Meroplankton community structure across oceanographic fronts along the South Brazil Shelf

Costa Brandao Manoela ^{1, 2, *}, Garcia Carlos A.E ^{3, 4}, Freire Andrea S. ¹

¹ Departamento de Ecologia e Zoologia, Universidade Federal de Santa Catarina, Campus Universitário, 88040-970 Florianópolis, Brazil

² Institut Français de Recherche pour l'Exploitation de la Mer, Centre de Bretagne, 1625 Route de Sainte-Anne, 29280 Plouzané, France

³ Instituto de Oceanografia, Universidade Federal do Rio Grande, Km 8 Avenida Itália, 96201-900 Rio Grande, Brazil

⁴ Programa de Pós-graduação em Oceanografia, Universidade Federal de Santa Catarina, Campus Universitário, 88040-970 Florianópolis, Brazil

* Corresponding author : Manoela C. Brandao, email address : manoelacb@yahoo.com.br

Abstract :

The influence of oceanographic fronts on the abundance and community composition of invertebrate larvae, mostly of benthic species, along nearly 2000 km of the southwestern Atlantic shelf (21–34°S) was investigated. Meroplankton was sampled through vertical hauls at 89 stations, distributed along 14 cross-shelf transects, during late spring 2010 and early summer 2011. Salinity and temperature were registered with a CTD/rosette system, which provided seawater for chlorophyll-a and nutrient concentrations estimations. Vertical profiles of temperature, salinity, chlorophyll-a and nutrients were used as proxies of the fronts. In addition, high-resolution thermosalinograph data were used to detect surface frontal features. Meroplankton abundance peaks were found at several fronts intersected by the ship, including upwelling zones, estuarine and plume fronts, a shelf-break front, and two cyclonic eddies. Furthermore, meroplankton abundance was also relatively higher at small-scale thermal and/or saline surface fronts observed along the shelf. Such increases in meroplankton abundance are likely to be ascribed to high nutrient input and primary production. Distinct taxa of invertebrate larvae occurred at different types of fronts, besides the coastal realm, which was virtually dominated by decapod, cirripede and bivalve larvae. Small-scale shelf fronts presented high abundances of decapod and gastropod larvae, for instance, while larvae of polychaetes were the most frequent in the estuarine front of Patos Lagoon section.

Highlights

► The sampling area covers 13° of latitudinal gradient and up to 400 km offshore. ► Large-scale maps of abundance of the main meroplanktonic groups are presented. ► Physical, chemical and biological variables are analyzed together. ► Influence of distinct types of fronts on meroplankton abundance is discussed.

Keywords : Frontal zones, Invertebrate larvae, Water mass, Large-scale variability, Southwestern Atlantic

44 **1. Introduction**

45

In neritic pelagic ecosystems, meroplanktonic larvae of benthic invertebrates comprise
a large portion of the zooplanktonic community (Shanks et al., 2002; Hidalgo et al., 2014).
Despite the limited ability to move, these larvae are capable of controlling their vertical
position in the water column (Morgan, 2014). This behavior in conjunction with physical
processes will determine whether larvae are exported, retained, or concentrated in specific
locations (Cowen et al., 2000).

52 As a transitional area between the coastal zone and the ocean, continental shelves 53 include water masses of different physical/chemical characteristics and, consequently, a series 54 of frontal zones (Munk et al., 2003). Oceanographic fronts are regions of larger-than-average 55 horizontal gradients of water properties such as temperature, salinity, and density (Joyce, 56 1983). These confluences of oceanographic processes of contrasting features are usually 57 characterized by high biological productivity (Le Fèvre, 1987; Acha et al., 2015), due to 58 nutrient entrainment, primary/secondary production or aggregation (e.g. Munk et al., 2003; 59 Acha et al., 2015; Hidalgo et al., 2014). In addition, frontal zones are generally assumed to 60 maximize diversity due to the convergence of species inhabiting different water masses (e.g. Acha et al., 2004). 61

Tropical and subtropical oceanic regions usually have a permanent thermocline, which prevents the mixing of surface and nutrient-rich deep waters, thus presenting low productivity and planktonic biomass (Nybakken, 1997). Accordingly, the continental margin of the South Brazil Shelf (SBS) (21–34°S) is predominantly oligotrophic, depicted by the strong influence of the Tropical Water (TW) driven by the Brazil Current (Brandini, 2006). However, a series of quasi-permanent or episodic oceanographic processes disrupt the vertical stability of the water column significantly increasing the availability of nutrients in the upper layers (Acha et al., 2004; Gaeta and Brandini, 2006). While others, such as the Plata Plume and the Patos
Lagoon estuarine front, do not disrupt stratification, but inject nutrients in the area (Acha et
al., 2004).

72 Among the processes that mainly increase the biological productivity in the SBS, it is 73 worth mentioning the wind-driven coastal and shelf-break upwelling of the South Atlantic 74 Central Water (SACW), as seen, for instance, in the inner shelf of Cape São Tomé (21°S), 75 Cape Frio (23°S), and Cape Santa Marta Grande (28°S) (Castro and Miranda, 1998; Möller et al., 2008; Campos et al., 2013). The biological activity is also enhanced by freshwater 76 discharges of several estuaries along the coast, and largely by the Río de la Plata (35-36°S) 77 78 and Patos Lagoon (32°S), which transports nutrient-rich waters northwards (Ciotti et al., 1995; 79 Acha et al., 2004). Furthermore, the Subtropical Shelf Front (STSF), a density-compensated 80 thermohaline subsurface front, which occurs over the shelf close to 32°S, creates a region with 81 high nutrient input, primary production, copepod and ichthyoplankton abundance (Piola et al., 82 2000; Muelbert et al., 2008; Acha et al., 2020). Additionally, the SBS is subject to episodic 83 instabilities, such as eddies (Ito et al., 2016) and meanders of the Brazil Current (Lorenzzetti 84 et al., 2009), which influence the distribution patterns of nutrients and planktonic organisms 85 (Brandini, 2006).

86 Worldwide, the role of oceanographic fronts on the distribution of pelagic larvae has 87 been investigated, for instance, along the coast (Belgrano et al., 1995), at topographically 88 generated fronts (Shanks et al., 2002), across rings (Villar et al., 2015), and at estuarine fronts 89 (Ayata et al., 2011), where fronts were responsible for aggregating, transporting, mixing or 90 separating specific assemblages. However, much more focus has been given for its influence 91 on holoplankters and fish larvae (e.g. Flint et al., 2002; Bakun, 2006; Ohman et al., 2012), 92 where elevated plankton abundance, as well as faunal transitions, have been attributed to the 93 presence of the fronts.

This study investigates the potential role of oceanographic fronts on the dynamics of the spring/summer benthic invertebrate larvae community along one of the Large Marine Ecosystems (LMEs), the South Brazil Shelf. The specific goals were (i) to identify frontal sites and their influence on the meroplankton abundance and composition, and (ii) to investigate associations between frontal types and distinct taxonomic groups of larvae. The central hypothesis is that meroplanktonic high abundance patches occur at frontal sites, due to nutrient and chlorophyll-*a* entrainment, and/or physical entrapment of these larvae.

101

102 **2. Materials and methods**

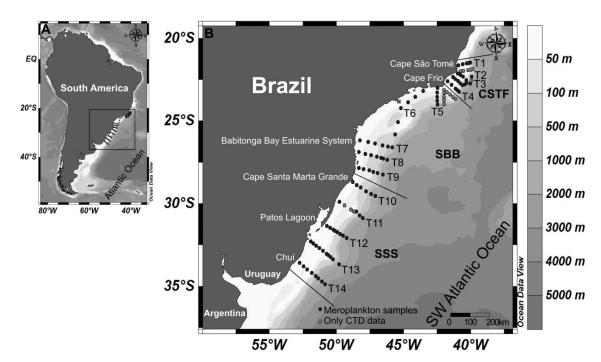
104 *2.1. Study area*

105

The South Brazil Shelf (SBS) extends from 22°S to 34°S along the South American southeast coast (Heileman and Gasalla, 2008) (Fig. 1A). The width of the continental shelf varies according to the latitude, being narrower in the northern than in the southern area (Fig. 1B). The continental slope is more pronounced in the northern region.

The SBS is often divided into three latitudinal subareas (Fig. 1B): (i) the Cape São Tomé-Cape Frio region (CSTF); (ii) the Southern Brazilian Bight (SBB) located between Cape Frio and Cape Santa Marta Grande (CSM); and (iii) the Southern Subtropical Shelf (SSS) between CSM and Río de la Plata. In the northern portion of SBS, the CSTF is mainly characterized by seasonal coastal upwelling, while SBB and SSS are dominated by the strong influence of less saline waters derived from the Río de la Plata and Patos Lagoon, which are stronger in winter and spring (Burrage et al., 2008; Möller et al., 2008).

117



118

Figure 1. (A) Geographic location of the study area. (B) Position of the sampling stations along the 14
cross-shelf transects (T). Subareas: CSTF = Cape São Tomé-Cape Frio; SBB = Southern Brazilian
Bight; SSS = Southern Subtropical Shelf. Plankton samples were collected at 89 stations (black
circles) out of 107 stations. Grey circles represent stations where only CTD data was available.

123

124 2.2. Sampling and laboratory procedures

Oceanographic cruises were conducted between Chuí (34°S) and Cape São Tomé 126 127 (21°S) on board of the R. V. Cruzeiro do Sul (owned by the Brazilian Navy). To cover the 128 entire sampling area three consecutive legs were carried out, with the first and second 129 occurring in austral late spring (December 06 to 14 and 17 to 22, 2010) and the third finishing 130 in early summer (January 04 to 11, 2011) in the CSTF region. The positions of CTD stations 131 were strategically selected to intersect several shelf fronts that could be seen on satellite 132 images. Prior to the cruises, high-resolution (~ 1 km) ocean colour and thermal infrared 133 satellite images were analysed for choosing locations of CTD stations. Vertical profiles of 134 temperature, salinity, fluorescence and dissolved oxygen were recorded at 107 stations 135 distributed at 17 cross-shelf transects using a SeaBird CTD (conductivity, temperature and 136 depth) profiler casts (Fig. 1B). During the cruises, continuous measurements of sea surface (~ 137 5 m) temperature and salinity were made by a well-calibrated thermosalinograph. CDT 138 measurements were only considered at depths greater than 10 m.

139 In addition, water samples were collected at selected depths (3 or 5 m, maximum 140 fluorescence depth and base of the thermocline) to determine chlorophyll-a and nutrient 141 concentrations with 5-L Niskin bottles. Water was filtered on board and chlorophyll-a 142 concentrations were determined by spectrophotometry using the approach detailed in 143 Strickland and Parsons (1972). Ammonia and phosphate concentrations were determined by 144 colorimetric analyses using a portable spectrophotometer, while nitrite, nitrate, and silicate 145 were analyzed using Flow Injection Analysis. Nutrient analysis followed the processing 146 recommendations in Aminot and Chaussepied (1983).

Plankton samples (89, black circles in Fig. 1B) were collected at 14 out of 17 crossshelf transects through vertical tows from the maximum fluorescence depth up to the surface in deep-water stations, from 10 m above the bottom when the water column was homogenous and from about 10 m depth at shallow stations (up to 20 m).

A WP2 net with a 0.5-m diameter mouth and 200-µm mesh equipped with a flowmeter (General Oceanics) was used for sampling planktonic organisms, through vertical tows at a speed of about 2 knots. All samples were immediately fixed and preserved in 4% buffered seawater-formaldehyde solution. The maximum fluorescence depth ranged from 7 to 125 m, and the plankton sampling depth ranged from 12 to 130 m. The distance of sampling locations from the coast ranged from 7 to 418 km. Local depths varied from 15 to 2,800 m, thus covering coastal, shelf and oceanic waters.

158 Invertebrate larvae were counted and sorted from all 89 samples. In a few coastal 159 stations, larvae were counted out of 1/2 or 1/4 fractions of the samples due to high abundances, and values further extrapolated. Larvae were identified into major taxonomic
groups, under stereomicroscope, according to Smith (1977), Boltovskoy (1981) and Young
(2001).

163

164 *2.3. Data analysis*

165

Larval counts were standardized to number of individuals per 100 m³ to calculate the relative abundance (RA) of each taxon. The frequency of occurrence (FO) was also calculated.

A potential temperature–salinity (T–S) diagram was built for the studied area on Ocean Data View (Schlitzer, 2009). Water masses were determined based on thermohaline indexes in Miranda (1985), Castro and Miranda (1998), Piola et al. (2000), and Möller et al. (2008). Temperature and salinity obtained with the thermosalinograph were used as proxies to detect fronts in the study area (Chaigneau and Morrow, 2002). Chlorophyll-*a* and nutrient data were also examined to detect responses to the fronts and some profiles are presented.

In addition to *in situ* data, we used 8-days satellite images of sea surface chlorophyll-*a* concentration for the dates of each leg of the oceanographic cruises, with a 4 km spatial resolution from the MODIS Aqua sensor. Images were obtained through the Giovanni/NASA web site (Berrick et al., 2009).

179 In order to verify the distribution of the most frequent taxa in relation to environmental 180 variables, a distance-based Redundancy Analysis (dbRDA) was conducted, using the Bray-181 Curtis index for similarity between samples. Biological data were Hellinger-transformed to 182 reduce the wide disparity in magnitude between taxa abundances (Legendre and Gallagher, 183 2001). Only the most frequent taxa were considered (> 10%). In order to avoid collinearity of 184 explanatory variables, we applied a variance inflation factor (VIF) and removed collinear 185 variables. A cut-off VIF value of 10 was applied to get the final set of covariates (Zuur et al., 186 2009). The dbRDA and additional tests were performed in R (R Foundation for Statistical 187 Computing), with the 'vegan' and 'HH' packages (Oksanem et al., 2013; Heiberger, 2013).

Additionally, a variance partitioning estimated from the dbRDA allowed to assess the relative amount of variance of the meroplankton abundance into components explained solely by effects of environmental or spatial variables, components explained by combined effects of environmental and spatial variables, and finally unexplained components (Borcard et al., 192 1992). The spatial variables used were selected from Brandão et al., 2015, where principal coordinates of neighbor matrices (PCNM) were applied in order to identify the most predominant spatial patterns (Borcard et al., 2004). Package 'vegan' (Oksanem et al., 2013)
was used for variation partitioning.

Mean values of temperature, salinity, chlorophyll-*a* and nutrient concentrations, calculated from the surface down to the plankton sampling depth, were used in the dbRDA. The oxygen vertical gradient was calculated using the surface oxygen value and the value at the bottom of the oxycline, as well as respective depths. These parameters were used in the dbRDA to characterize the environmental scenarios where meroplankton was distributed.

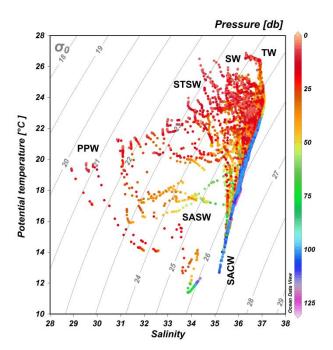
- 201
- 202 **3. Results**
- 203

204 3.1. Physical and biological features of the fronts

205

The T–S diagram in the range 0-130 m (maximum plankton sampling depth) showed the presence of six water masses (Fig. 2): Tropical Water (TW), Shelf Water (SW), Subtropical Shelf Water (STSW), Plata Plume Water (PPW), South Atlantic Central Water (SACW), and Subantarctic Shelf Water (SASW).

210



211

Figure 2. Potential temperature-salinity diagram for the first 130 m of all stations along the South Brazil Shelf during late spring 2010 and early summer 2011. TW = Tropical Water; SW = Shelf Water; STSW = Subtropical Shelf Water; PPW = Plata Plume Water; SACW = South Atlantic Central Water; SASW = Subantarctic Shelf Water.

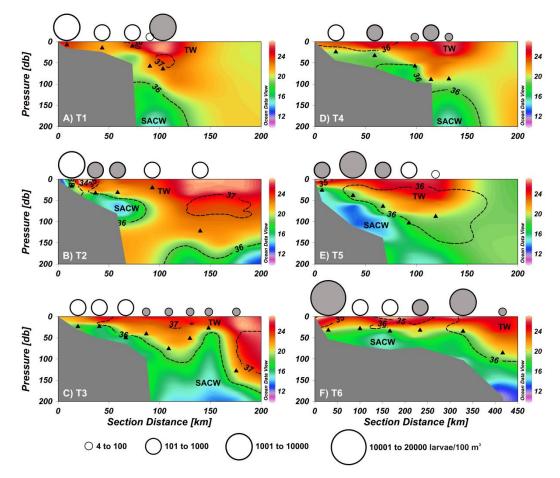
The cross-shelf distribution of temperature and salinity are shown vertically until 200 m depth (Fig. 3) and through high-resolution surface data (Fig. 4), together with meroplankton abundance.

- 220 The salty and warm Tropical Water (TW) was the dominant water mass in the surface 221 layer over the slope area in the entire region (Fig. 3). Great variability was observed regarding 222 the depth of the cool nutrient-rich South Atlantic Central Water (SACW). The isotherms of 223 20°C and 18.5°C define the upper limit below which the SACW dominates the bottom layers 224 over the shelf in the CSTF (Fig. 3A-F) and SBB/SSS (Fig. 3G-N), respectively. The 225 strongest onshore intrusions of the SACW were observed at T2, T5 (Cape Frio), T9 and T10 226 (Cape Santa Marta Grande) (Fig. 3B, E, I and J), depicting the upwelling in subsurface 227 waters. In the Southern portion, the estuarine plume, especially represented by the low-228 salinity Plata Plume Water (PPW), was observed over the shelf from T12 to T14 (Fig. 3L–N), 229 occupying a larger area along T13 (Patos Lagoon) and T14. In the southmost transect, the 230 Subantarctic Shelf Water (SASW) was present below 30 m (Fig. 3N).
- 231 Pelagic larvae of benthic invertebrates were present in all samples, with mean 232 abundance of $1,350 \pm 320$ larvae/100 m³. The highest abundance patches of meroplankton 233 were found in the upwelling zone of Cape Santa Marta Grande (T10) (19,250 larvae/100 m³), 234 in the estuarine front at Patos Lagoon's mouth (T13) (13,550 larvae/100 m³), in the 235 Subtropical Shelf Front (STSF), located at T14 (12,430 larvae/100 m³), and in the coastal realm of T6 (10,400 larvae/100 m³) (Fig. 3 and Fig. 4). At T10 and T6, meroplankton 236 237 abundance peaks were coincident with the upwelling of the SACW (Fig. 3F and J). At T10, 238 the area subject to the upwelling front was characterized by high concentrations of nitrate and 239 phosphate (Fig. 5F and G). In the area of the STSF, the plankton tow coincided with the zone 240 of a sharp change in salinity and temperature (Fig. 3N and Fig. 4N) due to intrusion of SASW 241 into the area. In some cases, meroplankton high abundance patches over the oceanic waters 242 coincided with nocturnal plankton hauls, as seen in offshore stations of T1, T6 and T11 (Fig. 243 3A, F and K).
- Several small-scale surface fronts were identified along the shelf (Fig. 4), and their putative influence on the meroplankton abundance varied throughout the region. For instance, along T12, the increase in meroplankton abundance (3-fold) seems to be associated with the quick increase in salinity ($\Delta S \sim 3.0$) and temperature ($\Delta T \sim 2^{\circ}C$) between stations (Fig. 4L). In addition, at T8, the front ($\Delta S \sim 1.0$; $\Delta T \sim 1^{\circ}C$) also coincided with an increase in meroplankton abundance (2-fold) (Fig. 4H). And at T1 the increase in meroplankton relative abundance (3-fold) was coincident with the thermohaline front ($\Delta S \sim 0.7$; $\Delta T \sim 4^{\circ}C$) (Fig.

4A). On the other hand, despite the thermal front ($\Delta T \sim 3^{\circ}C$) observed at T2, meroplankton abundance was virtually the same between stations (Fig. 4B).

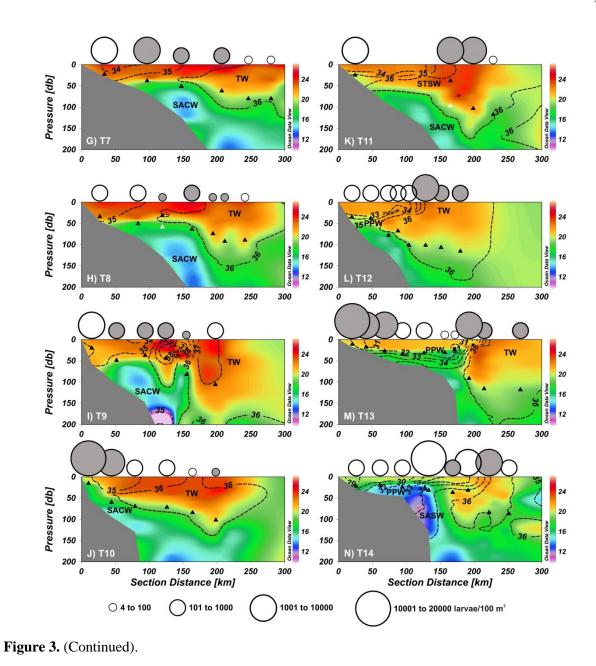
253 During the cruises, the ship crossed two cyclonic eddies over the shelf zone (Ito et al., 254 2016). The first cyclonic eddy-like structure was found in the offshore area of T14, whereas 255 the second was identified offshore T3 (Fig. 4C and N). The cyclonic eddy at T14 section was 256 smaller and weaker than at T3 section (Ito et al., 2016). At T3, the cyclonic vortex was strong 257 enough for upwelling of the SACW from deep layers up to about 50 m deep as well as the 258 maximum chlorophyll-a depth (see Fig. 3C and Fig. 5A), and for aggregating nutrients (Fig. 259 5B–D). Despite plankton sampling was carried out in shallower waters (~ 30 m), a slightly 260 increase in meroplankton relative abundance was observed (Fig. 4C and Fig. 5A).

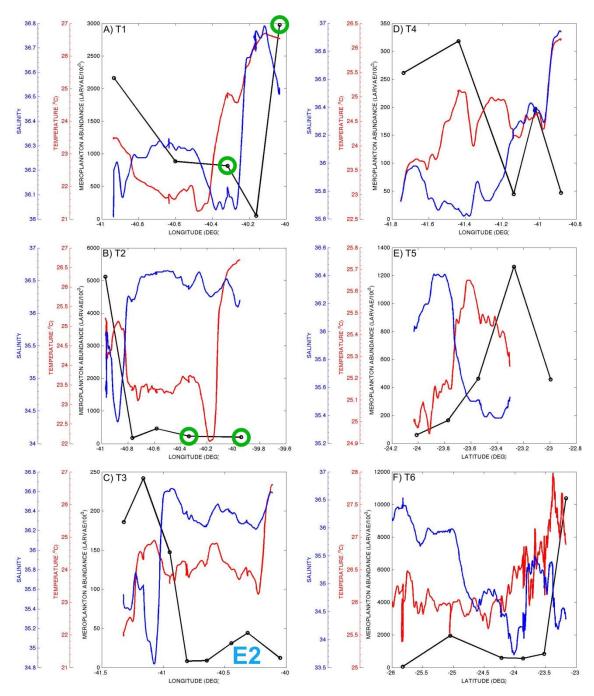
261 High relative abundances of meroplankton were mainly associated with the coastal 262 zone, seen in several transects (Fig. 4A, B, H, I and K). In the mouth of Patos Lagoon (T12) 263 an estuarine and a plume front were observed (Fig. 4M), and meroplankton patches were 264 observed until almost 200 km from the coast, where chlorophyll-a, phosphate and silicate 265 concentrations were also high down to 100 m deep (Fig. 5I, K and L). At T14, high abundances of meroplankton were observed in two stations, one related with the presence of 266 267 the STSF (Fig. 3N), where high concentrations of nitrate and phosphate were observed (Fig. 268 5N and O); and the other associated with the occurrence of the eddy (Fig. 3N and Fig. 4N), 269 depicted by a chlorophyll-*a* bloom (Fig. 5M).



271

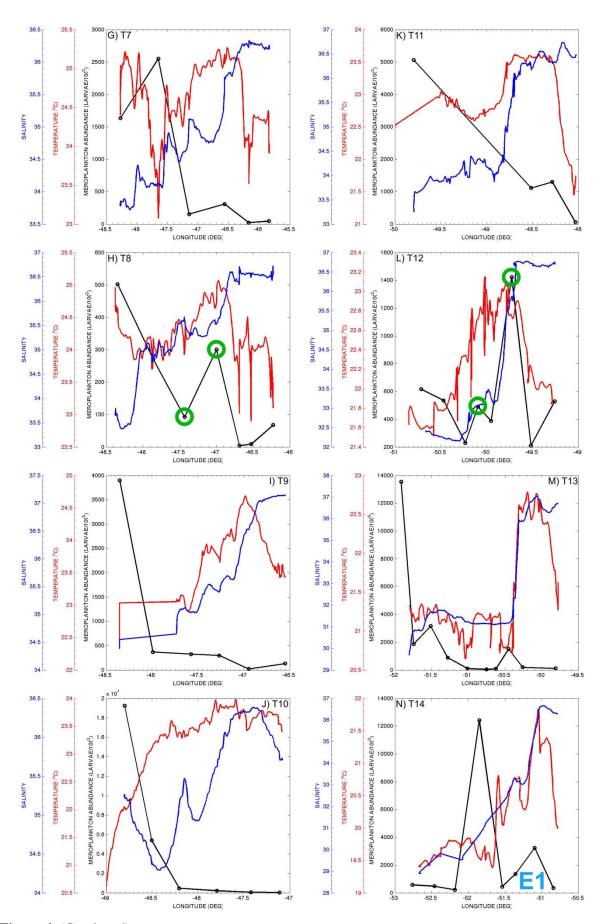
Figure 3. Cross-shelf distributions of temperature (°C) (colors), salinity (contour lines) and meroplankton abundance (larvae/100 m³) (circles) for the transects along the South Brazil Shelf. Circles filled in gray represent stations conducted at night. Black triangles indicate the plankton sampling depth, which was coincident or below (from 1 to 10 m) the maximum fluorescence depth. White triangles indicate the maximum fluorescence depth in stations where it was below plankton sampling depth (T8 and T11).

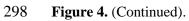


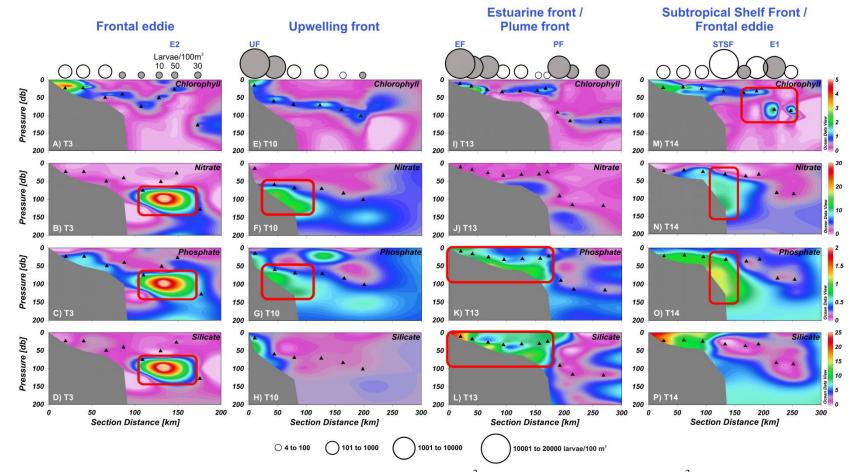


290

Figure 4. Surface salinity and temperature variability obtained from the thermosalinograph and meroplankton abundance (larvae/100 m³) by each transect along the South Brazil Shelf. Note that for T5 and T6 the sections are exceptionally shown by latitude, from South to North. Eddies position (E1 and E2) according to Ito et al. (2016). Green circles indicate the variation in meroplankton abundance and the corresponding limits of the shelf fronts.







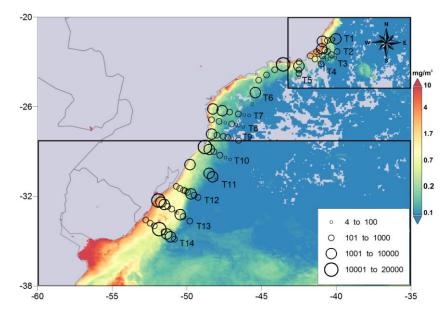
299 300

Figure 5. Cross-shelf distributions of meroplankton abundance (larvae/100 m³) (circles), chlorophyll-a (mg/m³), nitrate (µM), phosphate (µM) 301 and silicate (µM) at T3, T10, T13 and T14. Circles filled in gray represent stations conducted at night. Black triangles indicate the plankton 302 sampling depth. Eddies position (E1 and E2) according to Ito et al. (2016). UF = upwelling front; EF = estuarine front; PF = plume front; STSF = 303 Subtropical Shelf Front. Red rounded rectangles indicate chlorophyll-*a* or nutrient peaks in the area of the fronts.

305 Overall, the surface chlorophyll-*a* concentration and the meroplankton abundance 306 were high all along the continental shelf with maximum values in inshore waters. It was also 307 high over the entire shelf of the southernmost transects (Fig. 6), under the influence of the 308 PPW (Fig. 3L–N).

309

304



310

Figure 6. Meroplankton abundance (black circles) (larvae/100 m³) along the South Brazil Shelf during December 2010 and January 2011. In the background, 8-days satellite images of chlorophyll-*a* concentration are shown for the three legs of the cruise. Leg 1 corresponds to the bottom rectangle, sampled in December 06 to 14, 2010. Leg 2 occurred in December 17 to 22, 2010, during which transects 6 to 9 have been sampled. Leg 3 took place in January 04 to 11, 2011, represented by the upper rectangle.

317

318 3.2. Meroplankton community composition across the fronts

319

Larvae belonging to eleven phyla were found in the area. Among the groups, decapod larvae were the most frequent, while gastropod larvae were the most abundant, followed by larvae of polychaetes and bivalves, with these four groups accounting together for 80% of total larval abundance (Table 1). Besides these groups, cirripedes, holothurians and ophiuroids also presented relatively high mean larval abundance comparing to the others (~ 50 larvae/100 m³).

Polychaete larvae presented up to 100 larvae/100 m³ in most samples. Abundance hotspots of these larvae were observed in the coastal stations of T13 and T10 (~ 7,000 and 4,000 larvae/100 m³, respectively), as well as at T14, near the STSF area (~ 5,000 larvae/100

m³) (Fig. 7A) (Table 2). In turn, molluscan larvae were found in most samples with 329 abundances of up to 1,000 larvae/100 m³. Their highest abundances (~ 8,000 larvae/100 m³) 330 331 occurred in the nearshore stations of T6 (mostly gastropods) and T10 (mostly bivalves), and 332 at T14, at the same station of polychaete larval peak (mostly gastropods) (Fig. 7B; Fig. 8). Crustacean larvae showed a clear pattern of decrease in abundance towards the ocean (Fig. 333 334 7C), with the highest values in the coast of T2 and T13 transects (decapods and cirripedes) (~ 3,300 and 2,500 larvae/100 m³, respectively) (Fig. 7C; Fig. 8). For the detailed distribution of 335 decapod larvae in the area, see Brandão et al. (2015). Echinoderm larvae were found in high 336 337 abundances (up to 4,600 larvae/100 m³) only in a few coastal stations located southward 27°S (Fig. 7D), among which ophiuroids and holothurians were the most representative (Fig. 8; 338 339 Table 1).

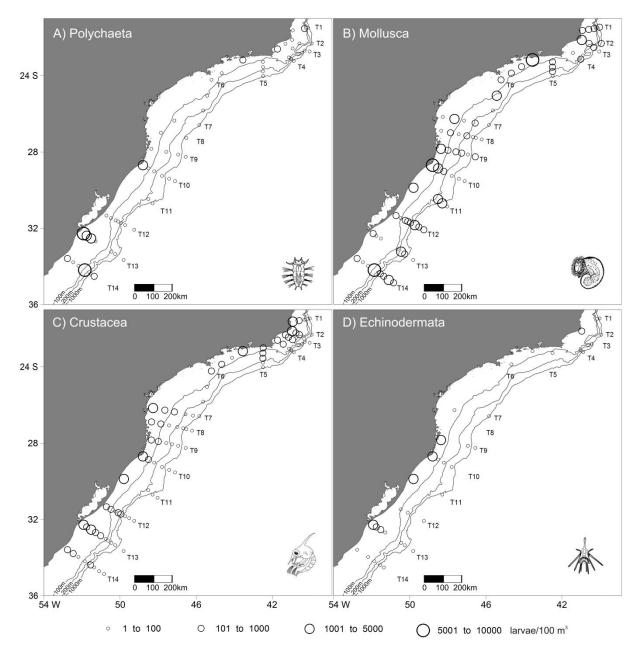


Figure 7. Distribution of larval abundance (larvae/100 m3) of: (A) Polychaeta, (B) Mollusca, (C)
Crustacea and (D) Echinodermata in the 89 stations sampled along the South Brazil Shelf.

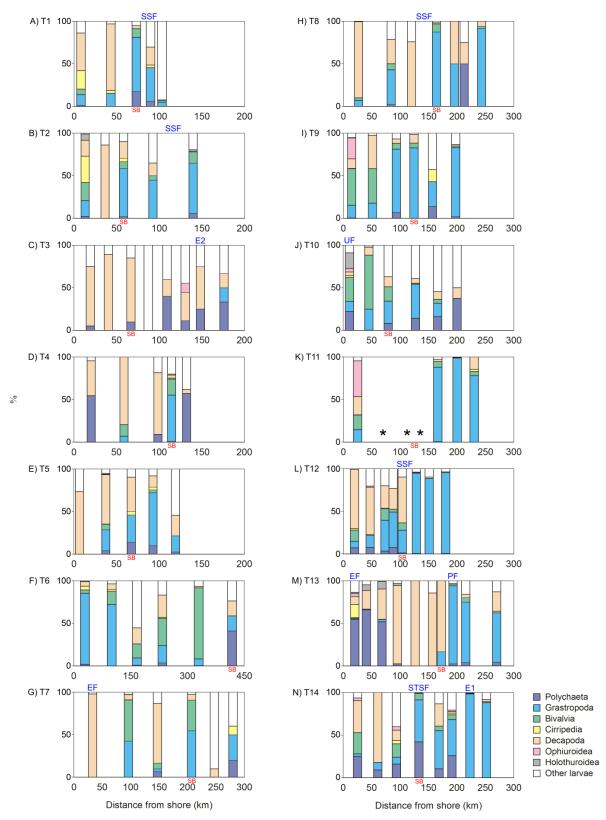




Figure 8. Cross-shelf relative abundance (%) of the main meroplankton groups found along the South
Brazil Shelf, showing the approximate location of the shelf break (SB) and fronts. SSF = small-scale
front; EF = estuarine front; PF = plume front; UF= upwelling front; STSF = Subtropical Shelf Front.
Eddies position (E1 and E2) according to Ito et al. (2016). Asterisks indicate stations without plankton
sampling.

352 Regarding the influence of environmental variables on the distribution of the most 353 frequent meroplankton groups, the first and second axes of the distance-based Redundancy 354 Analysis (dbRDA) ordination accounted together for 65.5% of the constrained variance (Fig. 355 9). Axis 1 represented mainly the cross-shelf gradient. It was positively correlated with 356 chlorophyll-a concentration and oxygen stratification, and negatively with bathymetry, 357 distinguishing neritic from oceanic assemblages. Larvae of gastropods showed an association 358 with offshore conditions. In contrast, crustacean larvae appeared in association with 359 chlorophyll-rich oxygen-stratified coastal waters (Fig. 9).

Axis 2 represented the nutrient-rich waters, especially in ammonia, silicate and phosphate, in opposition to the nutrient-poor waters, characterized by high salinity and temperature. This separation seems to be associated with the contrasting conditions between the PPW, which occupies the neritic waters in the south, and the Tropical Water (TW), dominant over the slope. Larvae of polychaetes and echinoderms were strongly associated with the estuarine plume waters (Fig. 9).

366

351

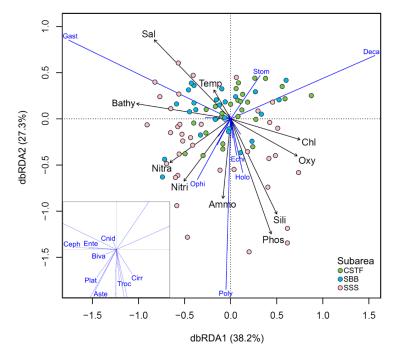


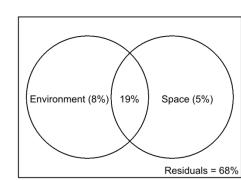
Figure 9. Distance-based Redundancy Analysis (dbRDA) ordination for meroplankton composition in
relation to environmental variables. Triplot with explanatory variables, taxa and samples (colored by
latitudinal subarea). Taxa: Gast = Gastropoda; Stom = Stomatopoda; Deca = Decapoda; Ophi =
Ophiuroidea; Echi = Echinoidea; Holo = Holothuroidea; Poly = Polychaeta; Ceph = Cephalopoda;
Ente = Enteropneusta; Cnid = Cnidaria; Biva = Bivalvia; Plat = Platyhelminthes; Aste = Asteroidea;
Troc = Trocophores; Cirr = Cirripedia. Explanatory variables: Bathy = bathymetry; Sal = salinity;

374 Temp = temperature; Nitra = nitrate; Nitri = nitrite; Ammo = ammonia; Phos = phosphate; Sili =
375 silicate; Oxy = oxygen; Chl = chlorophyll-*a*. Subareas: CSTF = Cape São Tomé-Cape Frio; SBB =
376 Southern Brazilian Bight; SSS = Southern Subtropical Shelf.

377

In summary, the coastal realm was generally dominated by larvae of decapods, cirripedes and bivalves. In addition, gastropod and decapod larvae were the most representative larvae in small-scale shelf fronts (Table 2). In turn, larvae of polychaetes were the dominant in the estuarine front of Patos Lagoon section (T13) (Fig. 8, Fig. 9 and Table 2). In the upwelling front of Cape Santa Marta (T10), larvae of a higher number of taxa occurred, being mostly represented by Bivalvia, Gastropoda and Polychaeta (Fig. 8 and Table 2).

Variance partitioning indicated that most of the variation in invertebrate larval distributions is due to unexplained or stochastic variance. Of the explained portion, variation was mainly due to the combined effect of the environmental and spatial structure of the hydrological environment which accounted for 19% of the total variation. Environmental variables alone explained 8%, while spatial variation retained 5% (Fig. 10).



390

389

Figure 10. Venn diagram showing the results of the variation partitioning procedure.

392

393 **4. Discussion**

394

395 The present findings provide information on the distribution of pelagic larvae of 396 several benthic invertebrate taxa and its relationship with oceanographic fronts within a wide 397 latitudinal range in the South Brazil Shelf. The coast-ocean gradient was the most striking 398 feature on the distribution of meroplankton along the area during spring 2010 and summer 399 2011. Meroplankton abundances were higher in the nearshore stations irrespectively of the 400 latitudinal hydroclimatic scenarios. The contrast between the coastal water masses, rich in 401 nutrients, and the dominant Tropical Water (TW) over the outer shelf and slope, seems to 402 greatly contribute to this general pattern. Besides, the surface chlorophyll-a concentration

distribution was also coincident with the meroplankton abundance gradient. Similar coastocean gradients have also observed for meroplankton in the Southern Ocean (Thatje et al.,
2003) or along a Patagonian fjord (Meerhoff et al., 2014), assigned mainly to nutrient and
chlorophyll-*a* inputs.

407 In addition to the cross-shelf gradient, meroplankton highest values were recorded at 408 specific zones under the influence of distinct frontal systems. It is worth highlighting the 409 frontal upwelling at Cape Santa Marta (CSM), where meroplankton presented a remarkable 410 20-fold increase in relation to the neighboring stations. Coastal upwelling events in this area 411 were observed most likely in response to NE winds (Möller et al., 2008), with strong 412 intrusions of the South Atlantic Central Water (SACW). Thus, physical and chemical 413 gradients are formed along shore between the upwelled water and the advected Coastal Water 414 (CW) (Brandini et al., 2018). This promotes an increase in the productivity and food 415 availability, particularly of diatoms (Brandini et al., 2014), ensuring the availability of food 416 for the larvae. High abundances of bacterioplankton, fish eggs and larvae were also found in 417 the area under the influence of CSM coastal upwelling (Fontes et al., 2018; Macedo-Soares et 418 al., 2014). In addition, high biomass values of invertebrates were also observed in the area, 419 notably gastropods, bivalves and polychaetes (Amaral and Rossi-Wongtschowski, 2004).

420 The estuarine and plume fronts at Patos Lagoon transect also presented an increase in 421 meroplankton abundance, of more than 10-fold. Plata Plume Water (PPW) displays a wide 422 range of physical and biogeochemical properties, as chlorophyll-a and silicates (Ito et al., 423 2016), which reflect elevated nutrient availability due to both respiratory processes and 424 subantarctic water mass contributions (Acha et al., 2004). Indeed, it has been shown to sustain 425 high chlorophyll-a concentrations and consequently high phytoplankton biomass (Ciotti et al., 426 1995; Möller et al., 2008), thus providing planktotrophic larvae with abundant food resources. 427 Regarding the silicates, the high concentrations could be due also to river discharges, since 428 continental freshwaters are characterized by high concentrations of silicates (Ciotti et al., 429 1995). In addition, river plumes and associated fronts also act as physical barriers for the 430 dispersal of pelagic larvae, which may aid to concentrate or retain them in their vicinities 431 (Largier, 2003).

It is also worth to point out the strong influence of the Subtropical Shelf Front (STSF) on the meroplankton abundance, presenting an increase of nearly 60-fold, coincident with a sharp change in temperature and salinity. The increase in meroplankton abundance is likely to be associated with high nutrient input (nitrate and phosphate) and primary production. High abundances of copepod and ichthyoplankton have also been registered in the area (Muelbert et al., 2008; Acha et al., 2018). This front could be compared to other shelf-break fronts, such as
in the Middle Atlantic Bight (Marra et al., 1990) or in the northeastern North Sea (Munk et
al., 1995), where increases were observed in phytoplankton concentration and fish larvae
abundance, respectively. Consistently, the perturbations caused by the front bring turbid,
nutrient-rich water into clearer water, making it more productive than elsewhere.

442 Close inspection of satellite thermal and color images show that the ship crossed two cyclonic vortices (Ito et al., 2016). Although only a slightly increase in meroplankton relative 443 444 abundance was observed in the corresponding areas. The occurrence of cyclonic eddies in the 445 Southern Brazilian Bight enriches nutrients at the bottom layer of the euphotic zone; therefore, 446 regenerated production, a common feature of these oligotrophic waters, is temporarily 447 replaced by new production, in which the nitrogen compound is primarily nitrate (Metzler et 448 al., 1997; Ito et al., 2016). The episodic occurrence of vortices may indeed enhance the 449 primary production and, consequently, zooplanktonic community in the area (Acha et al., 450 2004). In addition, eddy systems may act as retention areas for neritic invertebrate larvae, as 451 seen in the shelf of Gran Canaria in the NW Africa (Landeira et al., 2009), and in the Gulf 452 Stream (Anderson and Robinson, 2001), as well as for holoplankton, as seen in the North 453 Pacific (Mackas et al., 2005). For fish larvae, both physical trapping and biological attraction 454 to food contribute to the retention of fish larvae in eddies in the North Pacific (Chang et al., 455 2018).

456 Several other small-scale thermal and saline shelf fronts were detected during the 457 studied period, primarily in the surface and, secondarily, in the subsurface waters. Although 458 the extent of their influence was not consistent, varying between no influence to a 3-fold 459 increase, in most of the areas, higher values of meroplankton abundance were found in the 460 frontal zones, supporting the hypothesis.

Variance partitioning highlighted that the variation in meroplankton abundances was mainly explained by the combined effect of the geographical space and the environmental conditions. In fact, it has been shown that the interaction between the hydrological environment and spatial structure plays a major role on the distribution of meroplankton, especially in estuaries and coastal areas (Ayata et al., 2011; Brandão et al., 2015), meaning that the environmental conditions alone could have a negligible effect, and that larvae are often trapped by hydrological structures and fronts (Shanks et al., 2002).

468 Regarding the meroplankton community composition, certain groups occurred in 469 association with determined types of frontal systems. Most coastal fronts, coincident with 470 highest surface concentrations of chlorophyll-*a*, were dominated by decapod, cirripede, 471 gastropod and bivalve larvae. Previous studies focusing on the benthic megafauna community 472 showed that decapods (mainly crabs Portunus spinicarpus and Hepatus pudibundus) and 473 gastropods (mainly Buccinanops gradatum) dominate the upwelling region off Cape Frio 474 during spring (Léo and Pires-Vanin, 2006). In addition, in the CSTF region and from 475 Babitonga Bay (T7) to CSM (T10) the highest phytoplanktonic densities were observed in the 476 coastal stations, among which, diatoms were the dominant (Becker et al., 2018; Brandini et 477 al., 2014; Moser et al., 2014). In addition, smaller size plankton fractions, including 478 microzooplankton, were found in high abundances associated with coastal processes in 479 Babitonga Bay and CSM sections during the same cruises of the present study (Becker et al., 480 2018). These resources constitute the main prey items on the diet of bivalve veligers, cirripede 481 nauplii and early crab zoeae (Raby et al., 1994; Turner et al., 2001; Sulkin and McKeen, 482 1999). Accordingly, gastropods were the most abundant larvae found in a transect in front of 483 Cape Frio area, with higher values at coastal stations (Yoshinaga et al., 2010).

The estuarine front of Patos Lagoon section was dominated by larvae of polychaetes. An association between polychaete larvae and estuarine plumes has been reported in other nearshore environments, mainly associated with the adult species habitats (e.g. Shanks et al., 2002; Ayata et al., 2011). In fact, inside the Patos Lagoon, polychaetes that live on sediments in unvegetated shoals constitute one of the main food resources for birds, juvenile decapod crustaceans, and fishes (Bemvenuti, 1997).

490 The observed patterns of invertebrate larvae distribution are also very likely to be 491 influenced by the benthic megafauna community, which are the source for the meroplankton 492 community. In fact, we observed consistencies between the distributions of larvae and the 493 benthic community, based on the literature. Larvae of cephalopods showed an association 494 with the offshore waters. The squid Illex argentinus, the cephalopod mostly captured and one 495 of the main deep-sea demersal fishing resources in Brazil and Argentina, presents its highest 496 densities in the shelf break area, where this species is known to spawn (Rossi-Wongtschowski 497 et al., 2006; Vidal et al., 2010).

A biodiversity survey of the benthic community in the continental shelf and slope of the South Brazil Shelf also found that Gastropoda, Bivalvia, Polychaeta, Crustacea and Ophiuroidea were among the most abundant and frequent taxa (Amaral and Rossi-Wongtschowski, 2004). Higher abundances of organisms were associated with sandy and muddy substrata, both mainly present in the shelf until the 200 m bathymetry (Amaral and Rossi-Wongtschowski, 2004), where higher abundances of meroplankton were also found.

507 The findings of relationships between hydrological structures and invertebrate larvae 508 communities along the South Brazil Shelf add to the accumulating evidence that the frontal 509 systems play an important role in the plankton community, and consequently in the benthic 510 community (Acha et al., 2015). The coast-ocean gradient is a pervasive feature shaping the 511 distribution of the meroplankton community, influenced by several aspects, including distance to the coast and chlorophyll-a concentration. In addition, coastal areas with highly stochastic 512 513 processes, such as frontal systems, due to the confluence of oceanographic processes of 514 contrasting origin, promote enrichment of the biological productivity. The results support the 515 hypothesis that an increase in meroplankton abundance is observed in the frontal systems 516 present during the spring/summer in the South Brazil Shelf, such as in the Subtropical Shelf 517 Front, in surface thermal and saline fronts, and in the estuarine front derived from the Río de 518 la Plata and Patos Lagoon estuarine front. In addition, the highest concentrations of larvae 519 coincided with the strongest upwelling event present in the studied period, reinforcing the 520 importance of the SACW intrusions to enhance biological production in the coastal euphotic 521 zones (Moser et al., 2014). Different communities of larvae were observed in association with 522 waters derived from the Río de la Plata and Patos Lagoon estuarine front, the oceanic waters, 523 and the coastal waters, indicating the influence of the oceanographic regime in the 524 composition of the meroplanktonic assemblages. The fronts in the South Brazil Shelf perform 525 a fundamental role in enriching nutrients at the euphotic zone, thus fertilizing the generally 526 oligotrophic waters. Here it is shown that a change is also triggered in the plankton 527 community of invertebrate larvae, which responds with an increased abundance in relation to 528 its surrounding areas. Fronts play an important role in phytoplankton production and carbon 529 export. The carbon biomass produced in fronts may be exported downwards, fueling deeper 530 pelagic and benthic communities, which represents an important pathway in the global carbon 531 cycle (Brandini et al., 2018). In addition, frontal systems in a generally oligotrophic shelf, as 532 the South Brazil Shelf, are of paramount importance for sustaining and influencing the length 533 of the food webs (Acha et al., 2015), and thus, for providing ecosystem services (Martinetto et 534 al., 2020). Therefore, a more precise understanding of the effect of frontal systems on 535 integrated ecosystems community is pivotal, and should also be investigated in the light of 536 possible climate change effects.

537

538 Acknowledgments

This research was funded by The Brazilian Council for Science and Technological 539 540 Development (CNPq) and the Brazilian Ministry of Science and Technology (MCT). M.C. 541 Brandão received a PhD scholarship from the Coordination for the Improvement of Higher Education Personnel (CAPES). A.S. Freire benefited from CNPg grants (312644/2013-2 and 542 543 311994/2016-4). The authors thank the Brazilian Navy and all crew from the R. V. Cruzeiro 544 do Sul for their logistical support during fieldwork, the researches that helped conducting the field work and analysis in the laboratory, and A. Ciotti and R. Pollery, who provided 545 546 chlorophyll-*a* and nutrient data, respectively. 547 548 References 549 550 Acha E.M., Mianzan H.W., Guerrero R.A., Favero M., Bava J. (2004) Marine fronts at the continental shelves of 551 austral South America physical and ecological Processes. Journal of Marine Systems, 44, 83-105.

552

555

Acha E.M., Ehrlich M.D., Muelbert J.H., Pájaro M., Bruno D., Machinandiarena L., Cadaveira M. (2018)
Ichthyoplankton associated to the frontal regions of the Southwestern Atlantic. In: Hoffmeyer M., Sabatini M.E.,
Brandini F., Calliari D., Santinelli N.H. (Eds.) *Plankton Ecology of the Southwestern Atlantic, From Subtropical to the Subantarctic Realm.* Springer, Berlin, pp. 219–244.

560

563

569

Anderson L.A., Robinson A.R. (2001) Physical and biological modeling in the Gulf Stream region Part II.
Physical and biological processes. *Deep-Sea Research I*, 48, 1139–1168.

572

Ayata S., Stolba R., Comtet T., Thiébaut E. (2011) Meroplankton distribution and its relationship to coastal
 mesoscale hydrological structure in the northern Bay of Biscay (NE Atlantic). *Journal of Plankton Research*, 33,

- 575 1193–1211.
- 576

^{Acha E.M., Piola A., Iribarne O., Mianzan E. (2015)} *Ecological Processes at Marine Fronts: Oases in the Ocean.* Springer, Cham, 68 pp.

<sup>Acha E.M., Viñasa M.D., Derisio C., Alemany D., Piola A.R. (2020) Large-scale geographic patterns of pelagic
copepods in the southwestern South Atlantic.</sup> *Journal of Marine Systems*, 204, 103281.

Amaral A.C.Z., Rossi-Wongtschowski C.L.D.B. (2004) Biodiversidade Bentônica da Região Sudeste-Sul do
 Brasil – Plataforma Externa e Talude Superior. Instituto Oceanográfico – USP, São Paulo, 216 pp.

<sup>566
567</sup> Aminot A., Chaussepied, M. (1983) *Manuel des analyses chimiques en milieu marin*. Centre national pour
568 l'exploitation des océans, Brest, 395 pp.

- 577 Bakun A. (2006) Fronts and eddies as key structures in the habitat of marine fish larvae: opportunity, adaptive 578 response and competitive advantage. *Scientia Marina*, **70**, 105–122.
- 579
- Becker E.C., Garcia C.A.E., Freire A.S. (2018) Mesozooplankton distribution, especially copepods, according to
 water masses dynamics in the upper layer of the Southwestern Atlantic shelf (26°S to 29°S). *Continental Shelf Research*, 166, 10–21.
- 583
- Belgrano A., Legendre P., Dewarumez J., Frontier S. (1995) Spatial structure and ecological variations of
 meroplankton on the French-Belgian coast of the North Sea. *Marine Ecology Progress Series*, 128, 43–50.
- 586
- 587 Bemvenuti C.E. (1997) Unvegetated intertidal flats and subtidal bottoms. In: Seeliger U., Odebrecht C., Castello
- J.P. (Eds.) Subtropical Convergence environments: The coast and the sea in the warm temperate southwestern
 Atlantic. Springer, New York, pp. 78–82.
- 590
- Berrick S.W., Leptoukh G., Farley J.D., Rui H. (2009) Giovanni: a web service workflow-based data
 visualization and analysis system. *IEEE Transactions on Geoscience and Remote Sensing*, 47, 106–113.
- 593
- Boltovskoy D. (1981) Atlas del Atlantico Sudoccidental y métodos de trabajo con el zooplancton marino.
 INIDEP, Mar del Plata, 964 pp.
- 596
- Borcard D., Legendre P., Drapeau P. (1992) Partialling out the Spatial Component of Ecological Variation. *Ecology*, 73, 1045–1055.
- 599
- Borcard D., Legendre P., Avois-Jacquet C., Tuomisto H. (2004) Dissecting the spatial structure of ecological
 data at multiple scales. *Ecology*, 85, 1826–1832.
- 602
- 603 Brandão M.C., Garcia C.A.E., Freire A.S. (2015) Large-scale spatial variability of decapod and stomatopod 604 larvae along the South Brazil Shelf. *Continental Shelf Research*, **107**, 11–23.
- 605
- Brandini F.P. (2006) Zooplankton and ichthyoplankton distribution on the southern Brazilian shelf: an overview. *Scientia Marina*, **70**, 189–202.
- 608
- 609 Brandini F.P., Nogueira Jr. M., Simião M., Codina J.C.U., Noernberg M.A. (2014) Deep chlorophyll maximum
- and plankton community response to oceanic bottom intrusions on the continental shelf in the South Brazilian
 Bight. *Continental Shelf Research*, **89**, 61–75.
- 612
- 613 Brandini F.P., Tura P.M., Santos Pedro P.G.M. (2018) Ecosystem responses to biogeochemical fronts in the
- 614 South Brazil Bight. *Progress in Oceanography*, **164**, 52–62.
- 615

616	Burrage D., Wesson J., Martinez C., Pérez T., Möller Jr. O., Piola A. (2008) Patos Lagoon outflow within the		
617	Río de la Plata plume using an airborne salinity mapper: Observing an embedded plume. Continental Shelf		
618	<i>Research</i> , 28 , 1625–1638.		
619			
620	Campos P.C., Möller Jr. O.O., Piola A.R., Palma E.D. (2013) Seasonal variability and coastal upwelling near		
621	Cape Santa Marta (Brazil). Journal of Geophysical Research, 118, 1–14.		
622			
623	Castro B.M., Miranda L.B. (1998) Physical oceanography of Western Atlantic continental shelf located between		
624	4°N and 34°S. In: Robinson A.R., Brink K.H. (Eds.) The Sea. John Wiley and Sons, New York, pp. 209–251.		
625			
626	Chaigneau A., Morrow R. (2002) Surface temperature and salinity variations between Tasmania and Antarctica,		
627	1993–1999. Journal of Geophysical Research, 107, 1–8.		
628			
629	Chang Y.K., Miyazawa Y., Béguer-Pon M., Han Y.S., Ohashi K., Sheng, J. (2018). Physical and biological roles		
630	of mesoscale eddies in Japanese eel larvae dispersal in the western North Pacific Ocean. Scientific Reports, 8,		
631	5013.		
632			
633	Ciotti A.M., Odebrecht C., Fillmann G., Möller Jr. O.O. (1995) Freshwater outflow and Subtropical		
634	Convergence influence on phytoplankton biomass on the Southern Brazilian continental shelf. Continental Shelf		
635	<i>Research</i> , 15 , 1737–1756.		
636			
637	Cowen R.K., Lwiza K.M., Sponaugle S., Paris, C.B., Olson D.B. (2000) Connectivity of marine populations:		
638	open or closed? Science, 287, 857–859.		
639			
640	Flint M.V., Sukhanova I.N., Kopylov A.I., Poyarkov S.G., Whitledge T.E. (2002) Plankton distribution		
641	associated with frontal zones in the vicinity of the Pribilof Islands. Deep-Sea Research II, 49, 6069-6093.		
642			
643	Fontes M.L.S., Berri A., Carvalho M., Fonseca A.L.O., Antônio R.V., Freire A.S. (2018) Bacterioplankton		
644	abundance and biomass stimulated by water masses intrusions over the Southern Brazilian Shelf (between		
645	25°57'S and 29°24'S). Continental Shelf Research, 164, 28–36.		
646			
647	Gaeta S.A., Brandini F.P. (2006) Produção primária do fitoplâncton na região entre Cabo de São Tomé (RJ) e o		
648	Chuí (RS). In: Rossi-Wongschowski C.L.D.B., Madureira L.S. (Eds.) O Ambiente Oceanográfico da Plataforma		
649	Continental e do Talude na Região Sudeste-Sul do Brasil. EDUSP, São Paulo. pp. 219–264.		
650			
651	Heiberger R.M. (2013) HH: Statistical Analysis and Data Display. R package version 2.3-37. Available:		
652	http://CRAN.R-project.org/package=HH. Accessed 2015 Aug 31.		
653			

- Heileman S., Gasalla M.A. (2008) South Brazil Shelf LME Chapter XVI. In: Sherman K., Hempel G. (Eds).
 The UNEP Large Marine Ecosystems Report: A Perspective on Changing Conditions in LMEs of the World's
- 656 *Regional Seas.* United Nations Environment Programme, Nairobi, 723–734.
- 657
- Hidalgo M., Reglero P., Álvarez-Berastegui D., Torres A.P., Álvarez I., Rodriguez J.M., Carbonell A., Zaragoza
- N., Tor A., Goñi R., Mallol S., Balbín R., Alemany F. (2014) Hydrographic and biological components of the
 seascape structure the meroplankton community in a frontal system. *Marine Ecology Progress Series*, 505, 65–
 80.
- 662
- Ito R.G., Garcia C.A.E, Tavano V.M. (2016) Net sea-air CO₂ fluxes and modelled pCO₂ in the southwestern
 subtropical Atlantic continental shelf during spring 2010 and summer 2011. *Continental Shelf Research*, 119,
 665 68–84.
- 666
- Joyce T.M. (1983) Varieties of ocean fronts. In: Stern M.E., Mellor F.K. (Eds.) *Baroclinic instability and ocean fronts*. Woods Hole Oceanographic Institution, Woods Hole, pp. 59.
- 669
- Landeira J.M., Lozano-Soldevilla F., Hernández-León S., Barton E.D. (2009) Horizontal distribution of
 invertebrate larvae around the oceanic island of Gran Canaria: the effect of mesoscale variability. *Scientia Marina*, 73, 757–767.
- 673
- Largier J.L. (2003) Considerations in estimating larval dispersal distances from oceanographic data. *Ecological Applications*, 13, 71–89.
- 676
- 677 Le Fèvre J. (1987) Aspects of the biology of frontal systems. *Advances in Marine Biology*, **23**, 163–299.
- 678
- 679 Legendre P., Gallagher E. (2001) Ecologically meaningful transformations for ordination of species data.
 680 *Oecologia*, **129**, 271–280.
- 681
- Léo F.C.D., Pires-Vanin A.M.S. (2006) Benthic megafauna communities under the influence of the South
 Atlantic Central Water intrusion onto the Brazilian SE shelf: A comparison between an upwelling and a nonupwelling ecosystem. *Journal of Marine Systems*, **60**, 268–284
- 685
- Lorenzzetti J.A., Stech J.L., Filho W.L.M., Assireu A.T. (2009) Satellite observation of Brazil Current inshore
 thermal front in the SW South Atlantic: space/time variability and sea surface temperatures. *Continental Shelf*
- 688 *Research*, **29**, 2061–2068.
- 689
- 690 Macedo-Soares L.C.P., Garcia C.A.E., Freire A.S., Muelbert J.H. (2014) Large-scale ichthyoplankton and water
- 691 mass distribution along the South Brazil Shelf. *Plos One*, **9**, e91241.
- 692

Pacific Eddy of coastal origin: II. Mechanisms of eddy colonization by and retention of offshore species. Deep Sea Research II, 52, 1011–1035. Marra J., Houghton R.W., Garside, C. (1990) Phytoplankton growth at the shelf-break front in the Middle Atlantic Bight. Journal of Marine Research, 48, 851-868. Martinetto P., Alemany D., Botto F., Mastrángelo M., Falabella V., Acha E.M., Antón G., Bianchi A., Campagna C., Cañete G. et al. (2020) Linking the scientific knowledge on marine frontal systems with ecosystem services. Ambio 49, 541-556. Meerhoff E., Tapia F.J., Castro L.R. (2014) Spatial structure of the meroplankton community along a Patagonian fjord – The role of changing freshwater inputs. Progress in Oceanography, 129, 125–135. Metzler P.M., Gilbert P.M., Gaeta S.A., Lublan J.M. (1997) New and regenerated production in South Atlantic off Brazil. Deep Sea Research I, 44, 363-384. Miranda L.B. (1985) Forma de correlação TS de massas d'água das regiões costeira e oceânica entre o Cabo de São Tomé (RJ) e a Ilha de São Sebastião (SP), Brasil. Boletim do Instituto Oceanográfico da USP, 33, 105–119. Möller Jr. O.O., Piola A.R., Freitas A.C., Campos E.J.D. (2008) The effects of river discharge and seasonal winds on the shelf off southeastern South America. Continental Shelf Research, 28, 1607–1624. Morgan S. (2014) Behaviorally mediated larval transport in upwelling systems. Advances in Oceanography, , 1–1. Moser G.A.O., Takanohashi R.A., Braz M.C., Lima D.T., Kirsten F.V., Guerra J.V., Fernandes A.M., Pollery R.C.G. (2014) Phytoplankton spatial distribution on the Continental Shelf off Rio de Janeiro, from Paraíba do Sul River to Cabo Frio. *Hydrobiologia*, **728**, 1–21. Muelbert J.H., Acha M., Mianzan H., Guerrero R., Reta R., Braga E.S., Garcia V.M.T., Berasategui A., Gomez-Erachef M., Ramírez F. (2008) Biological, physical and chemical properties at the Subtropical Shelf Front Zone in the SW Atlantic Continental Shelf. Continental Shelf Research, 28, 1662–1673. Munk P., Larsson P.O., Danielsen D., Moksness E. (1995) Larval and small juvenile cod Gadus morhua concentrated in the highly productive areas of a shelf break front. Marine Ecology Progress Series, 125, 21–30. Munk P., Hansen B.W., Nielsen T.G., Thomsen H.A. (2003) Changes in plankton and communities across hydrographic fronts off West Greenland. Journal of Plankton Research, 25, 815-830.

Mackas D., Tsurumi M., Galbraith M., Yelland D. (2005) Zooplankton distribution and dynamics in a North

733	Nybakken J.W. (1997) Marine Biology: An Ecological Approach. Addison Wesley Longman, Menlo Park, 481
734	pp.
735	
736	Ohman M.D., Powell J.R., Picheral M., Jensen D.W. (2012) Mesozooplankton and particulate
737	matter responses to a deep-water frontal system in the southern California Current System. Journal of Plankton
738	Research, 34 , 815–827.
739	
740	Oksanen J., Blanchet F.G., Kindt R., Legendre P., Minchin P.R., O'Hara R.B., Simpson G.L., Solymos P.,
741	Stevens M.H.H., Wagner H. (2013) Vegan: Community Ecology Package. R package version 2.0-7. Available:
742	http://CRAN.R-project.org/package=vegan. Accessed 2015 Aug 31.
743	
744	Piola A.R., Campos E.J.D., Möller Jr. O.O., Charo M., Martinez C. (2000) Subtropical Shelf Front off eastern
745	South America. Journal of Geophysical Research, 105, 6565–6578.
746	
747	R Core Team (2018). R: A language and environment for statistical computing. R Foundation for Statistical
748	Computing, Vienna, Austria. https://www.R-project.org/
749	
750	Raby D., Lagadeuc Y., Dodson J.J., Mingelbier M. (1994) Relationship between feeding and vertical distribution
751	of bivalve in stratified and mixed waters. Marine Ecology Progress Series, 103, 275-284.
752	
753	Rossi-Wongtschowski C.L.D.B., Ávila-da-Silva A.O., Cergole M.C. (2006) Análise das Principais Pescarias
754	Comerciais da Região Sudeste-Sul do Brasil: Dinâmica Populacional das Espécies em Explotação - II. Instituto
755	Oceanográfico – USP, São Paulo, 96 pp
756	
757	Schlitzer R. (2009) Ocean Data View, htt://odv.awi.de.
758	
759	Shanks A.L., Largier J., Brink L. (2002) Observations on the distribution of meroplankton during a downwelling
760	event and associated intrusion of the Chesapeake Bay estuarine plume. Journal of Plankton Research, 24, 391-
761	416.
762	
763	Smith D.L. (1977) A Guide to Marine Coastal Plankton and Invertebrate Larvae. Kendall/Hunt Publishing
764	Company, California, 221 pp.
765	
766	Strickland J.D., Parsons T.R. (1972) A practical handbook of seawater analysis. Fisheries Research Board of
767	Canada, Ottawa, 310 pp.
768	
769	Sulkin S.D., McKeen G.L. (1999) The significance of feeding history on the value of heterotrophic
770	microzooplankton as prey for larval crabs. Marine Ecology Progress Series, 186, 219-225.
771	

772	Thatje S., Schnack-Schiel S., Arntz W.E. (2003) Developmental trade-offs in Subantarctic meroplankton				
773	communities and the enigma of low decapod diversity in high southern latitudes. Marine Ecology Progress				
774	Series, 260 , 195–207.				
775					
776	Turner J.T., Levinsen H., Nielsen T.G., Hansen B.W. (2001) Zooplankton feeding ecology: grazing on				
777	phytoplankton and predation on protozoans by copepod and barnacle nauplii in Disko Bay, West Greenland				
778	Marine Ecology Progress Series, 221 , 209–219.				
779					
780	Vidal E.A.G., Haimovici M., Hackbart V.C.S. (2010) Distribution of paralarvae and small juvenile cephalopods				
781	in relation to primary production in an upwelling area off southern Brazil. <i>ICES Journal of Marine Science</i> , 67,				
782 782	1346–1352.				
783 784	Villar E., Farrant G.K., Follows M., Garczarek L., Speich S., Audic S., Bittner L., Blanke B., Brum J.R., Brunet				
785	C. et al. (2015) Environmental characteristics of Agulhas rings affect interocean plankton transport. <i>Science</i> ,				
786	348 , 1261447–1261447.				
787					
788	Yoshinaga M.Y., Sumida P.Y.G., Silveira I.C.A., Ciotti A.M., Gaeta S.A., Pacheco L.F.C.M., Koettker A.G.				
789	(2010) Vertical distribution of benthic invertebrate larvae during an upwelling event along a transect off the				
790	tropical Brazilian continental margin, Journal of Marine Systems, 79, 124–133.				
791					
792	Young C.M. (2001) Atlas of Marine Invertebrate Larvae. Academic Press, San Diego, 626 pp.				
793					
794	Zuur A.F., Ieno E.N., Walker N.J., Saveliev A.A., Smith G. (2009) Mixed effects models and extensions in				
795	ecology with R. Springer, New York, 574 pp.				
796					
797					
798					
799					
800					
801					
802					
803					
804					
805					
806					
807					
808					
809					

810 Tables

Table 1. Mean abundance, relative abundance (RA) and frequency of occurrence (FO) of the

Phylum/Subphylum	Таха	Mean (larvae/100 m ³)	RA (%)	FO (%)	
Porifera	Porifera	0.02	< 0.01	1	
Cnidaria	Cnidaria	2.26	0.17	17	
Platyhelminthes	Platyhelminthes	2.90	0.21	17	
Nemertea	Nemertea	4.80	0.35	8	
Annelida	Polychaeta	243.68	17.99	69	
Mollusca	Gastropoda	422.49	31.18	78	
	Bivalvia	212.30	15.67	61	
	Cephalopoda	1.92	0.14	21	
Arthropoda/Crustacea	Cirripedia	59.40	4.38	36	
	Stomatopoda	11.91	0.88	55	
	Decapoda	197.65	14.59	96	
Sipuncula	Sipuncula	3.58	0.26	9	
Phoronida	Phoronida	2.02	0.15	9	
Echinodermata	Ophiuroidea	49.19	3.63	16	
	Asteroidea	6.77	0.50	28	
	Holothuroidea	51.95	3.83	13	
	Echinoidea	6.83	0.50	11	
Hemichordata	Enteropneusta	31.31	2.31	11	
-	Trocophores	13.61	1.00	28	
-	Unidentified	30.26	2.23	48	
	TOTAL	1,354.86	100.00		

813 meroplanktonic larvae sampled along the South Brazil Shelf.

815 **Table 2.** Types of fronts intersected over the South Brazil Shelf during late spring 2010 and 816 early summer 2011, response in the abundance of meroplankton (increase in relation to values 817 surrounding stations), and dominant taxa found. STSF = Subtropical Shelf Front. Front types 818 were classified according to Acha et al., 2004 and Acha et al., 2015.

Tuonacat	Front type	Meroplankton	Dominant groups	
Transect		(~ fold higher)		
T1 – Cape São Tomé	Small-scale	3	Decapoda / Gastropoda	
T2 – Feia Lagoon 2	Small-scale	-	Gastropoda	
T3 – Campos Bight	Eddie	4	Decapoda / Polychaeta	
T4 – Cape Frio 1	Coastal	5	Decapoda / Polychaeta	
T5 – Cape Frio 4	Coastal	3	Decapoda	
T6 – Ilhabela Island	Coastal	10	Gastropoda	
T7 – Babitonga Bay	Estuarine	10	Decapoda / Bivalvia / Gastropoda	
T8 – Itajaí River	Small-scale	2	Decapoda / Gastropoda	
T9 – Santa Catarina Island	Coastal	8	Bivalvia / Ophiuroidea	
T10 – Cape Santa Marta	Upwelling	20	Bivalvia / Gastropoda / Polychaeta	
T11 – Tramandaí	Coastal	5	Ophiuroidea	
T12 – Mostardas	Small-scale	3	Decapoda / Gastropoda	
T13 – Patos Lagoon	Estuarine	14	Polychaeta / Cirripedia / Ophiuroidea	
	Plume	10	Gastropoda	
T14 – Chuí	STSF	60	Polychaeta / Gastropoda	
	Eddie	20	Gastropoda	