NOAA Atlas NESDIS 84

WORLD OCEAN ATLAS 2018
Volume 4: Dissolved Inorganic Nutrients (phosphate, nitrate and nitrate+nitrite, silicate)

Silver Spring, MD
July 2019

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National Centers for Environmental Information (NCEI)

Silver Spring, Maryland
July 2019

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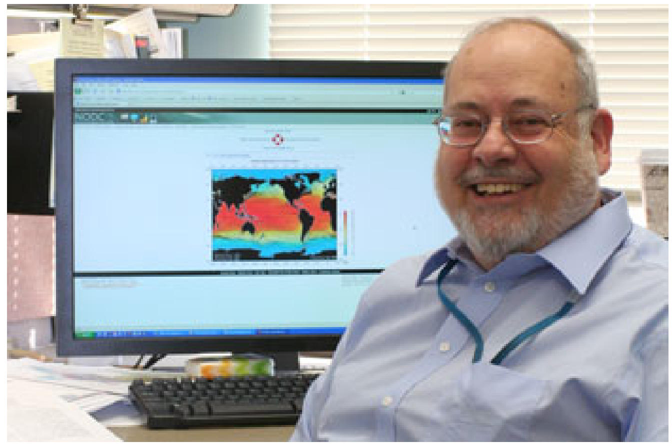
To Sydney (Syd) Levitus



Syd exemplifies the craft of careful, systematic inquiry of the large-scale distributions and low-frequency variability from seasonal-to-decadal time scales of ocean properties. He was one of the first to recognize the importance and benefits of creating objectively analyzed climatological fields of measured ocean variables including temperature, salinity, oxygen, nutrients, and derived fields such as mixed layer depth. Upon publishing *Climatological Atlas of the World Ocean* in 1982, he distributed this work without restriction, an act not common at the time. This seminal atlas moved the oceanographic diagnostic research from using hand-drawn maps to using objectively analyzed fields of ocean variables.

With his NODC Ocean Climate Laboratory (OCL) colleagues, and unprecedented cooperation from the U.S. and international ocean scientific and data management communities, he created the *World Ocean Database (WOD)*; the world's largest collection of ocean profile data that are available internationally without restriction. The *World Ocean Atlas (WOA)* series represents the gridded objective analyses of the WOD and these fields have also been made available without restriction.

The WOD and WOA series are used so frequently that they have become known generically as the "Levitus Climatology". These databases and products enable systematic studies of ocean variability in its climatological context that were not previously possible. His foresight in creating WOD and WOA has been demonstrated by their widespread use over the years. Syd has made major contributions to the scientific and ocean data management communities. He has also increased public understanding of the role of the oceans in climate. He retired in 2013 after 39 years of distinguished civil service. He distilled the notion of the synergy between rigorous data management and science; there are no shortcuts.

All of us at the Ocean Climate Laboratory would like to dedicate this atlas to Syd, his legacy, vision, and mentorship.

The OCL team members

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List of Acronyms

Acronym	Expanded Term
APB	Autonomous Pinniped Bathythermograph
BAMS	Bulletin of the American Meteorological Society
CSV	Comma-Separated Value
CTD	Conductivity Temperature Depth
DBT	Drifting Bathythermograph
DOC	Department of Commerce
DOE	Department of Energy
DRB	Drifting Buoy
ENSO	El Niño-Southern Oscillation
ERL	Earth Research Laboratory
ETOPO2	Earth Topography 2 arc minute
EVR	Extended Vertical Resolution
GIS	Geographic Information System
GLD	Glider
GMT	Greenwich Mean Time, or Generic Mapping Tools
GODAR	Global Ocean Data Archaeology and Rescue
GTSPP	Global Temperature-Salinity Profile Program
IAPSO	International Association for the Physical Sciences of the Oceans
IOC	International Oceanographic Commission
IODE	International Oceanographic Data Exchange
IRI	International Research Institute for Climate and Society

Acronym	Expanded Term
JAMSTEC	Japan Agency for Marine-Earth Science and Technology
JPOTS	Joint Panel on Oceanographic Tables and Standards
LDEO	Lamont-Doherty Earth Observatory
MAST	Marine Science and Technology
MBT	Mechanical Bathythermograph
MEDAR	Mediterranean Data Archeology and Rescue
MRB	Moored Buoy
NAO	North Atlantic Oscillation
NASA	National Aeronautics and Space Administration
NATO	North Atlantic Treaty Organization
NCEI	National Centers for Environmental Information
NESDIS	National Environmental Satellite, Data, and Information Service
NOAA	National Oceanic and Atmospheric Administration
NODC	National Ocean Data Center
OCL	Ocean Climate Laboratory
ODV	Ocean Data View
PFL	Profiling Float
PIRATA	Prediction and Research Moored Array in the Tropical Atlantic
PSS	Practical Salinity Scale
RAMA	Research Moored Array for African-Asian-Australian Monsoon Analysis and Prediction
RDML	Rear Admiral
SST	Sea Surface Temperature
SUR	Surface
TAO/TRITON	Tropical Atmosphere Ocean moored buoy array
TSK	Tsurumi-Seiki Company
UNESCO	United Nations Educational, Scientific and Cultural Organization
UOR	Undulating Oceanographic Recorder
USA	United States of America
USN	United States Navy
WDC	World Data Center
WOA	World Ocean Atlas
WOD	World Ocean Database
XBT	Expendable Bathythermograph
XCTD	Expendable Conductivity Temperature Depth

Preface

The World Ocean Atlas 2018 (WOA18) is the latest in a line of oceanographic analyses of subsurface ocean variables at standard depths extending back to the groundbreaking *Climatological Atlas of the World Ocean* (Levitus, 1982). The WOA has been published semi-regularly since 1994, with versions in 1998, 2001, 2005, 2009, 2013, and now 2018. Previous iterations of the WOA have proven to be of great utility to the oceanographic, climate research, geophysical, and operational environmental forecasting communities. The oceanographic variable analyses are used as boundary and/or initial conditions in numerical ocean circulation models and atmosphere-ocean models, for verification of numerical simulations of the ocean, as a form of "sea truth" for satellite measurements such as altimetric observations of sea surface height, for computation of nutrient fluxes by Ekman transport, and for planning oceanographic expeditions among others.

WOA18 includes analyses on a one-degree grid for all variables and on a quarter-degree grid for temperature and salinity. Since WOA13, the ocean variable analyses are produces on 102 depth levels from the surface to 5,500 m (previously 33 levels within the same depth limits). Ocean data and analyses of data at higher vertical resolution than previously available are needed to document the variability of the ocean, including improving diagnostics, understanding, and modeling of the physics of the ocean.

In the acknowledgment section of this publication, we have expressed our view that creation of global ocean profile and plankton databases and analyses are only possible through the cooperation of scientists, data managers, and scientific administrators throughout the international scientific community.

A pre-release version of WOA18 was made available in September, 2018. The final version of WOA18 was released in July, 2019. In the interim community feedback and our own work has led to changes in the temperature atlas in particular. Animal mounted pinniped temperature profiles have been added as a data source improving coverage in some high latitude areas. A different Expendable Bathythermograph (XBT) correction (Cheng et al., 2014) has been employed. These changes are detailed below. Also, the XBTs were doubly corrected in the pre-release version. The Levitus correction was applied after another correction had been applied (Cheng et al., 2014). This error led to an ocean which was less than 0.1°C cooleer in the pre-release WOA18 as compared to the final WOA18 for the most affected decades (1975-84, 1985-94, 1995-2004) in the upper 400m with smaller differences below. The 1981-2010 climate normal for temperature is slightly cooler (< 0.05°C) in the final WOA18 than in the pre-release WOA18 due to inadvertent double-weighting of the 2001-2010 decade in the pre-release version.

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Acknowledgments

This work was made possible by a grant from the NOAA Climate and Global Change Program, which enabled the establishment of a research group at the National Oceanographic Data Center (now the National Centers for Environmental Information – NCEI). The purpose of this group is to prepare research quality oceanographic databases, as well as to compute objective analyses of, and diagnostic studies based on, these databases. Support is now from base funds and from the NOAA Climate Program Office.

The data on which this atlas is based are in *World Ocean Database 2018* and are distributed on-line by NCEI. Many data were acquired as a result of the IOC/IODE *Global Oceanographic Data Archaeology and Rescue* (GODAR) project, and the IOC/IODE *World Ocean Database* project (WOD).

The WOD is a composite of publicly available ocean profile data, both historical and recent. We acknowledge the scientists, technicians, and programmers who have collected and processed data, those individuals who have submitted data to national and regional data centers as well as the managers and staff at the various data centers. We are working on a more substantive and formalized way to acknowledge all those who have collected and contributed to oceanographic measurements, which were used to calculate the fields in the WOA. Until we have such a system in place, we direct the reader's attention to lists of primary investigators, institutions, and projects, which contributed data (codes can be used to locate data in the World Ocean Database). We also thank our colleagues at the NCEI. Their efforts have made this and similar works possible.

We dedicate this work to Carla Coleman who always contributed with a smile and was taken from us too soon.

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ABSTRACT

This atlas consists of a description of data analysis procedures and horizontal maps of climatological distribution fields of dissolved inorganic nutrients (phosphate, nitrate and nitrate+nitrite, and silicate) at selected standard depth levels of the World Ocean on a one-degree latitude-longitude grid. The aim of the maps is to illustrate large-scale characteristics of the distribution of these nutrients. The oceanographic data fields used to generate these climatological maps were computed by objective analysis of all scientifically quality-controlled historical nutrient data in the *World Ocean Database 2018*. Maps are presented for climatological composite periods (annual, seasonal, monthly, seasonal and monthly difference fields from the annual mean field, and the number of observations) at 102 standard depths. We also provide estimates of the basin-scale uncertainty of the WOA18 nutrient objectively analyzed annual fields.

1. INTRODUCTION

The distribution and concentration of dissolved inorganic nutrients in the world ocean is affected by biochemical (*e.g.*, marine production, respiration, and oxidation or remineralization of labile dissolved and particulate organic matter) and physical processes (*e.g.,* water mass renewal, advection, and mixing). By "nutrients" in this atlas, we mean chemically reactive dissolved inorganic nitrate or nitrate and nitrite (N+N), ortho-phosphate or phosphate, and ortho-silicic acid or silicate (expressed in units of micro-mole per kilogram, μmol kg^{-1}).

This atlas is part of the *World Ocean Atlas 2018* (WOA18) series. The WOA18 series includes analysis for dissolved inorganic nutrients (this atlas), temperature (Locarnini *et al.*, 2019), salinity (Zweng *et al.*, 2019), and dissolved oxygen (Garcia *et al.*, 2019a, b). This atlas presents annual, seasonal, and monthly climatologies and related statistical fields for dissolved inorganic nutrients. Climatologies in this atlas are defined as mean oceanographic fields at selected standard depth levels based on the objective analysis of historical oceanographic profiles and select surface-only data. A profile is defined as a set of measurements of samples collected at discrete depths taken as an instrument such as a rosette CTD package drops or rises vertically in the water column.

This atlas includes an objective analysis of all scientifically quality-controlled historical nutrient measurements available in the *World Ocean Database 2018* (WOD18; Boyer *et al.*, 2019; Garcia *et al.*, 2019c). We present data analysis procedures and horizontal maps showing annual, seasonal, and monthly climatologies and related statistical fields at selected standard depth levels between the surface and the ocean bottom to a maximum depth of 5500 m. The complete set of maps, statistical and objectively analyzed data fields, and documentation are all available on-line.

All of the nutrient climatologies use all available quality control data regardless of year of observation. Most of the data used in this atlas were collected after 1960. The annual climatology was calculated using all data regardless of the month in which the observation was made. Seasonal climatologies were calculated using only data from the defined season (regardless of year). The seasons are here defined as follows. Winter is defined as the months of January, February, and March. Spring is defined as April, May, and June. Summer is defined as July, August, and September. Fall is defined as October, November, and December. Monthly climatologies were calculated using data only from the given month regardless of the day of the month in which the observation was made.

The dissolved inorganic nutrient data used in this atlas are available from the National Centers for Environmental Information (NCEI) and World Data Service for Oceanography (WDS-Oceanography). Large volumes of oceanographic data have been acquired as a result of the fulfillment of several data management projects including:

a) the Intergovernmental Oceanographic Commission (IOC) Global Oceanographic Data Archaeology and Rescue (GODAR) project (Levitus *et al.*, 2005);

b) the IOC World Ocean Database project (WOD);

c) the IOC Global Temperature Salinity Profile project (GTSPP) (IOC, 1998).

The WOD18 nutrient data used in the WOA18 have been analyzed in a consistent, objective manner on a one-degree latitude-longitude grid at standard depth levels from the surface to a maximum depth of 5500m. The procedures for "all-data" climatologies are identical to those used in the *World Ocean Atlas 2013* (WOA13) series (Garcia *et al.*, 2013 a, b). Slightly different procedures were followed in earlier analyses (Levitus, 1982; *World Ocean Atlas 1994* series [WOA94, Levitus *et al*., 1994; Levitus and Boyer, 1994 a, b; Conkright *et al*., 1994]). WOA18 uses 102 standard depth levels.

Objective analyses shown in this atlas are constrained by the nature of the historical nutrient data base (data are non-uniform in space, time, and data quality), characteristics of the objective analysis techniques, and the grid used. These limitations and characteristics are discussed below.

Since the publication of WOA13, substantial amounts of additional historical and modern nutrient data have become available. However, even with these additional data, we are still hampered in a number of ways by a lack of oceanographic data and their quality. The overall precision and uncertainty in the nutrient data have improved over time. Because of the lack of nutrient data coverage in time and space, we are forced to examine the annual cycle by compositing all data regardless of the year of observation. In some geographic areas, quality control is made difficult by the limited number of nutrient data collected in these areas. Data may exist in an area for only one season, thus precluding any representative annual analysis. In some areas there may be a reasonable spatial distribution of quality-controlled data points on which to base an analysis, but there may be only a few (perhaps only one) data values in each one-degree latitude-longitude square.

This atlas is divided into sections. We begin by describing the data sources and data distribution (Section 2). Then we describe the general data processing procedures (Section 3), the results (Section 4), summary (Section 5), and future work (Section 6). Global horizontal maps for the each of the nutrients at each individual depth levels for each time period are available on-line.

2. DATA AND DATA DISTRIBUTION

Data sources and quality control procedures are briefly described below. For further information on the data sources used in WOA18 refer to the *World Ocean Database 2018* (WOD18, Boyer *et al.*, 2019). The quality control procedures used in preparation of these analyses are described by Garcia *et al.* (2019a).

2.1. Data sources

Historical oceanographic data used in this atlas were obtained from the NCEI/WDS-Oceanography archives and include all data gathered as a result of the GODAR and WOD projects. The nutrient data used in this atlas were typically obtained by means of analysis of serial (discrete) samples using manual or continuous flow analysis photometric methods. We refer to the discrete water sample dataset in WOD18 as Ocean Station Data (OSD). Typically, each profile in the OSD dataset consists of 1 to 36 water samples collected at various depths between the surface and the ocean bottom using Nansen or Niskin samplers. Garcia *et al.* (2018a, c) describes the quality control procedures used in preparation of these analyses.

To understand the procedures for taking individual oceanographic observations and constructing climatological fields, definition of the terms "standard level data" and "observed level data" are necessary. We refer to the actual measured value of an oceanographic variable *in situ* (Latin for in place) as an "observation", and to the depth at which such a measurement was made as the "observed level depth". We refer to such data as "observed level data". Before the development of oceanographic instrumentation that measure at high frequencies along the vertical profile, oceanographers often attempted to make measurements at selected "standard levels" in the water column. Sverdrup *et al.* (1942) presented the suggestions of the International Association of Physical Oceanography (IAPSO) as to which depths oceanographic measurements should be made or interpolated to for analysis. Historically the World Ocean Atlas used a modified version of the IAPSO standard depths. However, with the increased global coverage of high depth resolution instrumentation, such as profiling floats, WOA has extended the standard depth levels from 33 to 102. The new standard depth levels include the original depth levels presented up to WOA09, but have tripled the resolution in the upper 100 meters, more than doubled the depth resolution of the upper 1000 meters, and almost three and a half times the resolution for overall depth levels. For many purposes, including preparation of the present climatologies, observed level data are interpolated to standard depth levels if observations did not occur at the desired standard depths (see section 3.1 for details). The levels at which the nutrient climatologies were calculated are given in Table 1. Table 2 shows the depths of each standard depth level. Section 3.1 discusses the vertical interpolation procedures used in our work.

2.2. Data quality control

Quality control of the nutrient data is a major task, the difficulty of which is directly related to lack of data and metadata (for some areas) upon which to base statistical checks. Consequently we applied certain empirical criteria (see sections 2.2.1 through 2.2.4), and as part of the last processing step, subjective judgment was used (see sections 2.2.5 and 2.2.6). Individual data, and in some cases entire profiles or all profiles for individual cruises, have been flagged and not used because these data produced features that were subjectively judged to be non-representative or questionable. As part of our work, we have made available WOD18 which contains both observed levels profile data and standard depth level profile data with various quality control flags applied. The flags mark individual nutrient measurements or entire profiles which were not used in the next step of the procedure, either interpolation to standard depth levels for observed level data or calculation of statistical means in the case of standard depth level data.

Our knowledge of the variability of the world ocean based on the instrumental record now includes a greater appreciation and understanding of the ubiquity of eddies, rings, and lenses in some parts of the world ocean as well as inter-annual and inter-decadal variability of water mass properties associated with modal variability of the atmosphere such as the North Atlantic Oscillation, Pacific Decadal Oscillation, and El Niño Southern Ocean Oscillation. Therefore, we have simply added quality control flags to the nutrient data, not eliminating them from the WOD18. In addition, some data values include the originator's quality flags (*e.g.*, World Ocean Circulation Experiment, CLIVAR repeat hydrography). Thus, individual investigators can make their own decision regarding the representativeness of the data. Investigators studying the distribution of features such as eddies will be interested in those data that we may regard as unrepresentative for the preparation of the analyses shown in this atlas.

2.2.1. Duplicate elimination

Because data are received from many sources, sometimes the same data set is received at NODC/WDC more than once but with slightly different time and/or position and/or data values, and hence are not easily identified as duplicate stations. Therefore, to eliminate the repetitive data values our databases were checked for the presence of exact and near exact replicates using eight different criteria. The first checks involve identifying stations with exact position/date/time and data values; the next checks involve offsets in position/date/time. Profiles identified as duplicates in the checks with a large offset were individually verified to ensure they were indeed duplicate profiles. In summary, we eliminated all but one profile from each set of replicate profiles at the first step of our data processing.

2.2.2. Range and gradient checks

Range checking (*i.e.*, checking whether individual nutrient concentration values are within preset minimum and maximum values as a function of depth and ocean region) was performed on all data values as a first quality control check to flag and withhold from further use the relatively few values that were grossly outside expected oceanic concentration ranges. Range checks were prepared for individual oceanic regions. A check as to whether excessive vertical gradients occur in the data has been performed for each nutrient variable in WOD18 both in terms of positive and negative gradients. We flagged and not used values that exceeded these gradients. Garcia *et al.* (2018 a,c) detail the quality control procedures.

2.2.3. Statistical checks

Statistical checks were performed as follows. All data for each nutrient variable (irrespective of year), at each standard depth level, were averaged within five-degree latitude-longitude squares to produce a record of the number of observations, mean, and standard deviation in each square. Statistics were computed for the annual, seasonal, and monthly compositing periods. Below 50 m depth, if data were more than three standard deviations from the mean, the data were flagged and withheld from further use in objective analyses. Above 50 m depth, a five-standard-deviation criterion was used in five-degree squares that contained any land area. In selected five-degree squares that are close to land areas, a four-standard-deviation check was used. In all other squares a three-standard-deviation criterion was used for the 0-50 m depth layer. For standard depth levels situated directly above the ocean bottom, a four-standard-deviation criterion was used.

The reason for the relatively weaker standard deviation criterion in coastal and near-coastal regions is the exceptionally large variability in the coastal five-degree square statistics for some variables. Frequency distributions of some variables in some coastal regions are observed to be skewed or bimodal. Thus to avoid flagging possibly good data in highly variable environments, the standard deviation criteria were broadened.

For each nutrient variable, the total number of measurements in each profile, as well as the total number of nutrient observations exceeding the standard deviation criterion, were recorded. If more than two nutrient observations in a profile were found to exceed the standard deviation criterion, then the entire profile was flagged. This check was imposed after tests indicated that surface data from particular casts (which upon inspection appeared to be erroneous) were being flagged but deeper data were not. Other situations were found where erroneous nutrient data from the deeper portion of a cast were flagged, while near-surface data from the same cast were not flagged because of larger natural variability in surface layers. One reason for this was the decrease of the number of nutrient observations with depth and the resulting change in sample statistics. The standard-deviation check was applied twice to the data set for each compositing period.

In summary, first the five-degree square statistics were computed, and the data flagging procedure described above was used to provide a preliminary data set. Next, new five-degree-square statistics were computed from this preliminary data set and used with the same statistical check to produce a new, “clean” data set. The reason for applying the statistical check twice was to flag (and withhold from further use), in the first round, any grossly erroneous or non-representative data from the data set that would artificially increase the variances. The second check is then more effective in identifying smaller, but non-representative, observations.

2.2.4. Subjective flagging of data

The nutrient data were averaged by one-degree squares for input to the objective analysis program. After initial objective analyses were computed, the input set of one-degree means still contained questionable data contributing to unrealistic distributions, yielding intense bull's-eyes or spatial gradients. Examination of these features indicated that some of them were due to profiles from particular oceanographic cruises. In such cases, data from an entire cruise were flagged and withheld from further use by setting a flag on each profile from the cruise. In other cases, individual profiles or measurements were found to cause these features and were flagged.

2.2.5. Representativeness of the data

Another quality control issue is nutrient data representativeness. The general paucity of data forces the compositing of all historical nutrient data to produce “climatological” fields. In a given one-degree square, there may be data from a month or season of one particular year, while in the same or a nearby square there may be data from an entirely different year. If there is large interannual variability in a region where scattered sampling in time has occurred, then one can expect the analysis to reflect this. Because the observations are scattered randomly with respect to time, except for a few limited areas (*i.e.*, time series stations such as Hawaii

Ocean Time Series, Bermuda Atlantic Time Series, CARIACO), the results cannot, in a strict sense, be considered a true long-term climatological average.

We present smoothed analyses of historical means, based (in certain areas) on relatively few observations. We believe, however, that useful information about the oceans can be gained through our procedures and that the large-scale features are representative of the real ocean. We believe that, if a hypothetical global synoptic set of ocean data (temperature, salinity, dissolved oxygen, nutrients, *etc*) existed and one were to smooth these data to the same degree as we have smoothed the historical means overall, the large-scale features would be similar to our results. Some differences would certainly occur because of interannual-to-decadal-scale variability.

The nutrient observations diminish in number with increasing depth. In the upper ocean, the all-data annual mean distributions are reasonable for representing large-scale features, but for the seasonal and monthly periods, the data base is inadequate in some regions. With respect to the deep ocean, in some areas the distribution of observations may be adequate for some diagnostic computations but inadequate for other purposes. If an isolated deep basin or some region of the deep ocean has only one observation, then no horizontal gradient computations are meaningful. However, useful information is provided by the observation in the computation of other quantities (*e.g.*, a volumetric mean over a major ocean basin).

3. DATA PROCESSING PROCEDURES

3.1. Vertical interpolation to standard levels

Vertical interpolation of observed depth level data to standard depth levels followed procedures in JPOTS Editorial Panel (1991). These procedures are in part based on the work of Reiniger and Ross (1968). Four observed depth level values surrounding the standard depth level value were used, two values from above the standard level and two values from below the standard level. The pair of values furthest from the standard level is termed “exterior” points and the pair of values closest to the standard level are termed “interior” points. Paired parabolas were generated via Lagrangian interpolation. A reference curve was fitted to the four data points and used to define unacceptable interpolations caused by “overshooting” in the interpolation. When there were too few data points above or below the standard level to apply the Reiniger and Ross technique, we used a three-point Lagrangian interpolation. If three points were not available (either two above and one below or vice-versa), we used linear interpolation. In the event that an observation occurred exactly at the depth of a standard level, then a direct substitution was made. Table 2 provides the range of acceptable distances for which observed level data could be used for interpolation to a standard level.

3.2. Methods of analysis

3.2.1. Overview

An objective analysis scheme of the type described by Barnes (1964) was used to produce the fields shown in this atlas. This scheme had its origins in the work of Cressman (1959). In *World Ocean Atlas 1994* (WOA94), the Barnes (1973) scheme was used. This required only one “correction” to the first-guess field at each grid point in comparison to the successive correction method of Cressman (1959) and Barnes (1964). This was to minimize computing time used in the processing. Barnes (1994) recommends a return to a multi-pass analysis when computing time is not an issue. Based on our own experience we agree with this assessment. The single pass analysis, used in WOA94, caused an artificial front in the Southeastern Pacific Ocean in a data sparse area (Anne Marie Treguier, personal communication). The analysis scheme used in generating WOA98, WOA01, WOA05, WOA13, and WOA18 analyses uses a three-pass "correction" which does not result in the creation of this artificial front.

Inputs to the analysis scheme were one-degree square means of data values at standard levels (for time period and variable being analyzed), and a first-guess value for each square. For instance, one-degree square means for our annual analysis were computed using all available data regardless of date of observation. For July, we used all historical July data regardless of year of observation.

Analysis was the same for all standard depth levels. Each one-degree latitude-longitude square value was defined as being representative of its square. The 360x180 gridpoints are located at the intersection of half-degree lines of latitude and longitude. An influence radius was then specified. At those grid points where there was an observed mean value, the difference between the mean and the first-guess field was computed. Next, a correction to the first-guess value at all gridpoints was computed as a distance-weighted mean of all gridpoint difference values that lie within the area around the gridpoint defined by the influence radius. Mathematically, the correction factor derived by Barnes (1964) is given by the expression:

$$C_{i,j} = \frac{\sum_{s=1}^{n} W_s Q_s}{\sum_{s=1}^{n} W_s} \qquad (1)$$

in which:

(i,j) - coordinates of a gridpoint in the east-west and north-south directions respectively;

$C_{i,j}$ - the correction factor at gridpoint coordinates (i,j);

n - the number of observations that fall within the area around the point i,j defined by the influence radius;

Q_s - the difference between the observed mean and the first-guess at the S^{th} point in the influence area;

$W_s = e^{-\frac{Er^2}{R^2}}$ (for $r \le R$; $W_s = 0$ for $r > R$);

r - distance of the observation from the gridpoint;

R - influence radius;

$E = 4$.

The derivation of the weight function, W_s, will be presented in the following section. At each gridpoint we computed an analyzed value $G_{i,j}$ as the sum of the first-guess, $F_{i,j}$, and the correction $C_{i,j}$. The expression for this is

$$G_{i,j} = F_{i,j} + C_{i,j} \qquad (2)$$

If there were no data points within the area defined by the influence radius, then the correction was zero, the first-guess field was left unchanged, and the analyzed value was simply the first-guess value. This correction procedure was applied at all gridpoints to produce an analyzed field. The resulting field was first smoothed with a median filter (Tukey, 1974; Rabiner *et al.*, 1975) and then smoothed with a five-point smoother of the type described by Shuman (1957) (hereafter referred as five-point Shuman smoother). The choice of first-guess fields is important and we discuss our procedures in section 3.2.5.

The analysis scheme is set up so that the influence radius, and the number of five-point smoothing passes can be varied with each iteration. The strategy used is to begin the analysis with a large influence radius and decrease it with each iteration. This technique allows us to analyze progressively smaller scale phenomena with each iteration.

The analysis scheme is based on the work of several researchers analyzing meteorological data. Bergthorsson and Doos (1955) computed corrections to a first-guess field using various techniques: one assumed that the difference between a first-guess value and an analyzed value at a gridpoint was the same as the difference between an observation and a first-guess value at a nearby observing station. All the observed differences in an area surrounding the gridpoint were then averaged and added to the gridpoint first-guess value to produce an analyzed value. Cressman (1959) applied a distance-related weight function to each observation used in the correction in order to give more weight to observations that occur closest to the gridpoint. In addition, Cressman introduced the method of performing several iterations of the analysis scheme using the analysis produced in each iteration as the first-guess field for the next iteration. He also suggested starting the analysis with a relatively large influence radius and decreasing it with successive iterations so as to analyze smaller scale phenomena with each pass.

Sasaki (1960) introduced a weight function that was specifically related to the density of observations, and Barnes (1964, 1973) extended the work of Sasaki. The weight function of Barnes (1964) has been used here. The objective analysis scheme we used is in common use by the mesoscale meteorological community. Several studies of objective analysis techniques have been made. Achtemeier (1987) examined the "concept of varying influence radii for a successive corrections objective analysis scheme." Seaman (1983) compared the "objective analysis accuracies of statistical interpolation and successive correction schemes." Smith and Leslie (1984) performed an "error determination of a successive correction type objective analysis scheme." Smith *et al.* (1986) made "a comparison of errors in objectively analyzed fields for uniform and non-uniform station distribution."

3.2.2. Derivation of Barnes (1964) weight function

The principle upon which the Barnes (1964) weight function is derived is that "the two-dimensional distribution of an atmospheric variable can be represented by the summation of an infinite number of independent harmonic waves, that is, by a Fourier integral representation". If *f(x,y)* is the variable, then in polar coordinates (r,θ), a smoothed or filtered function *g(x,y)* can be defined:

$$g(x,y) = \frac{1}{2\pi}\int_0^{2\pi}\int_0^{\infty} \eta f\,(x + r\cos\theta, y + r\sin\theta)\, d(\frac{r^2}{4K})d\theta \tag{3}$$

in which r is the radial distance from a gridpoint whose coordinates are *(x,y)*. The weight function is defined as

$$\eta = e^{-\frac{r^2}{4K}} \tag{4}$$

which resembles the Gaussian distribution. The shape of the weight function is determined by the value of K, which relates to the distribution of data. The determination of K follows. The weight function has the property that

$$\frac{1}{2\pi}\int_0^{2\pi}\int_0^{\infty} \eta d\left(\frac{r^2}{4K}\right) d\theta = 1 \tag{5}$$

This property is desirable because in the continuous case (3) the application of the weight function to the distribution *f(x,y)* will not change the mean of the distribution.

However, in the discrete case (1), we only sum the contributions to within the distance *R*. This introduces an error in the evaluation of the filtered function, because the condition given by (5) does not apply. The error can be pre-determined and set to a reasonably small value in the following manner. If one carries out the integration in (5) with respect to θ, the remaining integral can be rewritten as

$$\int_0^R \eta d\left(\frac{r^2}{4K}\right) + \int_R^\infty \eta d\left(\frac{r^2}{4K}\right) = 1 \quad (6)$$

Defining the second integral as ε yields

$$\int_0^R e^{-\frac{r^2}{4K}} d\left(\frac{r^2}{4K}\right) = 1 - \varepsilon \quad (7)$$

Integrating (7), we obtain

$$\varepsilon = e^{-\frac{R^2}{4K}} \quad (7a)$$

Taking the natural logarithm of both sides of (7a) leads to an expression for *K*,

$$K = R^2 / 4E \quad (7b)$$

where $E \equiv -\ln \varepsilon$

Rewriting (4) using (7b) leads to the form of weight function used in the evaluation of (1). Thus, choice of *E* and the specification of *R* determine the shape of the weight function. Levitus (1982) chose *E*=4 which corresponds to a value of ε of approximately 0.02. This choice implies with respect to (7) the representation of more than 98 percent of the influence of any data around the gridpoint in the area defined by the influence radius *R*. This analysis (WOA18) and previous analyses (WOA94, WOA98, WOA01, WOA05, WOA13) used *E*=4.

Barnes (1964) proposed using this scheme in an iterative fashion similar to Cressman (1959). Levitus (1982) used a four-iteration scheme with a variable influence radius for each pass. WOA94 used a one-iteration scheme. WOA98, WOA01, WOA05, WOA09, WOA13, and WOA18 employed a three-iteration scheme with a variable influence radius.

3.2.3. Derivation of Barnes (1964) response function

It is desirable to know the response of a data set to the interpolation procedure applied to it. Following Barnes (1964) and reducing to one-dimensional case we let

$$f(x) = A\sin(\alpha x) \quad (8)$$

in which $\alpha = 2\pi/\lambda$ with λ being the wavelength of a particular Fourier component, and substitute this function into equation (3) along with the expression for η in equation (4). Then

$$g(x) = D[A\sin(\alpha x)] = Df(x) \quad (9)$$

in which *D* is the response function for one application of the analysis and defined as

$$D = e^{-\left(\frac{\alpha R}{4}\right)^2} = e^{-\left(\frac{\pi R}{2\lambda}\right)^2}$$

The phase of each Fourier component is not changed by the interpolation procedure. The results of an analysis pass are used as the first-guess for the next analysis pass in an iterative fashion. The relationship between the filtered function *g(x)* and the response function after *N* iterations as derived by Barnes (1964) is

$$g_N(x) = f(x)D\sum_{n=1}^{N}(1-D)^{n-1} \quad (10)$$

Equation (10) differs trivially from that given by Barnes. The difference is due to our first-guess field being defined as a zonal average, annual mean, seasonal mean, or monthly mean, whereas Barnes used the first application of the analysis as a first-guess. Barnes (1964) also showed that applying the analysis scheme in an iterative fashion will result in convergence of the analyzed field to the observed data field. However, it is not desirable to approach the observed data too closely, because at least seven or eight gridpoints are needed to represent a Fourier component.

The response function given in (10) is useful in two ways: it is informative to know what Fourier components make up the analyses, and the computer programs used in generating the analyses can be checked for correctness by comparison with (10).

3.2.4. Choice of response function

The distribution of O_2 observations (see appendices) at different depths and for the different averaging periods, are not regular in space or time. At one extreme, regions exist in which every one-degree square contains data and no interpolation needs to be performed. At the other extreme are regions in which few if any data exist. Thus, with variable data spacing the average separation distance between gridpoints containing data is a function of geographical position and averaging period. However, if we computed and used a different average separation distance for each variable at each depth and each averaging period, we would be generating analyses in which the wavelengths of observed phenomena might differ from one depth level to another and from one season to another. In WOA94, a fixed influence radius of 555 kilometers was used to allow uniformity in the analysis of all variables. For the present analyses (as well as for WOA13, WOA09, WOA98, and WOA01), a three-pass analysis, based on Barnes (1964), with influence radii of 892, 669 and 446 km was used for the 1° analysis.

Inspection of (1) shows that the difference between the analyzed field and the first-guess field values at any gridpoint is proportional to the sum of the weighted-differences between the observed mean and first-guess at all gridpoints containing data within the influence area.

The reason for using the five-point Shuman smoother and the median smoother is that our data are not evenly distributed in space. As the analysis moves from regions containing data to regions devoid of data, small-scale discontinuities may develop. The five-point Shuman and median smoothers are used to eliminate these discontinuities. The five-point Shuman smoother does not affect the phase of the Fourier components that comprise an analyzed field.

The response function for the analyses presented in the WOA18 series is given in Table 4 and in Figure 1. For comparison purposes, the response function used by Levitus (1982), WOA94, and others are also presented. The response function represents the smoothing inherent in the objective analysis described above plus the effects of one application of the five-point Shuman smoother and one application of a five-point median smoother. The effect of varying the amount of smoothing in North Atlantic sea surface temperature (SST) fields has been quantified by Levitus (1982) for a particular case. In a region of strong SST gradient such as the Gulf Stream, the effect of smoothing can easily be responsible for differences between analyses exceeding 1.0°C.

To avoid the problem of the influence region extending across land or sills to adjacent basins, the objective analysis routine employs basin “identifiers” to preclude the use of data from adjacent basins. Table 5 lists these basins and the depth at which no exchange of information between basins is allowed during the objective analysis of data, *i.e.*, “depths of mutual exclusion.” Some regions are nearly, but not completely, isolated topographically. Because some of these nearly isolated basins have water mass properties that are different from surrounding basins, we have chosen to treat these as isolated basins as well. Not all such basins have been identified because of the complicated structure of the sea floor. In Table 5, a region marked with an (*) can interact with adjacent basins except for special areas such as the Isthmus of Panama.

3.2.5. First-guess field determination

There are gaps in the data coverage and, in some parts of the world ocean, there exist adjacent basins whose water mass properties are individually nearly homogeneous but have distinct basin-to basin differences. Spurious features can be created when an influence area extends over two basins of this nature (basins are listed in Table 6). Our choice of first-guess field attempts to minimize the creation of such features. To maximize data coverage and best represent global variability, a set of "time-indeterminant" climatologies were produced as a first-guess for each set of decadal climatologies. The time-indeterminant climatologies used the first-guess field procedures developed for earlier versions of WOA: To provide a first-guess field for the "all-data" annual analysis at any standard level, we first zonally averaged the observed temperature data in each one-degree latitude belt by individual ocean basins. The annual analysis was then used as the first-guess for each seasonal analysis and each seasonal analysis was used as a first-guess for the appropriate monthly analysis if computed.

We then reanalyzed the temperature data using the newly produced analyses as first-guess fields described as follows and as shown in Figure 3. A new annual mean was computed as the mean of the twelve monthly analyses for the upper 1500 m, and the mean of the four seasons below 1500 m depth. This new annual mean was used as the first-guess field for new seasonal analyses. These new seasonal analyses in turn were used to produce new monthly analyses. This procedure produces slightly smoother means.

These time-indeterminant monthly mean objectively analyzed temperature fields were used as the first-guess fields for each "decadal" monthly climatology. Likewise, time-indeterminant seasonal and annual climatologies were used as first-guess fields for the seasonal and annual decadal climatologies.

We recognize that fairly large data-void regions exist, in some cases to such an extent that a seasonal or monthly analysis in these regions is not meaningful. Geographic distribution of observations for the "all-data" annual periods (see appendices) is excellent for the upper layers of the ocean. By using an "all-data" annual mean, first-guess field regions where data exist for only one season or month will show no contribution to the annual cycle. By contrast, if we used a zonal average for each season or month, then, in those latitudes where gaps exist, the first-guess field would be heavily biased by the few data points that exist. If these were anomalous data in some way, an entire basin-wide belt might be affected.

One advantage of producing "global" fields for a particular compositing period (even though some regions are data void) is that such analyses can be modified by investigators for use in modeling studies.

For the time-indeterminant quarter-degree first-guess field, the one-degree time-indeterminant field was also used. Each of the sixteen quarter-degree boxes enclosed used the one-degree time-indeterminant value as a first-guess, thereby projecting the one-degree climatology onto the quarter-degree grid. In those areas where there was no one-degree value due to land or bottom mask, the statistical mean for the entire basin at the given depth was used. This first-guess field was then used to calculate time-indeterminant quarter-degree field. The time indeterminant quarter-degree field was then used for each quarter-degree decadal climatological mean.

3.3. Choice of objective analysis procedures

Optimum interpolation (Gandin, 1963) has been used by some investigators to objectively analyze oceanographic data. We recognize the power of this technique but have not used it to produce analyzed fields. As described by Gandin (1963), optimum interpolation is used to analyze synoptic data using statistics based on historical data. In particular, second-order statistics such as correlation functions are used to estimate the distribution of first order parameters such as means. We attempt to map most fields in this atlas based on relatively sparse data sets. By necessity we must composite all data regardless of year of observation, to have enough data to produce a global, hemispheric, or regional analysis for a particular month, season, or even yearly. Because of the paucity of data, we prefer not to use an analysis scheme that is based on second order statistics. In addition, as Gandin has noted, there are two limiting cases associated with optimum interpolation. The first is when a data distribution is dense. In this case, the choice of interpolation scheme makes little difference. The second case is when data are sparse. In this case, an analysis scheme based on second order statistics is of questionable value. For additional information on objective analysis procedures see Thiebaux and Pedder (1987) and Daley (1991).

3.4. Choice of spatial grid

The analyses that comprise WOA18 have been computed using the ETOPO2 (Earth Topography 2 arc minute) land-sea topography to define ocean depths at each gridpoint (ETOPO2, 2006). From the ETOPO2 land mask, a quarter-degree land mask was created based on ocean bottom depth and land criteria. If sixteen or more 2-minute square values out of a possible forty-nine in a one-quarter-degree box were defined as land, then the quarter-degree gridbox was defined to be land. If no more than two of the 2-minute squares had the same depth value in a quarter-degree box, then the average value of the 2-minute ocean depths in that box was defined to be the depth of the quarter-degree gridbox. If ten or more 2-minute squares out of the forty-nine had a common bottom depth, then the depth of the quarter-degree box was set to the most common depth value. The same method was used to go from a quarter-degree to a one-degree resolution. In the one-degree resolution case, at least four points out of a possible sixteen (in a one-degree square) had to be land in order for the one-degree square to remain land and three out of sixteen had to have the same depth for the ocean depth to be set. These criteria yielded a mask that was then modified by:

1. Connecting the Isthmus of Panama;
2. Maintaining an opening in the Straits of Gibraltar and in the English Channel;
3. Connecting the Kamchatka Peninsula and the Baja Peninsula to their respective continents.

The one-degree mask was created from the quarter-degree mask instead of directly from ETOPO2 in order to maintain consistency between the quarter-degree and one-degree masks.

4. RESULTS

The on-line figures for this atlas include seven types of horizontal maps representing annual, seasonal, and monthly spatial distribution of analyzed data and data statistics as a function of selected standard depth levels for dissolved inorganic nutrients over one-degree latitude-longitude grid:

a) Objectively analyzed climatology fields. Grid boxes for which there were less than three values available in the objective analysis defined by the influence radius are denoted by a white "+" symbol.

b) Statistical mean fields. Grid boxes for which there were less than three values available in the objective analysis defined by the influence radius are denoted by a white "+" symbol.

c) Data distribution fields for the number of observations in each grid box used in the objective analysis binned into 1 to 2, 3-5, 6-10, 11-30, 31-50 and greater than 51 observations.

d) Standard deviation fields binned into several ranges depending on the depth level. The maximum value of the standard deviation is shown on the map.

e) Standard error of the mean fields binned into several ranges depending on the depth level.

f) Difference between observed and analyzed fields binned into several ranges depending on the depth level.

g) Difference between seasonal/monthly temperature fields and the annual mean field.

h) The number of mean values within the radius of influence for each grid box was also calculated. This is not represented as stand-alone maps, but the results are used on a) and b) maps (see above) to mark the grid boxes with less than three mean values within the radius of influence. These calculations are available as data files.

The maps are arranged by composite time periods (annual, seasonal, month). Table 5 describes all available maps and data fields. We note that the complete set of all climatological maps (in color), objectively analyzed fields, and associated statistical fields at all standard depth levels shown in Table 2, as well as the complete set of data fields and documentation, are available on-line. The complete set of data fields and documentation are available on-line as well.

All of the figures use consistent symbols and notations for displaying information. Continents are displayed as light-grey areas. Coastal and open ocean areas shallower than the standard depth level being displayed are shown as solid gray areas. The objectively analyzed fields include the nominal contour interval used. In addition, these maps may include in some cases additional contour lines displayed as dashed black lines. All of the maps were computer drafted using Generic Mapping Tools (Wessel and Smith, 1998).

We describe next the computation of annual and seasonal fields (section 4.1) and available objective and statistical fields (section 4.2).

4.1. Computation of annual and seasonal fields

After completion of all of our analyses we define a final annual analysis as the average of our twelve monthly mean nutrient fields in the upper 800 m of the ocean (Figure 2). Our final seasonal analysis is defined as the average of monthly analyses in the upper 800 m of the ocean.

4.2. Available statistical fields

Table 5 lists all objective and statistical fields calculated as part of WOA18. Climatologies of oceanographic variables and associated statistics described in this document, as well as global figures of the same can be obtained online.

The sample standard deviation in a grid box was computed using:

$$s = \sqrt{\frac{\sum_{n=1}^{N}(x_n - \bar{x})^2}{N-1}} \qquad (11)$$

in which x_n= the n^{th} data value in the grid box, $\bar{x}$=mean of all data values in the grid box, and N= total number of data values in the grid box. The standard error of the mean was computed by dividing the standard deviation by the square root of the number of observations in each grid box.

In addition to statistical fields, the land/ocean bottom mask and basin definition mask are available on-line. A user could take the standard depth level data from WOD18 with flags and these masks, and recreate the WOA18 fields following the procedures outlined in this document. Explanations and data formats for the data files are found under documentation on the WOA18 webpage.

4.3. Obtaining WOA18 fields on-line

The objective and statistical data fields can be obtained online in different digital formats at the WOA18 webpage. The WOA18 fields can be obtained in ASCII format (WOA native and comma separated value [CSV]) and Network Common Data Form (NetCDF) through our WOA18 webpage. For users interested in specific geographic areas, the World Ocean Atlas Select (WOAselect) selection tool can be used to designate a subset geographic area, depth, and oceanographic variable to view, and optionally download, climatological means or related statistics in shapefile format which is compatible with GIS software such as ESRI ArcMap. WOA18 includes a digital collection of "JPEG" images of the objective and statistical fields. In addition, WOA18 can be obtained in Ocean Data View (ODV) format. WOA18 will be available through other online locations as well. WOA98, WOA01, WOA05, WOA09, and WOA13 are presently served through the IRI/LDEO Climate Data Library with access to statistical and objectively analyzed fields in a variety of digital formats.

To provide an estimate of the uncertainty of the WOA18 nutrient climatologies, we compared the WOA18 and the Global Ocean Data Analysis Project version 2 (GLODAPv2, Olsen *et al.*, 2016) gridded fields for phosphate, nitrate, and silicate. GLODAPv2 does not have seasonal or monthly nutrient gridded fields. The results suggest that the differences between the two annual mean climatologies are relatively small below about 500 m depth at the global and basin scale (60°N-60°S). The global average difference (± 1 standard deviation) of WOA18 minus GLODAPv2 gridded fields below 500 m depth is -0.02±0.07 μmol kg^{-1} for phosphate, -0.22±0.95 μmol kg^{-1} for nitrate, and -0.3±3.8 μmol kg^{-1} for silicate (Table 6). The average differences are generally less than or comparable to the current estimated long-term measurement precision of these nutrients by standard colorimetric methods using Continuous Flow Analysis. The data do not show a significant systematic depth offset. Above 500 m depth, there are significant measurable differences. This is expected because of larger variability in the upper ocean. WOA18 is based on a much representative larger spatial and monthly nutrient data coverage than GLODAPv2. WOA18 contains all of the data used in the creation of GLODAPv2.

5. SUMMARY

In the preceding sections we have described the results of a project to objectively analyze all historical nutrient data in WOD18. We desire to build a set of climatological analyses that are identical in all respects for all variables in WOA18 including relatively data sparse variables such as dissolved inorganic nutrients. This provides investigators with a consistent set of analyses to work with.

One advantage of the analysis techniques used in this atlas is that we know the amount of smoothing by objective analyses as given by the response function in Table 3 and Figure 1. We believe this to be an important function for constructing and describing a climatology of any parameter. Particularly when computing anomalies from a standard climatology, it is important that the field be smoothed to the same extent as the climatology, to prevent generation of spurious anomalies simply through differences in smoothing. A second reason is that purely diagnostic computations require a minimum of seven or eight grid points to represent any Fourier component. Higher order derivatives might require more data smoothing.

We have created objectively analyzed fields and data sets that can be used as a "black box." We emphasize that some quality control procedures used are subjective. For those users who wish to make their own choices, all the data used in our analyses are available both at standard depth levels as well as observed depth levels. The results presented in this atlas show some features that are suspect and may be due to non-representative data that were not flagged by the quality control techniques used. Although we have attempted to eliminate as many of these features as possible by flagging the data which generate these features, some obviously could remain. Some may eventually turn out not to be artifacts but rather to represent real features, not yet capable of being described in a meaningful way due to lack of data. The views, findings, and any errors in this document are those of the authors.

6. FUTURE WORK

Our analyses will be updated when justified by additional water column nutrient observations. As more oceanographic nutrient data are received at NCEI/WDS-Oceanography, we will also be able to extend the seasonal and monthly nutrient analysis to deeper levels.

Additional nutrient collected by automated sensors will likely improve the results and help provide additional observational constraints on observed inter-annual to decadal-scale biochemical variability. For example, recent advances in chemical sensors in Biogeochemical Argo profiling floats enable measurements of nitrate (Johnson *et al.*, 2017a, b). Merging and integrating nutrient data collected by classical photometric methods, using continuous flow analysis, with other sensor-based nutrient observing systems will likely improve the results and provide additional observational constraints on observed inter-annual to decadal-scale changes. Each of these different nutrient observing systems add much additional data coverage and have different data uncertainties and calibrations that must be reconciled before combining into an internally consistent climatology.

We are encouraged by the potential acquisition of much additional high-quality oceanographic observations through recently adopted complementary global projects such as the Global Ocean Observing System (GOOS) 2030 Strategy and the United Nations Decade of Ocean Science for Sustainable Development (2021-2030). GOOS is sponsored by the Intergovernmental Oceanographic Commission of UNESCO, the World Meteorological Organization (WMO), the United Nations Environment Programme (UNEP), and the International Science Council (ISC). Expansion of the current global ocean observing system will enable the creation of more robust climatologies that span shorter climatological time-periods of interest (e.g., inter-annual to decadal).

Creating WOA18 relies on the unrestricted and timely open access and use of oceanographic observations collected worldwide. One country cannot afford the observational system needed to monitor the entire Earth; and thus, open access and use of observations is essential for formulating informed science-based societal-relevant strategies for sustainable ocean use and respond to environmental challenges. The developing research-quality climatologies such as WOA nutrients serve as reliable science-based baselines from which to estimate low frequency regional to global variability.

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Table 1. Descriptions of climatologies for each nutrient variable in WOA18.

The climatologies have been calculated based on bottle data (OSD) from WOD13. The standard depth levels are shown in Table 2.

Oceanographic variable	Depths for annual climatology	Depths for seasonal climatology	Depths for monthly climatology
Nitrate (N+N), Phosphate, and Silicate	0-5500 m (102 levels)	0-800 m (43 levels)	0-800 m (43 levels)

Table 2. Acceptable distances (m) for defining interior (A) and exterior (B) values used in the Reiniger-Ross scheme for interpolating observed level data to standard levels.

Standard Level #	Standard Depths (m)	A	B	Standard Level #	Standard Depths (m)	A	B
1	0	50	200	52	1250	200	400
2	5	50	200	53	1300	200	1000
3	10	50	200	54	1350	200	1000
4	15	50	200	55	1400	200	1000
5	20	50	200	56	1450	200	1000
6	25	50	200	57	1500	200	1000
7	30	50	200	58	1550	200	1000
8	35	50	200	59	1600	200	1000
9	40	50	200	60	1650	200	1000
10	45	50	200	61	1700	200	1000
11	50	50	200	62	1750	200	1000
12	55	50	200	63	1800	200	1000
13	60	50	200	64	1850	200	1000
14	65	50	200	65	1900	200	1000
15	70	50	200	66	1950	200	1000
16	75	50	200	67	2000	1000	1000
17	80	50	200	68	2100	1000	1000
18	85	50	200	69	2200	1000	1000
19	90	50	200	70	2300	1000	1000
20	95	50	200	71	2400	1000	1000
21	100	50	200	72	2500	1000	1000
22	125	50	200	73	2600	1000	1000
23	150	50	200	74	2700	1000	1000
24	175	50	200	75	2800	1000	1000

Standard Level #	Standard Depths (m)	A	B	Standard Level #	Standard Depths (m)	A	B
25	200	50	200	76	2900	1000	1000
26	225	50	200	77	3000	1000	1000
27	250	100	200	78	3100	1000	1000
28	275	100	200	79	3200	1000	1000
29	300	100	200	80	3300	1000	1000
30	325	100	200	81	3400	1000	1000
31	350	100	200	82	3500	1000	1000
32	375	100	200	83	3600	1000	1000
33	400	100	200	84	3700	1000	1000
34	425	100	200	85	3800	1000	1000
35	450	100	200	86	3900	1000	1000
36	475	100	200	87	4000	1000	1000
37	500	100	400	88	4100	1000	1000
38	550	100	400	89	4200	1000	1000
39	600	100	400	90	4300	1000	1000
40	650	100	400	91	4400	1000	1000
41	700	100	400	92	4500	1000	1000
42	750	100	400	93	4600	1000	1000
43	800	100	400	94	4700	1000	1000
44	850	100	400	95	4800	1000	1000
45	900	200	400	96	4900	1000	1000
46	950	200	400	97	5000	1000	1000
47	1000	200	400	98	5100	1000	1000
48	1050	200	400	99	5200	1000	1000
49	1100	200	400	100	5300	1000	1000
50	1150	200	400	101	5400	1000	1000
51	1200	200	400	102	5500	1000	1000

Table 3. Response function of the objective analysis scheme as a function of wavelength for WOA18 and earlier analyses. Response function is normalized to 1.0.

Wavelength[1]	**Levitus (1982)**	**WOA94**	**WOA98, 01, 05, 09, 13, 18**
360ΔX	1.000	0.999	1.000
180ΔX	1.000	0.997	0.999
120ΔX	1.000	0.994	0.999
90ΔX	1.000	0.989	0.998
72ΔX	1.000	0.983	0.997
60ΔX	1.000	0.976	0.995
45ΔX	1.000	0.957	0.992
40ΔX	0.999	0.946	0.990
36ΔX	0.999	0.934	0.987
30ΔX	0.996	0.907	0.981
24ΔX	0.983	0.857	0.969
20ΔX	0.955	0.801	0.952
18ΔX	0.923	0.759	0.937
15ΔX	0.828	0.671	0.898
12ΔX	0.626	0.532	0.813
10ΔX	0.417	0.397	0.698
9ΔX	0.299	0.315	0.611
8ΔX	0.186	0.226	0.500
6ΔX	3.75×10^{-2}	0.059	0.229
5ΔX	1.34×10^{-2}	0.019	0.105
4ΔX	1.32×10^{-3}	2.23×10^{-3}	2.75×10^{-2}
3ΔX	2.51×10^{-3}	1.90×10^{-4}	5.41×10^{-3}
2ΔX	5.61×10^{-7}	5.30×10^{-7}	1.36×10^{-6}

[1]For ΔX = 111 km, the meridional separation at the Equator.

Table 4. Basins defined for objective analysis and the shallowest standard depth level for which each basin is defined.

#	Basin[1]	Standard Depth Level	#	Basin[1]	Standard Depth Level
1	Atlantic Ocean	1*	30	North American Basin	29
2	Pacific Ocean	1*	31	West European Basin	29
3	Indian Ocean	1*	32	Southeast Indian Basin	29
4	Mediterranean Sea	1*	33	Coral Sea	29
5	Baltic Sea	1	34	East Indian Basin	29
6	Black Sea	1	35	Central Indian Basin	29
7	Red Sea	1	36	Southwest Atlantic Basin	29
8	Persian Gulf	1	37	Southeast Atlantic Basin	29
9	Hudson Bay	1	38	Southeast Pacific Basin	29
10	Southern Ocean	1*	39	Guatemala Basin	29
11	Arctic Ocean	1	40	East Caroline Basin	30
12	Sea of Japan	1	41	Marianas Basin	30
13	Kara Sea	8	42	Philippine Sea	30
14	Sulu Sea	10	43	Arabian Sea	30
15	Baffin Bay	14	44	Chile Basin	30
16	East Mediterranean	16	45	Somali Basin	30
17	West Mediterranean	19	46	Mascarene Basin	30
18	Sea of Okhotsk	19	47	Crozet Basin	30
19	Banda Sea	23	48	Guinea Basin	30
20	Caribbean Sea	23	49	Brazil Basin	31
21	Andaman Basin	25	50	Argentine Basin	31
22	North Caribbean	26	51	Tasman Sea	30
23	Gulf of Mexico	26	52	Atlantic Indian Basin	31
24	Beaufort Sea	28	53	Caspian Sea	1
25	South China Sea	28	54	Sulu Sea II	14
26	Barents Sea	28	55	Venezuela Basin	14
27	Celebes Sea	25	56	Bay of Bengal	1*
28	Aleutian Basin	28	57	Java Sea	6
29	Fiji Basin	29	58	East Indian Atlantic Basin	32

[1]Basins marked with a "*" can interact with adjacent basins in the objective analysis.

Table 5. Statistical fields calculated as part of WOA18 ("√"denotes field was calculated and is publicly available).

Statistical field	One-degree Field Calculated	Five-degree Statisctics calculated
Objectively analyzed climatology	√	
Statistical mean	√	√
Number of observations	√	√
Seasonal (monthly) climatology minus annual climatology	√	
Standard deviation from statistical mean	√	√
Standard error of the statistical mean	√	√
Statistical mean minus objectively analyzed climatology	√	
Number of mean values within radius of influence	√	

Table 6. Nominal depth average phosphate, nitrate+nitrite, and silicate (µmol/kg) differences (± 1 standard deviation) of the GLODAPv2 minus WOA18 for 1-degree objectively analyzed fields (60°N-60°S).

Variable	Depth range (m)	Atlantic	Pacific	Indian	Global
Phosphate	0-500	0.02±0.11	0.01±0.13	-0.01±0.12	0.01±0.12
	500-5500	-0.01±0.05	-0.02±0.07	-0.02±0.08	-0.02±0.07
Nitrate + Nitrite	0-500	0.0±1.6	-0.2±1.9	-0.3±1.7	0.2±1.8
	500-5500	-0.2±0.7	-0.2±1.0	-0.2±1.0	-0.2±1.0
Silicate	0-500	0.1±3.1	1.0±3.5	0.9±3.0	0.8±3.6
	500-5500	-0.5±2.7	0.8±4.1	-0.2±3.9	-0.3±3.8

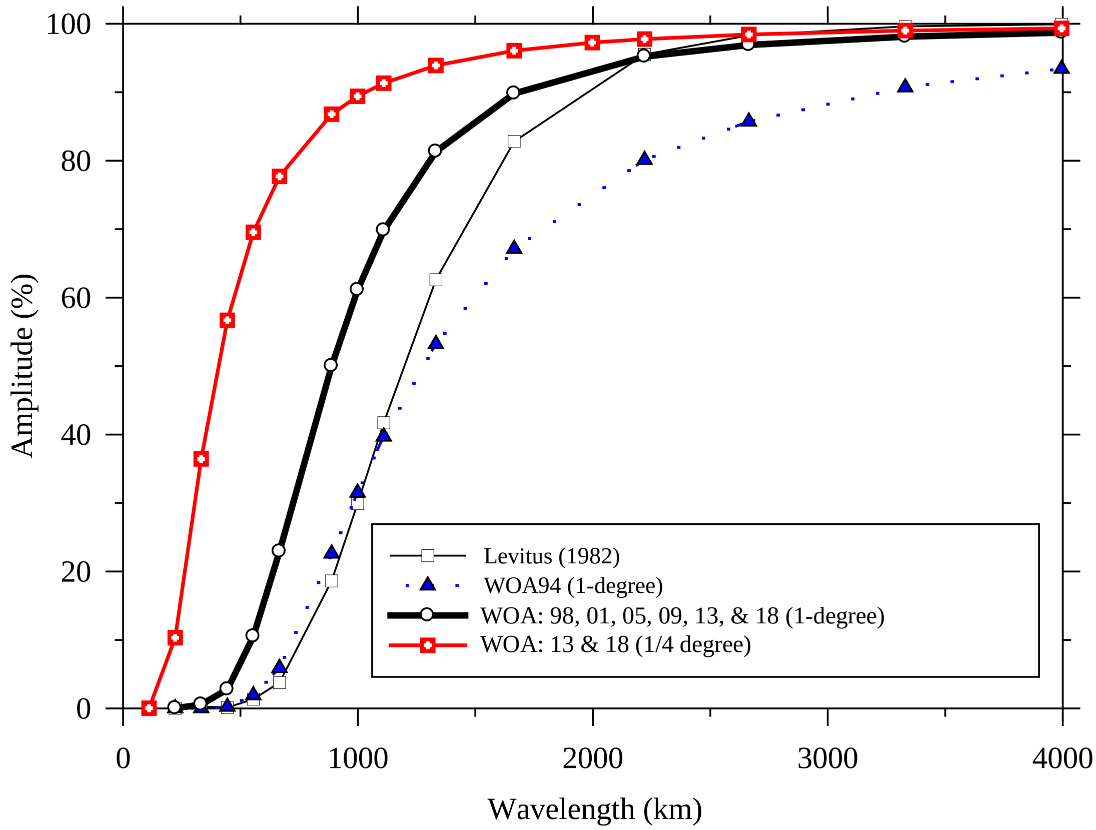

Figure 1. Response function of the WOA18, WOA13, WOA05, WOA01, WOA98, WOA94, and Levitus (1982) objective analysis schemes.

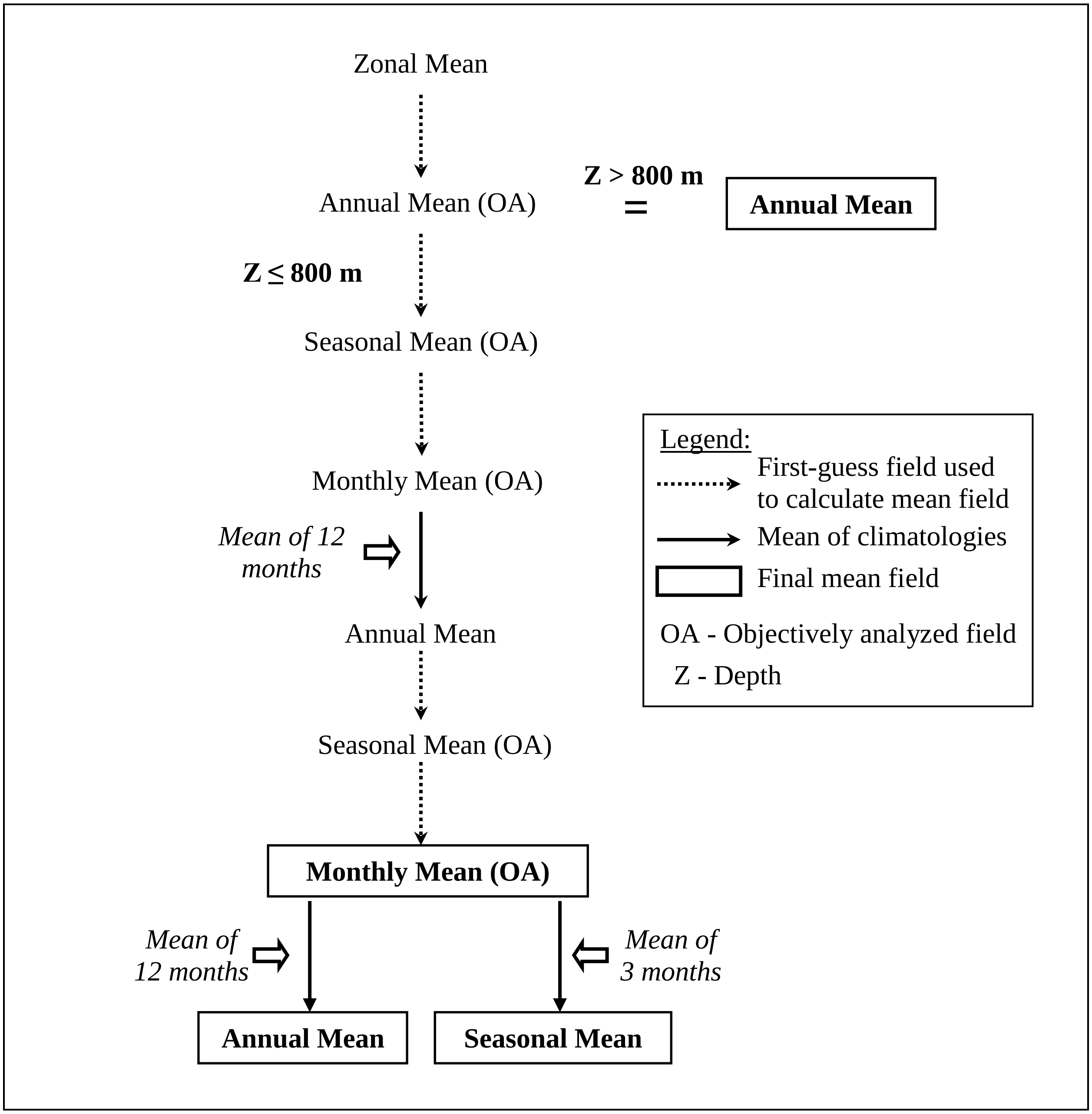

Figure 2. Scheme used in computing annual, seasonal, and monthly objectively analyzed means for phosphate, silicate, and nitrate.