



1 **Impact of dust addition on the metabolism of Mediterranean**  
2 **plankton communities and carbon export under present and**  
3 **future conditions of pH and temperature**

4 Frédéric Gazeau<sup>1</sup>, France Van Wambeke<sup>2</sup>, Emilio Marañón<sup>3</sup>, María Pérez-Lorenzo<sup>3</sup>, Samir  
5 Alliouane<sup>1</sup>, Christian Stolpe<sup>1</sup>, Thierry Blasco<sup>1</sup>, Nathalie Leblond<sup>4</sup>, Birthe Zäncker<sup>5,6</sup>, Anja  
6 Engel<sup>6</sup>, Barbara Marie<sup>7</sup>, Julie Dinasquet<sup>7,8</sup>, Cécile Guieu<sup>1</sup>

7 <sup>1</sup> Sorbonne Université, CNRS, Laboratoire d'Océanographie de Villefranche, LOV, 06230  
8 Villefranche-sur-Mer, France

9 <sup>2</sup> Aix-Marseille Université, Université de Toulon, CNRS/INSU, IRD, Mediterranean Institute of  
10 Oceanography (MIO), UM 110, 13288, Marseille, France

11 <sup>3</sup> Department of Ecology and Animal Biology, Universidade de Vigo, 36310 Vigo, Spain

12 <sup>4</sup> Sorbonne Université, CNRS, Institut de la Mer de Villefranche, IMEV, 06230 Villefranche-sur-  
13 Mer, France

14 <sup>5</sup> The Marine Biological Association of the UK, PL1 2PB Plymouth, United Kingdom

15 <sup>6</sup> GEOMAR Helmholtz Centre for Ocean Research, Kiel, Germany

16 <sup>7</sup> CNRS, Sorbonne Université, Laboratoire d'Océanographie Microbienne, LOMIC, F-66650  
17 Banyuls-sur-Mer, France

18 <sup>8</sup> Scripps Institution of Oceanography, University of California San Diego, USA

19 Correspondence to: Frédéric Gazeau ([f.gazeau@obs-vlfr.fr](mailto:f.gazeau@obs-vlfr.fr))

20 Keywords: Mediterranean Sea; Atmospheric deposition; Plankton community metabolism;

21 Carbon export; Ocean acidification; Ocean warming



22

## 23 **Abstract**

24           Although atmospheric dust fluxes from arid as well as human-impacted areas represent a  
25 significant source of nutrients to surface waters of the Mediterranean Sea, studies focusing on the  
26 evolution of the metabolic balance of the plankton community following a dust deposition event  
27 are scarce and none were conducted in the context of projected future levels of temperature and  
28 pH. Moreover, most of the experiments took place in coastal areas. In the framework of the  
29 PEACETIME project, three dust-addition perturbation experiments were conducted in 300-L  
30 tanks filled with surface seawater collected in the Tyrrhenian Sea (TYR), Ionian Sea (ION) and  
31 in the Algerian basin (FAST) onboard the R/V “Pourquoi Pas?” in late spring 2017. For each  
32 experiment, six tanks were used to follow the evolution of chemical and biological stocks,  
33 biological activity and particle export. The impacts of a dust deposition event simulated at their  
34 surface were followed under present environmental conditions and under a realistic climate  
35 change scenario for 2100 (ca. + 3 °C and -0.3 pH units). The tested waters were all typical of  
36 stratified oligotrophic conditions encountered in the open Mediterranean Sea at this period of the  
37 year, with low rates of primary production and a metabolic balance towards net heterotrophy.  
38 The release of nutrients after dust seeding had very contrasting impacts on the metabolism of the  
39 communities, depending on the station investigated. At TYR, the release of new nutrients was  
40 followed by a negative impact on both particulate and dissolved <sup>14</sup>C-based production rates,  
41 while heterotrophic bacterial production strongly increased, driving the community to an even  
42 more heterotrophic state. At ION and FAST, the efficiency of organic matter export due to  
43 mineral/organic aggregation processes was lower than at TYR likely related to a lower  
44 quantity/age of dissolved organic matter present at the time of the seeding. At these stations, both



45 the autotrophic and heterotrophic community benefited from dust addition, with a stronger  
46 relative increase in autotrophic processes observed at FAST. Our study showed that the potential  
47 positive impact of dust deposition on primary production depends on the initial composition and  
48 metabolic state of the investigated community. This potential is constrained by the quantity of  
49 nutrients added in order to sustain both the fast response of heterotrophic prokaryotes and the  
50 delayed one of primary producers. Finally, under future environmental conditions, heterotrophic  
51 metabolism was overall more impacted than primary production, with the consequence that all  
52 integrated net community production rates decreased with no detectable impact on carbon  
53 export, therefore reducing the capacity of surface waters to sequester anthropogenic CO<sub>2</sub>.



54

## 55 **1. Introduction**

56 Low Nutrient Low Chlorophyll (LNLC) areas represent 60% of the global ocean surface  
57 area (Longhurst et al., 1995) and, although phytoplankton production there is limited by the  
58 availability of nitrogen, phosphorus and iron, it accounts for 50% of global carbon export  
59 (Emerson et al., 1997). Atmospheric dust fluxes from arid as well as anthropogenic sources  
60 represent a significant source of these nutrients to surface waters in these regions and as such  
61 could play a significant role in stimulating primary production (e.g. Bishop et al., 2002; Guieu et  
62 al., 2014b; Jickells and Moore, 2015), potentially increasing the efficiency of the biological  
63 pump in the sequestration of atmospheric CO<sub>2</sub>. However, as heterotrophic prokaryotes have been  
64 shown to outcompete phytoplankton during nutrient addition experiments (e.g. Guieu et al.,  
65 2014a; Mills et al., 2008; Thingstad et al., 2005), dust deposition could induce even stronger  
66 enhancements of heterotrophic bacterial production and/or respiration rates thereby reducing net  
67 atmospheric CO<sub>2</sub> drawdown and the potential for carbon export outside the euphotic zone (Guieu  
68 et al., 2014b). Indeed, several experiments conducted in the Atlantic Ocean and in the  
69 Mediterranean Sea have shown a fast and dominant effect of dust additions on heterotrophic  
70 bacterioplankton metabolism (Herut et al., 2005, 2016; Lekunberri et al., 2010; Marañón et al.,  
71 2010; Pulido-Villena et al., 2008, 2014). However, to the best of our knowledge, no study  
72 focused on the evolution of the metabolic balance of the plankton community after such a dust  
73 event in the open sea. The metabolic balance (or net community production, NCP) is defined as  
74 the difference between gross primary production (GPP) of autotrophic organisms and community



75 respiration (CR) of both autotrophic and heterotrophic organisms, revealing the capacity of a  
76 system to sequester carbon via the biological pump.

77         The Mediterranean Sea is a perfect example of LNLC regions and receives anthropogenic  
78 aerosols originating from industrial and domestic activities from all around the basin and other  
79 parts of Europe and pulses of natural inputs from the Sahara (e.g. Bergametti et al., 1989;  
80 Desboeufs et al., 2018). These atmospheric depositions, mostly in the form of pulsed inputs  
81 (Loÿe-Pilot and Martin, 1996), provide new nutrients (Guieu et al., 2010; Kouvarakis et al.,  
82 2001; Markaki et al., 2003; Ridame and Guieu, 2002) to the surface waters with fluxes that are of  
83 the same order of magnitude as riverine inputs (Powley et al., 2017). These significant nutrient  
84 enrichments likely support primary production especially during the stratification period (Bonnet  
85 et al., 2005; Ridame and Guieu, 2002), however no clear correlation between dust and ocean  
86 color have been evidenced from long series of satellite observations (Guieu and Ridame, 2020).  
87 This raises the question on which compartment (autotrophic or heterotrophic) benefits the most  
88 from these transient relieves in nutrient limitation.

89         In response to ocean warming and increased stratification, LNLC areas are expected to  
90 expand in the future (Irwin and Oliver, 2009; Polovina et al., 2008) due to lower nutrient supply  
91 from sub-surface waters (Behrenfeld et al., 2006). Furthermore, dust deposition could increase in  
92 the future due to desertification (Moulin and Chiapello, 2006), although so far the trend for  
93 deposition remains uncertain because the drying of the Mediterranean basin might also induce  
94 less wet deposition over the basin (Laurent et al., 2021). Nevertheless, whether the fluxes  
95 increase or not in the coming decades and centuries, new nutrients from atmospheric sources will  
96 play an important role in a surface mixed layer even more stratified and isolated from the deeper  
97 nutrient-rich layer. The question remains on how plankton metabolism and carbon export would



98 respond in a warmer and more acidified ocean. Indeed, with an average annual anthropogenic  
99 CO<sub>2</sub> uptake, during the period 2010 to 2019, of  $2.5 \pm 0.6$  GtC (~22.9% of anthropogenic  
100 emissions; Friedlingstein et al., 2020), the oceans substantially contribute towards slowing down  
101 the increase in atmospheric CO<sub>2</sub> concentrations, and therefore towards limiting terrestrial and  
102 ocean warming. However, this massive CO<sub>2</sub> input induces global changes in seawater chemistry  
103 referred to as “ocean acidification” because increased CO<sub>2</sub> concentration lowers seawater pH  
104 (i.e. increases its acidity).

105         Although the response of plankton metabolism to ocean warming has been shown to be  
106 highly dependent on resource availability (Lewandowska et al., 2014), both for heterotrophic  
107 bacteria (Lopez-Urrutia and Moran, 2007) and phytoplankton (Marañón et al., 2018), it has been  
108 suggested that ocean warming will substantially weaken the ocean CO<sub>2</sub> sink in the future as a  
109 consequence of stronger increase in remineralization than in photosynthesis processes, following  
110 the metabolic theory of ecology (MTE; Brown et al., 2004; Gillooly et al., 2001). Ocean  
111 acidification alone has been shown to exert no or very limited influence on plankton metabolism  
112 in the Mediterranean Sea (Maugendre et al., 2017a; Mercado et al., 2014). To the best of our  
113 knowledge, only Maugendre et al. (2015) studied the combined impact of ocean warming and  
114 acidification on plankton metabolism in the Mediterranean Sea. They found a very limited  
115 impact of ocean acidification on the plankton community and a positive impact of warming on  
116 small phytoplankton species (e.g. Cyanobacteria) with a potential decrease of the export and  
117 energy transfer to higher trophic levels. Nevertheless, that study was conducted under nutrient  
118 depleted conditions and there is still a need to assess the combined impact of warming and  
119 acidification on the metabolic balance of plankton communities in this region, following a  
120 transient relief in nutrient availability (Maugendre et al., 2017b).



121           So far there has been no attempt to evaluate the evolution of plankton metabolism and  
122 carbon export following atmospheric deposition in the context of future levels of temperature  
123 and pH. Such experiments were conducted in the frame of the PEACETIME project (ProcEss  
124 studies at the Air-sEa Interface after dust deposition in the MEditerranean sea; [http://peacetime-](http://peacetime-project.org/)  
125 [project.org/](http://peacetime-project.org/)) during the cruise on board the R/V “Pourquoi Pas?” in May/June 2017 (Guieu et al.,  
126 2020). The project aimed at extensively studying and parameterizing the chain of processes  
127 occurring in the Mediterranean Sea after atmospheric deposition, especially of Saharan dust, and  
128 to put them in perspective of on-going environmental changes. During this cruise, three  
129 perturbation experiments were conducted in 300-L tanks filled with surface seawater collected in  
130 the Tyrrhenian Sea (TYR), Ionian Sea (ION) and in the Algerian basin (FAST; Fig. 1). Six tanks  
131 were used to follow the evolution of chemical and biological stocks, biological activity and  
132 export, following a wet dust deposition event simulated at their surface, both under present  
133 environmental conditions and following a realistic climate change scenario for 2100 (ca. + 3 °C  
134 and -0.3 pH units; IPCC, 2013). A companion paper presents the general setup of the  
135 experiments and the impacts of dust under present and future environmental conditions on  
136 nutrients and biological stocks (Gazeau et al., 2020). Here, we focus on the impacts of dust  
137 seeding on plankton metabolism (e.g. primary production, heterotrophic prokaryote production)  
138 and carbon export.



139

## 140 **2. Material and Methods**

### 141 **2.1. General set-up**

142           The general set-up of the experiments is fully detailed in Gazeau et al. (2020). Briefly,  
143 three experiments were performed at the long duration stations TYR, ION and FAST during the  
144 Peacetime cruise onboard R/V “Le Pourquoi Pas?” (Fig. 1). During these experiments (3 to 4  
145 days each), seawater was incubated in 300-L tanks (Fig. S1) installed in a temperature-controlled  
146 container, in which the irradiance spectrum and intensity can be finely controlled and in which  
147 future ocean acidification and warming conditions can be fully reproduced. The tanks were made  
148 of high-density polyethylene (HDPE) and were trace-metal free in order to avoid contaminations,  
149 with a height of 1.09 m, a diameter of 0.68 m, a surface area of 0.36 m<sup>2</sup> and a volume of 0.28 m<sup>3</sup>.  
150 The conical base of the tanks was equipped with a sediment trap that was left open during the  
151 duration of the experiments and removed at the end. The experimental protocol comprised two  
152 unmodified control tanks (C1 and C2), two tanks enriched with Saharan dust (D1 and D2) and  
153 two tanks enriched with Saharan dust and maintained simultaneously under warmer (+ 3 °C) and  
154 acidified (-0.3 pH unit) conditions (G1 and G2). At the three stations, tanks were always filled at  
155 the end of the day before the start of the experiments: TYR (17/05/2017), ION (25/05/2017) and  
156 FAST (02/06/2017). The tanks were filled by means of a large peristaltic pump (Verder© VF40  
157 with EPDM hose, flow of 1200 L h<sup>-1</sup>) collecting seawater below the base of the boat (depth of ~  
158 5 m), used to supply continuously surface seawater to a series of instruments during the entire  
159 campaign. While filling the tanks, seawater was sampled for the measurements of selected  
160 parameters (sampling time = t-12h). After filling the tanks, seawater was slowly warmed



161 overnight using 500 W heaters, controlled by temperature-regulation units (COREMA©), in G1  
162 and G2 to reach an offset of + 3 °C. <sup>13</sup>C-bicarbonate was added to all tanks at 4:00 am (all times  
163 in local time) and G1 and G2 were acidified by addition of CO<sub>2</sub>-saturated filtered (0.2 μm)  
164 seawater (~1.5 L in 300 L; collected when filling the tanks at each station) at 4:30 am to reach a  
165 pH offset of -0.3. Sampling for many parameters took place prior to dust seeding (sampling time  
166 = t<sub>0</sub>). Dust seeding was performed between 7:00 and 9:00 in tanks D1, D2, G1 and G2. The same  
167 dust analog was used and the same dust flux was simulated as for the DUNE 2009 experiments  
168 described in Desboeufs et al. (2014). To mimic a realistic wet flux event of 10 g m<sup>-2</sup>, 3.6 g of this  
169 analog dust were quickly diluted into 2 L of ultrahigh-purity water (UHP water; 18.2 MΩ cm<sup>-1</sup>  
170 resistivity), and sprayed at the surface of the tanks using an all-plastic garden sprayer (duration =  
171 30 min). Depending on the considered parameter or process, seawater sampling was conducted 1  
172 h (t<sub>1h</sub>), 6 h (t<sub>6h</sub>), 12 h (t<sub>12h</sub>), 24 h (t<sub>24h</sub>), 48 h (t<sub>48h</sub>) and 72 h (t<sub>72h</sub>) (+ 96 h = t<sub>96h</sub> for station  
173 FAST) after dust addition. Acid-washed silicone tubes were used for transferring the water  
174 collected from the tanks to the different vials or containers.

## 175 **2.2. Stocks**

### 176 **2.2.1. Dissolved and particulate organic carbon**

177 The concentration of dissolved organic carbon (DOC) was determined from duplicate 10  
178 mL GF/F (pre-combusted, Whatman) filtered subsamples that were transferred to pre-combusted  
179 glass ampoules, acidified with H<sub>3</sub>PO<sub>4</sub> (final pH = 2) and sealed. The sealed glass ampoules were  
180 stored in the dark at room temperature until analysis at the Laboratoire d'Océanographie  
181 Microbienne (LOMIC). DOC measurements were performed on a Shimadzu© TOC-V-CSH



182 (Benner and Strom, 1993). Prior to injection, DOC samples were sparged with CO<sub>2</sub>-free air for 6  
183 min to remove inorganic carbon. Sample (100 µL) were injected in triplicate and the analytical  
184 precision was 2%. Standards were prepared with acetanilid.

185 Seawater samples for measurements of particulate organic carbon concentrations (POC; 2  
186 L) were taken at t-12h, t<sub>0</sub>, t<sub>12h</sub>, t<sub>24h</sub>, t<sub>48h</sub> and t<sub>72h</sub> (or t<sub>96h</sub> for station FAST), filtered on pre-  
187 combusted GF/F membranes, dried at 60 °C and analyzed at the Laboratoire d'Océanographie de  
188 Villefranche (LOV, France) following decarbonation with a drop of HCl 2N, on an elemental  
189 analyzer coupled with an isotope ratio mass spectrometer (EA-IRMS; Vario Pyrocube-Isoprime  
190 100, Elementar©).

### 191 **2.2.2. Total hydrolysable carbohydrates and amino acids**

192 For total hydrolysable carbohydrates and amino acids, samples were taken at t<sub>0</sub>, t<sub>6h</sub>,  
193 t<sub>24h</sub>, t<sub>48h</sub> and t<sub>72h</sub> at all stations. For total hydrolysable carbohydrates (TCHO) > 1 kDa,  
194 samples (20 mL) were filled into pre-combusted glass vials (8 h, 500 °C) and stored at -20 °C  
195 pending analysis. Prior to analysis, samples were desalted with membrane dialysis (1 kDa  
196 MWCO, Spectra Por) at 1 °C for 5 h. Samples were subsequently hydrolyzed for 20 h at 100 °C  
197 with 0.8 M HCl final concentration followed by neutralization using acid evaporation (N<sub>2</sub>, for 5  
198 h at 50 °C). TCHO were analysed at GEOMAR using high performance anion exchange  
199 chromatography with pulsed amperometric detection (HPAEC-PAD), on a Dionex ICS 3000 ion  
200 chromatography system following the procedure of Engel and Händel (2011). Two replicates per  
201 TCHO sample were analyzed.

202 For total hydrolysable amino acids (TAA), samples (5 mL) were filled into pre-  
203 combusted glass vials (8 h, 500 °C) and stored at -20 °C. Samples were hydrolyzed at 100 °C for



204 20 h with 1 mL 30% HCl (Suprapur®, Merck) added to 1 mL of sample, and neutralized by acid  
205 evaporation under vacuum at 60 °C in a microwave. Samples were analyzed by high  
206 performance liquid chromatography (HPLC) using an Agilent 1260 HPLC system following a  
207 modified version of established methods (Dittmar et al., 2009; Lindroth and Mopper, 1979).  
208 Separation of 13 amino acids with a C18 column (Phenomenex Kinetex, 2.6 µm, 150 x 4.6 mm)  
209 was obtained after in-line derivatization with o-phthalaldehyde and mercaptoethanol. A  
210 gradient with solvent A containing 5 % acetonitrile (LiChrosolv, Merck, HPLC gradient grade)  
211 in sodium dihydrogenphosphate (Suprapur®, Merck) buffer (pH 7.0) and solvent B being  
212 acetonitrile was used for analysis. A gradient from 100% solvent A to 78% solvent A was  
213 produced in 50 min. Two replicates per TAA sample were analyzed.

### 214 **2.2.3. Transparent exopolymer particles**

215 Samples for transparent exopolymer particles (TEP) were taken at t<sub>0</sub>, t<sub>24h</sub> and t<sub>72h</sub> at all  
216 stations. The abundance and area of TEP were microscopically measured following the  
217 procedure given in Engel (2009). Samples of 10-50 mL were directly filtered under low vacuum  
218 (< 200 mbar) onto a 0.4 µm Nucleopore membrane (Whatman©) filter, stained with 1 mL Alcian  
219 Blue solution (0.2 g l<sup>-1</sup> w/v) for 3 s and rinsed with MilliQ water. Filters were mounted on  
220 Cytoclear© slides and stored at -20 °C until analysis. Two filters per sample with 30 images each  
221 were analyzed using a Zeiss Axio Scope.A1 (Zeiss©) and an AxioCam MRc (Zeiss©). The  
222 pictures with a resolution of 1388 x 1040 pixels were saved using AxioVision LE64 Rel. 4.8  
223 (Zeiss©). All particles larger than 0.2 µm<sup>2</sup> were analyzed. ImageJ© and R were subsequently  
224 used for image analysis (Schneider, Rasband and Eliceiri 2012, R Core Team, 2014). Filters  
225 prepared with 10 mL MilliQ water instead of samples served as a blank. The carbon content of  
226 TEP (TEP-C) was estimated after Mari (1999) using the size-dependent relationship:



$$227 \quad TEP-C = a \sum_i n_i r_i^D \quad (1)$$

228 with  $n_i$  being the number of TEP in the size class  $i$  and  $r_i$  being the mean equivalent spherical  
229 radius of the size class. The constant  $a = 0.25 * 10^{-6}$  ( $\mu\text{g C}$ ) and the fractal dimension of  
230 aggregates  $D = 2.55$  were used as proposed by Mari (1999). To relate to organic carbon  
231 concentration in seawater, data for TEP-C are given as  $\mu\text{mol L}^{-1}$ .

## 232 **2.3. Processes**

### 233 **2.3.1. Dissolved and particulate $^{14}\text{C}$ incorporation rates**

234 The photosynthetic production of particulate ( $< 0.2\text{-}2 \mu\text{m}$  and  $> 2 \mu\text{m}$  size fractions) and  
235 dissolved organic matter was determined from samples taken at  $t_0$ ,  $t_{24\text{h}}$ ,  $t_{48\text{h}}$  and  $t_{72\text{h}}$  (or  $t_{96\text{h}}$   
236 at station FAST) with the  $^{14}\text{C}$ -uptake technique. From each tank, four polystyrene bottles (70  
237 mL; three light and one dark bottles) were filled with sampled seawater and amended with 40  
238  $\mu\text{Ci}$  of  $\text{NaH}^{14}\text{CO}_3$ . Bottles were incubated for 8 h in two extra 300 L tanks maintained under  
239 similar light and temperature regimes than in the experimental tanks (ambient temperature for  
240 C1, C2, D1 and D2 and ambient temperature + 3 °C for G1 and G2). Incubations were  
241 terminated by sequential filtration of the sample through polycarbonate filters (pore sizes 2  $\mu\text{m}$   
242 and 0.2  $\mu\text{m}$ , 47 mm diameter) using low-pressure vacuum. Filters were exposed for 12 h to  
243 concentrated HCl fumes to remove non-fixed, inorganic  $^{14}\text{C}$ , and then transferred to 4 mL plastic  
244 scintillation vials to which 3.5 mL of scintillation cocktail (Ultima Gold XR, Perkin Elmer©)  
245 were added. For the measurement of dissolved primary production, a 5 mL aliquot of each  
246 sampling bottle was filtered, at the end of incubation, through a 0.2  $\mu\text{m}$  polycarbonate filter (25  
247 mm diameter). This filtration was conducted, under low-pressure vacuum, in a circular filtration



248 manifold that allows the recovery of the filtrate into 20 mL scintillation vials. The filtrates were  
249 acidified with 200  $\mu\text{L}$  of 50% HCl and maintained in an orbital shaker for 12 h. Finally, 15 mL  
250 of liquid scintillation cocktail was added to each sample. All filter and filtrate samples were  
251 measured onboard in a liquid scintillation counter (Packard© 1600 TR).  $^{14}\text{C}$ -based production  
252 rates (PP; in  $\mu\text{g C L}^{-1} \text{ h}^{-1}$ ) were calculated as:

$$253 \quad \text{PP} = C_{\text{T}} \times \left( \frac{\text{DPM}_{\text{sample}} - \text{DPM}_{\text{dark}}}{\text{DPM}_{\text{added}} \times t} \right) \quad (2)$$

254 where  $C_{\text{T}}$  is the concentration of total dissolved inorganic carbon ( $\mu\text{g C L}^{-1}$ ),  $\text{DPM}_{\text{sample}}$  and  
255  $\text{DPM}_{\text{dark}}$  are the radioactivity counts in the light and dark bottle, respectively,  $\text{DPM}_{\text{added}}$  is the  
256 radioactivity added to each sample, and  $t$  is the incubation time (h).

257 The percentage extracellular release (PER%) was calculated as:

$$258 \quad \text{PER}\% = \frac{\text{PPd}}{\text{PPd} + \text{PPp}} \times 100 \quad (3)$$

259 where PPd refers to  $^{14}\text{C}$ -based dissolved production and PPp refers to  $^{14}\text{C}$ -based particulate  
260 production (sum of  $< 2$  and  $> 2$   $\mu\text{m}$  size fractions).

### 261 **2.3.2. Integrated $^{13}\text{C}$ incorporation**

262 Addition of  $^{13}\text{C}$ -bicarbonate ( $\text{NaH}^{13}\text{CO}_3$  99%; Sigma-Aldrich©) was performed in each  
263 tank before  $t_0$  in order to increase the isotopic level ( $\delta^{13}\text{C}$  signature) of the dissolved inorganic  
264 carbon pool to ca. 350‰. We followed with time the evolution of the  $\delta^{13}\text{C}$  signature in dissolved  
265 inorganic carbon ( $\delta^{13}\text{C}-C_{\text{T}}$ ), dissolved organic carbon ( $\delta^{13}\text{C}-\text{DOC}$ ) and particulate organic carbon  
266 pools ( $\delta^{13}\text{C}-\text{POC}$ ). For the analysis of the actual  $\delta^{13}\text{C}-C_{\text{T}}$ , 60 mL of sampled seawater (at  $t-12\text{h}$ ,  
267  $t_0$ ,  $t_{12\text{h}}$ ,  $t_{24\text{h}}$ ,  $t_{48\text{h}}$  and  $t_{72\text{h}}$ ; +  $t_{96\text{h}}$  at station FAST) was gently transferred to glass vials



268 avoiding bubbles. Vials were sealed after being poisoned with 12  $\mu\text{L}$  saturated  $\text{HgCl}_2$  and stored  
269 upside-down at room temperature in the dark pending analysis. At the University of Leuven, a  
270 helium headspace (5 mL) was created in the vials and samples were acidified with 2 mL of  
271 phosphoric acid ( $\text{H}_3\text{PO}_4$ , 99%). Samples were left to equilibrate overnight to transfer all  $\text{C}_T$  to  
272 gaseous  $\text{CO}_2$ . Samples were injected in the carrier gas stream of an EA-IRMS (Thermo©  
273 EA1110 and Delta V Advantage), and data were calibrated with NBS-19 and LSVEC standards  
274 (Gillikin and Bouillon, 2007).

275 At the same frequency than for  $\delta^{13}\text{C}-\text{C}_T$ , samples for  $\delta^{13}\text{C}-\text{DOC}$  were filtered online (see  
276 above), transferred to 40 mL pre-cleaned borosilicate amber EPA vials with septa caps (PTFE-  
277 lined silicone) and stored in the dark pending analysis at the Ján Veizer Stable Isotope  
278 Laboratory (Ottawa, Canada).

279 At t-12h, t0, t12h, t24h, t48h and t72h (or t96h at station FAST), the  $\delta^{13}\text{C}-\text{POC}$  was  
280 obtained based on the same measurements as described above for POC, on a an elemental  
281 analyzer coupled with an isotope ratio mass spectrometer (EA-IRMS; Vario Pyrocube-Isoprime  
282 100, Elementar©).

283 Carbon isotope data are expressed in the delta notation ( $\delta$ ) relative to Vienna Pee Dee  
284 Belemnite (VPDB) standard. The carbon isotope ratio was calculated as:

$$285 \quad R_{\text{sample}} = \left( \frac{\delta^{13}\text{C}_{\text{sample}}}{1000} + 1 \right) \times R_{\text{VPDB}} \quad (4)$$

286 with  $R_{\text{VPDB}} = 0.011237$ .



### 287 **2.3.2. Community metabolism (oxygen light-dark method)**

288 At the same frequency as for  $^{14}\text{C}$  incorporation, from each tank, a volume of 2 L was  
289 sampled in plastic bottles and distributed in 15 biological oxygen demand (BOD; 60 mL)  
290 borosilicate bottles. Five BOD bottles were immediately fixed with Winkler reagents (initial  $\text{O}_2$   
291 concentrations), five BOD bottles were incubated in the dark for the measurement of community  
292 respiration (CR) in two incubators maintained respectively at ambient temperature for C1, C2,  
293 D1 and D2 and at ambient temperature + 3 °C for G1 and G2. Additionally, five BOD bottles  
294 were incubated for the measurement of net community production (NCP) in the same tanks as  
295 described above for  $^{14}\text{C}$ -incorporation. Upon completion of the incubations (24 h), samples were  
296 fixed with Winkler reagents. Within one day,  $\text{O}_2$  concentrations were measured using an  
297 automated Winkler titration technique with potentiometric endpoint detection. Analyses were  
298 performed on board with a Metrohm© Titrand 888 and a redox electrode (Metrohm© Au  
299 electrode). Reagents and standardizations were similar to those described by Knap et al. (1996).  
300 NCP and CR were estimated by regressing  $\text{O}_2$  values against time, and CR was expressed as  
301 negative values. Gross primary production (GPP) was calculated as the difference between NCP  
302 and CR. The combined standard errors were calculated as:

$$303 \quad \text{SE}_{xy} = \sqrt{\text{SE}_x^2 + \text{SE}_y^2} \quad (5)$$



### 304 **2.3.4. Heterotrophic prokaryotic production and** 305 **ectoenzymatic activities**

306 At all sampling times, heterotrophic bacterial production (BP, *sensu stricto* referring  
307 to heterotrophic prokaryotic production) was determined onboard using the microcentrifuge  
308 method with the <sup>3</sup>H- leucine (<sup>3</sup>H-Leu) incorporation technique to measure protein production  
309 (Smith and Azam, 1992). The detailed protocol is in Van Wambeke et al. (2020b). Briefly,  
310 triplicate 1.5 mL samples and one blank were incubated in the dark for 1-2 h in two  
311 thermostated incubators maintained respectively at ambient temperature for C1, C2, D1 and  
312 D2 and at ambient temperature +3 °C for G1 and G2. Incubations were ended by the addition  
313 of TCA to a final concentration of 5%, followed by three runs of centrifugation at 16000 g  
314 for 10 min. Pellets were rinsed with TCA 5% and ethanol 80%. A factor of 1.5 kg C mol  
315 leucine<sup>-1</sup> was used to convert the incorporation of leucine to carbon equivalents, assuming no  
316 isotopic dilution (Kirchman et al., 1993).

317 Ectoenzymatic activities were measured fluorometrically, using fluorogenic model  
318 substrates that were L-leucine-7-amido-4-methyl-coumarin (Leu-MCA) and 4  
319 methylumbelliferyl – phosphate (MUF-P) to track aminopeptidase activity (LAP) and  
320 alkaline phosphatase activity (AP), respectively (Hoppe, 1983). Stocks solutions (5mM)  
321 were prepared in methycellosolve and stored at -20 °C. Release of the products of LAP and  
322 AP activities, MCA and MUF, were followed by measuring increase of fluorescence (exc/em  
323 380/440 nm for MCA and 365/450 nm for MUF, wavelength width 5 nm) in a  
324 VARIOSCAN LUXmicroplate reader calibrated with standards of MCA and MUF solutions.  
325 For measurements, 2 mL of unfiltered samples from the tanks were supplemented with 100



326  $\mu\text{L}$  of a fluorogenic substrate solution diluted so that different concentrations were  
327 dispatched in a black 24-well polystyrene plate in duplicate (0.025, 0.05, 0.1, 0.25, 0.5, 1  $\mu\text{M}$   
328 for MUF-P, 0.5, 1, 5, 10, 25  $\mu\text{M}$  for MCA-leu). Incubations were carried out in the same  
329 thermostatically controlled incubators than those used for BP and reproducing temperature  
330 levels in the experimental tanks. Incubations lasted up to 12 h long with a reading of  
331 fluorescence every 1 to 2 h, depending on the intended activities. The rate was calculated  
332 from the linear part of the fluorescence versus time relationship. Boiled-water blanks were  
333 run to check for abiotic activity. From varying velocities obtained, we determined the  
334 parameters  $V_m$  (maximum hydrolysis velocity) and  $K_m$  (Michaelis-Menten constant which  
335 reflects enzyme affinity for the substrate) by fitting the data using a non-linear regression on  
336 the following equation:

$$337 \quad V = V_m \times \frac{S}{K_m + S} \quad (6)$$

338 where  $V$  is the hydrolysis rate and  $S$  the fluorogenic substrate concentration added.

### 339 **2.3.5. Inorganic and organic material export**

340 At the end of each experiment ( $t_{72\text{h}}$  for TYR and ION and  $t_{96\text{h}}$  for FAST, after artificial  
341 dust seeding), the sediment traps were removed, closed and stored with formaldehyde 4%. Back  
342 in the laboratory, after the swimmers were removed, the samples were rinsed to remove the salts  
343 and then freeze-dried. The total amount of material collected was first weighted to measure the  
344 total exported flux. Several aliquots were then weighted to measure the following components:  
345 total carbon and organic carbon, lithogenic and biogenic silicates and calcium. Total carbon was  
346 measured on an elemental analyzer coupled with an isotope ratio mass spectrometer (EA-IRMS;



347 Vario Pyrocube-Isoprime 100, Elementar©). Particulate organic carbon (POC) was measured in  
348 the same way after removing inorganic carbon by acidification with HCl 2N. Particulate  
349 inorganic carbon (PIC) was obtained by subtracting particulate organic carbon from particulate  
350 total carbon. Calcium concentrations were measured by ICP-OES (Inductively Coupled Plasma -  
351 Optic Emission Spectrometry; Perkin-Elmer© Optima 8000) on acid digested samples (the  
352 organic matrix was removed by HNO<sub>3</sub> while the mineral aluminosilicate matrix was eliminated  
353 with HF). Biogenic silica (BSi) and Lithogenic silica (LSi) were measured by colorimetry  
354 (Analytikjena© Specor 250 plus spectrophotometer) after a NaOH/HF digestion, respectively  
355 (Mosseri et al., 2005). The carbonate fraction of the exported material was determined from  
356 particulate calcium concentrations (%CaCO<sub>3</sub> = 5/2 x (%Ca). The organic matter fraction was  
357 calculated as 2 x (%POC). The lithogenic fraction was calculated as [total mass – (organic matter  
358 + CaCO<sub>3</sub> + opal) and was very comparable to the lithogenic fraction calculated from LSi (taking  
359 Si concentration in dust analog used for seeding from Desboeufs et al., 2014; ca. 11.9%). In the  
360 controls, the amount of material exported was low and the entire content of the traps was filtered  
361 in order to measure total mass and organic matter mass fluxes.

362

## 363 **2.4. Data processing**

364 All metabolic rates were integrated over the duration of the experiments using trapezoidal  
365 integrations and the relative changes (in %) in tanks D and G as compared to the controls  
366 (average between C1 and C2) were computed following:

$$367 \text{ Relative change} = \left( \frac{\text{Rate}_{\text{Treatment}} - \text{Rate}_{\text{Controls}}}{\text{Rate}_{\text{Controls}}} \right) \times 100 \quad (7)$$

368 Where Rate<sub>Treatment</sub> is the integrated rate measured in treatments D and G (D1, D2, G1 or G2) and  
369 Rate<sub>Controls</sub> is the averaged integrated rates between the duplicate controls (treatment C). Daily



370 rates of  $^{14}\text{C}$ -based production were computed from hourly rates assuming a 14 h daylight period.  
371 As incubations performed from samples taken at  $t_0$  (before dust addition) do not represent what  
372 happened in the tanks between  $t_0$  and  $t_{24\text{h}}$ , as a first assumption, we considered a linear  
373 evolution between these rates and those measured from samples at  $t_{24\text{h}}$ , and recomputed an  
374 average value for the time interval  $t_0 - t_{24\text{h}}$ . At FAST, no incubations were performed for  $^{14}\text{C}$   
375 incorporation and oxygen metabolism between  $t_{72\text{h}}$  and  $t_{96\text{h}}$ , again an average rate between  
376 rates measured from samples taken at  $t_{48\text{h}}$  and  $t_{96\text{h}}$  was used for this time interval. Since  
377 bacterial respiration rates were not measured, bacterial growth efficiency (BGE, expressed as a  
378 percentage) was estimated based on BP (carbon units) and community respiration (CR, oxygen  
379 units). As BP was determined more often than CR during the first 48 h, hourly BP rates were  
380 integrated using trapezoidal integrations during the time period when CR was measured. We  
381 assumed that heterotrophic prokaryotes were responsible for 70% of CR (BR/CR ratio; Lemée et  
382 al., 2002) and used a respiratory quotient (RQ) of 0.8 (del Giorgio and Williams, 2005),  
383 following the equation:

$$384 \quad \text{BGE} = \left( \frac{\text{BP}}{\text{CR} \times \frac{\text{BR}}{\text{CR}} \text{ratio} \times \text{RQ} + \text{BP}} \right) \times 100 \quad (8)$$

385 When BP varied following an exponential growth, we calculated growth rates ( $\mu_{\text{BP}}$ ) from linear  
386 least square regression of  $\ln$  BP rates versus time.



## 387 **3. Results**

### 388 **3.1. Initial conditions**

389 Initial conditions in terms of the chemical and biological standing stocks measured while  
390 filling the tanks at the three stations are fully described in Gazeau et al. (2020). Briefly, the three  
391 experiments were conducted with surface seawater collected during stratified oligotrophic  
392 conditions typical of the open Mediterranean Sea at this period of the year (Table 1). Nitrate +  
393 nitrite ( $\text{NO}_x$ ) concentrations were maximal at station FAST with a  $\text{NO}_x$  to dissolved inorganic  
394 phosphate (DIP) molar ratio of  $\sim 4.6$ . Very low  $\text{NO}_x$  concentrations were observed at stations  
395 TYR and ION ( $14$  and  $18 \text{ nmol L}^{-1}$ , respectively). DIP concentrations were the highest at station  
396 TYR ( $17 \text{ nmol L}^{-1}$ ) and the lowest at the most eastern station (ION,  $7 \text{ nmol L}^{-1}$ ). Consequently,  
397 the lowest  $\text{NO}_x$ :DIP ratio was measured at TYR ( $0.8$ ), compared to ION and FAST ( $2.8$  and  $4.6$ ,  
398 respectively). Silicate ( $\text{Si}(\text{OH})_4$ ) concentrations were similar at TYR and ION ( $\sim 1 \mu\text{mol L}^{-1}$ ) and  
399 the lowest at FAST ( $\sim 0.6 \mu\text{mol L}^{-1}$ ). Both POC and DOC concentrations were the highest at  
400 station TYR ( $12.9$  and  $72.2 \mu\text{mol L}^{-1}$ , respectively) and the lowest at FAST ( $6.0$  and  $69.6 \mu\text{mol}$   
401  $\text{L}^{-1}$ , respectively). Very low and similar concentrations of chlorophyll *a* were measured at the  
402 three stations ( $0.063$  -  $0.072 \mu\text{g L}^{-1}$ ). Phytoplankton communities at stations TYR and ION were  
403 dominated by Prymnesiophytes followed by Cyanobacteria, while, at station FAST, the  
404 phytoplanktonic community was clearly dominated by photosynthetic prokaryotes. At all three  
405 stations, the proportion of pigments representative of larger species was very small ( $< 5\%$ ;  
406 Gazeau et al., 2020). Heterotrophic prokaryotes were the most abundant at station FAST ( $6.15 \times$   
407  $10^5 \text{ cells mL}^{-1}$ ) and the least abundant at station ION ( $2.14 \times 10^5 \text{ cells mL}^{-1}$ ).



408 Relatively similar  $^{14}\text{C}$ -based particulate production rates were measured at the start of the  
409 experiments ( $t_0$ ) in the control tanks (C1 and C2) at station ION and FAST (ca.  $0.014 - 0.015 \mu\text{g}$   
410  $\text{C L}^{-1} \text{h}^{-1}$ ). At both stations, ca. 80% of the production was attributed to larger ( $> 2 \mu\text{m}$ ) cells and  
411 the percentage of extracellular release (%PER) did not exceed 45%. Lower rates were estimated  
412 at station TYR (total particulate production of  $0.08 \mu\text{g C L}^{-1} \text{h}^{-1}$ ) from which 87.5% was due to  
413 large cells  $> 2 \mu\text{m}$ . A larger amount of  $^{14}\text{C}$  incorporation was released as dissolved organic  
414 matter at station TYR compared to the two other stations (PER ca. 60%). Metabolic balance  
415 derived from oxygen measurements showed that, at all three stations, the community was net  
416 heterotrophic with a higher degree of heterotrophy at station TYR (NCP were  $-1.9, -0.2, -0.8$   
417  $\mu\text{mol O}_2 \text{L}^{-1} \text{d}^{-1}$  at TYR, ION and FAST, respectively, as measured in the controls from seawater  
418 sampled at  $t_0$ ). CR and GPP rates were respectively the highest and the lowest at station TYR  
419 compared to the other two stations. Finally, BP rates were the highest at station FAST ( $35.8 \text{ ng C}$   
420  $\text{L}^{-1} \text{h}^{-1}$ ), intermediate at ION ( $26.1 \text{ ng C L}^{-1} \text{h}^{-1}$ ) and the lowest at TYR ( $21.3 \text{ ng C L}^{-1} \text{h}^{-1}$ ).

### 421 **3.2. Changes in biological stocks**

422 DOC concentrations showed a general increasing trend during the three experiments and  
423 a large variability between duplicates (Fig. 2). This variability appeared as soon as 1 h after dust  
424 seeding ( $t_{1h}$ ) while the range of variation at  $t_0$  (before dust seeding) was rather moderate  
425 (difference between minimal and maximal values in all tanks of  $1.3, 6.2$  and  $4.3 \mu\text{mol C L}^{-1}$  at  
426 station TYR, ION and FAST, respectively). As a consequence of this variability, no clear impact  
427 of dust seeding (D) could be highlighted at station TYR and FAST. Indeed, DOC concentrations  
428 in the two duplicates (D1 and D2) were higher than values in the controls (C1 and C2) in only  
429 33% of the samples along the experiments (after dust seeding). In contrast, at station ION, DOC  
430 concentrations appeared impacted by dust seeding as higher concentrations were almost



431 systematically (83% of the time after dust seeding) measured for this treatment as compared to  
432 control tanks at the same time. At all stations, this impact was somewhat exacerbated under  
433 conditions of temperature and pH projected for 2100 (G1 and G2) as DOC concentrations were  
434 almost all the time higher in these tanks than in control tanks (83 - 100% of the samples after  
435 dust seeding, depending on the station).

436 Total hydrolysable carbohydrates and amino acids concentrations along the three  
437 experiments are shown in Fig. S2. TCHO concentrations were quite variable between tanks  
438 before dust seeding ( $t_0$ ; 649 - 954, 569 - 660 and 600 - 744  $\text{nmol L}^{-1}$  at station TYR, ION and  
439 FAST, respectively) and no visible impact of the treatments were visible at station TYR (TCHO  
440 tended to decrease everywhere). In contrast, at station ION and FAST, values in dust amended  
441 tanks increased and appeared higher than in control tanks towards the end of the experiments  
442 although the large variability between duplicates tended to mask this potential effect. An impact  
443 of dust seeding was much clearer for TAA concentrations that showed larger increases  
444 throughout the three experiments in tanks D1 and D2 as compared to control tanks, this effect  
445 being exacerbated for warmer and acidified tanks (G1 and G2). The ratio between TAA and  
446 DOC concentrations (Fig. 2) showed increasing trends in tanks D and G during all three  
447 experiments with a clear distinction between treatments at the end of the experiments ( $G > D >$   
448 C). The strongest increase was observed at station FAST in tanks G where final values were  
449 above 3%.

450 Particulate organic carbon (POC) concentrations strongly decreased at all stations  
451 between  $t-12\text{h}$  and  $t_0$ , this decrease being the largest at station TYR where concentrations  
452 dropped from 25.7 to 9.6 - 13.2  $\mu\text{mol C L}^{-1}$  (Fig. 3). After dust seeding, POC concentrations did  
453 not show clear temporal trends for the three experiments although a slight general increase could



454 be observed at station FAST. Furthermore, no impact of dust seeding and warming/acidification  
455 could be observed for this parameter. While concentrations of transparent exopolymer particles  
456 (TEP-C) were rather constant through time in control tanks at the three stations, a large increase  
457 was observed in dust-amended tanks (D and G) with TEP-C reaching values up to  $\sim 2 \mu\text{mol C L}^{-1}$   
458 in tank G1 at station TYR after 24 h (i.e.  $\sim 17\%$  of POC concentration, Fig. 3). In all cases except  
459 for tank G2 at station ION, TEP-C further decreased towards the end of the experiments although  
460 concentrations remained well above those observed in the controls. As the variability between  
461 duplicated tanks G was rather high, no impact of warming/acidification on TEP dynamics could  
462 be highlighted at the three stations.

### 463 **3.3. Changes in metabolic rates**

464  $^{14}\text{C}$ -based particulate production rates as measured during the different time intervals at  
465 the three stations were low in control tanks (maximal total particulate production of  $0.34 \mu\text{g L}^{-1}$   
466  $\text{h}^{-1}$  at station FAST) and did not show any particular temporal dynamics (Fig. 4). In these tanks,  
467 the vast majority of particulate production was attributed to cells above  $2 \mu\text{m}$  (65 - 89%). The  
468 percentage of extracellular release (%PER) was overall maximal at station TYR and minimal at  
469 station FAST with a tendency to decrease with time at the three stations although large variations  
470 were observed between duplicates.

471 Dust addition alone did not have any clear positive impact on all  $^{14}\text{C}$ -based rates at  
472 station TYR, with even an observable decrease in production rates from larger cells ( $> 2 \mu\text{m}$ )  
473 compared to the controls. In contrast, at this station, dust seeding under warmer and acidified  
474 conditions (tanks G) had a positive effect on particulate production rates, this effect being  
475 particularly visible for cells  $< 2 \mu\text{m}$  and to a lesser extent on dissolved production with a general



476 decrease of %PER. An important discrepancy between the duplicates of treatment G was  
477 observable at the end of the experiment with much larger rates measured in tank G2.

478 In contrast to station TYR, an enhancement effect of dust addition was clearly visible at  
479 station ION where all rates increased towards the end of this experiment reaching a maximal  
480 total particulate production of  $0.6 - 0.7 \mu\text{g L}^{-1} \text{h}^{-1}$  in tanks D1 and D2. Since this positive effect  
481 was similar between small and larger cells, dust addition alone had no effect on the partitioning  
482 of production at this station, with cells  $> 2 \mu\text{m}$  representing  $\sim 80\%$  of total production. Although  
483 being also positively impacted and increasing with time, dissolved production appeared less  
484 sensitive than particulate production leading to an overall decrease of %PER at this station  
485 following dust addition. These positive impacts of dust seeding on  $^{14}\text{C}$ -based particulate  
486 production rates were even more visible at this station under warmer and acidified conditions  
487 (tanks G) with maximal rates more than doubled compared to those measured under present  
488 conditions of temperature and pH ( $1.5 - 1.6 \mu\text{g L}^{-1} \text{h}^{-1}$ ). Dust seeding under warmer and acidified  
489 conditions had a slight impact on the partitioning of particulate production at this station with  
490 smaller cells benefiting the most from these conditions. %PER remained between 20 and 30%.

491 At station FAST, similarly to station ION, total particulate production rates were clearly  
492 enhanced by dust addition (tanks D) reaching maximal values during the incubation time interval  
493 t48 - 56h. No clear increase was observed for total particulate production on the next incubation  
494 (t96 - 120h) while production rates of cells larger than  $2 \mu\text{m}$  increased and rates of smaller cells  
495 decreased. However at FAST, in contrast to station ION, there was much less impact of  
496 warming/acidification on all measured rates although rates measured on smaller cells ( $< 2 \mu\text{m}$ )  
497 did not decrease at the end of the experiment as observed under present environmental



498 conditions. %PER under both present conditions of temperature and pH (tanks D) decreased  
499 during this experiment reaching values lower than in the controls and in tanks G.

500 The initial enrichment of the tanks in  $^{13}\text{C}$ -bicarbonate led to an increase in the  $^{13}\text{C}$   
501 signature of dissolved inorganic carbon ( $\delta^{13}\text{C-C}_T$ ) of above 300‰, with generally lower values  
502 measured in warmer and acidified tanks (G; Fig. S3). After this initial enrichment,  $\delta^{13}\text{C-C}_T$  levels  
503 decreased linearly in all tanks. At stations TYR and ION, the isotopic signature of dissolved  
504 organic carbon ( $\delta^{13}\text{C-DOC}$ ; Fig. S3) increased with time, although these increases were rather  
505 low and limited to  $\sim 4\%$  over the course of the experiments. In contrast to station TYR, at ION,  
506 an enhanced incorporation of  $^{13}\text{C}$  into DOC was visible after 24 h in tanks D and G in  
507 comparison to control tanks. A similar observation was done at station FAST, especially at the  
508 end of the experiment, although much more variability was observed at this station.

509 The incorporation of  $^{13}\text{C}$  onto particulate organic carbon ( $\delta^{13}\text{C-POC}$ ) is shown in Fig. 5.  
510 At all stations,  $\delta^{13}\text{C-POC}$  increased with time but reached lower enrichment levels at station  
511 TYR as compared to ION and FAST. At this station, incorporation rates appeared smaller in  
512 dust-amended tanks under present environmental conditions (tanks D). As for  $^{14}\text{C}$ -based  
513 production rates, an important discrepancy was observed between duplicates under future  
514 conditions of temperature and pH (tanks G) with much higher final  $\delta^{13}\text{C-POC}$  at the end of the  
515 experiment in tank G2. At station ION, enrichment levels obtained at the end of the experiment  
516 were more important in dust-amended tanks reaching maximal levels of 73‰ in tank G2 at t72h.  
517 This enhancement effect was even more visible at station FAST with maximal enrichment levels  
518 of 146‰ (tank D2 at t96h). Since no sampling occurred at t72h, these enrichment levels cannot  
519 be directly compared to what was measured at station TYR and ION. However, by interpolating



520 values at t72h assuming a linear increase between these time intervals, enrichment levels  
521 appeared similar although slightly higher for tanks D between station ION and FAST.  
522 NCP rates as measured using the O<sub>2</sub> light-dark method showed that, under control  
523 conditions, the communities remained the vast majority of the time throughout the three  
524 experiments in a net heterotrophic state (NCP < 0; Fig. 6). This was especially true at station  
525 TYR where the lowest NCP rates were measured. At this station, dust addition whether under  
526 present or future conditions of temperature and pH did not switch the community towards net  
527 autotrophy but even drove the community towards a stronger heterotrophy. This was related to  
528 the fact that while gross primary production rates were not positively impacted, community  
529 respiration increased in tanks D and G. At station ION, dust addition alone (tanks D) led to a  
530 switch from net heterotrophy to net autotrophy after two days of incubation due to a stronger  
531 positive effect of dust on GPP than on CR. Under future environmental conditions (tanks G), the  
532 same observation was made with higher NCP and GPP rates than in tanks D. CR rates reacted  
533 quickly to these forcing factors in tanks G and initially (first incubation) drove the community  
534 towards a much stronger heterotrophy as compared to the other tanks. Finally, at station FAST,  
535 similarly to what was observed at ION, the community became autotrophic after two days of  
536 incubation in dust amended tanks as, although both GPP and CR were positively impacted by  
537 dust addition, this impact was less important for CR. Warming and acidification had a limiting  
538 impact on this enhancement, with a lower final NCP in tanks G compared to tanks D, a  
539 difference that can be related to an absence of effects of these environmental stressors on GPP  
540 while CR clearly increased at higher temperature and lower pH.

541 While BP remained constant or gradually increased in control tanks depending on the  
542 station, a clear and quick fertilization effect was observable following dust addition (treatment D



543 and G) at all stations (Fig. 7). At station TYR, BP rates sharply increased to reach maximal  
544 values at t24h, with an even stronger increase observed under warmer and acidified conditions  
545 (tanks G). After this initial increase, rates slightly decreased towards the end of the experiment.  
546 This fertilization effect appeared less important at station ION where lower maximal rates were  
547 obtained after 24 h as compared to station TYR. Nevertheless, the same observations can be  
548 made, namely, 1) higher rates were measured under future temperature and pH levels and 2) after  
549 this initial sharp increase, rates gradually decreased towards the end of the experiment especially  
550 in tanks G. At station FAST, a much stronger effect of warming/acidification was observed with  
551 an important increase of BP in tanks G until 24 or 48 h post-seeding, depending on the duplicate.  
552 A sharp decline was observed for this treatment until the end of the experiment although rates  
553 remained higher than those measured in tanks C and D. The impact of dust addition under  
554 present environmental conditions (tanks D) was somehow more limited than at the other stations  
555 with a gradual increase until t72h with maximal rates ~ 40 - 100% higher than rates measured in  
556 the controls. However, BP increased exponentially between t0 and t12h in all tanks including  
557 controls, and in all experiments (Table 2). The growth rate of BP ( $\mu_{BP}$ ) in control tanks was the  
558 highest at TYR, intermediate at ION and the lowest at FAST.  $\mu_{BP}$  increased significantly in all  
559 dust amended tanks compared to controls. Under future environmental scenarios,  $\mu_{BP}$  tended to  
560 increase compared to treatment D but with a variable relative change.

561 BGE increased in dust amended tanks under present environmental conditions (treatment  
562 D) at TYR and ION, while no changes were detectable at station FAST due to a strong  
563 discrepancy between control duplicates and overall higher BGE at this station in the controls  
564 (Table 3). In contrast, warming and acidification exerted the strongest effect at station FAST  
565 with a doubling of BGE between treatment G and D. Although an increase in BGE was also



566 observed at the two other stations in treatment G as compared to present environmental  
567 conditions (treatment D), this increase was more limited (ca. 1 to 1.4-fold increase).

568         The alkaline phosphatase Vm (AP Vm) increased in all experiments after dust seeding,  
569 with amplified effects in G treatments (Fig. S4). Note that AP Vm increased also in the controls  
570 at TYR and FAST. In contrast, leucine aminopeptidase Vm (LAP vm) showed succession of  
571 peaks instead of continuously increasing (Fig. S4). It was higher in dust alone treatment (D) as  
572 compared to the controls at TYR and FAST. A larger variability between duplicates at ION  
573 prevents such an observation. At all stations, maximum velocities were measured under future  
574 environmental conditions (G). Vm being possibly influenced by enzyme synthesis but also by the  
575 number of cells inducing such enzymes, we computed also specific AP Vm per heterotrophic  
576 bacterial cell (Fig. 7). Specific AP Vm slightly increased during all experiments in controls and  
577 dust-amended tanks (D) with no visible differences between these treatments, a clear over-  
578 expression of this enzyme was observed under warmer and more acidified conditions (treatment  
579 G) especially at station FAST where velocities were enhanced by a ~8-fold at t96h.

### 580         **3.4. Inorganic and organic material export**

581         Both total mass and organic matter fluxes, as measured from analyses of the sediment  
582 traps at the end of each experiment, were extremely low under control conditions (Fig. 8).  
583 Additions of dust in tanks D and G led to a strong increase in both fluxes with a large variability  
584 between the duplicates of treatment D at ION. No clear changes between tanks maintained under  
585 present and future conditions of temperature and pH could be highlighted.



## 586 4. Discussion

### 587 4.1. Initial conditions of the tested waters and evolution in 588 controls

589 As discussed in the companion paper from Gazeau et al. (2020), the three sampling  
590 stations were typical of stratified oligotrophic conditions encountered in the open Mediterranean  
591 Sea in late spring / early summer. DOC concentrations at the start of the experiments were in the  
592 same range ( $60 - 75 \mu\text{mol C L}^{-1}$ ) as those measured from samples collected in surface waters  
593 using clean sampling procedures (Van Wambeke et al., 2020b), revealing no contamination  
594 issues from our sampling device. TAA concentrations as measured in the tanks at  $t_0$  were also  
595 consistent with measurements from surface water samples (Van Wambeke et al., 2020b) with an  
596 average across stations and treatments of  $254 \pm 36 \text{ nmol L}^{-1}$  (Fig. S2). In contrast, TCHO  
597 appeared higher at  $t_0$  (average across stations and treatments of  $681 \pm 98 \text{ nmol L}^{-1}$ ) than  
598 concentrations based on clean *in situ* sampling (average of  $595 \pm 43 \text{ nmol L}^{-1}$ ; Van Wambeke et  
599 al., 2020b). The decrease in POC concentrations between pumping ( $t-12\text{h}$ ) and  $t_0$  for the three  
600 experiments, especially at station TYR (likely linked to higher initial concentrations), was likely  
601 a consequence of sedimentation of senescent cells and/or fecal pellets in our experimental  
602 systems, which are designed to evaluate the export of matter thanks to their conical shape. TEP  
603 concentrations were not quantified at  $t-12\text{h}$  and therefore there is no possibility to evaluate if  
604 sedimentation of these particles occurred before  $t_0$  in our tanks. At  $t_0$ , larger and more abundant  
605 TEP were measured at station TYR compared to the two other stations (data not shown) leading  
606 to a larger contribution of TEP carbon content (TEP-C) to POC concentrations (Fig. 3).



607 As a consequence of a very low availability in inorganic nutrients, TChl $a$  and  $^{14}\text{C}$ -based  
608 production rates were very low, all typical of oligotrophic conditions. Nano- and micro-  
609 phytoplanktonic cells ( $> 2 \mu\text{m}$ ) contributed most of the  $^{14}\text{C}$ -based particulate production ( $\sim 80\%$ ),  
610 as found also on several on-deck incubations at the three stations (on average  $73 \pm 6\%$ ; Marañón  
611 et al., 2020). %PER values were also very similar to those measured during these on-deck  
612 incubations ( $\sim 40\text{-}45\%$ ; see Marañón et al., 2020). This suggests no significant impact of our  
613 experimental protocol on rates and partitioning of  $^{14}\text{C}$ -based production rates (i.e. sampling from  
614 the continuous seawater supply, delay of 12 h before initial measurements, artificial light etc.).  
615 The low values of chlorophyll stocks as well as of  $^{14}\text{C}$ -based production rates are consistent with  
616 previous estimates based on direct measurements, satellite observations and modelling  
617 approaches in the same areas in late spring / early summer (e.g. Bosc et al., 2004; Lazzari et al.,  
618 2016; Moutin and Raimbault, 2002).

619 The metabolic balance was in favor of net heterotrophy at all stations at the start of the  
620 experiments (NCP  $< 0$ ). Net heterotrophy in the open Mediterranean sea at this period of the year  
621 has been reported by Regaudie-de-Gioux et al. (2009) and Christaki et al. (2011) in agreement  
622 with our measurements at  $t_0$  in control tanks (Table 1). The lowest NCP and the highest CR rates  
623 were measured at station TYR, suggesting that the autotrophic plankton community was not very  
624 active at this station. This was confirmed by the  $^{14}\text{C}$ -based particulate production rates, which  
625 were about half the ones measured at the other two stations. The community at TYR was most  
626 likely relying on regenerated nutrients, as shown by the highest levels of ammonium ( $\text{NH}_4^+$ )  
627 measured at the start of this experiment (Gazeau et al., 2020). As discussed in Guieu et al.  
628 (2020), a dust deposition event took place several days before the arrival of the vessel in this  
629 area, likely on May 10-12. This dust event was confirmed by inventory of particulate aluminium



630 in the water column at several stations of the Tyrrhenian Sea including TYR, 6 to 9 d after the  
631 event (Matthieu Bressac, pers. comm.). This dust deposition likely stimulated phytoplankton  
632 growth and POC accumulation shortly after the deposition and consequently increased the  
633 abundance of herbivorous grazers (copepods) and attracted carnivorous species (Feliú et al.,  
634 2020), subsequently driving the community towards a net heterotrophic state that characterized  
635 the initial condition of the experiment at this station. The optimal conditions for BP growth at  
636 this station were also confirmed by the highest  $\mu_{BP}$  growth rates obtained among the three  
637 experiments (Table 2; 0.06 - 0.07 h<sup>-1</sup>) in controls tanks.

638 The two other stations, although both also showing a slight net heterotrophic state, were  
639 clearly different from each other in terms of initial biological stocks and metabolic rates. Indeed,  
640 whereas TChla and abundances of pico- and nano-autotrophic cells (flow cytometry counts;  
641 Gazeau et al., 2020) were higher at FAST compared to ION, the autotrophic community was not  
642 more efficient at fixing carbon at this station, as shown by similar initial <sup>14</sup>C-based production  
643 rates. In contrast, both heterotrophic prokaryotic abundances and BP were much higher at station  
644 FAST as compared to ION, leading to initial higher CR and lower NCP. At ION, the initial NCP  
645 closer to metabolic balance further suggests a tight coupling between heterotrophic prokaryotes  
646 and phytoplankton at this station, as discussed by Dinasquet et al. (2021).

647 For most of the chemical and biological stocks (e.g. nutrients, pigments etc.) presented in  
648 Gazeau et al. (2020), no major changes took place during the three experiments under control  
649 conditions. Here, we further show that DOC, POC as well as TEP concentrations did not exhibit  
650 strong changes during the experiments. For DOC, large variability between the duplicates (C1  
651 and C2) potentially masked an increase towards the end of the experiments. The same holds true  
652 for autotrophic metabolic rates, as <sup>14</sup>C-based particulate production rates showed no marked



653 variations during the three experiments, although a slight increase was visible at FAST until  
654 t48h. The communities at the three stations remained heterotrophic under the nutrient-limited  
655 conditions in the controls. However, heterotrophic prokaryotes probably benefited from initial  
656 inputs of available organic matter issued from other stressed eukaryotic organisms and/or POC  
657 decay between t-12h and t0, which could be due to both sedimentation and degradation. This was  
658 reflected in the progressive increase of BP, their variable initial growth rates ( $\mu_{BP}$  ranged from  
659 0.02 to 0.06 h<sup>-1</sup> in control tanks according to the experiment) as well as increasing TAA/DOC  
660 ratios at the three stations. Finally, an initial increase of BP during incubations is generally  
661 described and classically attributed to a bottle effect, which favours large, fast-growing bacteria  
662 and often induces mortality of some phytoplankton cells (Calvo-Díaz et al., 2011; Ferguson et  
663 al., 1984; Zobell and Anderson, 1936)

## 664 **4.2. Impact of dust addition under present environmental** 665 **conditions**

666 The addition of nitrogen and phosphorus in the experimental tanks through dust seeding  
667 (+ 11 to + 11.6  $\mu\text{mol L}^{-1}$  and + 22 to + 30.8  $\text{nmol L}^{-1}$  for  $\text{NO}_x$  and DIP, respectively, in dust  
668 enriched, i.e. D1 and D2, versus controls; Gazeau et al., 2020) had very contrasting impacts on  
669 the metabolism of the communities, depending on the station. At TYR, surprisingly, the relieving  
670 of nutrient limitation had a negative impact on <sup>13</sup>C incorporation as well as on both particulate  
671 and dissolved <sup>14</sup>C-based production rates (as seen by the relative changes compared to the  
672 control presented in Fig. 9). These observations are fully corroborated by the observed relative  
673 decrease in GPP in these tanks (D1 and D2) relative to controls and by the negative impact of  
674 dust-addition on TChla concentrations as discussed by Gazeau et al. (2020). Integrated <sup>14</sup>C-



675 incorporation rates converted to P (using a C:P molar ratio of 245:1 determined in the particulate  
676 organic matter in surface waters of the Northwestern Mediterranean Sea during stratification;  
677 Tanaka et al., 2011) showed that phytoplankton P requirements in treatment D ( $\sim 2 \text{ nmol P L}^{-1}$ )  
678 were much lower than the release of DIP through dust addition at this station (+ 20.4 to + 24.6  
679  $\text{nmol P L}^{-1}$ ; Gazeau et al., 2020). This suggests that the observed strong decrease of DIP at this  
680 station following dust addition was due to an utilization by the heterotrophic compartment.  
681 Indeed, in contrast to the autotrophic compartment, both heterotrophic prokaryotic abundances  
682 (Gazeau et al., 2020) and BP (this study, Fig. 9) showed that heterotrophic prokaryotes reacted  
683 quickly and strongly to the increase in DIP availability. Integrated BP increased by almost 400%  
684 in tanks D1 and D2 as compared to controls (Fig. 9). Such relative increases of BP surpassing by  
685 far the observed relative increases of CR suggest a much more efficient utilization of resources  
686 by heterotrophic prokaryotes in this treatment (i.e. BGE increased by 200% as compared to the  
687 controls; Fig. 9). As such, at this station, the addition of dust drove the community to an even  
688 more heterotrophic state. Such absence of response of the autotrophic community despite the  
689 input of new N and P from simulated wet deposition was never observed in dust enrichment  
690 experiments performed in the Mediterranean Sea (Guieu and Ridame, 2020). To the best of our  
691 knowledge, it is the first time that a negative effect of dust addition is experimentally  
692 demonstrated on the metabolic balance. The apparent utilization of nutrients, especially DIP  
693 (Gazeau et al., 2020), by heterotrophic prokaryotes was extremely fast, starting right after dust  
694 addition and driving DIP concentrations back to control levels at the end of the experiment  
695 (t72h). While heterotrophic prokaryotic abundances increased until the end of the experiment,  
696 BP rates increased exponentially during the first 24h, and then BP reached a plateau.  
697 Heterotrophic prokaryotes appeared limited by nutritive resources although DIP concentrations



698 were not yet back to their initial level and no relative increase of the AP Vm per cell compared to  
699 the control was observed in these tanks. Independent nutrient experiments showed a direct  
700 stimulation of BP in the dark after addition of DIP (Van Wambeke et al., 2020b), suggesting a  
701 great competition with phytoplankton for DIP utilization at TYR. After 24 h, abundances of  
702 heterotrophic prokaryotes continued to increase while BP stabilized, suggesting a less extent of  
703 lysis and viral control than in the other experiments (abundances of heterotrophic nanoflagellates  
704 decreased; Dinasquet et al., 2021). This limitation of BP was potentially a consequence of  
705 relatively less available access to labile DOC sources, as <sup>14</sup>C-based production rates decreased  
706 relative to the controls at t24h and t48h although BP increased by 200 - 800%. The very tight  
707 coupling between phytoplankton and bacteria at all stations investigated was further confirmed  
708 by the absence of an important <sup>13</sup>C incorporation into DOC (Fig. S3).

709 At stations ION and FAST, in contrast to TYR, both the autotrophic and heterotrophic  
710 community benefited from dust addition relative to the controls (Fig. 9). Interestingly, while the  
711 relative increase in integrated autotrophic processes (GPP and all <sup>14</sup>C-based production rates)  
712 was more important at FAST than at ION, the opposite was observed for BP. Estimated BGE  
713 values even suggest an absence of response to dust addition at station FAST compared to the  
714 controls. The different (relative) responses of BP at the two stations could be partly explained by  
715 the dynamics of BP in the control tanks as no clear pattern could be observed at ION while a  
716 continuous increase was observed at FAST. As shown by Gazeau et al. (2020), at FAST,  
717 abundances of heterotrophic prokaryotes were much higher at the start of the experiment, further  
718 increased until t48h and then declined until the end of the experiment.

719 We can rule out a potential limitation of BP from DIP availability at station FAST as DIP  
720 levels remained much higher in tanks D than in the controls (Gazeau et al., 2020). Furthermore,



721 the amount of maximum DIP reached before its decline compared to TYR and ION showed a  
722 less important direct DIP uptake, suggesting that communities were not as much P limited at  
723 FAST compared to the other stations at the start of the experiment. Finally, no increase of  
724 specific AP Vm was observed in these tanks as compared to the controls (Fig. 7), suggesting no  
725 particular additional needs for AP synthesis per unit cell following dust addition. A potential  
726 explanation resides in the competition between heterotrophic bacteria and phytoplankton for DIP  
727 utilization. At station ION, P requirements of the autotrophic community were low compared to  
728 the initial input of DIP following dust seeding ( $\sim 9 \text{ nmol P L}^{-1}$  as compared to an input of + 22 to  
729 + 23.3  $\text{nmol P L}^{-1}$ ; Gazeau et al., 2020). In contrast, at FAST, the autotrophic community  
730 consumed a much larger proportion of the initial DIP input ( $\sim 25 \text{ nmol P L}^{-1}$  as compared to an  
731 input of 30.8 - 31.3  $\text{nmol P L}^{-1}$ ) and phytoplankton appeared as a winner for the utilization of  
732 DIP towards the end of the experiment at this station. It seems that heterotrophic bacteria and  
733 phytoplankton were more in a steady state of equilibrium and less stressed at the start of the  
734 experiment at FAST, i.e. phytoplankton abundances showed no decrease between t-12h and t0  
735 and BP did not increase as much as during the other two experiments, suggesting a strong  
736 predation pressure ( $\mu_{BP}$  was the lowest of the three experiments: ca.  $0.02 \text{ h}^{-1}$  in the controls).

737 The explanation for the observed differential responses of the autotrophic community at  
738 the two stations (FAST > ION) is not evident and further complicated by the fact that the  
739 sampling strategy differed between the two stations (i.e. no sampling at t72h, replaced by a  
740 sampling at t96h). It is however unlikely that this different sampling strategy was responsible for  
741 the different changes in computed integrated autotrophic rates at the two stations. As a maximal  
742 increase in nano-eukaryote abundance was observed at t72h at FAST (followed by a drastic  
743 reduction at t96h; Gazeau et al., 2020), excluding this sampling point in the calculation of



744 autotrophic metabolic rates would most likely have led to an underestimation of these rates rather  
745 than an overestimation. Furthermore, a similar partitioning of  $^{14}\text{C}$ -based production rates  
746 throughout the two experiments did not provide clear insights on which size-group benefited the  
747 most at station FAST compared to ION. Two non-exclusive explanations could be proposed: (1)  
748 as mentioned above, a less important immediate consumption of DIP by heterotrophic bacteria  
749 leading to a higher availability of new DIP for phytoplankton growth at FAST (+ 31 vs + 22 to +  
750 23  $\text{nmol L}^{-1}$  at FAST and ION, respectively; Gazeau et al., 2020) along with (2) the presence of a  
751 potentially more active community at the start of the experiment at FAST with a much higher  
752 contribution from smaller cells (i.e. pico-eukaryotes, *Synechococcus*; Gazeau et al., 2020) that  
753 are well known to be better competitors for new nutrients and that were less stressed at the start  
754 of the experiments (e.g. Moutin et al., 2002).

755         During both experiments at ION and FAST, communities switched from net heterotrophy  
756 to net autotrophy between 48 and 72 h following dust addition (Fig. 6), leading to a positive  
757 integrated NCP at both stations (Fig. 9). This is an important observation since, to the best of our  
758 knowledge, the present study constitutes the first investigation of the community metabolism  
759 response to dust addition. However, it is important to discuss the timing of such a switch in  
760 community metabolism. Since heterotrophic prokaryotes reacted faster than autotrophs to the  
761 relief of nutrient limitation (i.e. BP already increased by 150-500% at t24 h, while  $^{14}\text{C}$ -based  
762 production rates increased only after 48-72 h), NCP was first lower (and negative) in the dust-  
763 amended tanks as compared to the controls. Marañón et al. (2010) and Pulido-Villena (2008,  
764 2014) have already reported on a much faster response of the heterotrophic prokaryote  
765 community to dust enrichment in the central Atlantic Ocean and Mediterranean Sea,  
766 respectively. As DIP concentrations at the completion of their 48 h incubations did not differ



767 from that in the controls, it is unlikely that primary production rates and consequently NCP  
768 would have further increased. In contrast, during our experiments, DIP concentrations in dust-  
769 amended tanks (D) reached initial levels only after 72 h at TYR and ION and remained far above  
770 ambient levels at FAST until the end of the experiment (t96h). During the PEACETIME cruise,  
771 high frequency sampling of CTD casts allowed following the evolution of biogeochemical  
772 properties and fluxes before and after wet dust deposition that took place in the area around  
773 FAST on June 3-5 (Van Wambeke et al., 2020a). As in our experiment, a rapid increase in BP  
774 was responsible for the observed *in situ* decline in DIP concentrations in the mixed layer  
775 following the rain with no detectable changes in primary production (Van Wambeke et al.,  
776 2020a). The intensity of the wet deposition event that was simulated during our experiments was,  
777 by far, more important, but still representative of a realistic scenario (Bonnet and Guieu, 2006;  
778 Loÿe-Pilot and Martin, 1996; Ternon et al., 2010).

779         The most intriguing result concerning the export of inorganic and organic matter is that  
780 these fluxes were maximal at the end of the experiment at TYR in the dust-amended tanks  
781 despite the fact that <sup>14</sup>C-based production was relatively low and not enhanced by dust addition.  
782 Based on previous studies (Bressac et al., 2014; Louis et al., 2017; Ternon et al., 2010), organic  
783 matter export was most likely mainly due to the formation of organic-mineral aggregates  
784 triggered by the introduced lithogenic particles (referred thereafter to as POC<sub>litho</sub>). Indeed, Louis  
785 et al. (2017) showed that such an aggregation process occurs within 1 h after dust deposition.  
786 These authors further demonstrated the key role of TEP as the conversion of dissolved organic  
787 matter (DOM) to POC was mediated by TEP formation/aggregation activated by the introduction  
788 of dust. As TEP concentrations were only measured on two occasions after seeding with the first  
789 measurement occurring at t24h, ), it prevents studying in detail the dynamics of these particles.



790 Nevertheless, it is very likely that the sharp decrease of TEP abundances (data not shown)  
791 between  $t_{24h}$  and  $t_{72h}$  was related to  $POC_{litho}$  export. The coefficient linking  $POC_{litho}$  to  $Litho_{flux}$   
792 (i.e. the mass of sedimented particles) measured here (0.02) is consistent with values reported for  
793 other experiments conducted in the Mediterranean Sea (Louis et al., 2017).

794 Even though  $^{14}C$ -based production rates were enhanced in the dust-amended tanks at  
795 stations ION and FAST, the amount of POC exported at the end of these experiments remained  
796 lower than at TYR, with fluxes  $\sim 10\text{-}20 \text{ mg C m}^{-2} \text{ d}^{-1}$ . It must be stressed that not all the  
797 lithogenic material introduced in the tanks was recovered after 4 (and 5) days, with the highest  
798 percentage ( $\sim 30\%$ ) being found at TYR, indicating that the tested waters at this station had a  
799 better capacity to aggregate dust. This efficiency to export  $POC_{litho}$  more rapidly at TYR  
800 compared to ION and FAST was likely due to the age and quantity of dissolved organic matter  
801 present at the time of the seeding (Bressac and Guieu, 2013). At TYR, impacted by a strong dust  
802 event several days before the experiment started (see above), the likely stimulation of the  
803 autotrophs after this *in situ* event should have been followed by the production of a fresh and  
804 abundant DOM, comparable to the “post-bloom situation” in Bressac and Guieu (2013).

### 805 **4.3. Impact of dust addition under future environmental** 806 **conditions**

807 Warming and/or acidification had a clear impact on most evaluated stocks and metabolic  
808 rates. Gazeau et al. (2020) have already discussed temperature/pH mediated changes in nutrient  
809 uptake rates and autotrophic community composition in these experiments. The difference in the  
810 relative response of plankton communities to dust addition under present and future conditions of  
811 temperature and pH was highly dependent on the sampling station (Fig. 9). At all stations,  $^{14}C$ -



812 based particulate production rates were enhanced under future conditions as compared to those  
813 measured under present environmental conditions (treatment D) although this pattern was not  
814 observed for  $^{13}\text{C}$  incorporation into POC at stations ION and FAST. At ION, no differences  
815 could be detected and at FAST an even lower  $^{13}\text{C}$ -enrichment was measured at the end of the  
816 experiment. These contrasting patterns between  $^{14}\text{C}$ -uptake rates and  $^{13}\text{C}$ -enrichment of POC are  
817 likely explained by the fact that the latter covered the whole experimental period (including dark  
818 periods) and represents net community carbon production while  $^{14}\text{C}$ -based rates were measured  
819 over 8 h incubations in the light, providing an estimate in between gross and net carbon  
820 production.

821 Similarly, the heterotrophic compartment was more stimulated, as BP rates increased  
822 strongly at all stations under this treatment compared to treatment D. The relatively smaller  
823 increase in CR rates, compared to BP, leading to higher BGE suggests a better utilization of  
824 resources by heterotrophic prokaryotes under future environmental conditions. Overall, CR was  
825 more impacted than GPP, with the consequence that all integrated NCP rates decreased under  
826 future environmental conditions compared to present conditions (treatment D). At station TYR,  
827 as discussed previously, dust addition under present conditions did not lead to a switch from net  
828 heterotrophy to net autotrophy. This pattern was even more obvious under warmer/acidified  
829 conditions, with a larger decrease in integrated NCP at this station. The decrease of integrated  
830 NCP at station FAST relative to controls, as well as the smaller increase of all  $^{14}\text{C}$ -based  
831 production rates relative to those observed at station ION must be taken with caution. As already  
832 discussed, the fact that for these processes ( $\text{O}_2$  metabolism and  $^{14}\text{C}$ -incorporation), no samples  
833 were taken at FAST at  $t_{72\text{h}}$  when maximal cell abundances were recorded for all autotrophic  
834 groups (pico- and nano-eukaryotes, autotrophic bacteria) must have artificially led to an



835 underestimation of these integrated metabolic rates. The question of the timing appeared even  
836 more preponderant under warmer/acidified conditions, especially at station FAST, where the  
837 very important increase in BP led to a full consumption of DIP before t48h (Gazeau et al., 2020)  
838 and drove the community towards a strong heterotrophy. The metabolic balance further switched  
839 to a slight autotrophy at t72h when heterotrophic bacterial activity appeared limited by nutrient  
840 availability.

841 Both elevated partial pressure of CO<sub>2</sub> ( $p\text{CO}_2$ ) and warming are major global change  
842 stressors impacting marine communities. Elevated  $p\text{CO}_2$  may directly facilitate oceanic primary  
843 production through enhanced photosynthesis (Hein and Sand-Jensen, 1997; Riebesell et al.,  
844 2007) although the effects appear to be species- and even strain-specific (e.g. Langer et al.,  
845 2009). Warming affects organisms by enhancing their metabolic rates (Brown et al., 2004;  
846 Gillooly et al., 2001). Although recent studies suggest large differences in temperature sensitivity  
847 between phytoplankton taxa (Chen and Laws, 2017) and no significant overall difference  
848 between algae and protozoa (Wang et al., 2019), mineralization rates are usually believed to be  
849 more impacted by warming than primary production rates, potentially leading to a decline in net  
850 oceanic carbon fixation (Boscolo-Galazzo et al., 2018; Garcia-Corral et al., 2017; Lopez-Urrutia  
851 and Moran, 2007; Regaudie-de-Gioux and Duarte, 2012) and carbon export efficiency (Cael et  
852 al., 2017; Cael and Follows, 2016). Overall, our experimental set-up did not allow discriminating  
853 warming from acidification effects, precluding an evaluation of their potential individual  
854 impacts. Nevertheless, we could speculate to which extent a 3 °C warming and a doubling of  
855 CO<sub>2</sub> can explain some of the observed differences between D and G (for instance, a 2-fold  
856 increase in <sup>14</sup>C-based production rates at ION). For photosynthesis, meta-analysis studies  
857 indicate minor effects of  $p\text{CO}_2$  on most investigated species (Kroeker et al., 2013; Mackey et al.,



858 2015). Recent studies show a strong, although species-dependent, temperature sensitivity of  
859 phytoplankton growth (Chen and Laws, 2017; Wang et al., 2019), suggesting that a 3 °C  
860 warming could explain most of the increased carbon fixation in G compared to D. With respect  
861 to NCP, our results are in line with the general view and suggest a weakening of the so-called  
862 fertilization effect of atmospheric deposition in the coming decades.

863 In contrast, we did not observe an additional impact of future environmental conditions  
864 on the export of organic matter after dust addition as, at each station, this export was of the same  
865 order of magnitude for treatments D and G. This result is in agreement with the findings of a  
866 similar experiment in coastal Mediterranean waters that considered only pH change (Louis et al.,  
867 2017) but stands in contrast with the findings of Müren et al. (2005) who showed a clear  
868 decrease in sedimentation following a 5 °C warming in the Baltic Sea. Only a few studies have  
869 addressed the combined effect of both temperature and pH changes on aggregation processes and  
870 export but none considered dust as the particulate phase. These studies, focused mainly on the  
871 formation of TEP, were inconclusive on the impact of these combined factors (Passow and  
872 Carlson, 2012, and references therein). As the potential effect of warming and acidification on  
873 biogenic carbon export was certainly, over the rather restricted duration of the experiments,  
874 insignificant as compared to the large amount of carbon exported through the lithogenic pump,  
875 observations over longer temporal scales are probably required to ascertain the interactive effects  
876 of these stressors in the coming decades.



## 877 **5. Conclusion**

878           Although the three experiments were conducted under rather similar conditions in terms  
879 of nutrient availability and chlorophyll stock of the tested seawater, contrasting responses were  
880 observed following the simulation of a wet dust deposition event. Under present conditions of  
881 temperature and pH, at the site where the community was the most heterotrophic (TYR), no  
882 positive impact of new nutrients could be observed on autotrophs, while a fast and strong  
883 response of heterotrophic bacteria drove the metabolic balance towards an even more  
884 heterotrophic state. The situation was different at the two other stations where a more active  
885 autotrophic community responded quickly to the relief in nutrient limitation, driving the  
886 community to an autotrophic state at the end of these experiments. In all tested waters, an overall  
887 faster response of the heterotrophic prokaryote community, as compared to the autotrophic  
888 community, was observed after new nutrients were released from dust. Phytoplankton could  
889 benefit from nutrient inputs, only if the amount released from dust was enough to sustain both  
890 the fast bacterial demand and the delayed one of phytoplankton. As our experimental protocol  
891 consisted in simulating a strong, although realistic, wet dust deposition, further work should  
892 explore at which flux a wet dust deposition triggers an enhancement of net community  
893 production and therefore increases the capacity of the surface oligotrophic ocean to sequester  
894 atmospheric CO<sub>2</sub>. This question, of the utmost importance in particular for modelling purposes,  
895 should be answered through future similar experiments as the ones considered in our study but  
896 following a gradient approach of dust fluxes. As a consequence of a stronger sensitivity of  
897 heterotrophic prokaryotes to temperature and/or pH, the ongoing warming and acidification of  
898 the surface ocean will result in a decrease of the dust fertilization of phytoplankton in the coming  
899 decades and a weakening the CO<sub>2</sub> sequestration capacity of the surface oligotrophic ocean.



## 900 **Data availability**

901 All data and metadata will be made available at the French INSU/CNRS LEFE CYBER database  
902 (scientific coordinator: Hervé Claustre; data manager, webmaster: Catherine Schmechtig).  
903 INSU/CNRS LEFE CYBER (2020)

## 904 **Author contributions**

905 FG and CG designed and supervised the study. All authors participated in sample analyses. FG  
906 wrote the paper with contributions from all authors.

## 907 **Financial support**

908 This study is a contribution to the PEACETIME project (<http://peacetime-project.org>), a joint  
909 initiative of the MERMEX and ChArMEX components supported by CNRS-INSU, IFREMER,  
910 CEA, and Météo-France as part of the programme MISTRALS coordinated by INSU.  
911 PEACETIME was endorsed as a process study by GEOTRACES and is a contribution to IMBER  
912 and SOLAS International programs. PEACETIME cruise (<https://doi.org/10.17600/17000300>).  
913 The project leading to this publication has received funding from European FEDER Fund under  
914 project 1166-39417. The research of EM and MPL was supported by the Spanish Ministry of  
915 Science, Innovation and Universities through project POLARIS (Grant No. PGC2018-094553B-  
916 I00) and by European Union's H2020 research and innovation programme through project  
917 TRIATLAS (Grant No. 817578). JD was funded by a Marie Curie Actions-International  
918 Outgoing Fellowship (PIOF-GA-2013-629378).

919



## 920 **Acknowledgments**

921 The authors thank the captain and the crew of the RV “Pourquoi Pas ?” for their professionalism  
922 and their work at sea. Céline Ridame and Kahina Djaoudi are thanked for their help during  
923 sampling, Sophie Guasco and Marc Garel for their help in ectoenzymatic measurements onboard.



## 924 **References**

- 925 Behrenfeld, M. J., O'Malley, R. T., Siegel, D. A., McClain, C. R., Sarmiento, J. L., Feldman, G.  
926 C., Milligan, A. J., Falkowski, P. G., Letelier, R. M. and Boss, E. S.: Climate-driven  
927 trends in contemporary ocean productivity, *Nature*, 444(7120), 752–755, 2006.
- 928 Benner, R. and Strom, M.: A critical evaluation of the analytical blank associated with DOC  
929 measurements by high-temperature catalytic oxidation, *Marine Chemistry*, 41(1), 153–  
930 160, [https://doi.org/10.1016/0304-4203\(93\)90113-3](https://doi.org/10.1016/0304-4203(93)90113-3), 1993.
- 931 Bergametti, Gi., Dutot, A.-L., Buat-Ménard, P., Losno, R. and Remoudaki, E.: Seasonal  
932 variability of the elemental composition of atmospheric aerosol particles over the  
933 Northwestern Mediterranean, *Tellus B: Chemical and Physical Meteorology*, 41(3), 353–  
934 361, <https://doi.org/10.3402/tellusb.v41i3.15092>, 1989.
- 935 Bishop, J. K. B., Davis, R. E. and Sherman, J. T.: Robotic Observations of Dust Storm  
936 Enhancement of Carbon Biomass in the North Pacific, *Science*, 298(5594), 817–821,  
937 <https://doi.org/10.1126/science.1074961>, 2002.
- 938 Bonnet, S. and Guieu, C.: Atmospheric forcing on the annual iron cycle in the western  
939 Mediterranean Sea: A 1-year survey, *Journal of Geophysical Research: Oceans*, 111(C9),  
940 <https://doi.org/10.1029/2005JC003213>, 2006.
- 941 Bonnet, S., Guieu, C., Chiaverini, J., Ras, J. and Stock, A.: Effect of atmospheric nutrients on the  
942 autotrophic communities in a low nutrient, low chlorophyll system, *Limnology and*  
943 *Oceanography*, 50(6), 1810–1819, <https://doi.org/10.4319/lo.2005.50.6.1810>, 2005.
- 944 Bosc, E., Bricaud, A. and Antoine, D.: Seasonal and interannual variability in algal biomass and  
945 primary production in the Mediterranean Sea, as derived from 4 years of SeaWiFS  
946 observations, *Global Biogeochemical Cycles*, 18(1),



- 947 <https://doi.org/10.1029/2003GB002034>, 2004.
- 948 Boscolo-Galazzo, F., Crichton, K. A., Barker, S. and Pearson, P. N.: Temperature dependency of  
949 metabolic rates in the upper ocean: A positive feedback to global climate change?, *Global*  
950 *and Planetary Change*, 170, 201–212, <https://doi.org/10.1016/j.gloplacha.2018.08.017>,  
951 2018.
- 952 Bressac, M. and Guieu, C.: Post-depositional processes: What really happens to new atmospheric  
953 iron in the ocean’s surface?, *Global Biogeochemical Cycles*, 27(3), 859–870,  
954 <https://doi.org/10.1002/gbc.20076>, 2013.
- 955 Bressac, M., Guieu, C., Doxaran, D., Bourrin, F., Desboeufs, K., Leblond, N. and Ridame, C.:  
956 Quantification of the lithogenic carbon pump following a simulated dust-deposition event  
957 in large mesocosms, *Biogeosciences*, 11(4), 1007–1020, [https://doi.org/10.5194/bg-11-](https://doi.org/10.5194/bg-11-1007-2014)  
958 1007-2014, 2014.
- 959 Brown, J. H., Gillooly, J. F., Allen, A. P., Savage, V. M. and West, G. B.: Toward a Metabolic  
960 Theory of Ecology, *Ecology*, 85(7), 1771–1789, <https://doi.org/10.1890/03-9000>, 2004.
- 961 Cael, B. B. and Follows, M. J.: On the temperature dependence of oceanic export efficiency,  
962 *Geophysical Research Letters*, 43(10), 5170–5175,  
963 <https://doi.org/10.1002/2016GL068877>, 2016.
- 964 Cael, B. B., Bisson, K. and Follows, M. J.: How have recent temperature changes affected the  
965 efficiency of ocean biological carbon export?, *Limnology and Oceanography Letters*,  
966 2(4), 113–118, <https://doi.org/10.1002/lol2.10042>, 2017.
- 967 Calvo-Díaz, A., Díaz-Pérez, L., Suárez, L. Á., Morán, X. A. G., Teira, E. and Marañón, E.:  
968 Decrease in the Autotrophic-to-Heterotrophic Biomass Ratio of Picoplankton in  
969 Oligotrophic Marine Waters Due to Bottle Enclosure, *Appl. Environ. Microbiol.*, 77(16),



- 970 5739–5746, <https://doi.org/10.1128/AEM.00066-11>, 2011.
- 971 Chen, B. and Laws, E. A.: Is there a difference of temperature sensitivity between marine  
972 phytoplankton and heterotrophs?, *Limnology and Oceanography*, 62(2), 806–817,  
973 <https://doi.org/10.1002/lno.10462>, 2017.
- 974 Christaki, U., Van Wambeke, F., Lefevre, D., Lagaria, A., Prieur, L., Pujo-Pay, M.,  
975 Grattepanche, J.-D., Colombet, J., Psarra, S., Dolan, J. R., Sime-Ngando, T., Conan, P.,  
976 Weinbauer, M. G. and Moutin, T.: Microbial food webs and metabolic state across  
977 oligotrophic waters of the Mediterranean Sea during summer, *Biogeosciences*, 8(7),  
978 1839–1852, <https://doi.org/10.5194/bg-8-1839-2011>, 2011.
- 979 Desboeufs, K., Leblond, N., Wagener, T., Bon Nguyen, E. and Guieu, C.: Chemical fate and  
980 settling of mineral dust in surface seawater after atmospheric deposition observed from  
981 dust seeding experiments in large mesocosms, *Biogeosciences*, 11(19), 5581–5594,  
982 <https://doi.org/10.5194/bg-11-5581-2014>, 2014.
- 983 Desboeufs, K., Bon Nguyen, E., Chevaillier, S., Triquet, S. and Dulac, F.: Fluxes and sources of  
984 nutrient and trace metal atmospheric deposition in the Northwestern Mediterranean,  
985 *Atmospheric Chemistry and Physics*, 18(19), 14477–14492, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-18-14477-2018)  
986 [18-14477-2018](https://doi.org/10.5194/acp-18-14477-2018), 2018.
- 987 Dinasquet, J., Bigeard, E., Gazeau, F., Marañón, E., Ridame, C., Van Wambeke, F.,  
988 Obernosterer, I. and Baudoux, A.-C.: Impact of dust enrichment on the microbial food  
989 web under present and future conditions of pH and temperature, *Biogeosciences*  
990 *Discussions*, 2021.
- 991 Dittmar, T., Cherrier, J. and Ludwichowski, K.-U.: The analysis of amino acids in seawater, in  
992 *Practical Guidelines for the Analysis of Seawater*, edited by O. Wurl, pp. 67–77, CRC



- 993 Press Taylor & Francis Group, Boca Raton, FL., , 2009.
- 994 Emerson, S., Quay, P., Karl, D., Winn, C., Tupas, L. and Landry, M.: Experimental  
995 determination of the organic carbon flux from open-ocean surface waters, *Nature*,  
996 389(6654), 951–954, <https://doi.org/10.1038/40111>, 1997.
- 997 Engel, A.: Determination of marine gel particles, in *Practical Guidelines for the Analysis of*  
998 *Seawater*, edited by O. Wurl, pp. 125–142, CRC Press Taylor & Francis Group, Boca  
999 Raton, FL., , 2009.
- 1000 Engel, A. and Händel, N.: A novel protocol for determining the concentration and composition  
1001 of sugars in particulate and in high molecular weight dissolved organic matter (HMW-  
1002 DOM) in seawater, *Marine Chemistry*, 127(1), 180–191,  
1003 <https://doi.org/10.1016/j.marchem.2011.09.004>, 2011.
- 1004 Feliú, G., Pagano, M., Hidalgo, P. and Carlotti, F.: Structure and function of epipelagic  
1005 mesozooplankton and their response to dust deposition events during the spring  
1006 PEACETIME cruise in the Mediterranean Sea, *Biogeosciences*, 17, 5417–5441,  
1007 <https://doi.org/10.5194/bg-17-5417-2020>, 2020.
- 1008 Ferguson, R. L., Buckley, E. N. and Palumbo, A. V.: Response of marine bacterioplankton to  
1009 differential filtration and confinement., *Appl. Environ. Microbiol.*, 47(1), 49–55, 1984.
- 1010 Friedlingstein, P., O’Sullivan, M., Jones, M. W., Andrew, R. M., Hauck, J., Olsen, A., Peters, G.  
1011 P., Peters, W., Pongratz, J., Sitch, S., Le Quéré, C., Canadell, J. G., Ciais, P., Jackson, R.  
1012 B., Alin, S., Aragão, L. E. O. C., Arneeth, A., Arora, V., Bates, N. R., Becker, M., Benoit-  
1013 Cattin, A., Bittig, H. C., Bopp, L., Bultan, S., Chandra, N., Chevallier, F., Chini, L. P.,  
1014 Evans, W., Florentie, L., Forster, P. M., Gasser, T., Gehlen, M., Gilfillan, D., Gkritzalis,  
1015 T., Gregor, L., Gruber, N., Harris, I., Hartung, K., Haverd, V., Houghton, R. A., Ilyina,



1016 T., Jain, A. K., Joetzjer, E., Kadono, K., Kato, E., Kitidis, V., Korsbakken, J. I.,  
1017 Landschützer, P., Lefèvre, N., Lenton, A., Lienert, S., Liu, Z., Lombardozzi, D., Marland,  
1018 G., Metz, N., Munro, D. R., Nabel, J. E. M. S., Nakaoka, S.-I., Niwa, Y., O'Brien, K.,  
1019 Ono, T., Palmer, P. I., Pierrot, D., Poulter, B., Resplandy, L., Robertson, E., Rödenbeck,  
1020 C., Schwinger, J., Séférian, R., Skjelvan, I., Smith, A. J. P., Sutton, A. J., Tanhua, T.,  
1021 Tans, P. P., Tian, H., Tilbrook, B., van der Werf, G., Vuichard, N., Walker, A. P.,  
1022 Wanninkhof, R., Watson, A. J., Willis, D., Wiltshire, A. J., Yuan, W., Yue, X. and  
1023 Zaehle, S.: Global Carbon Budget 2020, *Earth System Science Data*, 12(4), 3269–3340,  
1024 <https://doi.org/10.5194/essd-12-3269-2020>, 2020.

1025 Garcia-Corral, L. S., Holding, J. M., Carrillo-de-Albornoz, P., Steckbauer, A., Pérez-Lorenzo,  
1026 M., Navarro, N., Serret, P., Gasol, J. M., Morán, X. A. G., Estrada, M., Fraile-Nuez, E.,  
1027 Benítez-Barrios, V., Agusti, S. and Duarte, C. M.: Temperature dependence of plankton  
1028 community metabolism in the subtropical and tropical oceans, *Global Biogeochemical*  
1029 *Cycles*, 31(7), 1141–1154, <https://doi.org/10.1002/2017GB005629>, 2017.

1030 Gazeau, F., Ridame, C., Van Wambeke, F., Alliouane, S., Stolpe, C., Irisson, J.-O., Marro, S.,  
1031 Grisoni, J.-M., De Liège, G., Nunige, S., Djaoudi, K., Pulido-Villena, E., Dinasquet, J.,  
1032 Obernosterer, I., Catala, P. and Guieu, C.: Impact of dust enrichment on Mediterranean  
1033 plankton communities under present and future conditions of pH and temperature: an  
1034 overview, *Biogeosciences Discussions*, <https://doi.org/10.5194/bg-2020-202>, 2020.

1035 Gillikin, D. P. and Bouillon, S.: Determination of  $\delta^{18}\text{O}$  of water and  $\delta^{13}\text{C}$  of dissolved inorganic  
1036 carbon using a simple modification of an elemental analyser-isotope ratio mass  
1037 spectrometer: an evaluation, *Rapid Communications in Mass Spectrometry*, 21(8), 1475–  
1038 1478, <https://doi.org/10.1002/rcm.2968>, 2007.



- 1039 Gillooly, J. F., Brown, J. H., West, G. B., Savage, V. M. and Charnov, E. L.: Effects of Size and  
1040 Temperature on Metabolic Rate, *Science*, 293(5538), 2248–2251,  
1041 <https://doi.org/10.1126/science.1061967>, 2001.
- 1042 del Giorgio, P. and Williams, P.: Respiration in Aquatic Ecosystems, Oxford University Press.  
1043 [https://oxford.universitypressscholarship.com/view/10.1093/acprof:oso/9780198527084.](https://oxford.universitypressscholarship.com/view/10.1093/acprof:oso/9780198527084.001.0001/acprof-9780198527084)  
1044 001.0001/acprof-9780198527084, last access: 22 January 2021, 2005.
- 1045 Guieu, C. and Ridame, C.: Impact of atmospheric deposition on marine chemistry and  
1046 biogeochemistry, in *Atmospheric Chemistry in the Mediterranean Region: Comprehensive Diagnosis and Impacts*, edited by F. Dulac, S. Sauvage, and E. Hamonou,  
1047 Springer, Cham, Switzerland, , 2020.
- 1049 Guieu, C., Loye-Pilot, M. D., Benyahya, L. and Dufour, A.: Spatial variability of atmospheric  
1050 fluxes of metals (Al, Fe, Cd, Zn and Pb) and phosphorus over the whole Mediterranean  
1051 from a one-year monitoring experiment: Biogeochemical implications, *Marine*  
1052 *Chemistry*, 120(1–4), 164–178, <https://doi.org/10.1016/j.marchem.2009.02.004>, 2010.
- 1053 Guieu, C., Ridame, C., Pulido-Villena, E., Bressac, M., Desboeufs, K. and Dulac, F.: Impact of  
1054 dust deposition on carbon budget: a tentative assessment from a mesocosm approach,  
1055 *Biogeosciences*, 11(19), 5621–5635, 2014a.
- 1056 Guieu, C., Aumont, O., Paytan, A., Bopp, L., Law, C. S., Mahowald, N., Achterberg, E. P.,  
1057 Marañón, E., Salihoglu, B., Crise, A., Wagener, T., Herut, B., Desboeufs, K., Kanakidou,  
1058 M., Olgun, N., Peters, F., Pulido-Villena, E., Tovar-Sanchez, A. and Völker, C.: The  
1059 significance of the episodic nature of atmospheric deposition to Low Nutrient Low  
1060 Chlorophyll regions, *Global Biogeochemical Cycles*, 28(11), 1179–1198,  
1061 <https://doi.org/10.1002/2014GB004852>, 2014b.



- 1062 Guieu, C., D’Ortenzio, F., Dulac, F., Taillandier, V., Doglioli, A., Petrenko, A., Barrillon, S.,  
1063 Mallet, M., Nabat, P. and Desboeufs, K.: Process studies at the air-sea interface after  
1064 atmospheric deposition in the Mediterranean Sea: objectives and strategy of the  
1065 PEACETIME oceanographic campaign (May–June 2017), *Biogeosciences*, 2020(17),  
1066 5563–5585, <https://doi.org/10.5194/bg-17-5563-2020>, 2020.
- 1067 Hein, M. and Sand-Jensen, K.: CO<sub>2</sub> increases oceanic primary production, *Nature*, 388(6642),  
1068 526–527, 1997.
- 1069 Herut, B., Zohary, T., Krom, M. D., Mantoura, R. F. C., Pitta, P., Psarra, S., Rassoulzadegan, F.,  
1070 Tanaka, T. and Frede Thingstad, T.: Response of East Mediterranean surface water to  
1071 Saharan dust: On-board microcosm experiment and field observations, *Deep Sea*  
1072 *Research Part II: Topical Studies in Oceanography*, 52(22), 3024–3040,  
1073 <https://doi.org/10.1016/j.dsr2.2005.09.003>, 2005.
- 1074 Herut, B., Rahav, E., Tsagaraki, T. M., Giannakourou, A., Tsiola, A., Psarra, S., Lagaria, A.,  
1075 Papageorgiou, N., Mihalopoulos, N., Theodosi, C. N., Violaki, K., Stathopoulou, E.,  
1076 Scoullou, M., Krom, M. D., Stockdale, A., Shi, Z., Berman-Frank, I., Meador, T. B.,  
1077 Tanaka, T. and Paraskevi, P.: The Potential Impact of Saharan Dust and Polluted  
1078 Aerosols on Microbial Populations in the East Mediterranean Sea, an Overview of a  
1079 Mesocosm Experimental Approach, *Front. Mar. Sci.*, 3,  
1080 <https://doi.org/10.3389/fmars.2016.00226>, 2016.
- 1081 Hoppe, H.-G.: Significance of exoenzymatic activities in the ecology of brackish water:  
1082 measurements by means of methylumbelliferyl-substrates, *Marine Ecology Progress*  
1083 *Series*, 11(3), 299–308, 1983.
- 1084 IPCC: *Climate Change, The Physical Science Basis.*, 2013.



- 1085 Irwin, A. J. and Oliver, M. J.: Are ocean deserts getting larger?, *Geophysical Research Letters*,  
1086 36, <https://doi.org/10.1029/2009gl039883>, 2009.
- 1087 Jickells, T. and Moore, C. M.: The Importance of Atmospheric Deposition for Ocean  
1088 Productivity, *Annual Review of Ecology, Evolution, and Systematics*, 46(1), 481–501,  
1089 <https://doi.org/10.1146/annurev-ecolsys-112414-054118>, 2015.
- 1090 Kirchman, D. L., Kemp, P., Sherr, B., Sherr, E. and Cole, J.: Leucine Incorporation as a Measure  
1091 of Biomass Production by Heterotrophic Bacteria, in *Handbook of Methods in Aquatic*  
1092 *Microbial Ecology*, pp. 509–512, CRC Press, <https://doi.org/10.1201/9780203752746-59>,  
1093 , 1993.
- 1094 Knap, A., Michaels, A., Close, A., Ducklow, H. and Dickson, A.: Protocols for the Joint Global  
1095 Ocean Flux Study (JGOFS) Core Measurements, UNESCO 1994., 1996.
- 1096 Kouvarakis, G., Mihalopoulos, N., Tselepides, A. and Stavrakakis, S.: On the importance of  
1097 atmospheric inputs of inorganic nitrogen species on the productivity of the Eastern  
1098 Mediterranean Sea, *Global Biogeochemical Cycles*, 15(4), 805–817,  
1099 <https://doi.org/10.1029/2001GB001399>, 2001.
- 1100 Kroeker, K. J., Kordas, R. L., Crim, R., Hendriks, I. E., Ramajo, L., Singh, G. S., Duarte, C. M.  
1101 and Gattuso, J. P.: Impacts of ocean acidification on marine organisms: quantifying  
1102 sensitivities and interaction with warming, *Global Change Biology*, 19(6), 1884–1896,  
1103 <https://doi.org/10.1111/gcb.12179>, 2013.
- 1104 Langer, G., Nehrke, G., Probert, I., Ly, J. and Ziveri, P.: Strain-specific responses of *Emiliania*  
1105 *huxleyi* to changing seawater carbonate chemistry, *Biogeosciences*, 6(11), 2637–2646,  
1106 <https://doi.org/10.5194/bg-6-2637-2009>, 2009.
- 1107 Laurent, B., Audoux, T., Bibi, M., Dulac, F. and Bergametti, G.: Mass deposition in the



- 1108 Mediterranean region, in *Atmospheric Chemistry in the Mediterranean Region:*  
1109 *Comprehensive Diagnosis and Impacts*, edited by F. Dulac, S. Sauvage, and E. Hamonou,  
1110 Springer, Cham, Switzerland, , 2021.
- 1111 Lazzari, P., Solidoro, C., Salon, S. and Bolzon, G.: Spatial variability of phosphate and nitrate in  
1112 the Mediterranean Sea: A modeling approach, *Deep Sea Research Part I: Oceanographic*  
1113 *Research Papers*, 108, 39–52, <https://doi.org/10.1016/j.dsr.2015.12.006>, 2016.
- 1114 Lekunberri, I., Lefort, T., Romero, E., Vázquez-Domínguez, E., Romera-Castillo, C., Marrasé,  
1115 C., Peters, F., Weinbauer, M. and Gasol, J. M.: Effects of a dust deposition event on  
1116 coastal marine microbial abundance and activity, bacterial community structure and  
1117 ecosystem function, *J Plankton Res*, 32(4), 381–396,  
1118 <https://doi.org/10.1093/plankt/fbp137>, 2010.
- 1119 Lemée, R., Rochelle-Newall, E., Van Wambeke, F., Pizay, M., Rinaldi, P. and Gattuso, J.:  
1120 Seasonal variation of bacterial production, respiration and growth efficiency in the open  
1121 NW Mediterranean Sea, *Aquat. Microb. Ecol.*, 29, 227–237,  
1122 <https://doi.org/10.3354/ame029227>, 2002.
- 1123 Lewandowska, A. M., Boyce, D. G., Hofmann, M., Matthiessen, B., Sommer, U. and Worm, B.:  
1124 Effects of sea surface warming on marine plankton, *Ecology Letters*, 17(5), 614–623,  
1125 <https://doi.org/10.1111/ele.12265>, 2014.
- 1126 Lindroth, P. and Mopper, K.: High performance liquid chromatographic determination of  
1127 subpicomole amounts of amino acids by precolumn fluorescence derivatization with o-  
1128 phthaldialdehyde, *Anal. Chem.*, 51(11), 1667–1674,  
1129 <https://doi.org/10.1021/ac50047a019>, 1979.
- 1130 Longhurst, A., Sathyendranath, S., Platt, T. and Caverhill, C.: An estimate of global primary



- 1131 production in the ocean from satellite radiometer data, *Journal of Plankton Research*,  
1132 17(6), 1245–1271, <https://doi.org/10.1093/plankt/17.6.1245>, 1995.
- 1133 Lopez-Urrutia, A. and Moran, X. A. G.: Resource limitation of bacterial production distorts the  
1134 temperature dependence of oceanic carbon cycling, *Ecology*, 88(4), 817–822, 2007.
- 1135 Louis, J., Pedrotti, M. L., Gazeau, F. and Guieu, C.: Experimental evidence of formation of  
1136 transparent exopolymer particles (TEP) and POC export provoked by dust addition under  
1137 current and high  $p\text{CO}_2$  conditions, *PLOS ONE*, 12(2), e0171980,  
1138 <https://doi.org/10.1371/journal.pone.0171980>, 2017.
- 1139 Lojze-Pilot, M. D. and Martin, J. M.: Saharan Dust Input to the Western Mediterranean: An  
1140 Eleven Years Record in Corsica, in *The Impact of Desert Dust Across the Mediterranean*,  
1141 edited by S. Guerzoni and R. Chester, pp. 191–199, Springer Netherlands, Dordrecht,  
1142 [https://doi.org/10.1007/978-94-017-3354-0\\_18](https://doi.org/10.1007/978-94-017-3354-0_18), , 1996.
- 1143 Mackey, K., Morris, J. J., Morel, F. and Kranz, S.: Response of photosynthesis to ocean  
1144 acidification, *Oceanography*, 25(2), 74–91, <https://doi.org/10.5670/oceanog.2015.33>,  
1145 2015.
- 1146 Marañón, E., Fernández, A., Mouriño-Carballido, B., Martínez-García, S., Teira, E., Cermeño,  
1147 P., Chouciño, P., Huete-Ortega, M., Fernández, E., Calvo-Díaz, A., Morán, X. A. G.,  
1148 Bode, A., Moreno-Ostos, E., Varela, M. M., Patey, M. D. and Achterberg, E. P.: Degree  
1149 of oligotrophy controls the response of microbial plankton to Saharan dust, *Limnology*  
1150 and *Oceanography*, 55(6), 2339–2352, <https://doi.org/10.4319/lo.2010.55.6.2339>, 2010.
- 1151 Marañón, E., Lorenzo, M. P., Cermeño, P. and Mouriño-Carballido, B.: Nutrient limitation  
1152 suppresses the temperature dependence of phytoplankton metabolic rates, *The ISME*  
1153 *Journal*, 12(7), 1836–1845, <https://doi.org/10.1038/s41396-018-0105-1>, 2018.



- 1154 Marañón, E., Van Wambeke, F., Uitz, J., Boss, E. S., Pérez-Lorenzo, M., Dinasquet, J.,  
1155 Haëntjens, N., Dimier, C. and Taillandier, V.: Deep maxima of phytoplankton biomass,  
1156 primary production and bacterial production in the Mediterranean Sea during late spring,  
1157 Biogeosciences Discussions, 1–28, <https://doi.org/10.5194/bg-2020-261>, 2020.
- 1158 Mari, X.: Carbon content and C:N ratio of transparent exopolymeric particles (TEP) produced by  
1159 bubbling exudates of diatoms, Marine Ecology Progress Series, 183, 59–71,  
1160 <https://doi.org/10.3354/meps183059>, 1999.
- 1161 Markaki, Z., Oikonomou, K., Kocak, M., Kouvarakis, G., Chaniotaki, A., Kubilay, N. and  
1162 Mihalopoulos, N.: Atmospheric deposition of inorganic phosphorus in the Levantine  
1163 Basin, eastern Mediterranean: Spatial and temporal variability and its role in seawater  
1164 productivity, Limnology and Oceanography, 48(4), 1557–1568,  
1165 <https://doi.org/10.4319/lo.2003.48.4.1557>, 2003.
- 1166 Maugendre, L., Gattuso, J.-P., Louis, J., de Kluijver, A., Marro, S., Soetaert, K. and Gazeau, F.:  
1167 Effect of ocean warming and acidification on a plankton community in the NW  
1168 Mediterranean Sea, ICES Journal of Marine Science: Journal du Conseil, 72(6), 1744–  
1169 1755, <https://doi.org/10.1093/icesjms/fsu161>, 2015.
- 1170 Maugendre, L., Gattuso, J.-P., Poulton, A. J., Dellisanti, W., Gaubert, M., Guieu, C. and Gazeau,  
1171 F.: No detectable effect of ocean acidification on plankton metabolism in the NW  
1172 oligotrophic Mediterranean Sea: Results from two mesocosm studies, Estuarine, Coastal  
1173 and Shelf Science, 186, 89–99, <https://doi.org/10.1016/j.ecss.2015.03.009>, 2017a.
- 1174 Maugendre, L., Guieu, C., Gattuso, J.-P. and Gazeau, F.: Ocean acidification in the  
1175 Mediterranean Sea: Pelagic mesocosm experiments. A synthesis, Estuarine, Coastal and  
1176 Shelf Science, 186, 1–10, <https://doi.org/10.1016/j.ecss.2017.01.006>, 2017b.



- 1177 Mercado, J. M., Sobrino, C., Neale, P. J., Segovia, M., Reul, A., Amorim, A. L., Carrillo, P.,  
1178 Claquin, P., Cabrerizo, M. J., León, P., Lorenzo, M. R., Medina-Sánchez, J. M.,  
1179 Montecino, V., Napoleon, C., Prasil, O., Putzeys, S., Salles, S. and Yebra, L.: Effect of  
1180 CO<sub>2</sub>, nutrients and light on coastal plankton. II. Metabolic rates, *Aquatic Biology*, 22,  
1181 43–57, <https://doi.org/10.3354/ab00606>, 2014.
- 1182 Mills, M. M., Moore, C. M., Langlois, R., Milne, A., Achterberg, E., Nachtigall, K., Lochte, K.,  
1183 Geider, R. J. and La, R. J.: Nitrogen and phosphorus co-limitation of bacterial  
1184 productivity and growth in the oligotrophic subtropical North Atlantic, *Limnology and*  
1185 *Oceanography*, 53(2), 824–834, <https://doi.org/10.4319/lo.2008.53.2.0824>, 2008.
- 1186 Mosseri, J., Quéguiner, B., Rimmelin, P., Leblond, N. and Guieu, C.: Silica fluxes in the  
1187 northeast Atlantic frontal zone of Mode Water formation (38°–45°N, 16°–22°W) in  
1188 2001–2002, *Journal of Geophysical Research: Oceans*, 110(C7),  
1189 <https://doi.org/10.1029/2004JC002615>, 2005.
- 1190 Moulin, C. and Chiapello, I.: Impact of human-induced desertification on the intensification of  
1191 Sahel dust emission and export over the last decades, *Geophysical Research Letters*,  
1192 33(18), <https://doi.org/10.1029/2006GL025923>, 2006.
- 1193 Moutin, T. and Raimbault, P.: Primary production, carbon export and nutrients availability in  
1194 western and eastern Mediterranean Sea in early summer 1996 (MINOS cruise), *Journal of*  
1195 *Marine Systems*, 33–34, 273–288, [https://doi.org/10.1016/S0924-7963\(02\)00062-3](https://doi.org/10.1016/S0924-7963(02)00062-3),  
1196 2002.
- 1197 Moutin, T., Thingstad, T. F., Wambeke, F. V., Marie, D., Slawyk, G., Raimbault, P. and  
1198 Claustre, H.: Does competition for nanomolar phosphate supply explain the  
1199 predominance of the cyanobacterium *Synechococcus*?, *Limnology and Oceanography*,



- 1200 47(5), 1562–1567, <https://doi.org/10.4319/lo.2002.47.5.1562>, 2002.
- 1201 Müren, U., Berglund, J., Samuelsson, K. and Andersson, A.: Potential Effects of Elevated Sea-  
1202 Water Temperature on Pelagic Food Webs, *Hydrobiologia*, 545(1), 153–166,  
1203 <https://doi.org/10.1007/s10750-005-2742-4>, 2005.
- 1204 Passow, U. and Carlson, C. A.: The biological pump in a high CO<sub>2</sub> world, *Marine Ecology*  
1205 *Progress Series*, 470, 249–271, 2012.
- 1206 Polovina, J. J., Howell, E. A. and Abecassis, M.: Ocean’s least productive waters are expanding,  
1207 *Geophysical Research Letters*, 35(3), <https://doi.org/10.1029/2007gl031745>, 2008.
- 1208 Pulido-Villena, E., Wagener, T. and Guieu, C.: Bacterial response to dust pulses in the western  
1209 Mediterranean: Implications for carbon cycling in the oligotrophic ocean, *Global*  
1210 *Biogeochemical Cycles*, 22(1), <https://doi.org/10.1029/2007gb003091>, 2008.
- 1211 Pulido-Villena, E., Baudoux, A.-C., Obernosterer, I., Landa, M., Caparros, J., Catala, P.,  
1212 Georges, C., Harmand, J. and Guieu, C.: Microbial food web dynamics in response to a  
1213 Saharan dust event: results from a mesocosm study in the oligotrophic Mediterranean  
1214 Sea, *Biogeosciences*, 11(19), 5607–5619, 2014.
- 1215 Regaudie-de-Gioux, A. and Duarte, C. M.: Temperature dependence of planktonic metabolism in  
1216 the ocean, *Global Biogeochemical Cycles*, 26(1), <https://doi.org/10.1029/2010GB003907>,  
1217 2012.
- 1218 Regaudie-de-Gioux, A., Vaquer-Sunyer, R. and Duarte, C. M.: Patterns in planktonic  
1219 metabolism in the Mediterranean Sea, *Biogeosciences*, 6(12), 3081–3089,  
1220 <https://doi.org/10.5194/bg-6-3081-2009>, 2009.
- 1221 Ridame, C. and Guieu, C.: Saharan input of phosphate to the oligotrophic water of the open  
1222 western Mediterranean Sea, *Limnology and Oceanography*, 47(3), 856–869, 2002.



- 1223 Riebesell, U., Schulz, K. G., Bellerby, R. G. J., Botros, M., Fritsche, P., Meyerhofer, M., Neill,  
1224 C., Nondal, G., Oschlies, A., Wohlers, J. and Zollner, E.: Enhanced biological carbon  
1225 consumption in a high CO<sub>2</sub> ocean, *Nature*, 450(7169), 545-U10, 2007.
- 1226 Smith, D. C. and Azam, F.: A simple, economical method for measuring bacterial protein  
1227 synthesis rates in seawater using 3H-leucine, *Marine Microbial Food Webs*, 6(2), 107–  
1228 114, 1992.
- 1229 Tanaka, T., Thingstad, T. F., Christaki, U., Colombet, J., Cornet-Barthaux, V., Courties, C.,  
1230 Grattepanche, J.-D., Lagaria, A., Nedoma, J., Oriol, L., Psarra, S., Pujo-Pay, M. and  
1231 Wambeke, F. V.: Lack of P-limitation of phytoplankton and heterotrophic prokaryotes in  
1232 surface waters of three anticyclonic eddies in the stratified Mediterranean Sea,  
1233 *Biogeosciences*, 8(2), 525–538, <https://doi.org/10.5194/bg-8-525-2011>, 2011.
- 1234 Ternon, E., Guieu, C., Loÿe-Pilot, M.-D., Leblond, N., Bosc, E., Gasser, B., Miquel, J.-C. and  
1235 Martín, J.: The impact of Saharan dust on the particulate export in the water column of  
1236 the North Western Mediterranean Sea, *Biogeosciences*, 7(3), 809–826,  
1237 <https://doi.org/10.5194/bg-7-809-2010>, 2010.
- 1238 Thingstad, T. F., Krom, M. D., Mantoura, R. F. C., Flaten, G. a. F., Groom, S., Herut, B., Kress,  
1239 N., Law, C. S., Pasternak, A., Pitta, P., Psarra, S., Rassoulzadegan, F., Tanaka, T.,  
1240 Tselepidis, A., Wassmann, P., Woodward, E. M. S., Riser, C. W., Zodiatis, G. and  
1241 Zohary, T.: Nature of Phosphorus Limitation in the Ultraoligotrophic Eastern  
1242 Mediterranean, *Science*, 309(5737), 1068–1071, <https://doi.org/10.1126/science.1112632>,  
1243 2005.
- 1244 Van Wambeke, F., Taillandier, V., Deboeufs, K., Pulido-Villena, E., Dinasquet, J., Engel, A.,  
1245 Marañón, E., Ridame, C. and Guieu, C.: Influence of atmospheric deposition on



- 1246 biogeochemical cycles in an oligotrophic ocean system, *Biogeosciences Discussions*, 1–  
1247 51, <https://doi.org/10.5194/bg-2020-411>, 2020a.
- 1248 Van Wambeke, F., Pulido, E., Dinasquet, J., Djaoudi, K., Engel, A., Garel, M., Guasco, S.,  
1249 Nunige, S., Taillandier, V., Zäncker, B. and Tamburini, C.: Spatial patterns of biphasic  
1250 ectoenzymatic kinetics related to biogeochemical properties in the Mediterranean Sea,  
1251 *Biogeosciences Discussions*, 1–38, <https://doi.org/10.5194/bg-2020-253>, 2020b.
- 1252 Wang, Q., Lyu, Z., Omar, S., Cornell, S., Yang, Z. and Montagnes, D. J. S.: Predicting  
1253 temperature impacts on aquatic productivity: Questioning the metabolic theory of  
1254 ecology’s “canonical” activation energies, *Limnology and Oceanography*, 64(3), 1172–  
1255 1185, <https://doi.org/10.1002/lno.11105>, 2019.
- 1256 Zobell, C. E. and Anderson, D. Q.: Observations on the multiplication of bacteria in different  
1257 volumes of stored sea water and the influence of oxygen tension and solid surfaces, *The*  
1258 *Biological Bulletin*, 71(2), 324–342, <https://doi.org/10.2307/1537438>, 1936.



## Tables

- 1259
- 1260 Table 1. Initial chemical and biological stocks as measured while filling the tanks (initial conditions in pumped surface water;
- 1261 sampling time: t-12h). NO<sub>x</sub>: nitrate + nitrite, DIP: dissolved inorganic phosphorus, Si(OH)<sub>4</sub>: silicate, POC: particulate organic carbon,
- 1262 DOC: dissolved organic carbon, TEP: transparent exopolymer particles, TChl<sub>a</sub>: total chlorophyll *a*. Values shown for <sup>14</sup>C
- 1263 incorporation rates, percentages of extracellular release (%PER) as well as for net community production (NCP), community
- 1264 respiration (CR) and gross primary production (GPP) were estimated from samples taken at t0 in the control tanks. For heterotrophic
- 1265 bacterial production (BP), rates were estimated from seawater sampled at t-12h.

Sampling station	TYR	ION	FAST
Coordinates (decimal)	39.34 N, 12.60 E	35.49 N, 19.78 E	37.95 N, 2.90 N
Bottom depth (m)	3395	3054	2775
Day and time of pumping (local time)	17/05/2017 17:00	25/05/2017 17:00	02/06/2017 21:00
Temperature (°C)	20.6	21.2	21.5
Salinity	37.96	39.02	37.07
Stocks			
NO <sub>x</sub> (nmol L <sup>-1</sup> )	14.0	18.0	59.0
DIP (nmol L <sup>-1</sup> )	17.1	6.5	12.9
Si(OH) <sub>4</sub> (μmol L <sup>-1</sup> )	1.0	0.96	0.64



POC ( $\mu\text{mol L}^{-1}$ )	12.9	8.5	6.0
DOC ( $\mu\text{mol L}^{-1}$ )	72.2	70.2	69.6
TEP ( $\times 10^6 \text{ L}^{-1}$ )	6.8	3.8	3.7
TChl <i>a</i> ( $\mu\text{g L}^{-1}$ )	0.063	0.066	0.072
Heterotrophic prokaryotes abundance ( $\times 10^5 \text{ cell mL}^{-1}$ )	4.79	2.14	6.15
Processes			
$^{14}\text{C}$ -based total particulate production ( $\mu\text{g C L}^{-1} \text{ h}^{-1}$ )	$0.08 \pm 0.03$	$0.14 \pm 0.04$	$0.15 \pm 0.04$
$^{14}\text{C}$ -based $> 2 \mu\text{m}$ particulate production ( $\mu\text{g C L}^{-1} \text{ h}^{-1}$ )	$0.07 \pm 0.02$	$0.11 \pm 0.02$	$0.11 \pm 0.02$
$^{14}\text{C}$ -based $< 2 \mu\text{m}$ particulate production ( $\mu\text{g C L}^{-1} \text{ h}^{-1}$ )	$0.01 \pm 0.01$	$0.04 \pm 0.02$	$0.05 \pm 0.01$
%PER	$60 \pm 20$	$45 \pm 3$	$32 \pm 23$
NCP ( $\mu\text{mol O}_2 \text{ L}^{-1} \text{ d}^{-1}$ )	$-1.9 \pm 0.3$	$-0.2 \pm 0.2$	$-0.8 \pm 0.9$
CR ( $\mu\text{mol O}_2 \text{ L}^{-1} \text{ d}^{-1}$ )	$-2.6 \pm 0.1$	$-1.2 \pm 0.5$	$-1.9 \pm 1.6$
GPP ( $\mu\text{mol O}_2 \text{ L}^{-1} \text{ d}^{-1}$ )	$0.7 \pm 0.4$	$1.1 \pm 0.3$	$1.1 \pm 0.7$
BP ( $\text{ng C L}^{-1} \text{ h}^{-1}$ )	11.6	15.2	34.6



1267 Table 2. Heterotrophic bacterial production (BP) growth rates ( $\mu_{BP}$  in  $h^{-1}$ ) estimated from the  
1268 exponential phase of BP growth, observable from at least four sampling points, between  $t_0$  and  
1269  $t_{12h}$ , during the three experiments (TYR, ION and FAST) in the six tanks (controls: C1, C2; dust  
1270 addition under present conditions of temperature and pH: D1, D2; dust addition under future  
1271 conditions of temperature and pH: G1 and G2). Values  $\pm$  SE are shown.

	$\mu_{BP}$		
	TYR	ION	FAST
C1	$0.076 \pm 0.025$	$0.042 \pm 0.007$	$0.020 \pm 0.003$
C2	$0.066 \pm 0.018$	$0.041 \pm 0.005$	$0.026 \pm 0.004$
D1	$0.117 \pm 0.008$	$0.095 \pm 0.020$	$0.089 \pm 0.014$
D2	$0.194 \pm 0.020$	$0.145 \pm 0.007$	$0.090 \pm 0.007$
G1	$0.164 \pm 0.020$	$0.126 \pm 0.011$	$0.124 \pm 0.005$
G2	$0.150 \pm 0.003$	$0.137 \pm 0.033$	$0.163 \pm 0.014$

1272



1273 Table 3. Estimated bacterial growth efficiency (BGE in %) during the course of the three  
1274 experiments (TYR, ION and FAST) in the six tanks (controls: C1, C2; dust addition under  
1275 present conditions of temperature and pH: D1, D2; dust addition under future conditions of  
1276 temperature and pH: G1 and G2). BGE was calculated based on integrated heterotrophic  
1277 bacterial production (BP) and community respiration (CR) rates by applying a bacterial  
1278 respiration to CR ratio of 0.7 and a respiratory quotient of 0.8 (see Material and Methods).

---

Bacterial growth efficiency (BGE)			
	TYR	ION	FAST
C1	11.1	9.8	15.4
C2	11.7	14.5	22.0
D1	31.8	21.0	17.3
D2	32.3	30.6	19.9
G1	39.3	35.2	37.6
G2	32.5	34.8	38.1

---

1279

1280



## 1281 **Figure caption**

1282 Fig. 1. Map showing the sampling stations in the Mediterranean Sea along the transect performed  
1283 onboard the R/V “Pourquoi Pas ?” during the PEACETIME cruise.

1284 Fig. 2. Dissolved organic carbon (DOC) concentrations and ratio between total hydrolysable  
1285 amino acids (TAA) and DOC concentrations measured in the six tanks (controls: C1, C2; dust  
1286 addition under present conditions of temperature and pH: D1, D2; dust addition under future  
1287 conditions of temperature and pH: G1 and G2) during the three experiments (TYR, ION and  
1288 FAST). The dashed vertical line indicates the time of seeding (after  $t_0$ ).

1289 Fig. 3. Particulate organic carbon (POC) concentrations and transparent exopolymer particle  
1290 carbon content (TEP-C) measured in the six tanks (controls: C1, C2; dust addition under present  
1291 conditions of temperature and pH: D1, D2; dust addition under future conditions of temperature  
1292 and pH: G1 and G2) during the three experiments (TYR, ION and FAST). The dashed vertical  
1293 line indicates the time of seeding (after  $t_0$ ).

1294 Fig. 4.  $^{14}\text{C}$ -based production rates ( $< 2 \mu\text{m}$  and  $> 2 \mu\text{m}$  size fractions, total particulate) estimated  
1295 from 8 h incubations on samples taken in the six tanks (controls: C1, C2; dust addition under  
1296 present conditions of temperature and pH: D1, D2; dust addition under future conditions of  
1297 temperature and pH: G1 and G2) during the three experiments (TYR, ION and FAST). The  
1298 percentage of extracellular release (%PER) is also shown.

1299 Fig. 5. Incorporation of  $^{13}\text{C}$  into particulate organic carbon ( $\delta^{13}\text{C}$ -POC) in the six tanks (controls:  
1300 C1, C2; dust addition under present conditions of temperature and pH: D1, D2; dust addition



1301 under future conditions of temperature and pH: G1 and G2) during the three experiments (TYR,  
1302 ION and FAST). The dashed vertical line indicates the time of seeding (after  $t_0$ ).

1303 Fig. 6. Net community production (NCP), community respiration (CR) and gross primary  
1304 production (GPP) rates estimated using the oxygen light-dark method (24 h incubations) on  
1305 samples taken in the six tanks (C1, C2, D1, D2, G1 and G2) during the three experiments (TYR,  
1306 ION and FAST).

1307 Fig. 7. Heterotrophic bacterial production rates (BP) and cell-specific maximum hydrolysis  
1308 velocity ( $V_m$ ) of the alkaline phosphatase (both over 1-2 h incubations) on samples taken in the  
1309 six tanks (C1, C2, D1, D2, G1 and G2) during the three experiments (TYR, ION and FAST).

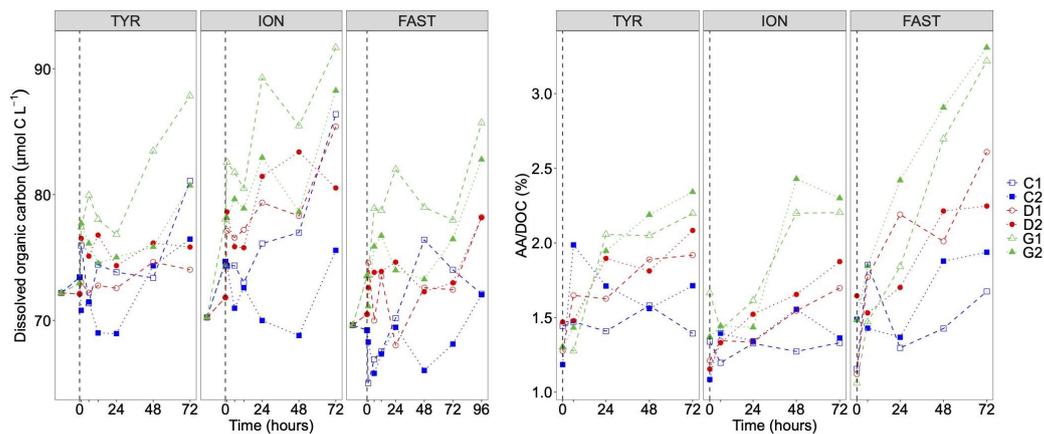
1310 Fig. 8. Total mass and organic matter fluxes measured in the sediment traps at the end of the  
1311 three experiments (TYR, ION and FAST) in the six tanks (C1, C2, D1, D2, G1 and G2).

1312 Fig. 9. Relative difference (%) between integrated rates measured in tanks D (D1, D2; dust  
1313 addition under present conditions of temperature and pH) and G (G1, G2; dust addition under  
1314 future conditions of temperature and pH) as compared to the controls (C1, C2) during the three  
1315 experiments (TYR, ION and FAST). Vertical boxes represent the range observed between the  
1316 two replicates per treatment.



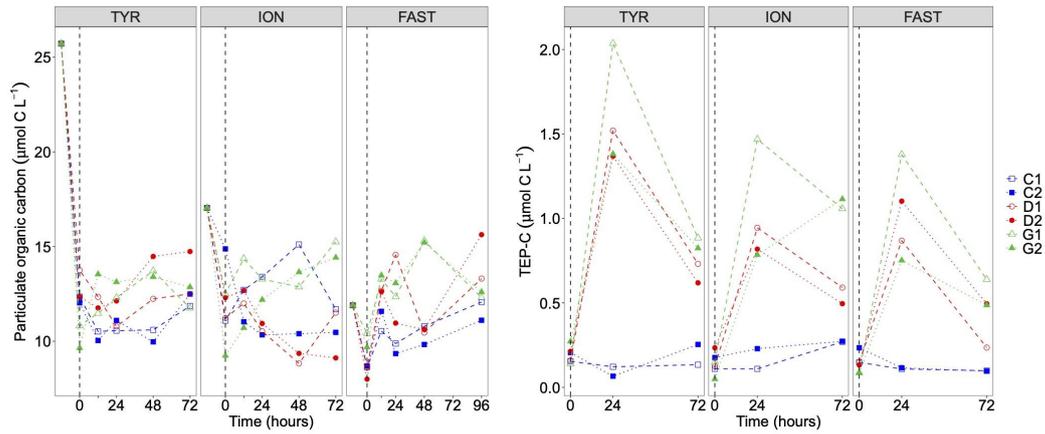
1317

1318 Fig. 1



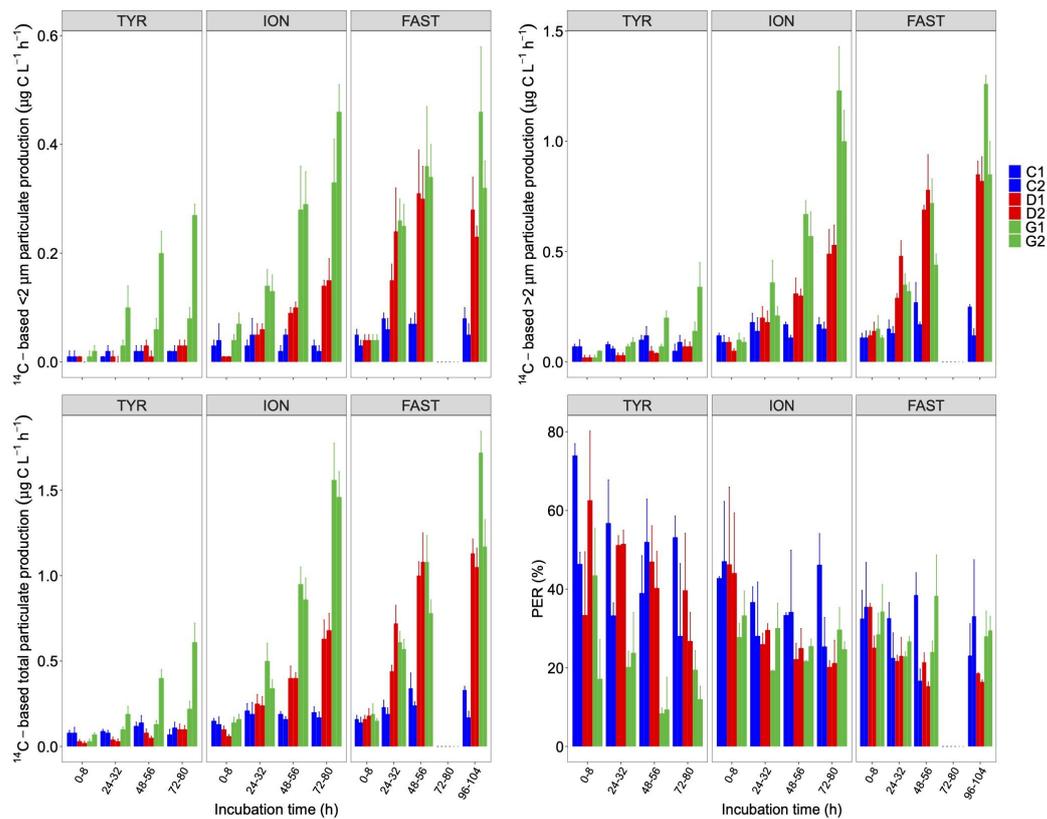
1319

1320 Fig. 2



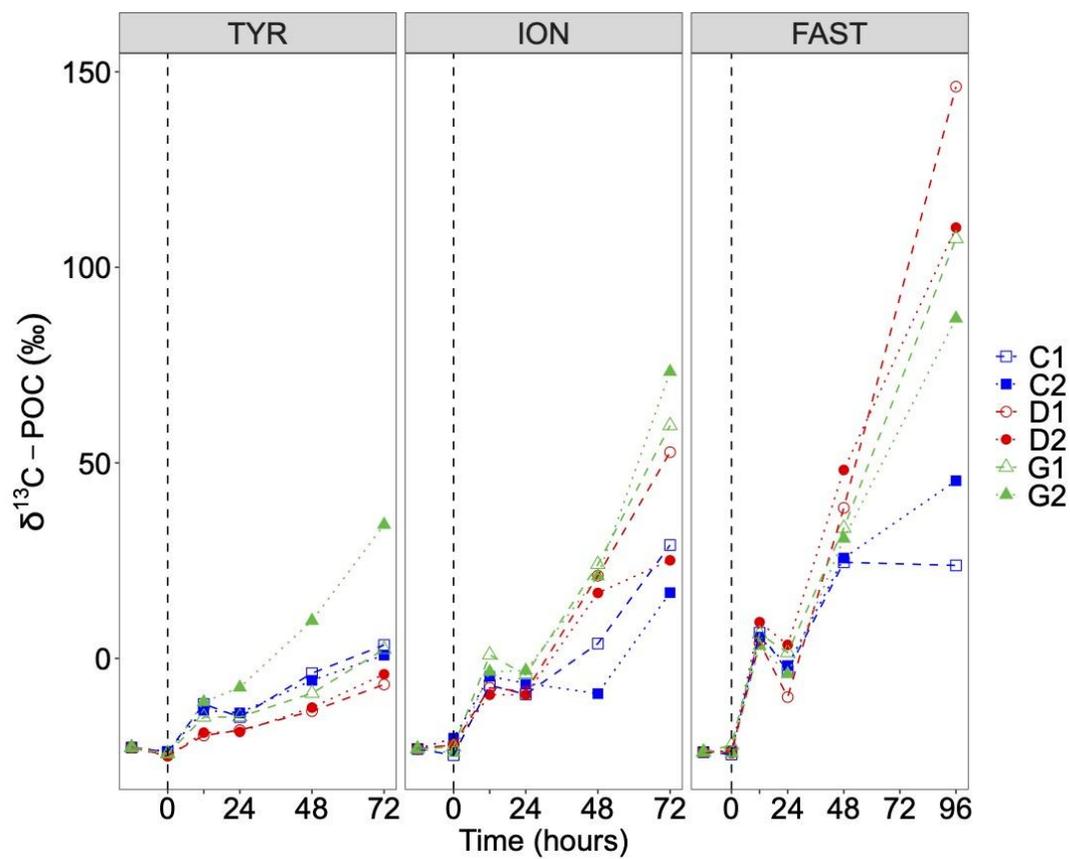
1321

1322 Fig. 3



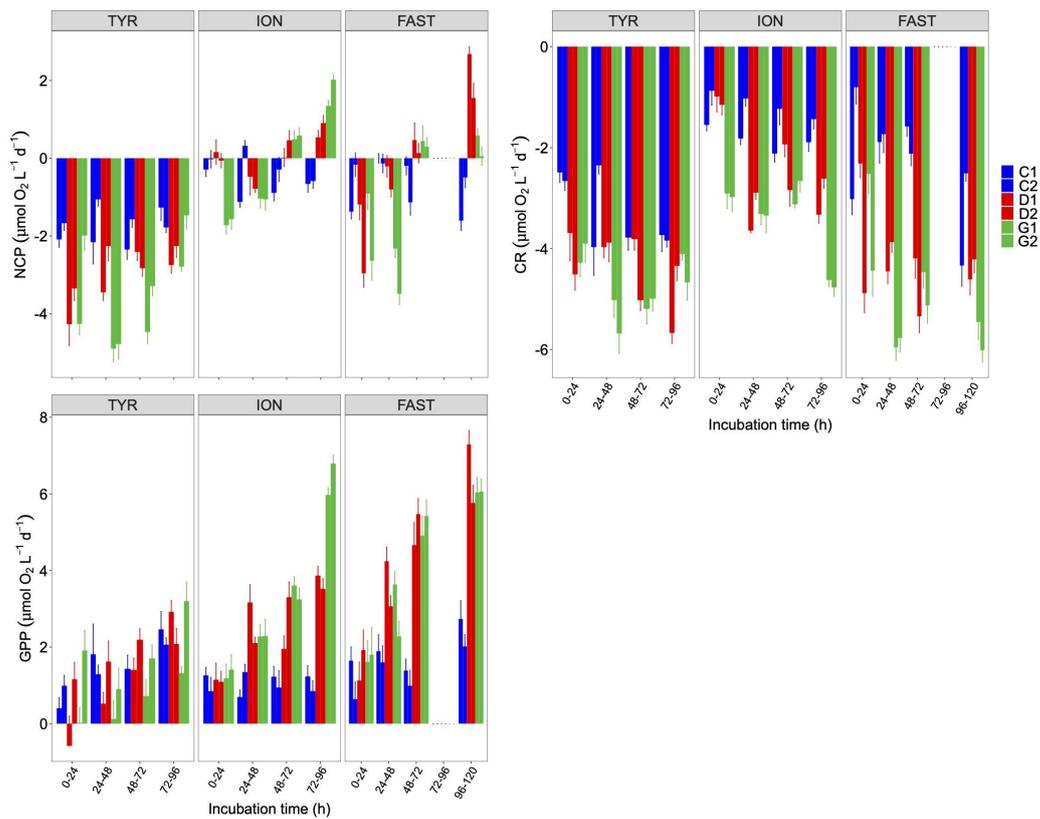
1323

1324 Fig. 4



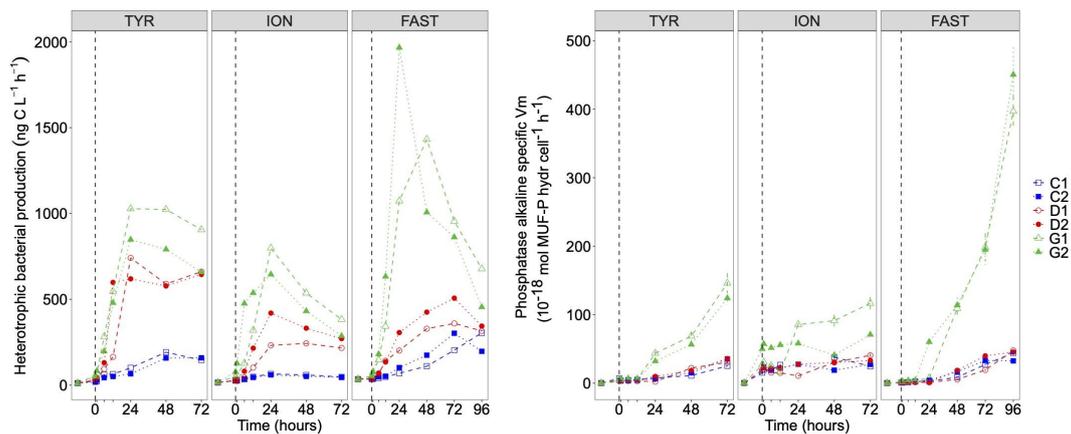
1325

1326 Fig. 5



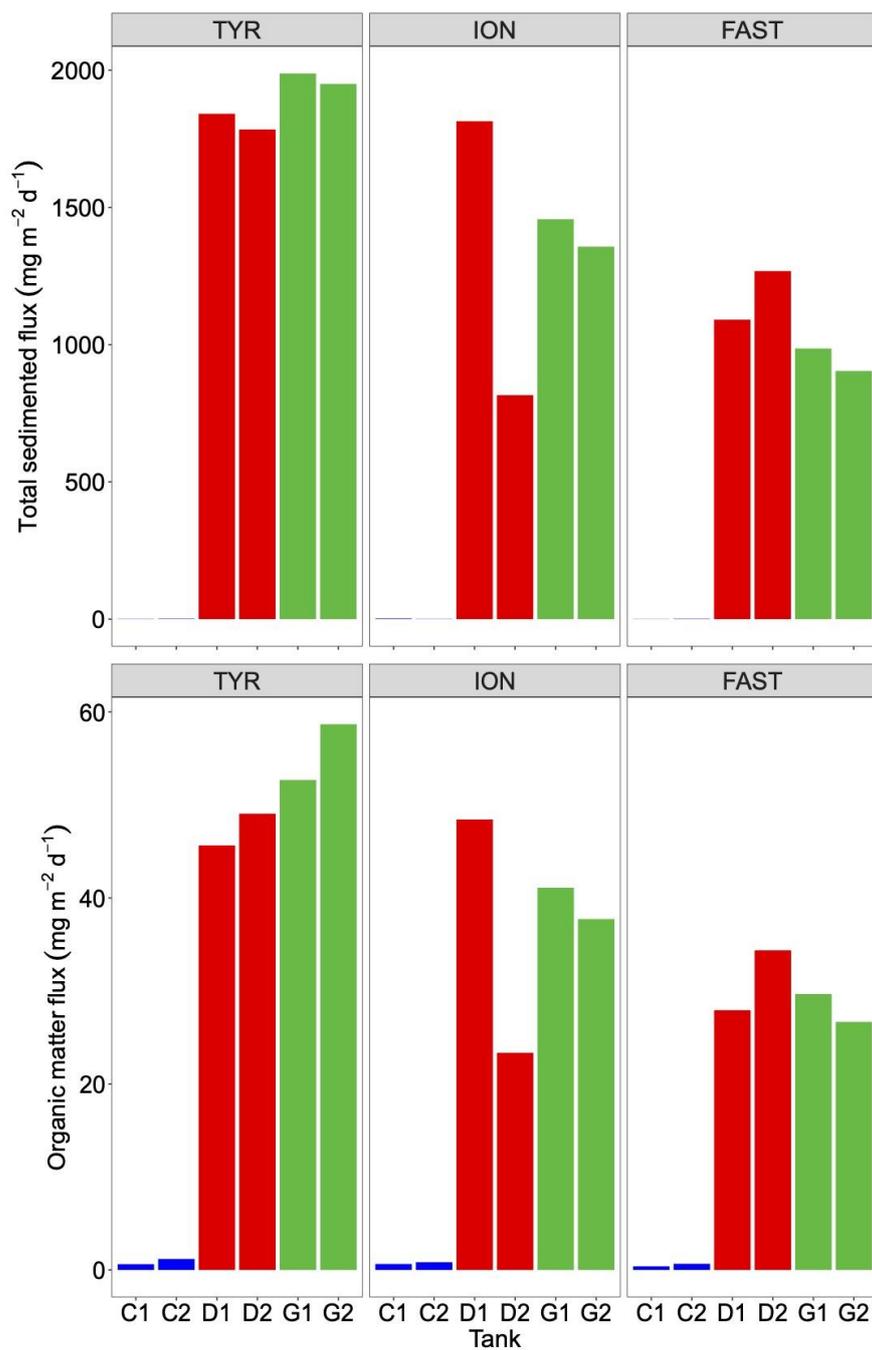
1327

1328 Fig. 6



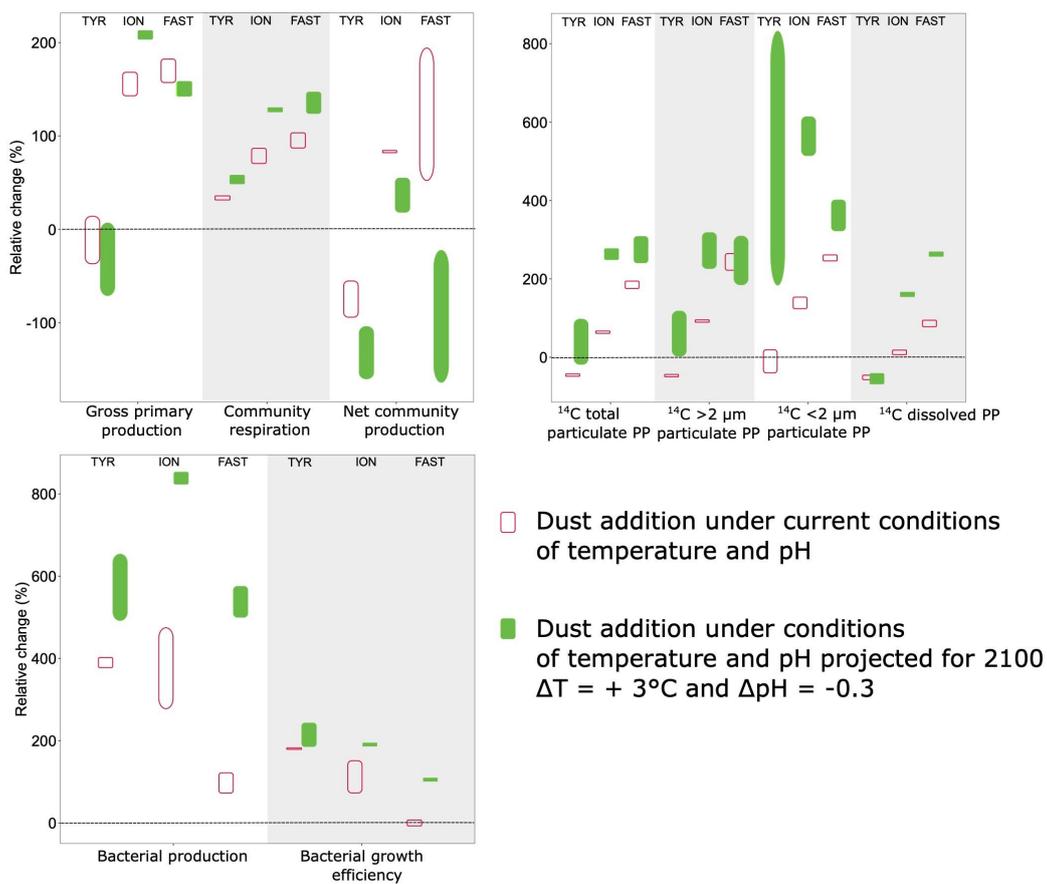
1329

1330 Fig. 7



1331

1332 Fig. 8



1333

1334 Fig. 9