests, and thus not visually controlled, are assessed by a second set of tests developed by CLS. The suspicious profiles are also visually controlled and manually flagged. Last, all the tested measurements are gathered in the CORA database with the updated flags.

160 A first quality check aims to complement the real time QC procedure with redundant tests with sharper threshold than NRT. The following tests are concerned.

Data-file consistency test

This test checks the obviously out of range position ($|\text{Lat}| > 90$ and $|\text{Lon}| > 180$ and out of range immersion ($\text{PRES} > 12000$ decibar and $\text{Depth} > 12000$ m or $\text{PRES} < -2.5$ decibar and $\text{DEPTH} < -2.5$ m). These tests are redundant 165 with the NRT checks and are designed to avoid any writing error in the CORA file. The few detections are visually checked.

Depth check, Stability Check, Vertical check

The depth check, stability check and vertical check have initially been developed by the UK Met-Office for the EN4 dataset validation. They have been added to the CORA validation framework after a collaborative comparison 170 of the two dataset validation methods with the UK Met-Office team. This study has shown that most of the profiles flagged in EN4 and not in CORA were detected by these three tests and that applying a visual control to the profiles detected in this way results in more accurate flags. The tests have been described in Ingleby and Huddleston, 2007. The stability test detects density inversions for profiles where both temperature and salinity are available. The density inversions with $0 > d\rho > -0.03$ kg.m³ are dismissed. Both temperature and salinity are visualized for 175 profiles with larger density inversion. Experience has shown however that most of the density inversions detected in this way are caused by small spikes in the salinity measurements, probably a consequence of anomalies in the conductivity measurement or alignment with temperature when estimating salinity. The spike test is designed to detect the temperature and salinity spikes and steps. It runs with a threshold of temperature and salinity variability varying from 5°C in surface to 1.5° C below 600 meter depth for temperature and from 1°PSU at surface and 180 0.2°PSU below 300 meter depth for salinity. These tests differ from the real time QC test since the trigger points are lower. They however sometimes create 'false positive' detection either by detecting the wrong point on a spurious profile or by detecting a correct measurement. A systematic PI visual flag selection is then performed on each of the detected profiles.

Level disorder and duplicated levels

185 The profiles with a non-monotonous PRES or DEPTH vector are detected and the PRES or DEPTH vector are flagged in order to be monotonous. This test has been requested by the CORA end users, the oceanographic



reanalysis community, to have a user friendly dataset to work with. Most of the detected profiles are indeed measurements with a very slow sinking speed near the surface, giving pressure vector inversion when exposed to the sea surface swell. Most of the detections are thus confined to the surface layer. Exceptions may however occur
190 in the case of black sea ARGO floats for which a recurrent problem of slow sinking speed is found at sub surface due to the low salinity level of the black sea. Last, “hedgehog” type profiles are detected by this test, which are often caused by transmission mistakes on Argo floats.

Global range

The global range test aims to detect obvious measurement mistakes. The Temperature measurements under -3°C
195 or over 43°C and the salinity measurements under 0 PSU or over 46 PSU are detected. This test has a very low detection rate, but it still detects some erroneous profiles each year. Most of them are profiles with a non-classical shape so that they avoid detection by redundant tests (Minmax test or climatological test). A recent example was an ARGO float grounded near Mogadicio, Somalia, measuring a temperature exceeding 43°C whereas the corresponding pressure was just above 0 decibar, so that the measurement avoided the other NRT and delayed time
200 mode tests confined to depths between 0 and 2000 m.

The following step of the CORA data validation is performed in the Coriolis datacenter to detect profiles diverging from the known ocean variability. Each temperature and salinity profile is compared with the minimum and maximum measured value reference profiles. Those profiles originate from reference fields on a gridded mesh with 1 degree resolution horizontal hexagonal cells of 20m thickness. The reference fields are the maximum and
205 minimum measured values on a set of 1.2 million ARGO profiles, vertically interpolated from the surface to 2000 m depth. The field coverage is reinforced, especially in the inner seas and in the Austral Ocean, badly covered by the ARGO network, by CTDs from the World Ocean Database and sea mammals measurements from the MEOP database. The CORA 5.2 measurements are compared to the minimum and maximum reference values of the corresponding cell and the upper and lower adjacent cells. The profiles containing measurements exceeding the
210 reference values are checked by an oceanographer. The minmax method is relaxed on the continental shelf since the minmax sampling is insufficient in the continental shelf zones. The temperature and salinity profiles measured on the continental shelf (depth < 1800m) are compared to a climatology field (ARIVO, Gaillard et al) to which is added or subtracted 10 times the climatological standard deviation field. Due to the lack of accuracy of the global climatologies near the coasts and on some shelves, the profiles lying above the 250 m isobath are not tested with
215 this method. See Gourrion et al. 2017 for further discussion on the minmax field.

A third validation is performed with the ISAS objective analysis tool, following the method developed by Gaillard et al, (2009). This validation test is similar to the validation performed by the Coriolis datacenter in near real time with the ARGO floats. The scope of the test is however extended to other profiles (XBT, CTD, etc...) and to ARGO profiles which have been updated in the Coriolis database too late to be part of the near real time validation.

220 A last set of delayed mode validation tests has been developed by the CLS research and development team and aims to complement the CNRS validation tests. These tests provide a sharper expertise on bias detection, spike detection and ocean variability in the continental shelf zones. These tests also aim to complement the Coriolis real time quality check tests for measurements directly included in the delayed mode dataset. The CLS tests are divided in two categories. A first set of tests are applied to detect small density inversions in the measurement profiles and
225 non-monotonous pressure/depth vector. These tests are reliable so the detected suspicious levels are automatically



flagged. A second set of tests is applied to detect smaller errors. These tests aim to detect unlikely extremes in temperature and salinity by comparing measurements to regional upper and lower bounds and World Ocean Atlas 2013 climatology. Tests are also applied on vertical density inversions, spikes and offsets with respect to the climatology.

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5. CORA 5.2 quality control results

5.1 Quality assessment - Influence of the QC flags on the estimated ocean variability

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The relevance of ocean climate studies strongly depends on the accuracy of ocean measurements. Systematic data errors might thus result in biasing the estimation of ocean state indicators such as the global ocean heat content, the global ocean freshwater content or the global mean steric height (Levitus et al. 2009). Furthermore, random measurement and data error may lead in overestimating ocean variability. Therefore, indirectly, one can assess the reliability of the global dataset by estimating the influence of the quality control on global metrics such as the ocean mean temperature and salinity and the associated variability.

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Two mappings of ocean temperature and salinity based on the CORA dataset measurements are calculated: a raw estimation (GOHCraw) which considers every measurement without taking the ocean flags and a flagged estimation (GOHCflg) which only consider the good and probably good QCs.

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Interpolated fields are calculated following the method presented by Forget and Wunch, 2007 that has the advantage of not biasing mean fields and not relying on specifying them. The global ocean is divided in 1° per 1° grid cells with 10 m vertical layers from the surface to 700 m depth. A first estimation of the mean parameter for a given month is given by calculating the mean of the temperature or the salinity data measured in a given cell. The variance field is estimated by taking the variance of the measurements located in a given cell, if the number of available measurements is greater than 4.

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A spatial weighting function is defined:

255

$$G(i, j) = e^{-\frac{(l_p - l(i, j))^2}{r_l^2} - \frac{(l_p - l(i, j))^2}{r_L^2}} \quad [1]$$

260

With r_l and r_L latitude and longitude decorrelation scales, both taken equal to 5°, and l_p and L_p the latitude and longitude of a grid point.

The combined mean is then:



$$\bar{T}(i, j) = \sum_p \frac{G_p(i, j) T_p n_p}{N(i, j)} \quad [2]$$

265 with:

$$N(i, j) = \sum_p G_p(i, j) n_p \quad [3]$$

The combined variance is estimated with a similar operator.

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$$\text{var}(T)(i, j) = \sum_p \frac{G_p(i, j) \text{var}(T_p) n_p}{N(i, j)} \quad [4]$$

With n_p the number of measurements available in the summed grid point and $N(i, j)$ the total number of measurements involved in the calculation of a grid point value.

275 The values of r_l and r_L are set to 5° longitude and latitude in order to include enough grid points with data in this averaging. To reduce the calculation time of the analysis, each $N(i, j)$ calculation is performed on a 20 per 20 grid points windows.

280 The objective analysis is performed at three steps of the global dataset. A first analysis is performed on a raw dataset, considering all available profile measurements. All the QC flags are considered good. A second analysis is performed on the same data profiles considering the QC available on NRT mode. A third one is performed on the same profiles considering the QC available on delayed time mode.

285 The ocean data coverage is sometimes insufficient to perform the monthly objective analysis on the whole ocean. Figure 5 shows an estimation of the ocean layer covered by the objective analysis as a percentage of the ocean layer surface between 60°N and 60°S . It shows that the ocean coverage is higher for temperature than for salinity objective analysis. The upper layers coverage are very close. It varies from 95% in 2005 to over 98% after 2012. The 1475-1525m depth layer departs from the others since it has a global coverage lower from the others, starting from 65% in January 2005. It converges to over 98% after 2014. A monthly variability is observed in the ARGO development period (2005-2010). It is probably caused by the ARGO profiles drift in the southern zones. This behavior lasts up to 2012 in the deeper layer.

295 Fig: 6 shows the percentage of good QCs in the NRT and CORA datasets compared to the RAW dataset. It shows that the proportion of good flags yearly tendencies are almost the same at all depth. Moreover, in any cases, the CORA and NRT differs from less than 0.5%. The proportion of good temperature flags varies from a minimum of 92% in 2006 to a plateau of about 98% after 2013. The 975-1025 m depth and 1475-1525 m depth layers depart from the others with 1 to 2 % lower rate between 2005 and 2013. Punctual decrease of good temperature flag rates are observed in late 2007, late 2012, late 2014 and in the beginning of 2016 for the surface and subsurface layers. These spikes are caused by a sharp increase in the number of profiles distributed from a tropical mooring from the RAMA network. These profiles are indeed first distributed in the Coriolis dataset as TESAC profiles transmitted from the GTSP. The profiles corresponding to tropical moorings are usually later replaced by the corresponding measurements transmitted by PMEL and the TESAC profiles are erased from the database. In this particular case,



the TESAC profiles had been retained and flagged as bad profiles instead. The yearly number of profiles in the RAW dataset is thus strongly increased but the corresponding number for the NRT and CORA dataset is not. The good salinity flag rate tendency is opposite to the good temperature flag rate, with a maximum of over 98% before 305 2010, then a decrease to a level of about 94% with a high interannual variability after 2011.

The mean 0-50m, 75-125 m, 275-325 m, 475-525 m, 975-1025 m and 1475-1525 m depth salinity standard deviations analyzed by the method (eq. 4) from 2005 to 2016 are shown on figure 7 and 8. The mean salinity 310 standard deviation is averaged between 60°N and 60°S for each dataset analysis. The comparison of the raw dataset analysis with the NRT analysis and the CORA analysis shows the gain in dataset quality resulting from the QC performed. In the raw dataset analysis, numerous random mistakes result in a high average salinity standard deviation. The raw dataset standard deviation is however lower in the early period at almost all levels, despite a rather high global level for the 475-525 m depth layer and a variability spike in late 2006 in the bottom layer. This 315 lower variability level is probably a consequence of the ARGO program development from 2005 to 2008, the low coverage in the southern oceans preventing the emergence of high level values. During this period, a large seasonal variability is present in the upper layers, varying from an order 0.2 PSU during winter to 0.4 PSU during summer, in the surface layer. The peaks in the ocean variability are thus correlated with peaks of ocean coverage (see figure 5). The objective analyses also have a higher proportion of ship-borne measurements, CTDs for instance, 320 essentially made during summer, compared to the autonomous measurements these years. We can thus assume that the lower number of profiles during winter does not allow to sample correctly the subsurface ocean fronts, leading to an underestimated winter time ocean variability. The increase in ARGO float data from 2005 to 2008 slowly decreased this bias in the ocean variability estimation. The raw dataset surface salinity standard deviation increases during the 2010-2016 period at all depth levels, with an order 0.6 PSU amplitude and spikes up to 1.2 325 PSU in 2010 in the surface layer, and spikes varying from 0.9 to 1.2 PSU in the other layers.

The NRT analysis is very close to the CORA analysis before 2008. This behavior is a consequence of the low number of measurements corresponding to this period collected or updated in the database after the validation of the last version of the CORA dataset. The NRT and the CORA datasets flags are indeed equal except if an updated 330 version of a profile is loaded in the database or if a new profile is loaded in the Coriolis database. On the other hand, large discrepancies between the NRT and the CORA datasets are recorded between 2009 and 2012 and between 2013 and 2016. Another fraction of the discrepancy between the NRT and the CORA error bars are caused by non-ARGO profiles updated in the Coriolis database without delayed time mode assessment. Most of these measurements are sea mammal profiles in the northern Pacific Ocean or mooring data, imported from the GTSP 335 TESAC messages with biased salinity sensors. Sea gliders with an order 5 to 10 PSU bias in salinity were also documented. Moreover, despite a lower number of profiles flagged, a few CORA flagged ARGO profiles have biases large enough to strongly increase the analyzed ocean variability. Some of the spikes in the NRT ocean variability documented in the upper layers, the late 2007 to begin of 2008 spike for instance, are observed in the surface layers since they are caused by biased instruments operating at surface to subsurface layers. Some other 340 spikes in the ocean variability, the 2009-2011 spike for instance, are caused by biased ARGO measurements and thus impact the ocean variability from the surface to 2000 m depth.



A striking feature is the corresponding spike visible in the NRT analysis and in the raw dataset analysis in late 2010, which suggest that major data errors have not been flagged in the dataset during the NRT validation. Further
345 exploration of this anomaly has shown that a fraction of the larger error bar in the NRT analysis is caused by an issue in the update of delayed time mode processed ARGO profiles. In a few cases when salinity measurements present large drifts, the ARGO PIs can decide that the salinity drift is too high to be adjusted. In these cases, the PI provides to the global DAC a delayed time version of the profiles with an adjusted temperature field, but with a practical salinity field filled with fillvalues and a salinity QC field filled with “4” values (bad measurement status). In some cases, the Coriolis data center had updated the profiles by getting the temperature adjusted field but without creating a salinity adjusted field. The available salinity field and QC in the Coriolis datacenter is therefore the original salinity field which might not have been flagged at “4”. In this study, a handful of these profiles, often associated with large salinity drifts (for instance salinity values on the order 20 PSU in the Indian Ocean) have produced large error bars in the NRT analysis fields. This issue will be soon resolved in the Coriolis
350 database.
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The CORA analysis salinity standard deviation varies slowly in time, with an order 0.15 PSU in the surface layer, an order 0.1 PSU in the 75m depth – 125m depth layer and an order 0.08 PSU in the 275-325m depth layer and below 0.05 PSU in the deeper layers. This behavior is a consequence of the delayed time mode validation process which reduces drastically the number of random mistakes in the dataset. This variability is probably a function of
360 the local data resolution, the oceanic variability and measurement errors. The slow variability of the CORA salinity standard deviation and its reasonable range suggests that remaining errors in the dataset will not have a large importance. Thus this product is likely to present a low error amplitude.

Figure 9 and 10 show time series of the mean temperature standard deviation of the CORA, NRT and RAW analysis. As anticipated, the mean temperature standard deviation time series is noisy and rather high in the RAW
365 dataset case. The mean amplitude varies almost linearly between 1.2 °C in the 0-50 m depth layer to 0.4°C in the 1475-1525 m depth layer, except for the 975-1025 m depth layer with a 1.2 °C spike. A striking feature is the decreasing mean temperature standard deviation amplitude in time for the RAW analysis. The reason of this behavior is rather unclear. One shall assume that the overall quality of the oceanographic in-situ temperature measurement improves because of improvements in the temperature sensor. On the other hand, it might also be
370 the decrease in the number of deployed XBTs in the 2010s that reduces the number of random errors in the dataset, since the XBT instruments have proved to produce erroneous measurements when they are not handled properly.

The NRT analysis and CORA analysis time series are rather close in all the analyzed layers, except for a 0.8 – 1°C spike in 2014-2015, detected in all the layer but the 1475-1525 m depth layer, and enhanced in the 475-525 m
375 depth layer and in the 975-1025m depth layer. Most of the temperature profiles flagged in the CORA delayed time mode validation process are indeed XBTs because the XBT instruments are more likely to fail (spikes or bias caused by a stretching of the XBT wire or a contact between the XBT wire and the ship hull), or bad estimation of the measurement depth. These measurements error, even if numerous (about 2000 profiles flagged between 2005 and 2016) often have a too small amplitude to induce a large bias between the NRT analysis and the CORA analysis. Most of the XBTs corrected during this period are moreover T-4 and Deep blue models. These models are usually not measuring in-situ temperature below 460 m depth and 760 m depth respectively. The CORA and
380 NRT variability amplitudes are thus almost the same in the deeper layers.



385 The CORA analysis variability has a mean amplitude of 0.85°C with a clear seasonal cycle of about 0.3°C. The
CORA analysis mean variability amplitude averages 0.95 °C in the 75-125 m depth layer, with a monthly
variability uncorrelated with the seasonal cycle. The seasonal cycle amplitude is null in the other layers, with a
CORA analysis mean variability amplitude of 0.6°C in the 275-325 m depth layer, 0.4°C in the 475-525 m depth
layer, 0.2°C in the 975-1025 m depth layer and 0.1°C in the 1475-1525 m depth layer. The higher frequency
variability decreases with depth and is almost null in the deeper layers as seen on figure 9 and 10. The noisy shape
of this high frequency variability is probably a result of ocean monthly variability and the changing locations of
390 the ocean profiles.

The 2014-2015 spike in the ocean variability, detected in all the layers except for the deeper one in the NRT
analysis, are caused by many XBT profiles. Most of those profiles are deployed in the Indian Ocean across a
transect linking the Gulf of Aden to Perth, Australia, corresponding to measurements performed by the ship of
opportunity program (Goni et al. 2009). The profiles have been extracted from the World Ocean Database and
395 have thus not been validated with the Coriolis real time validation framework. Many biases and spikes, probably
due to issues with the probes or with poor insolation of the XBT wire problems, have been flagged in delayed time
mode. The largest part of the upper layers spikes in the NRT and RAW analysis is a result of these erroneous
measurements. In addition to the usual issues with the XBT measurements, the profiles sometimes indicated
negative values at subsurface depth or temperature of 36.269°C at depth located above the maximum functioning
400 depth of the XBT (460 m depth for T-4 and T-6, 760 m depth for Deep Blue). These unrealistic values have not
been flagged after the extraction from the WOD dataset, resulting in exponential growth of the local amplitude of
temperature standard deviation in the RAW and NRT analysis in the 475-525 m depth and 975-1025 m depth
layers.

A closer look at the vertical profiles of the temperature and salinity mean variability (Figure 9) shows that the
405 CORA analysis temperature and salinity variability is far smaller than the RAW analysis and the NRT analysis
estimation. The depth variability of the temperature and salinity mean variability is moreover closer to the expected
oceanic variability, with a maximum ocean variability at the surface or close at sub surface with decreasing
variability below the ocean mixed layer depth. We don't think that the CORA dataset has been over flagged, which
could result in artificially decreasing the global ocean variability. Indeed, one should keep in mind that most of the
410 flags applied on these profiles are manually applied by physical oceanographers after receiving a detection alert.
The robustness of the CORA analysis variability time series, either for temperature and salinity, shows that this
hypothesis might be considered unlikely. The study by Forget and Wunsch, 2006 has moreover shown that the
analysis results were in good agreement with an 0.8 degree global model reanalysis estimate (Hill et al. 2007).

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6. Conclusion

The CORA dataset is an extensive dataset of temperature and salinity measurements. Efforts have been made to provide to the scientific community an information as close as possible from the physical measurement and to perform a strict quality control on all profiles. The CORA dataset indeed stands out from the EN4 dataset since the delayed time mode validation is based on automatic detections and systematic PI decision, reducing the number of mistaken bad flags. In addition to that, the profiles are not subsampled and the time series (TSGs and drifters) are distributed. It also stands out from the WOD dataset since all measurements within a profile are validated in delayed time mode, reducing the number of mistaken measurements.

Moreover, this study develops an innovative method to assess the overall quality of a dataset. This method shows the improvements of the dataset quality flags thanks to Coriolis real time QC and the CORA delayed time mode QC frameworks. This method however lacks a comparison with an analysis based on other datasets to ensure that the CORA validation framework is not constraining its description of the ocean variability by over flagging good measurements. This discussion shall be further pursued. Last, this method is based on the mapping of the Ocean variability. It is thus implicit that the ocean sampling is homogeneous and sufficient to perform a monthly analysis. These conditions are met at a global scale and for the ocean measurements from surface to 2000 m depth since the full deployment of the ARGO network. The ocean data coverage is however insufficient to have a global coverage before 2005 (see Fig.3 for the ocean basin data coverage ratio), especially at depth larger than 1000 m between 1990 and 2005 and at depth larger than 500 m before 1990, as seen on Fig.2. The method will thus have to be adapted to the ocean data coverage to provide a synoptic view of the dataset quality. Last, the CORA ocean dataset variability has proven to be rather homogeneous from 2005 to 2016. From this viewpoint, comparison of the ocean variability described by a dataset with a target variability may be an encouraging method for the detection of measurement errors. Further studies shall be conducted to develop this method.

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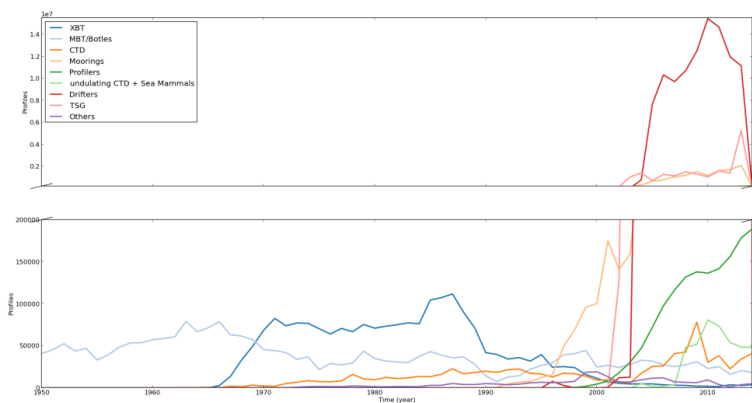
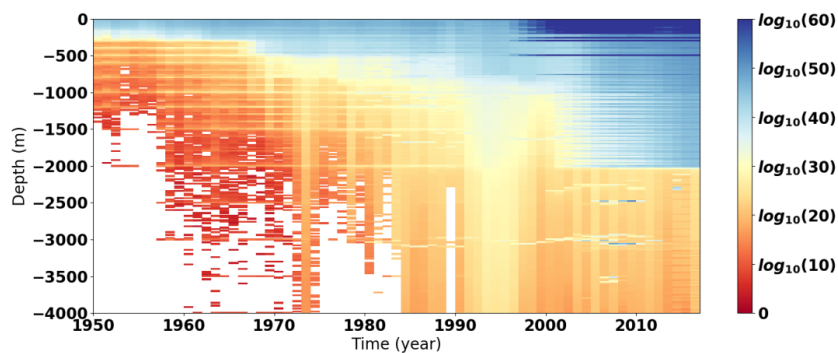


Figure 1: Yearly number of distributed profiles, sorted by instrument types.



535 Figure 2: Yearly number of measurements as a function of depth.

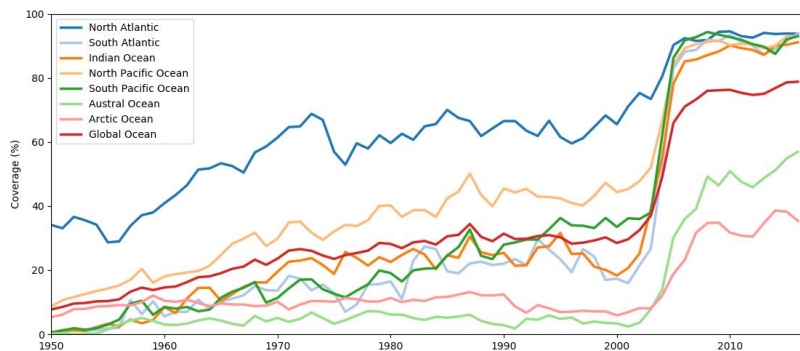
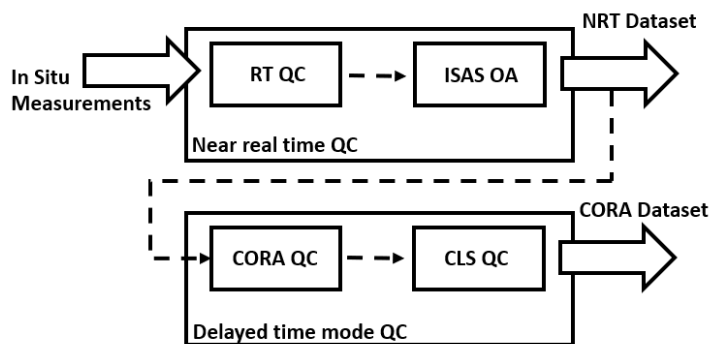
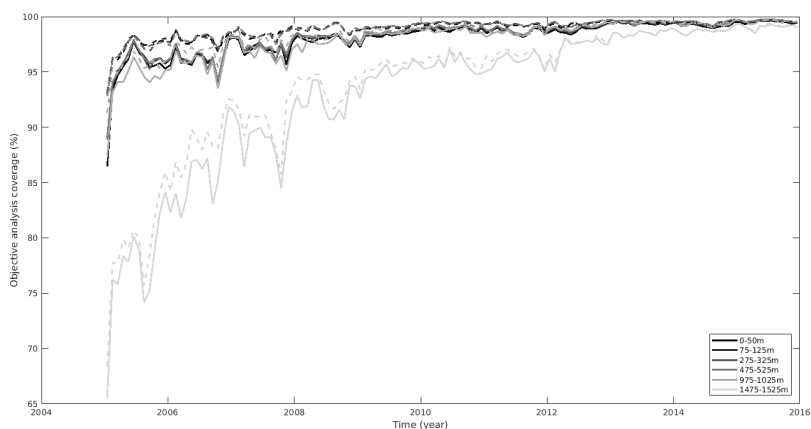


Figure 3: Yearly filling ratio of 3° latitude per 3° longitude gridded field of ocean basins



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Figure 4: Coriolis database validation process.



545 **Figure 5: Coverage of the Temperature and Salinity objective analysis for temperature (dashed line) and salinity (hard line) objective analysis.**

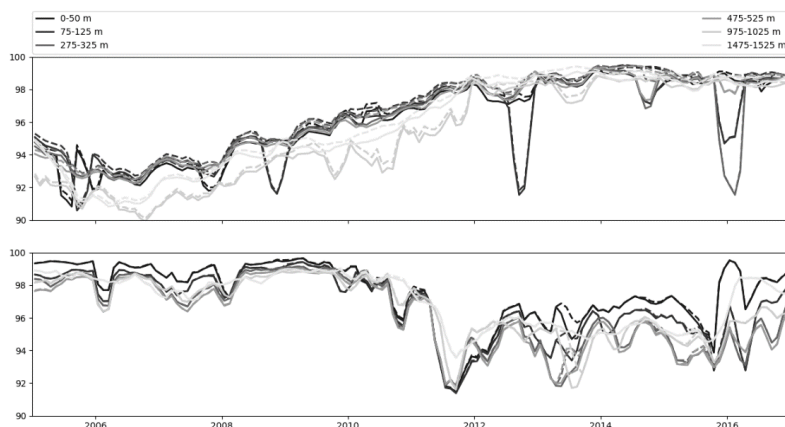
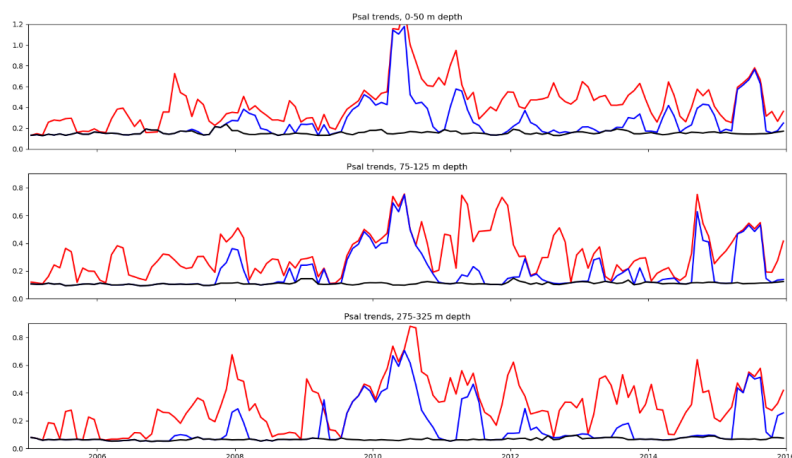
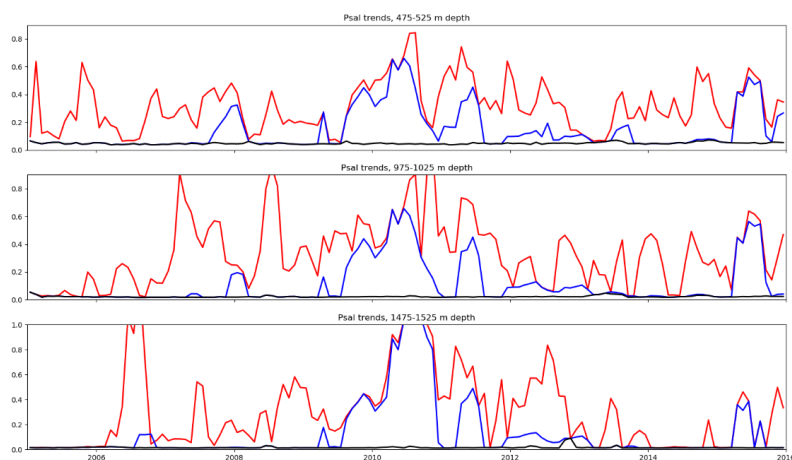


Figure 6: Percentage of good flags (flags 1 and 2) in the analyzed layers for the NRT dataset (hard line) and for the CORA dataset (dashed line). Upper panel for temperature, lower panel for salinity



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Figure 7: Mean salinity standard deviation in the 0-50m layer (top), 75-125 m depth layer (mid.) and 275-325 m depth layer (bot.). The raw dataset (red), NRT dataset (blue) CORA dataset (black) are represented.



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Figure 8: Mean salinity standard deviation in the 475-525 m depth layer (top), 975-1025 m depth layer (mid.) and 1475-1525 m depth layer (bot.) The raw dataset (red), NRT dataset (blue) CORA dataset (black) are represented.

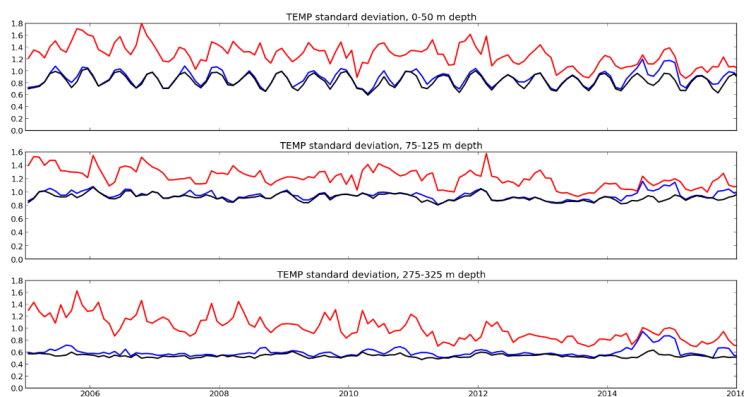


Figure 9: Mean temperature standard deviation in the 0-50m layer (top), 75-125 m depth layer (mid.) and 275-325 m depth layer (bot.). The raw dataset (red), NRT dataset (blue) CORA dataset (black) are represented.

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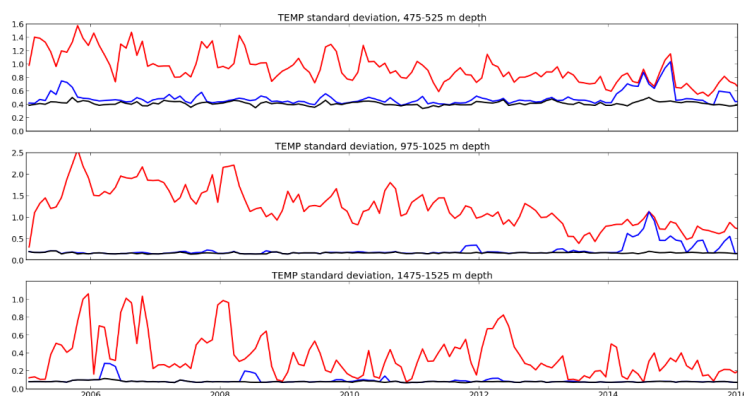


Figure 10: Mean temperature standard deviation in the 475-525 m depth layer (top), 975-1025 m depth layer (mid.) and 1475-1525 m depth layer (bot.). The raw dataset (red), NRT dataset (blue) CORA dataset (black) are represented.

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