A GC-SSIM-CRDS System: Coupling a Gas Chromatograph with a Cavity Ring-Down Spectrometer for Onboard Twofold Analysis of Molecular and Isotopic Compositions of Natural Gases during Ocean-Going Research Expeditions

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Abstract :

Carbon dioxide (CO2) and methane (CH4) are two climate-sensitive components of gases migrating within sediments and emitted into the water column on continental margins. They are involved in several key biogeochemical processes entering into the global carbon cycle. In order to perform onboard measurements of both the molecular and stable carbon isotope ratios (δ13C) of CH4 and CO2 of natural gases during oceanic cruises, we have developed a novel approach coupling gas chromatography (GC) with cavity ring-down spectroscopy (CRDS). The coupled devices are connected to a small sample isotope module (SSIM) to form a system called GC-SSIM-CRDS. Small volumes of natural gas samples (<1 mL) are injected into the GC using a headspace autosampler or a gas-tight syringe to separate the chemical components using a Shincarbon ST packed column and for molecular quantification by thermal conductivity detection (TCD). Subsequently, CO2 from the sample is trapped in a 7 mL loop at 32 °C before being transferred to the CRDS analyzer for sequential determination of the stable carbon isotope ratios of CH4 and CO2 in 24 min. The loop is an open column (without stationary phase). This approach does not require the use of adsorbents or cooling for the trapping step. Optimization of the separation step prior to analysis was focused on the influence of two key separation factors 1) the flow of the carrier gas and 2) the temperature of the oven. Our analytical system and the measurement protocol were validated on samples collected from gas seeps in the Sea of Marmara (Turkey). Our results show that the GC-SSIM-CRDS system provides a reliable determination of the molecular identification of CH4 and CO2 in complex natural gases, followed by the stable carbon isotope ratios of methane and carbon dioxide.

Graphical abstract



Highlights

▶ We present a coupled analytical system associating gas chromatography and cavity ring down spectroscopy (GC-SSIM-CRDS) for natural gas analyses. ▶ The analytical system allows fast onboard analyses of both molecular composition of natural gases and stable carbon isotope ratios of methane and carbon dioxide in 24 min. ▶ The GC-SSIM-CRDS was optimized and validated with samples from natural gas seeps. ▶ The GC-SSIM-CRDS is a decision-making tool suitable to refining sampling strategy of gases when characterizing seep areas and is a useful analytical system for scientific purposes.

Keywords: Coupled analytical technique, cavity ring-down spectroscopy, carbon dioxide, gas chromatography, methane, molecular composition, natural gases, onboard analysis, stable carbon isotope ratio

1 Introduction 39

| 39 | 1 Introduction |
|----|---|
| 40 | Methane (CH ₄) and carbon dioxide (CO ₂) are among the most abundant greenhouse |
| 41 | gases on Earth [14] and are strongly responsible for climate change. The major sources of |
| 42 | atmospheric methane are wetlands, landfills, livestock farming, hydrocarbon-field production, |
| 43 | as well as natural seepages from geologic structures including marine emissions. During the |
| 44 | 2008–2017 decade, the total emissions was estimated between 550 to 600 Tg CH_4 yr ⁻¹ with a |
| 45 | mean estimation of 7 Tg CH ₄ yr ⁻¹ for marine geological emissions [1–4]. Human activity is a |
| 46 | major source of carbon dioxide responsible for the imbalance of this compound in the |
| 47 | atmosphere. Global antropogenic fossil CO ₂ emissions was estimated at 38.0 Gt in 2019 [5]. |
| 48 | In marine environments, methane and carbon dioxide are also key molecules that are |
| 49 | encountered in hydrothermal systems and at cold seeps [6]. Their sources as well as their |
| 50 | mechanisms of formation are different from the processes that take place at these different |
| 51 | environments. These ecosystems are characterized by emissions of fluids rich in methane and |
| 52 | hydrogen sulphide that sustains the development of specific chemosynthetic communities [7– |

10]. Methane is either the product (methanogenesis) or the substrate (methanotrophy) of 53

biogeochemical redox reactions within both the sedimentary and the water columns [11–15]. 54 55 Oceanic methane can also be of abiotic origin, and may be generated from reduction of carbon dioxide or serpentinization processes [10–13]. 56

57 Although natural CO₂ seepages are found in sedimentary basins, it is mainly encountered in volcanic and hydrothermal areas [16–18]. In nature, the four main processes that 58 generate CO₂ are the decomposition of organic matter at low temperature, the oxidation of 59 hydrocarbons either via thermal reactions or by microbial sulfate reduction, the decomposition 60 of carbonates and degassing from the mantle [19]. In the case of hydrocarbon-derived CO₂, its 61 δ^{13} C is influenced by the abundance of the hydrocarbons and their isotopic signature under 62 63 reservoir conditions [20–22]. In sedimentary basins, positive or higher δ^{13} C values could be related to an inorganic source of CO₂ generated from the thermal breakdown of carbonate rocks 64 [22]. In contrast to methane, carbon dioxide seepages from sedimentary environments have 65 been less investigated [23]. 66

Earth scientists analyze molecular and isotopic composition of natural gases to interpret 67 their origin. The origin of natural gases containing (CH₄ or C1), ethane (C_2H_6 or C2), propane 68 (C₃H₈ or C₃) and carbon dioxide (CO₂), is commonly interpreted using binary genetic diagrams 69 of δ^{13} C-C1 versus C1/(C2 + C3) [24], δ^{13} C-C1 versus δ^{2} H-C1 [25,26] and δ^{13} C-C1 versus δ^{13} C-70 CO₂ [27]. A revision of these empirical diagrams has been done and it is based on more than 71 72 20000 samples [28]. It includes the genetic fields for primary microbial gases from CO_2 73 reduction and methyl-type fermentation, secondary microbial gases generated during petroleum biodegradation, thermogenic (ie thermal cracking of organic matter in deep-hydrocarbon 74 75 reservoirs) and abiotic gases. These diagrams are key tools to provide interpretations of gas origins. 76

The wide range of CO₂ and CH₄ sources and their potential mixing can lead to a broad 77 range of δ^{13} C signature; and at cold seeps characterized by a large number of gas seeps over 78

hundreds of square meters to square kilometers, themselves characterized by numerous gas
streams. Understanding gas transport and geochemical-transformation processes requires a
multifold sampling strategy to collect many gas samples over a wide range of molecular and
isotopic compositions.

Accordingly onboard measurement of δ^{13} C-CH₄ and δ^{13} C-CO₂, together with the 83 molecular composition of the gases, can be valuable as a decision-making tool on 84 oceanographic cruises to quickly characterize gas seep sites and also for scientific purposes to 85 define the gas emissions with the binary genetic diagrams δ^{13} C-C1 versus δ^{13} C-CO₂. For 86 instance, such information will help in targeting the more interesting samples and refining the 87 sampling resolution at gas emission sites, and thus avoiding sampling gas streams having close 88 89 molecular and isotopic compositions of CH₄ and CO₂ [29]. This was our primary incentive for the development of the GC-SSIM-CRDS. In the following paragraph, we describe the 90 specifications for the design of the GC-SSIM-CRDS and its detailed description, the 91 optimization of the method, and the evaluation of its accuracy and precision against well-proven 92 analytical methods. 93

Isotope-Ratio Mass Spectrometry (IRMS) is the most prevalent, reliable and accurate 94 analytical technique used to measure the δ^{13} C of C-bearing compounds, including methane and 95 carbon dioxide. Regardless of the compounds, the end product measured from this technique is 96 the relative abundance of ¹²CO₂ and ¹³CO₂ [30–32]. Accordingly, the analysis of compounds 97 98 other than carbon dioxide, e.g. methane, requires its conversion into CO₂ via a combustion step before entering into the mass detector. However, an Isotope-Ratio Mass Spectrometer is relative 99 100 expensive, requires significant space for its installation, in an environment devoid of strong vibrations, and therefore is not suitable for fast onboard analysis during oceanic expeditions. 101 102 cavity ring down spectroscopy (CRDS) has been proposed as an alternative technique to IRMS

[33–35], and is based on the absorption rate of a near-infrared laser light confined within a 103 104 closed optical cavity. The analyzer is compact and easy to operate, provides a fast and low-drift signal, and can be used both in laboratory and in field (e.g. shipboard analysis). The instrument 105 106 has been designed for continuous gas sampling through an optical cavity. The CRDS technique has been used for isotopic measurements of carbon dioxide, methane and water vapor at 107 atmospheric concentrations [33]. However, applications in the field of deep-sea fluid systems 108 109 like cold seeps to support the sampling strategy while providing reliable and accurate data remains scarce. The main issue to unlock such applications is the determination of the possible 110 influence of chemical compounds between each other and to understand how this affects the 111 measurement. 112

One of the major constraints related to CRDS applied to deep-sea environments is the 113 limitation in the volume of samples that can be collected and the heterogeneity in concentration. 114 Indeed, methane concentration in pore water ranges from nmol L^{-1} to mmol L^{-1} , while this 115 compound can represent more than 99 mol% of the total gas composition [16,36]. Similar 116 117 constraints hold true for carbon dioxide. These constraints have to be taken into consideration for the development of analytical methods, more particularly for the injection step of discrete 118 samples. In the case of oceanic samples, injection using manifold and syringe, or based on the 119 120 headspace technique are usually the most appropriate [37–39]. Moreover, previous studies have shown that analyses from CRDS instruments can exhibit bias in measurements due to 121 interference from other natural gases on the laser absorption. As an example, the presence of 122 H₂S can cause interference on δ^{13} C-CO₂ [40] which consequently leads to an overestimated 123 depletion in the measured δ^{13} C. The effect of air compounds (nitrogen, N₂; oxygen, O₂; argon, 124 Ar; water, H₂O; and helium, He) also affects significantly the determination of CO₂ and CH₄ 125 concentrations with the CRDS [38,39]. Other recent studies have shown biases induced from 126 ethane on the δ^{13} C-CH₄ [37,41]. Some of these studies proposed specific isotopic corrections 127

derived from a calibration step with known standards. Thus, cross sensitivity is measured by 128 creating a gas-dilution series to control the concentration of the compounds responsible for 129 interference. For instance, to correct the interference of ethane on δ^{13} C-CH₄, the concentration 130 of ethane needs to be precisely determined as there are cross-interferences with the 131 concentration of H₂O, CO₂ and CH₄. Often, multiple corrections are needed to obtain reliable 132 and accurate δ^{13} C-CH₄ values. According to the manufacturer of the CRDS analyzer used in 133 this study, methane at a concentration higher than 500 ppm could also be an interfering 134 compound on the isotopic measurement of δ^{13} C-CO₂ [42]. 135

Here, we present a coupled system consisting of a CRDS analyzer and a gas 136 chromatograph (GC) for upstream separation of the chemical compounds present in the 137 analysed gases followed by molecular analysis, and prior to the δ^{13} C analysis. This system 138 enables easy separation of CH₄ and CO₂ in a complex matrix of compounds commonly present 139 140 in natural gases using either a headspace device, a gas injector [43] or by applying direct manual injection with a syringe. A similar system coupling a gas chromatograph with a CRDS analyser 141 has been described once in the literature [44] to measure stable carbon isotope ratios of methane, 142 ethane and propane. The coupled devices are used 1) to directly and rapidly determine the 143 molecular composition of the natural gases, followed by 2) for fast analysis of the δ^{13} C of 144 methane and carbon dioxide. The influence of the amount of non-methane hydrocarbons, 145 carbon dioxide and hydrogen sulfide on the reliability of the method has been studied. 146

- 147
- 148 2 Materials and Methods

149 2.1 Description of the coupled GC-SSIM-CRDS system

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The system is composed of a modified gas chromatograph (GC) Agilent 6850 (Agilent
Technologies, Santa Clara, CA, USA) coupled with the SSIM-CRDS (Fig.1 and 2). The isotopic

analyzer used for this study is a G2201-i Cavity Ring-Down Spectrometer (CRDS) coupled 153 with a Small Sample Isotope Module (SSIM), both from Picarro®, USA. The CRDS G2201-i 154 can measure both δ^{13} C-CH₄ and δ^{13} C-CO₂. The SSIM is a sample management peripheral 155 designed to accommodate small gas volumes (< 20 mL) prior to injection into the analyzer. 156 This device is commonly used to inject gases via gastight syringes, samples Tedlar® bags or 157 directly by connecting a bottle onto it. The SSIM is composed of a 20 mL sample chamber, five 158 159 solenoid valves, an external vacuum pump, and an internal pressure sensor. It requires a pressurized supply of zero air dry gas. Every SSIM-CRDS analysis begins with two purge 160 cycles of the SSIM chamber using vacuum and zero air supply to remove any memory effects 161 between samples. Before and after each purge, the vacuum pump evacuates the SSIM chamber. 162 Once ready for injection, the SSIM chamber stays under vacuum and a gas sample can be 163 delivered into the chamber by vacuum transfer. Once the injection of the gas sample into the 164 165 SSIM chamber is complete, the Coordinator software from Picarro closes the solenoid valve in relation with the injection port and opens the one connected to the cavity to deliver the sample 166 into the CRDS cavity and measure the δ^{13} C-CH₄ and δ^{13} C-CO₂. This system allows an easy 167 way to accommodate and dilute samples. 168

The SSIM could also induces a dilution of the sample with zero air. For instance, if a 169 sample of 5 mL at atmospheric pressure is injected inside the SSIM, the pressure inside the cell 170 will be about 5/20*760 Torr. The zero air will be added inside the cell to allow a good transfer 171 of the gas to the CRDS. If the pressure of zero air is about 900 torr, this will create a dilution 172 of about 4.7. Further information on the use of the SSIM is described in supplementary material. 173 According to the manufacturer, isotopic measurements should be made in well-defined 174 ranges of methane concentrations. The so-called High Range (HR) is applied when the methane 175 concentration ranges between 10 and 1000 ppm, while the High Precision range (HP) is most 176 suitable for methane concentration ranging from 1.8 to 12 ppm. The GC has been modified to 177

meet the requirement for individual separation of CH₄ and CO₂ from natural gases after the 178 179 determination of their molecular composition. It is composed of an injector, a Shincarbon ST (80/100 mesh and 2 m x 3 mm) column, a µthermal conductivity detector (µTCD) and a 180 trapping loop of 7 mL. There are three possible ways to inject samples into the Agilent 6850 181 gas chromatograph using either a gas injector [43] of 5mL connected to the gas valve (Vi), a 182 gas-tight syringe to be inserted into the purge packed injector (PPI) or a headspace auto-sampler 183 184 connected to the PPI. Simultaneously, direct injection into the SSIM-CRDS can be performed using a gas-tight syringe (Fig.1 and 2). Environmental gas samples analyzed in this study were 185 injected directly inside the SSIM port by the use of a gas-tight syringe or by the use of the gas 186 187 injector to fill the 1 mL injection loop of the Vt valve. The amount of gas injected inside the injection loop was calculated using the ideal gas law and controlled by decreasing gently the 188 pressure of the gas inside the gas injector with the vacuum pump of the CRDS and by measuring 189 190 the pressure with a sensor.

The optimized analytical conditions applied to the GC are given in Table 1. Two softwares, Open lab from Agilent and Coordinator, were used for the GC and the SSIM-CRDS, respectively, to set and control the instrument parameters for analysis. A homemade subroutine was created for communication between the GC and SSIM-CRDS.

Although helium is commonly used as the carrier gas for natural gas analysis using GC- μ TCD as its thermal conductivity is much higher than that of hydrocarbons, it significantly affects the measurement of δ^{13} C-CO₂ with the CRDS due to collisional line broadening and Dicke narrowing effects [39]. Accordingly, nitrogen gas as a carrier has been chosen to limit these effects. The influence of nitrogen as a carrier gas on the influence of the GC performance was also studied and discussed.

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202 2.2 *Measurement principle*

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204 First, the molecular composition of the gas sample is determined from classical GCµTCD analysis. Next, the injected gas sample is transferred into a trapping loop installed on the 205 gas-switching 8-port valve (Vt). This valve is pneumatically activated, and the 7 mL loop which 206 is an open column is used to trap and store the CO₂ while the δ^{13} C-CH₄ measurement is taking 207 place. This system does not require the use of adsorbent or cooling for trapping CO₂. By 208 switching the Vt valve from one position to another, the loop can be connected to the injection 209 port of the SSIM device to deliver the gas sample under vacuum transfer. While a δ^{13} C-CH₄ 210 measurement starts, CO₂ is still retained on the column. Once CH₄ has been transferred to the 211 CRDS, the Vt valve is switched back to its initial position and wait for the elution of CO₂. Once 212 eluted, CO₂ is trapped and stored inside the loop until the opening of the SSIM. Finally the CO₂ 213 214 is injected inside the SSIM by vacuum transfer.

Valve Vt is regulated at 32 °C to avoid any ambient temperature variations and to maintain good repeatability of the trapping. Fig. 3 shows the flow path for trapping and the subsequent δ^{13} C analysis of methane and carbon dioxide by switching the valve from one position to another.

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220 2.3 *Reference Materials and calibration procedures*

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222 Certified δ^{13} C-CH₄ and δ^{13} C-CO₂ gas standards were used for the isotopic calibration 223 of the SIIM-CRDS analyzer alone and in the coupled configuration with the GC-SSIM-CRDS. 224 Four commercial certified δ^{13} C-CH₄ gas standards, supplied by Isometric Instruments (Victoria, 225 Canada), were used (Table 2). These gas standards are composed of methane diluted in 226 hydrocarbon-free air. The calibration of carbon dioxide was performed using four commercial

standards from -40.1‰ to -7.1‰ diluted into hydrocarbon-free air and with a given uncertainty of less than $\pm 0.5\%$ (Table 2).

The CRDS was first calibrated directly without the coupling system by injecting 5 229 replicates 5 mL of certified δ^{13} C-CH₄ gas standards inside the inlet port of the SSIM with the 230 help of gastight syringes (Hamilton) and Tedlar® sampling bags (Restek). Due to the 231 concentration of the CH₄ standards at 2500 ppm and the dilution occurring with the SSIM, the 232 volume of 5 mL of CH₄ was chosen to have approximately up to 500 ppm of CH₄ detected at 233 the CRDS. This allows to make molecular and isotopic measurements in the middle of the 234 concentration range of the HR mode. The same procedure was applied for the certified δ^{13} C-235 CO₂ gas standards by injecting 5 replicates of 500 µL of the standard at -7.12, -13.51 and -22.62 236 237 % and 50 µL of the standard at -40.1% to reach a CO₂ concentration of approximately 1000 ppm inside the CRDS. The δ^{13} C-CH₄ and δ^{13} C-CO₂ calibration with the coupling system was 238 achieved by the use of the headspace autosampler (Dani HSS86.50, italy) and a 1mL injection 239 240 loop. Compounds introduced inside the GC were trapped by the use of the loop trap and the Vt valve and then transfer to the SSIM-CRDS. The δ^{13} C-CH₄ certified standards were realized by 241 flushing 10 mL sealed vials with certified isotopic standard (table 2) at 500 mL min⁻¹. Only the 242 certified δ^{13} C-CO₂ standard at -40.1‰ was realized differently by injecting 1mL of the certified 243 244 standard (from a Tedlar® bag) inside vials previously flushed with zero air. As the Autosampler introduce a dilution inside the vial, because a pressurization step is needed for the filling of its 245 injection loop, it was not used to analyze the samples. The autosampler was used for the 246 automatization of the calibration procedure of the whole analytical line going from separation 247 of compounds, trapping of the gas of interest and isotopic analysis into the CRDS. 248

A preliminary test was made to check the stability of the response signal of the CRDS analyzer prior to analysis. Thus, 10 repeated injections of 5 mL of the δ^{13} C-CH₄ standard at -38.3‰ and 500 µL of the δ^{13} C-CO₂ standard at -23.9‰ were performed directly into the SSIM once the cavity reached its working temperature of 45 °C. This test was achieved with separate
runs for CH₄ and CO₂.

254 2.4 In Situ gas sampling

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256 Natural gas samples were collected during the Marsitecruise [16,29] in the sea of Marmara, at water depth ranging from -22 m to -1273 m, with a gas-bubble sampler (PEGAZ) 257 [36,45]. These samples were used for validating the system by comparing our measured δ^{13} C-258 CH₄ values with the results from Isolab b.v (The Netherlands) using a GC-C-IRMS. The 259 PEGAZ sampler was designed to collect gas bubbles and preserve the sample at the in situ 260 pressure. Aliquots of gas were then subsampled at lower pressure of 2 to 4 bars in 12 mL pre-261 evacuated vials from Labco[®] by connecting the PEGAZ sampler to a gas transfer system [46]. 262 The remaining gases were stored in metallic containers of 100, 200 or 1000 mL. 263

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265 2.5 Molecular gas analysis

The present study is mainly focus on the separation of CH₄ and CO₂ follow by δ^{13} C 266 analysis with a SSIM-CRDS. In order to achieve good isotopic measurement, it is necessary to 267 preliminary know the composition of gases from the samples to control the amount of gas 268 injected inside the GC to avoid overloading on the GC column or inside the CRDS. We have 269 chosen to work with samples with known concentrations of gases and previously analyzed. The 270 concentration of these gases are given in this study to give informations about the complexity 271 272 of the gas mixture from environmental samples. Molecular composition analysis was carried out both on-board and at the "Laboratoire des Cycles Géochimiques et ressources" (LCG) at 273 274 If remer. A gas chromatograph µGC R3000 from SRA equipped with a µTCD and a PoraPlot U 275 capillary column was used on-board to determine nitrogen, oxygen, methane to hexane and

carbon-dioxide concentrations. Hydrocarbons at low concentrations in the samples (< 1%-mol) were analyzed at the LCG using an Agilent 7890 A gas chromatograph equipped with a 32 m, 0.32 mm Porapak Q column. The uncertainty in the measurements was of \pm 2% for methane and carbon dioxide concentrations, and \pm 4% for the heavier hydrocarbons.

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281 2.6 Performance of the GC for separation and analysis of CH_4 and CO_2

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In order to optimize the separation of CH₄ and CO₂ with the nitrogen carrier gas, an 283 experimental design has been conducted. The idea was to have enough time between the two 284 peaks to allow enough time for the Vt valve to trap CH₄ and come back to its initial position 285 before trapping CO₂. The gas volume of each peak were studied. As the carrier gas flow and 286 the oven temperature have a direct influence on the resolution and peak shapes, they were 287 selected as factors for the experimental design. The Doehlert matrix [47] has been used on 288 several types of chromatography techniques. It is useful to optimize the separation between 289 290 species [48]. This matrix has been used in the separation of H₂, O₂, N₂ and CH₄ on Shincarbon ST column [49]. The application of the experimental design uses basically three steps: realized 291 the designed experiments, estimate the coefficients in a mathematics model, predicting the 292 responses and checking the model adequacy. The aim of this application was to find the best 293 compromise which allows a good separation between the peaks of CO₂ and CH₄ on the GC 294 regarding the carrier gas flow and the oven temperature. The number of experiments required 295 for the Doehlert experimental design (N) is given by N = n2 + n + n0, where n and n0 are 296 respectively the number of variables and n0 is the number of center points. In our case, the n0297 value was fixed at 3; thus, with two factors (gas flow and oven temperature). The number of 298 experiments was 9. The carrier gas flow was studied at five levels (6, 7, 8, 9, 10 mL min⁻¹) and 299

temperature at three levels (161, 180 and 199°C). Three responses were determinated; peaks width of the two gases and the difference of retention time between the compounds. Further information on the experimental design is provided in the supplementary material. All the calculations related to the experimental design were performed with the statistical software NemrodW. The significance of each coefficient was determined by the Student test. Hence, it is assumed that the residues are randomly distributed with a constant variance and the observed response data points follow normal distribution.

Once the parameters of the carrier gas flow and the oven temperature defined, the GC configuration was evaluated for separation of CO_2 and CH_4 in a complex gas mixture (1ml of $CO_2 5\%$ CLM-3783-10 + 1 mL of a gas mixture of 1% of C1 to C5 in nitrogen from Air Liquide) by injecting this 2 mL of this mixture inside the gas injector of the GC. The injected volume on the GC is 1mL. CO_2 and the C1-C5 mixture have also been injected separately for peaks identification.

A study was also made to determine the limits and compromises in term of gas concentrations in which the GC-SSIM-CRDS is able to manage for consecutive (CH_4 then CO_2) or individual runs. The limit of detection for both gases were evaluated using a mixture of CH_4 and CO_2 (2% in air balance from Messer) at different pressures and by using the 0.1 mL injection loop of Vi.

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319 2.7 Statistics used to assess the effect of concentration variation on the $\delta^{13}C$ -

320 *CH*₄

Method accuracy was evaluated by measuring bias and precision. The maximal bias (MB) was defined as being 0.8‰ as we believe that a precision of <1‰ allows to reliably estimate the gas sources or the level of mixing. The measurement of trueness was expressed as follows:

325
$$Bias(\%) = \frac{|z - ref|}{ref} * 100$$

With z being the average value obtained from five series of duplicate measurements and ref, the certified δ^{13} C-CH₄ value. Precision was estimated by the relative standard deviation calculated from the intra and inter-series variances. Finally, the method was validated in terms of accuracy control if it satisfies the following constraints:

$$ref + MB > z + 2s \qquad ref - MB < z - 2s$$

The bias is considered negligible to estimate the uncertainty (U) at k = 2 with normalized error (NE) < 2.

333
$$NE = \frac{|z - ref|}{\sqrt{s^2 + uref^2}} \qquad U = 2^* \sqrt{s^2 + uref^2}$$

With s being the standard deviation of the series, uref is the uncertainty of the reference value,and k the coverage factor in the uncertainty estimates for a 95% confidence interval.

336

337 3 Results and discussion

338 *3.1 CRDS stability and injection repeatability*

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Regarding the preliminary test, results show that the CH₄ and CO₂ concentration measurements from the ten consecutive injections remain relatively constant, with a mean value of 539 ppm and 624 ppm, and a relative standard deviation RSD of 0.7% and 0.6% for CH₄ and CO₂, respectively (see supplementary). While methane concentration at the detector remains stable over the entire injection series, the δ^{13} C-CH₄ varies significantly for the first five injections with an amplitude of 0.9‰ for the two tested standards T-iso1 (-38.3‰ for δ^{13} C-CH₄) and CLM-3781 (-22.7‰ for δ^{13} C-CO₂), and such behavior has already been pointed out

by the manufacturer [42] in case of hyphenated systems such as the G2121-i CRDS analyzer 347 348 coupled with a combustion module (CM-CRDS). In such a case, they recommended to calculate the standard deviation of six consecutive analyses, and then to check that the obtained value is 349 under a specific threshold depending on the instrument. The G2201-i instrument uses a 350 wavelength monitor with a line-locking algorithm to detect and analyzed the ${}^{13}C$ and ${}^{12}C$ of the 351 CH₄ or CO₂. This means that the CRDS analyzer needs a certain time to run and stabilize the 352 signal. As the instrument is working with the SSIM to deliver a small volume of gas inside the 353 cavity, the CRDS is then working in a closed system without continuously analyze ambient air 354 and detect CH₄ and CO₂ for its line-locking detection and positioning. Therefore, the first five 355 356 measurements were automatically discarded from our measurement series.

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358 3.2 Influence of the gas concentration on δ^{13} C-CH₄ and δ^{13} C-CO₂

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As described previously, the G2201-i analyzer was designed to measure the δ^{13} C-CH₄ and δ^{13} C-CO₂ at specific ranges of concentrations. Fig. 4 shows the δ^{13} C-CH₄ measurements in the concentration range recommended by the manufacturer, which is the HR mode (between 12 to 1000 ppm of methane). This test was achieved directly by the used of the SSIM-CRDS without the GC.

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Clearly, there is an influence of the methane concentration on δ^{13} C-CH₄ when moving towards the lowest concentration limits of the HR mode. For methane concentrations below 25 ppm, the measured δ^{13} C-CH₄ varies significantly, and several values fall outside the certified accuracy of the standards (in the range of -38.1 to -38.5‰). This may be explained by the socalled optical extinction due to scattering processes [50,51]. Nevertheless, these results are in

agreement with the precision provided by the manufacturer with a standard deviation (σ_p) of 371 0.55% for δ^{13} C-CH₄ [42]. As the GC-SSIM-CRDS was designed to analyze mainly methane-372 concentrated gas samples, its use at concentrations close to the lowest limit of the HR mode is 373 unlikely. In order to achieve reliable measurements, the isotopic analysis of methane was 374 carried out in HR mode with methane concentration higher than 25 ppm. The HR mode was 375 also assessed in the concentration range between 50 to 1000 ppm (Table 3) with a maximum 376 bias (MB) value of 0.8‰, the method is validated for all four isotopic standards and in a range 377 of 55 to 1011 ppm of CH₄ injected within the CRDS cavity. The averaged values of measured 378 δ^{13} C-CH₄ fall within the uncertainty range of the certified materials, and below the maximal 379 380 bias defined in this study.

Fig. 5 shows the variation of δ^{13} C-CO₂ with its concentration ([CO₂]) and for direct 381 injection inside the SSIM-CRDS. For each certified standard, there is a linear relationship 382 between δ^{13} C-CO₂ and 1/[CO₂]. The measured δ^{13} C-CO₂ is enriched in ¹³C with increasing 383 [CO₂]. Such physical effect was reported in the literature [39], and was explained as the 384 differences between ¹²CO₂ and ¹³CO₂ regarding adsorption/desorption process of the light 385 circulating through the cavity. The observed effects was described to give positive δ^{13} C bias 386 correlated with increasing CO₂ concentration. However, such linearity clearly shows that it is 387 possible to link the δ^{13} C with 1/[CO₂]. Comparison between measured and certified δ^{13} C-CO₂ 388 was carried out for different CO₂ concentrations (Fig.6). Previous works have shown the 389 possibility to linearize the δ^{13} C-CO₂ versus the variation of 1/[CO₂] with a GG2201-i analyzer 390 [52]. We then studied the influence of the variation of $[CO_2]$ on the typical plot use to correct 391 the δ^{13} C-CO₂ (ie δ^{13} C-CO₂ measured vs δ^{13} C-CO₂ certified) 392

Equation (1) was used to correct measured δ^{13} C-CO₂ values in accordance to certified values. The figure also shows the effect of [CO₂] variation on the correlation curves. Both the slope and the intercept could be linearized regarding the concentration (Fig. 7) to give

| 396 | respectively equation (1.1) and equation (1.2) . These equations are then used to obta | in the final | | | | | |
|-----|--|------------------------|--|--|--|--|--|
| 397 | equation (1.3) which is the δ^{13} C-CO ₂ corrected from [CO ₂] measured inside the analyzer for a | | | | | | |
| 398 | range of concentration in accordance with the specifications from Picarro (380-200 | 0 ppm) and | | | | | |
| 399 | for δ^{13} C-CO ₂ ranging from -7.2 to -40.1 ‰. | | | | | | |
| 400 | $\delta^{13}C = \delta^{13}C \text{ measured}^*A_N + B_N$ | (1) | | | | | |
| 401 | | | | | | | |
| 402 | with A_N the slope and B_N the intercept of the normalization correction. | | | | | | |
| 403 | | | | | | | |
| 404 | $A_N = a/[CO_2] + b$ | (1.1) | | | | | |
| 405 | | | | | | | |
| 406 | with a and b the slope and intercept for A_N , respectively. | | | | | | |
| 407 | | | | | | | |
| 408 | $B_{\rm N} = \alpha / [CO_2] + \beta$ | (1.2) | | | | | |
| 409 | | | | | | | |
| 410 | with α and β the slope and intercept for B_N , respectively. | | | | | | |
| 411 | | | | | | | |
| 412 | $\delta^{13}C = \delta^{13}C \text{ measured}^*(a/[CO_2] + b) + (\alpha/[CO_2] + \beta)$ | (1.3) | | | | | |
| 413 | | | | | | | |
| 414 | with a = 2.5814, b=1.0147, α = 2211.6 and β = -1.8389 and δ^{13} C measured (‰) and | [CO ₂] the | | | | | |
| 415 | CO ₂ concentration measured at the CRDS in ppm. | | | | | | |

417 3.3 Performance of the GC separation for trapping of CH_4 and CO_2 .

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The carrier flow of 6 mL min⁻¹ and the oven temperature of 184°C obtained from the 419 Doehlert optimization procedure allow a nice separation of CH₄ and CO₂ on the Shincarbon ST 420 column. These two conditions were selected to have enough time for the system to trap the CH₄ 421 422 and send it to the CRDS while the CO₂ was still retained and passing through the Shincarbon ST column. The volumes of each of the two gases were used as responses in the Doehlert matrix 423 by multiplying the peaks width with the total gas flow at the end of the GC system. Total gas 424 flow was given by the addition of the carrier gas flow of 6 mL min⁻¹ and the detector flow of 9 425 mL min⁻¹. With a carrier gas flow of 6 mL min⁻¹ and the oven temperature of 184°C, peaks 426 width were well constrained during the Doehlert optimization procedure giving 0.26 min for 427 CH₄ and 0.34 min for CO₂. 428

429 The GC system was evaluated for the separation of CH₄ and CO₂ in a mixture of C1 to 430 C5 compounds CH₄ and CO₂ are well separated from a matrix of gases from C2 to C5. Retention times of CH₄ and CO₂ are respectively of 3.0 and 4.1 min. It was not possible to elute and 431 separate gases > C3. C2 was eluted by the use of thermal and N_2 flow gradients presented on 432 Figure 8. To avoid any little accumulations of C3 to C5 (as these are present at lower 433 concentrations) which can progressively change the GC retentions of CH₄ and CO₂, it has been 434 chosen to regenerate the column during 1h at a carrier gas flow of 20 mL min⁻¹ and with an 435 oven temperature of 280 °C every 24 samples. No variation of retention times were observed. 436 The CRDS analyzer needs approximately 10 min to run a sample and so the global time of 437 438 analysis of CH₄ followed by CO₂ is about 24 min.

As previously mentioned, N₂ was chosen as a carrier gas thanks to its compatibility with
 CRDS. The main influence of choosing nitrogen as a carrier gas compare to helium is the lack

of contrast regarding the detection of CO_2 as they have similar thermal conductivity [53]. 441 442 According to the literature, at 20°C, the respective thermal conductivity of He, N₂, CH₄ and CO₂ are respectively of 150, 24, 35 and 17 mW mK⁻¹. In this configuration, the limits of 443 detection were evaluated for CH₄ and CO₂. The limits of detection for CH₄ and CO₂ with N₂ is 444 respectively of 6 ppm and 30 ppm. The limits of detection for CH₄ and CO₂ with Helium are 445 1.4 ppm and 2.1 ppm. As the isotopic measurements of δ^{13} C-CH₄ and CO₂ have to be operate 446 in a valid range of concentrations inside the CRDS, we recommend to first measure the 447 concentrations of gases and then to make isotopic measurements. The coupling system 448 developed in this study allows to control the amount of gas injected on the GC column easily 449 450 by the use of a gas injector. It is possible to decrease the amount of gas loaded on the injection loop by decreasing gently the pressure inside the gas injector by the use of the CRDS pump 451 directly connected onto it and a needle valve. The amount of a gas injected on the column has 452 453 a direct influence on the peak width and on the concentration detected at the CRDS. We have defined compromises in terms of maximum injected gas amounts by measuring the peaks 454 volumes of each gas and by checking the CRDS gas concentrations measured. Table 4 shows 455 the range of concentrations of CH₄ and CO₂ which have to be injected inside the GC-SSIM-456 CRDS system for appropriate δ^{13} C measurements with the CRDS. 457

458 The lowest limit of methane concentration which can be injected on the GC prior to δ 13C-CH₄ analysis on the CRDS is about 0.14% in HR mode. As the HP mode of the CRDS 459 use a concentration of CH₄ low as 1.8 ppm it is potentially possible to work with gas sample as 460 methane concentration as low as 0.010%. Compromises have been made by evaluating the 461 maximum amount of gases tolerated on the column which limit peaks width of the two 462 compounds and also correspond to the range of concentrations needed for valid isotopic 463 measurements. The most important factor was the defined range of concentration needed for 464 the CRDS. The range of concentrations which permit sequential isotopic analysis for CH₄ and 465

CO₂ is low. Due to this narrow range and to the high differences of concentrations between the 466 two gases in environmental samples, δ^{13} C-CH₄ and δ^{13} C-CO₂ were mostly measured 467 individually. The main limitation of the system is that a minimum concentration of 1.1% of 468 CO₂ is needed for the δ^{13} C-CO₂ measurement as the CRDS needs a minimum concentration of 469 380 ppm of CO₂. In perspective to improve the system, it is possible to increasing the trapping 470 volume to study the possibility of transferring more gas to the CRDS. A work on GC parameters 471 can also be done to increase the volume of the injection loop of the Vi valve to increase the 472 amount of gas injected for lower CO₂ concentration. The study of the influence of the gas 473 concentration of CO₂ on the δ^{13} C-CO₂ have shown a good linear relation between the 1/[CO₂] 474 and the δ^{13} C-CO₂. Some of these measurements were done at lower concentration than 380 475 ppm. For instance, the certified δ^{13} C-CO₂ standard of -7.2 ‰ was analyzed at a concentration 476 of 323 ppm with still a good linear relation between δ^{13} C-CO₂ and the inverted gas 477 478 concentration. This also means that it is possible to work to extend the limit of the CO₂ concentration needed for δ^{13} C-CO₂ measurements and to analyze lower concentrated gases with 479 the coupling system. 480

481 3.4 Calibration of the GC-SSIM-CRDS

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Fig. 9 shows good agreement between the calculated and certified δ^{13} C values for methane and carbon dioxide after correction. The linearity is highly satisfactory over the full δ^{13} C range explored, from -7.2‰ to -40‰ for CO₂ and -23.9 to -66.5‰ for δ^{13} C-CH₄. These results show the possibility to use a headspace autosampler for the automatization of the calibration procedure.

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489 3.5 Application to natural gas samples and comparison with GC-C-IRMS

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491 Gas samples collected at several natural gas seep sites of the Marmara sea and characterized by a large range of methane concentrations (between 66 and 99.6 mol% of total 492 gas) were analyzed with the GC-SSIM-CRDS to study the influence of compounds other than 493 methane and carbon dioxide on δ^{13} C of the two latter. The results were compared with those 494 obtained from GC-C-IRMS analysis [22]. There is good agreement on the measured δ^{13} C-CH₄ 495 between the two analytical techniques (Table 5), with a maximum RSD of 1.4%. The results 496 also show that the other compounds do not significantly affect the δ^{13} C-CH₄. For samples 497 denoted ND, it was not possible to measure the δ^{13} C-CO₂ as their CO₂ concentrations were 498 under 1.1%. Due to multiple assays and losses of samples, there was no more gases available 499 for the analysis of δ^{13} C-CH₄ for DV3PE3 and for the analysis of δ^{13} C-CO₂ for Marpegas 500 501 4/western high (samples denoted NA).

Table 6 compares the results of δ^{13} C-CO₂ measurements carried out with the non-coupled 502 SSIM-CRDS, the GC-SSIM-CRDS and the GC-C-IRMS. It also shows the concentration of 503 CO₂ and CH₄ in ppm measured with the CRDS. The direct analysis of the sample with the 504 SSIM-CRDS without the coupling with the GC leads to very different δ^{13} C-CO₂ values to those 505 506 obtained when using the GC-SSIM-CRDS or the GC-C-IRMS; except for the CO₂-rich sample 507 DV3-PE3. This sample is the only one where CO_2 is largely dominant and exhibits the lowest 508 methane concentration. The four other samples are mainly composed of methane at 509 concentrations ranging between 82.4 and 94.8 mol-%, and such high concentrations make for the unreliable analysis of δ^{13} C-CO₂ and clearly justify our coupling system GC-SSIM-CRDS. 510 Indeed, the separation step of the GC-SSIM-CRDS makes it possible to obtain reliable δ^{13} C-511 512 CO₂ measurements by avoiding the spectroscopic interferences due to high methane 513 concentration, with accuracy similar to what is achieved with the GC-C-IRMS (Fig. 11). In this example, the difference between the direct injection through the SSIM and the GC-SSIM-514

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| 515 | CRDS is that a high content of methane interferes in the measurement of δ^{13} C-CO ₂ even if this |
|-----|--|
| 516 | gas is defined in the range of 380 to 2000 ppm. By injecting the sample through the GC before, |
| 517 | it allows to remove this interference. |
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| 525 | 4 Conclusion |
| 526 | |
| 527 | We have designed and built a coupled system, the GC-SSIM-CRDS, enabling fast |
| 528 | measurement of stable carbon isotope ratios of methane and carbon dioxide, together with the |

molecular composition of natural gases. This system is transportable and easy-to-use onboard

during oceanic cruises. The GC-SSIM-CRDS separates the different components of a natural-

gas sample on a packed column, and measures its molecular composition with a thermal

conductivity detector. The gas components are sequentially trapped in a loop, avoiding

interference effects before being transferred to the SSIM-CRDS for the δ^{13} C determination of

534 CH₄ and CO₂.

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Several tests have been performed to evaluate the capability of the GC-SSIM-CRDS to accurately measure both isotopic and molecular compositions of natural gases characterized by a wide range of methane and carbon dioxide concentrations. The measured δ^{13} C-CH₄ and δ^{13} C-CO₂ are in agreement with the values obtained from the commonly applied and well-proven analytical method using a GC-C-IRMS.

Besides providing reliable and accurate molecular composition, δ^{13} C-CH₄ and δ^{13} C-CO₂ of natural gases, the GC-SSIM-CRDS can be useful during the exploration step of cruises devoted to the investigation of large-scale gas-emission sites on continental margins and

shelves. Indeed, in the context of climate change and its impact on the ocean, and considering the widespread occurrence of gas emissions on continental margins and shelves, the GC-SSIM-CRDS is a valuable tool for decision-making during sampling and onboard analysis. It would thus allow for smart sampling of the gas seeps by identifying the most relevant gas seeps to sample (the ones characterized by gas streams with highly heterogeneous molecular composition, δ^{13} C-CH₄ and/or δ^{13} C-CO₂), and would consequently lead to a better understanding of gas sources, transport processes and fates in the ocean.

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551 Author contributions

Christophe Brandily designed and built the coupled system, optimized the analytical method, 552 conducted data acquisition and analyses, and wrote the paper. Nolwenn Le Cuff designed and 553 built the coupled system, optimized the analytical method, and conducted data acquisition and 554 analyses. Alexis De Prunele contributed to the optimization of the analytical method for the 555 556 CRDS. Cécile Cathalot reviewed the paper. Claire Croguennec contributed to the optimization of the analytical method for the CRDS. Jean-Pierre Donval designed and built the coupled 557 system, optimized the analytical method, and conducted data acquisition and analyses. Vivien 558 Guyader designed and built the coupled system, optimized the analytical method, and 559 conducted data acquisition and analyses. Jean-Claude Caprais initiated the acquisition of the 560 561 CRDS analyzer and developed the methods for this instrument. Livio Ruffine led the Carnot Project, designed the coupled system, interpreted the data and wrote the original draft. Thanks 562 563 to Alison Chalm for proofreading the article.

564

565 **Declaration of competing interest**

The authors declare that they have no known competing financial interests or personalrelationships that could influence the work reported in this paper.

568

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- 574

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752753 Table captions

- 754
- Table 1. Analytical conditions applied to the GC and the Vi and Vt valves.

756

- Table 2. Commercial Certified δ^{13} C-CH₄ and δ^{13} C-CO₂ gas standards used for this study.
- Table 3. Statistic test of δ^{13} C-CH₄ with the [CH₄] (ppm) variations for the HR mode with z the
- average from 10 replicates, NE the normalized error, U the uncertainty, uref the uncertainty
- 760 from the reference value, MB the maximal bias and ref the isotopic certified value.
- 761

| \sim | | D | 10 | n | | |
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| U | un | | | p. | | |

| 762 | Table 4. Limits of concentration for the gas samples injected into the GC-SSIM-CRDS with a |
|-----|--|
| 763 | 1 mL injection loop and the gas injector |
| 764 | |
| 765 | Table 5. Comparison of δ^{13} C-CH ₄ and δ^{13} C-CO ₂ measured with the SSIM-CRDS and the GC- |
| 766 | C-IRMS for 16 natural gases. Gases compositions (% mol) were determined in previous |
| 767 | studies. |
| 768 | |
| 769 | Table 6. Comparison between the GC-SSIM-CRDS, the SSIM-CRDS and GC-C-IRMS |
| 770 | measurements of $\delta^{13}\text{C-CO}_2$ for 5 natural gas samples from the Marmara sea. Gases |
| 771 | compositions (% mol) were determined in previous studies. |
| 772 | |
| 773 | |
| 774 | |
| 775 | Figure captions |
| 776 | |
| 777 | Fig. 1. Schematic overview of the coupled system: Gas Chromatography 6850 Agilent |
| 778 | implemented with an 8-port valve, Vt, connected to SSIM-CRDS. A head-space sampler or a |
| 779 | gas tight syringe can be coupled to the GC using the purge packed inlet injector (PPI). |
| 780 | |
| 781 | Fig. 2. Photo of the GC-SSIM-CRDS. |
| 782 | |
| 783 | Fig. 3. Chromatogram of the trapping and subsequent $\delta^{13}C$ analysis of methane and carbon |
| 784 | dioxide by switching from one position to another. (a) Vt OFF: the CO ₂ is trapped in the loop |
| 785 | during the δ^{13} C-CH ₄ measurement; (b) Vt ON: the CO ₂ is delivered to the SSIM-CRDS for the |
| 786 | measurement of its δ^{13} C. |
| 787 | |

Fig. 4. Variation of δ^{13} C-CH₄ with CH₄ concentration (ppm) in the HR mode from 29 sample

injections of the standard at of -38.3‰ (+/- 0.2‰) of δ^{13} C-CH₄. σ_p represents the standard deviation reported by Picarro.

791

Fig. 5. Fluctuation of δ^{13} C-CO₂ with the reciprocal CO₂ concentration for each of the four certified standards.

794

Fig. 6. Influence of concentration of the injected CO₂ on δ^{13} C-CO₂ calibration.

796

Fig. 7. A_N and B_N plots versus $1/[CO_2]$.

798

Fig. 8. Chromatogram of CH_4 and CO_2 in a mixture of C1 to C5.

800

Fig. 9. Calibration curve obtained with the GC-SSIM-CRDS for δ^{13} C-CH₄ and δ^{13} C-CO₂ for 5 replications.

803

Fig. 10. Graph showing the good agreement on the δ^{13} C-CH₄ values measured with the SSIM-CRDS and the GC-C-IRMS for 16 natural gas samples.

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807

Fig. 11. Comparison between δ^{13} C-CO₂ values measured with the GC-SSIM-CRDS and the GC-C-IRMS for 4 natural gas samples. The concentrations of CO₂ and CH₄ in mol% were determined previously (not with this system).

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| Injection valve box | | | | | | | |
|---------------------|------------------------|-----------|--|--|--|--|--|
| Sample loop size | 0.1 and 1 mL | <u>c</u> | | | | | |
| Temperature | 32 °C | 0 | | | | | |
| Column | | | | | | | |
| Туре | Shincarbon (ST 80/100, | 2m x 3mm) | | | | | |
| Flow | 6 mL min ⁻¹ | | | | | | |
| Temperature | 184 °C | | | | | | |
| μΤCD | $\langle \rangle$ | | | | | | |
| Temperature | 250 °C | | | | | | |
| Reference flow | 9 mL min ⁻¹ | | | | | | |
| Valve | time | Position | | | | | |
| Vi | 0 min | off | | | | | |
| Vt | 3.2 min | on | | | | | |
| Vt | 3.6 min | off | | | | | |
| Vt | 4.3 min | on | | | | | |
| Vt | 13 min | off | | | | | |
| Vi | 13 min | on | | | | | |

| Gas | Standard | Origin | [gas] mol% in air balance | δ ¹³ C vs PDB ‰ | $\pm\sigma$ ‰ |
|--------|---------------------------------------|-----------------------------------|------------------------------|-------------------------------|---------------|
| | CLM-3783-10 High level Gas | | 5 | -7.12 | 0.4 |
| | CLM-9026-10 Mid level Gas | Cambridge isotope laboratories | 5 | -13.51 | 0.5 |
| CO_2 | CLM-3781-10 Baseline Calibrant Gas | | 5 | -22.62 | 0.2 |
| | UN1956 | AirGas (Air Liquide) | 50 | -40.1 | 0.3 |
| | L-iso1 | 2 | 0.25 | -66.5 | 0.2 |
| CU | B-iso1 | T | 0.25 | -54.5 | 0.2 |
| CH_4 | T-iso1 | Isometric instrument | 0.25 | -38.3 | 0.2 |
| | H-iso1 | | 0.25 | -23.9 | 0.2 |
| | Jonu | | | | |

| | | | | | | | | | _ |
|--|---------------|---------------|-------------|-------------|---------------------------|-----|--------|------|---|
| δ ¹³ C-CO ₂ ‰ | ref+uref ‰ | ref-uref ‰ | ref+MB ‰ | ref-MB ‰ | [CO ₂] ppm | U | z ‰ | NE | |
| | | | | | 55 | 0.4 | -66.44 | 0.33 | |
| -66.5 | -66.3 | -66.7 | -65.7 | -67.3 | 542 | 0.4 | -66.45 | 0.45 | |
| | | | | | 1015 | 0.4 | -66.34 | 0.80 | |
| | | | | | 55 | 0.4 | -54.45 | 0.25 | |
| -54.5 | -54.3 | -54.7 | -53.7 | -55.3 | 545 | 0.4 | -54.50 | 0.41 | |
| | | | | | 1023 | 0.4 | -54.38 | 0.55 | |
| | | | | | 55 | 0.4 | -38.31 | 0.05 | |
| -38.3 | -38.1 | -38.5 | -37.5 | 7.5 -39.1 | 533 | 0.4 | -38.40 | 0.76 | |
| | | | | | 1011 | 0.5 | -38.24 | 0.27 | |
| | | | | | 55 | 0.4 | -23.83 | 0.32 | |
| -23.9 | -23.7 | -24.1 | -23.1 | -24.7 | 541 | 0.4 | -23.80 | 0.55 | |
| | | | | | 1021 | 0.4 | -23.77 | 1.56 | |
| | | | | | | | | | - |

Table 4

| Table 4 | | | | |
|-----------------|---|---|--|--|
| Compound | Limit of detection (% mol) on the GC | Low gas concentration limit (% mol) to inject for isotopy | High gas concentration limit (% mol) to inject for isotopy | maximum amount of gas (nmol) to inject on the GC for isotopy |
| CH4 | 0.0006 | 0.14 for HR and 0.010 for HP | 5.4 | 2200 |
| CO ₂ | 0.0030 | 1.1 | 4.4 | 1800 |

| Cruise/Sample name | | % mol (previously determined) | | | | | δ ¹³ C ‰ | δ ¹³ C-CH ₄ ‰ | δ ¹³ C ‰ | δ ¹³ C-CO ₂ ‰ |
|--------------------|--------------------|-------------------------------|------|---|------|-----------------|---------------------|-------------------------------------|---------------------|-------------------------------------|
| | | methane | C2+ | CH4 GC Carbon Hydrogen CRDS dioxide sulfide | | CH4 GC- CRDS | Isolab | CO2 GC- CRDS | Isolab | |
| | MRS-DV1-PE02 | 98,71 | 0,49 | 0,55 | 0,25 | 0 | -44,1 | -43,5 | ND | ND |
| | MRS-DV1-PE03 | 98,65 | 0,85 | 0,14 | 0,36 | 0 | -53 | -53 | ND | ND |
| | MRS-DV2-PE02 | 82,39 | 9,01 | 8,6 | 0 | 0 | -43,5 | -44 | 28,3 | 28,4 |
| | MRS-DV3-PE01 | 94,76 | 3,44 | 1,59 | 0,21 | 0,18 | -52,1 | -52,3 | 11,8 | 12 |
| | MRS-DV3-PE06 | 94,48 | 3,39 | 1,95 | 0,18 | 0,1 | -52,1 | -52,2 | 17,6 | 16,9 |
| | MRS-DV3-PE09 | 99,53 | 0,13 | 0,11 | 0,22 | 0,07 | -58 | -58,4 | ND | ND |
| Marsite | MRS-DV4-PE02 | 99,79 | 0,01 | 0,1 | 0,1 | 0 | -64,1 | -63,8 | ND | ND |
| 2014 | MRS-DV4-PE07 | 99,67 | 0,01 | 0,09 | 0,24 | 0 | -65,8 | -66,1 | ND | ND |
| | MRS-DV4-PE08 | 99,53 | 0,02 | 0,09 | 0,36 | 0,02 | -65,8 | -66 | ND | ND |
| | DV3-PE3 | 2,01 | 9 | 97,64 | 0,2 | 0 | NA | -34 | -3,5 | -3,7 |
| | MRS-DV5-PE01 | 99,68 | 0,01 | 0,14 | 0,17 | 0 | -63,4 | -63,5 | ND | ND |
| | MRS-DV5-PE02 | 99,74 | 0,01 | 0,13 | 0,12 | 0 | -62,6 | -63,1 | ND | ND |
| | MRS-DV5-PE03 | 99,74 | 0,01 | 0,09 | 0,16 | 0 | -63,6 | -63,8 | ND | ND |
| | MRS-DV5-PE04 | 99,61 | 0,09 | 0,18 | 0,11 | 0,02 | -62 | -62,1 | ND | ND |
| Marpegas | s1 / Central High | 98,86 | 0,52 | 0,36 | 0,26 | NA | -44,2 | -44,4 | ND | ND |
| Marpegas | 4 / Western High | 90,90 | 5,2 | 3,9 | 0 | NA | -44,6 | -44,4 | NA | NA |
| Marpegas | 5 / Cinarcik Basin | 99,63 | 0,01 | 0,1 | 0,26 | NA | -64 | -64,1 | ND | ND |
| | | | | | | | | | | |
| | | | | | | | | | | |
| | | | | | | | | | | |
| | | | | | | | | | | |

| Sample | GC-SSIM- | IRMS | SSIM- | [CO ₂] | [CH4] | CO_2 | CH ₄ |
|----------|----------|-------|--------|--------------------|----------|--------|-----------------|
| | CKDS/ ‰ | %00 | CRDS ‰ | CRDS ppm | CRDS ppm | mol % | mol % |
| DV2-PE2 | +28.3 | +28.4 | +2806 | 592 | 8030 | 8.60 | 82.39 |
| DV3-PE3 | -3.5 | -3.7 | -3.78 | 1101 | 26.1 | 97.64 | 2.01 |
| DV3 PE01 | +11.8 | +12.0 | +11287 | 552 | 33463 | 1.59 | 94.76 |
| DV3 PE06 | +17.6 | +16.9 | +9606 | 548 | 26676 | 1.95 | 94.48 |





















 δ^{13} C (CH₄) ‰ GC-SSIM-CRDS



Highlights:

- We present a coupled analytical system associating gas chromatography and cavity ring down spectroscopy (GC-SSIM-CRDS) for natural gas analyses
- The analytical system allows fast onboard analyses of both molecular composition of natural gases and stable carbon isotope ratios of methane and carbon dioxide in 24 min.
- The GC-SSIM-CRDS was optimized and validated with samples from natural gas seeps
- The GC-SSIM-CRDS is a decision-making tool suitable to refining sampling strategy of gases when characterizing seep areas and is a useful analytical system for scientific purposes.

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Declaration of interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

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