

## Hydrology and small pelagic fish drive the spatio-temporal dynamics of springtime zooplankton assemblages over the Bay of Biscay continental shelf

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### Abstract :

As mesozooplankton is the preferential prey of small pelagic fish (SPF), environmentally-driven mesozooplankton dynamics can have critical effects on SPF population dynamics. Despite previous studies on SPF habitats' dynamics, hydrological landscapes and mesozooplankton dynamics in the Bay of Biscay (BoB), knowledge gaps persist at the BoB regional-scale pelagic ecology and in particular about the mesozooplankton assemblages and their long-term space-time patterns. Here, we present 16 years of spring mesozooplankton assemblage interannual spatial dynamics over the BoB continental shelf and we describe the correlations between the mesozooplankton space-time patterns and those in hydrology, primary producers and SPF. We gathered data originating from the PELGAS surveys (2004–2019) and remote sensing products. Mesozooplankton samples were collected with a 200- $\mu$ m mesh size WP2 net vertically towed from 100 m depth (or 5 m above the sea floor) to the surface. They were analysed with imaging and deep-learning tools and the biomass in 24 coarse taxonomic groups was calculated. Automated procedures for spatial gridding and missing data imputation enable the generation of yearly maps time series with the same spatial resolution across the pelagic ecosystem components and years. These comprehensive multivariate datasets were analysed with a multi-table method known as Multiple Factor Analyses to depict time-consistent spatial patterns in each ecosystem component and the temporal variability around them. Finally, the main time-consistent spatial patterns in the hydrology, primary producers and SPF ecosystem components were used as predictors in generalized linear models, to explain those in the mesozooplankton. Mesoscale coastal-offshore and north-south gradients were the main patterns observed in each of the pelagic ecosystem components studied. The spatial patterns in the mesozooplankton assemblage were stable, without any significant changes detected in the taxonomic composition nor its spatial structure over the studied period. Small copepods, gelatinous and meroplanktonic organisms characterised coastal areas. Euchaetidae and meroplanktonic crustaceans' larvae displayed higher biomass in the northern part of the BoB while Metridinidae, Cladocera,

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Appendicularia and Echinodermata had higher biomass in the southern part. Surface and bottom water temperature, salinity-related parameters, water column stratification and SPF biomasses were the variables that best explained the observed space-time patterns in the mesozooplankton communities.

### Highlights

► Taxonomically and spatially-resolved zooplankton dataset over 16 years in spring. ► Time-consistent zooplankton composition spatial structure with multi-table analysis. ► Explanation of zooplankton spatial patterns with gradients in the pelagic ecosystem. ► Zooplankton spatial pattern underpinned by taxonomic composition and size structure. ► River plumes, temperature gradients and pelagic fish correlated zooplankton pattern.

**Keywords** : Mesozooplankton, Imaging analysis, Time-consistent spatial patterns, Ecosystem structure, Spring habitat, Multi Factor Analysis, Bay of Biscay

# 78 1 Introduction

79           Mesozooplankton is the preferential prey of small pelagic fish (SPF, e.g. sardine  
80 and anchovy; Plounevez and Champalbert, 1999; Van Der Lingen, 2002; Bachiller and  
81 Irigoien, 2015; Fonseca et al., 2022). It is therefore a key trophic link that enables the  
82 transfer of matter and energy from the primary production to higher trophic levels in  
83 oceanic (Banse, 1995), upwelling ecosystems (Rykaczewski and Checkley, 2008) as  
84 well as in continental shelf ecosystems (Irigoien et al., 2008; Dessier et al., 2018;  
85 Noyon et al., 2022).

86           The short generation time of mesozooplankton and their sometimes non-linear  
87 response to environmental changes make them sensitive to environmental variability  
88 from small spatio-temporal scales (*i.e.* Romagnan et al., 2015, 2016) to long term,  
89 global scale climate change (Hays et al., 2005; Batchelder et al., 2012; González-Gil  
90 et al., 2015). Yet, the diversity of taxa, morphologies and size ranges often make  
91 difficult to sample, analyse and model this important ecosystem component at the  
92 appropriate spatio-temporal as well as biological resolutions (Mitra and Davis, 2010).  
93 Further, the diversity of mesozooplanktonic organisms' life cycles and habitat  
94 requirements translate into a variety of community dynamics that may require local to  
95 large scales, and populations to communities' studies for their understanding.

96           Numerous studies suggest a potentially critical effect of environmentally driven  
97 mesozooplankton dynamics on SPF population dynamics in various ecosystems  
98 worldwide (*i.e.* Cury et al., 2000; Van Der Lingen, 2002; Rykaczewski and Checkley,  
99 2008; Capuzzo et al., 2018). In the European waters, recent studies highlighted the  
100 probable bottom-up effects of mesozooplankton-mediated decline in SPF body  
101 condition and size in the Gulf of Lion (NW Mediterranean, Saraux et al., 2019). The  
102 authors hypothesized that a decadal shift towards smaller mesozooplanktonic species  
103 in the Mediterranean adversely affected the potential energy income of SPF (Queiros  
104 et al., 2019 but see Feuilloley et al., 2022). A similar decadal trend in the decline in  
105 SPF size at age has recently been evidenced over the 2000-2015 period (Doray et al.,  
106 2018b), throughout the French continental shelf of the Bay of Biscay (hereafter BoB),  
107 an open ocean bay delimited by the Spanish coast (south) and the French coast (east  
108 & north). However, a comprehensive explanatory study still needs to be set up. This  
109 study is a first step towards this goal, and aims at understanding the mesozooplankton

110 habitat distribution, and its correlations with variables describing the pelagic ecosystem  
111 components, from hydrology (water temperature, salinity, water column stratification  
112 index) to higher trophic levels (SPF).

113 A corpus of data and studies on the SPF habitats' patterns and dynamics exists  
114 in the BoB (Petitgas et al., 2018; Doray et al., 2018c, 2022), as well as for hydrological  
115 landscapes (Koutsikopoulos et al., 1998; Castaing et al., 1999; Planque et al., 2004;  
116 Guillaud et al., 2008). Distribution patterns of mesozooplankton have also been  
117 described but in the south-eastern area of the BoB only (Albaina and Irigoien, 2007a;  
118 Dessier et al., 2018) and in the adjacent Cantabrian shelf (Albaina and Irigoien, 2007b;  
119 Iriarte et al., 2022). Coastal waters show larger biomass and smaller organisms than  
120 offshore waters (Sourisseau and Carlotti, 2006; Vandromme et al., 2014), and are  
121 dominated by copepods and meroplanktonic organisms (Irigoien et al., 2008; Dessier  
122 et al., 2018). Mesoscale oceanographic structures *i.e.* river plumes and shelf-break  
123 fronts (Albaina and Irigoien, 2004, 2007b), hydrological features (*i.e.* water column  
124 stratification), winter mixing and microphytoplankton abundance were identified as  
125 drivers of mesozooplankton dynamics (González-Gil et al., 2015; Dessier et al., 2018).  
126 Nevertheless, important knowledge gaps persist at the scale of the BoB and small  
127 pelagic fish populations, as the spatial coverage of existing mesozooplankton studies  
128 was consistently focused on the south of the BoB (South of 46°30N) although the North  
129 is ecologically important *e.g.* as SPF spawning, feeding and nursery habitats (Bellier  
130 et al., 2007; ICES, 2010). In addition, currently available mesozooplankton space-time  
131 data series are often short: the longest spatially resolved series spans from 2003 to  
132 2013 (Dessier et al., 2018), making the detection of long term changes detected in  
133 many other ecological components (Chust et al., 2022) more difficult for the  
134 mesozooplankton. Therefore, important uncertainties remain about the patterns of  
135 mesozooplankton assemblages and their long-term space-time evolution, at the scale  
136 of the entire BoB, in the context of accelerating ecological changes due to the local  
137 effects of global warming (Chust et al., 2022).

138 Here, we present the first comprehensive space-time analysis of the  
139 mesozooplankton community over 16 years (2004 to 2019), in spring, over the BoB  
140 continental shelf. The spatial coverage of our study spans from the coast to the shelf  
141 break, and from the eastern Spanish coast to southern Brittany coastal waters,  
142 therefore encompassing the whole French BoB continental shelf (Fig. 1).

143 Mesozooplankton samples collected during the pelagic ecosystem survey PELGAS  
144 (Doray et al., 2018c) were analysed with imaging tools (Gorsky et al., 2010; Colas et  
145 al., 2018) to achieve a coarse but consistent taxonomic resolution throughout the  
146 series. The homogeneity and completeness in the data spatial resolution over time was  
147 guaranteed using a gridding smoothing procedure (Masse et al., 2018 section 3) and  
148 a missing data imputation technique (Josse and Husson, 2016) when necessary. The  
149 same data preparation procedure was applied to concomitant hydrology, primary  
150 producers and SPF data series originating from the PELGAS surveys and remote  
151 sensing data to generate a comprehensive dataset with the same space-time  
152 resolution across the pelagic ecosystem components. Our goal is to explore the  
153 variability in the spatial patterns over time in the BoB mesozooplankton assemblage  
154 and how they correlate with the other pelagic ecosystem components. The  
155 mesozooplankton component was studied here considering its estimated biomass to  
156 produce time series of maps that could be integrated in future zooplankton modelling  
157 studies. This new data could be integrated to marine food web studies, which are often  
158 based on biomass data. This spatial ecology approach is a first step in identifying  
159 potential bottom-up and top-down controls exerted on the BoB mesozooplankton  
160 assemblage, in a climate change context.

## 161 2 Materials and methods

### 162 2.1 The PELGAS survey

163 The PELGAS integrated survey takes place every year in spring since 2000, over  
164 the French continental shelf of the BoB. The aim of this survey is to assess small  
165 pelagic fish biomass and monitor the pelagic ecosystem, to inform ecosystem based  
166 fisheries management. Concomitantly with fish data, hydrology, phyto- and zoo-  
167 plankton samples and megafauna sightings (marine mammals and seabirds) are  
168 collected to build long-term spatially-resolved time series of the BoB pelagic  
169 ecosystem. The PELGAS sampling strategy combines en-route data collection (small  
170 pelagic fish and megafauna) during the day, with fixed points, depth-integrated  
171 hydrology and plankton sampling during the night (Fig. 1). Detailed PELGAS survey  
172 protocols can be found in Doray et al., 2018b and 2021.

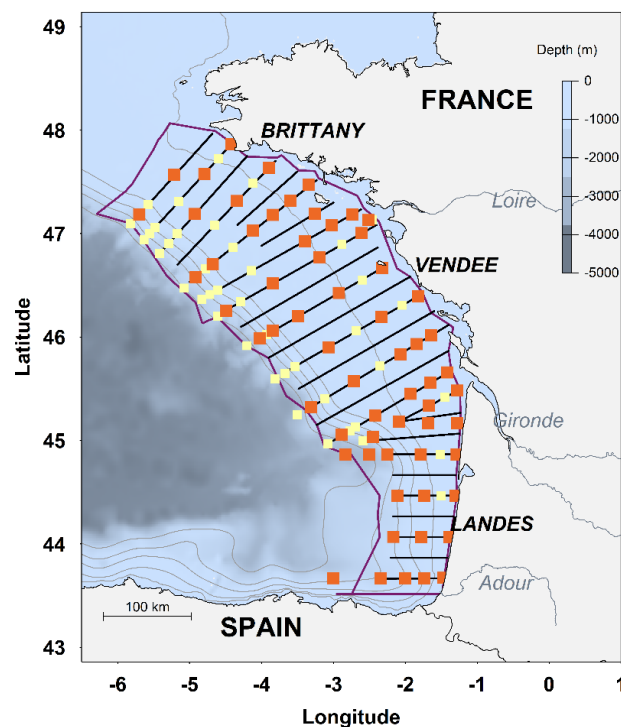


Fig. 1: PELGAS sampling scheme. Solid black lines: acoustic prospection transects for the evaluation of SPF biomass. Squares: Fixed-point water column sampling for hydrology parameters (light yellow) or hydrology parameters and plankton (orange) sampling. Note that the fixed-point geographical positions and sampling may change from year to year. Solid purple line: area considered for this study. Light grey lines: 100 to 500 m isobaths.

## 173 **2.2 Mesozooplankton samples collection**

174 Mesozooplankton samples were collected with vertical hauls of 200- $\mu$ m mesh size  
175 WP2 net, from 100 m depth (or 5 m above the sea floor) to the surface. In 2004 and  
176 2005, the sampling maximum depth was 200 m. Since 2014, the net has been  
177 equipped with a flowmeter (Hydrobios) to measure the sampled water volume. Before  
178 2014, the sampled water volume was estimated by multiplying the deployed cable  
179 length by the net opening surface (0.25 m<sup>2</sup>). From 2004 to 2019, the number of  
180 sampling points varied between 41 and 64 per year, due to adjustments in the sampling  
181 strategy and weather conditions. Between 2004 and 2016, 699 samples were collected  
182 and preserved in 4% formaldehyde on board for further analysis back on land. Between  
183 2017 and 2019, 190 samples were collected and were analysed live on board. Both  
184 sets of samples were analysed with imaging instruments.

## 185 **2.3 Mesozooplankton analyses and data**

186 The preserved samples (2004-2016) were digitized on land with the ZooScan, a  
187 waterproof flatbed scanner generating 16 bit gray-level high resolution images (pixel  
188 size: 10.56  $\mu$ m, 2400 dpi) (Gorsky et al., 2010). Prior to digitization, the samples were  
189 size-fractionated with a 1 mm sieve, into organisms > 1 mm and <1 mm size fractions,  
190 to avoid underestimation of large and rare objects due to subsequent subsampling.  
191 Then, both fractions were separately subsampled with a Motoda splitter, to obtain sub-  
192 samples containing 500-1500 objects. Each subsample was imaged after manual  
193 separation of objects on the scanning tray, to minimize the occurrence of touching  
194 objects (Vandromme et al., 2012). Remaining touching objects were manually digitally  
195 separated with a custom Zooprocess tool to ensure the quality of further identifications  
196 and counts. From 2017 onwards, the samples were analysed live on board with the  
197 ZooCAM, an in-flow imaging instrument (pixel size: 10.3  $\mu$ m) allowing the quasi real  
198 time analysis of samples (Colas et al., 2018). Samples were split as a single fraction  
199 with a Motoda splitter before being digitized. The agreement between the ZooScan  
200 and the ZooCAM in identifying similar communities and producing similar total  
201 mesozooplankton abundances and size distributions was demonstrated earlier (Colas  
202 et al., 2018). All raw images generated with both instruments were processed to obtain  
203 individual vignettes of each object digitized and associated morphological features,

204 including the size in pixels of each object. All individual planktonic vignettes were  
205 classified using Ecotaxa (Picheral et al., 2017), a dedicated online tool that combines  
206 a Random-Forest and a Convolutional Neural Network to achieve automatic  
207 identification, which is now a classic semi-automatic identification procedure. Each  
208 identified object, originating from both instruments, was then visually checked and  
209 explicitly validated or corrected when necessary, using Ecotaxa again. Eventually,  
210 2,135,401 objects were analysed and sorted into 24 broad taxonomic groups, to  
211 achieve a trade-off between the diversity of the taxa seen in the area, and the  
212 identification skills of experts who validated the automatic classification results. The  
213 taxonomic resolution is detailed in Table 1. Some taxonomic groups aggregate a  
214 diversity of species. For example, the group “Eumalacostraca” comprised adult forms  
215 of crustaceans such as Amphipods, Decapods and Isopods. The group “Meroplankton  
216 crustacean larvae” consisted of larval forms of meroplanktonic crustaceans, for  
217 example Cirripedia, Brachyours and Decapods zoe. The category “Crustacean nauplii”  
218 grouped the nauplii forms of holoplanktonic crustaceans, mostly copepods (Table 1).  
219 Non-living objects (i.e. detrituses, large aggregates, artefacts) and remaining multiple  
220 objects were excluded from the dataset.

221 Then, the biomass expressed in  $\mu\text{g}$  Dry Weight of each object was calculated from  
222 each object’s individual surface in pixels converted to  $\text{mm}^2$ , following Lehet and  
223 Hernández-León (2009) and Garijo and Hernández-León (2015) (Table 1). These  
224 equations are a common way to estimate biomasses of coarse zooplankton taxa and  
225 were used in zooplankton studies in the Southern Ocean (Stirnemann et al., 2021;  
226 Kerkar et al., 2022), tropical oceans and temperate environments (Marcolin et al.,  
227 2015; Garcia-Herrera et al., 2022; Giering et al., 2019; Makhlof Belkahia et al., 2021;  
228 Noyon et al., 2022) and at the global scale (Hernández-León et al., 2020; Siviadan et  
229 al., 2022), in open ocean as well as in shelf seas. The biomass of each taxonomic  
230 group was calculated at every sampling station, by summing the objects’ biomass  
231 within each taxonomic group, multiplied by the subsampling ratio and divided by the  
232 sampled water volume, to obtain the biomass expressed in  $\mu\text{g}$  Dry Weight. $\text{m}^{-3}$  for each  
233 taxa, at each sampling point. The organisms’ size considered in this study ranged  
234 between 0.3 and 3.4 mm ESD (large mesozooplankton). Data from both instruments  
235 were finally assembled to form a spatially-resolved dataset comprising 24  
236 mesozooplankton taxa over 16 years from 2004 to 2019.



Table 1: Mesozooplankton taxonomic groups used to characterise mesozooplankton spring community in the Bay of Biscay; the conversion equations of body area “A” (mm<sup>2</sup>) into dry weight (µg Dry Weight) were found in Lehette & Hernandez – Leon (2009) and Garijo & Hernandez – Leon (2015).

Organisms	Taxonomic groups	Code	Dry weight
Copepods	Acartiidae	Acart	$43.97 * A^{1.52}$
	Calanidae	Cal	
	Centropagidae	Centrop	
	Euchaetidae	Euch	
	Metridinidae	Metri	
	Temoridae	Temo	
	Other Calanoida	Calano	
	Cyclopoida	Cyclo	
	Harpacticoida	Harpact	
	Poecilostomatoida	Poecil	
Other crustaceans	Eumalacostraca	Eumal	
	Meroplankton crustacean larvae	Larv_mero	
	Crustacean nauplii	Npl_crust	
	Cladocera	Clad	
General euphausiids	Shrimp-like	Shrimp	$49.58 * A^{1.48}$
Chaetognaths	Chaetognaths	Chaeto	$23.45 * A^{1.19}$
Cnidarian: Siphonophores	Siphonophores	Sipho	$43.17 * A^{1.02}$
Other gelatinous plankton	Appendicularians	Append	$4.03 * A^{1.24}$
	Thaliacea	Thal	
	Other Cnidarians	Cnid	
Other plankton	Bivalvia larvae	Biv	$43.38 * A^{1.54}$
	Fish larvae	Actin	
	Echinodermata larvae	Echin	
	Thecosomata	Thec	

## 238 **2.4 Hydrology and primary producers data**

239 At-station vertical profiles were performed with a conductivity-temperature-depth  
240 (CTD, SeaBird SBE19 + V2) probe fitted with a fluorimeter and Niskin bottles, to  
241 conduct hydrological and phytoplankton sampling. Several water column structure  
242 descriptors were calculated from CTD casts, at each sampling station (Table 2): the  
243 surface salinity (mean between 2 and 7 m), the bottom temperature (value recorded at  
244 5 m above the seabed), the equivalent freshwater height, and three water column  
245 stratification indices: the potential energy deficit, the mixed layer maximum depth, and  
246 the pycnocline depth. The equivalent freshwater height is an index of river plume  
247 influence, which measures the freshwater height over the water column, considering a  
248 salinity reference value set at 35.5 psu. It better integrates the recent history and local  
249 influence of river run-offs and is less sensitive to vertical mixing than surface salinity.  
250 The deficit of potential energy is an index of water column stratification, defined as the  
251 energy necessary to homogenise the water density over the water column. A detailed  
252 description of these water column descriptors can be found in Huret et al. (2013).

253 Water samples collected with the Niskin bottles at the surface were filtered on board,  
254 stored at -80°C and further analysed by spectrophotometry back on land, to estimate  
255 the total and size-fractionated chlorophyll-a concentration (pico- (< 3 µm), nano- (3-20  
256 µm) and micro-phytoplankton (> 20 µm)). The water column integrated chlorophyll-a  
257 concentration was estimated using the fluorescence data from the vertical profiles  
258 which was calibrated with the actual measures of chlorophyll-a concentrations at three  
259 depths (surface, deep chlorophyll maximum, and below the thermocline).

260 Sea surface temperatures (SST) and surface chlorophyll-a concentrations from  
261 remote sensing data were downloaded from <https://marine.copernicus.eu/> (named as  
262 SST\_ATL\_SST\_L4\_REP\_OBSERVATIONS\_010\_026 and  
263 OCEANCOLOUR\_ATL\_CHL\_L4\_REP\_OBSERVATIONS\_009\_098, respectively).  
264 The downloaded SST data products were derived from AVHRR sensors of NOAA  
265 satellites daily products, and interpolated to fill in missing data due to the clouds,  
266 following Saulquin and Gohin (2010). Surface chlorophyll-a concentration data were  
267 derived from several sensors (SeaWiFS, MERIS, MODIS) and processed following  
268 Gohin (2011).

269 Table 2: Parameters used to describe the hydrological conditions and the primary producers  
 270 structure in the Bay of Biscay, in spring.

Ecosystem component	Parameter		Code	Units	Time series	Source
Hydrology	Bottom temperature		BTemp	°C	2004 - 2019	PELGAS surveys
	Potential energy deficit		PotEDef	Kg.m <sup>-1</sup> .m <sup>-2</sup>		
	Equivalent freshwater height		EqFH	meters		
	Mixed layer depth		MLD			
	Pycnocline depth		PycnD			
	Surface salinity		SSal	psu		
	Surface temperature		STemp	°C		Satellite data
Primary producers	Size-fractionated surface chlorophyll-a concentration	< 3 µm	Chl3	mg.m <sup>-3</sup>	2009 - 2019	PELGAS surveys
		3 - 20 µm	Chl-3-20			
		> 20 µm	Chl20			
	Total surface chlorophyll-a concentration		TotChl	mg.m <sup>-2</sup>		
	Integrated chlorophyll-a		IntChl			
	Chlorophyll-a maximum concentration depth		ChIMD			
	Chlorophyll-a concentration		satChl			mg.m <sup>-2</sup>

271 **2.5 Small pelagic fish biomass**

272 Small pelagic fish acoustic densities were recorded at 10 knots during daytime,  
 273 along transects using a calibrated monobeam echosounder operating at 38 kHz.  
 274 Midwater trawl hauls were adaptively performed to identify to species the echotraces  
 275 and provide their length, weight and age composition. Acoustic and biotic trawl data  
 276 were combined using the standard methodology described in Doray et al. (2021) to  
 277 derive biomass estimates per species and 5 cm length classes, within one nautical  
 278 mile (1852 m) long Elementary Sampling Units along the survey track. Biomass  
 279 estimates were expressed in tonnes per nautical miles square (tonnes.nm<sup>-2</sup>). The SPF  
 280 species appearing in at least 50 % of the 16 PELGAS surveys considered (2004-2019)  
 281 were selected to characterise the fish component. Selected species included boarfish  
 282 (*Capros aper*), Atlantic chub mackerel (*Scomber colias*), Atlantic mackerel (*Scomber*

283 *scombrus*), Atlantic horse mackerel (*Trachurus trachurus*), Mediterranean horse  
 284 mackerel (*Trachurus mediterraneus*), blue whiting (*Micromesistius poutassou*),  
 285 European anchovy (*Engraulis encrasicolus*), European sardine (*Sardina pilchardus*)  
 286 and sprat (*Sprattus sprattus*) (Table 3).

Table 3: Time series of small pelagic fish species considered by 5 cm length classes.

SPF species	Code	Length class	Time series (n years)
Boarfish ( <i>Capros aper</i> )	CAPR-APE	(10, 15]	2004 - 2019 (9)
European anchovy ( <i>Engraulis encrasicolus</i> )	ENGR-ENC	(5, 10]	2006 - 2019 (8)
		(10, 15]	2004 - 2019 (16)
		(15, 20]	
Blue whiting ( <i>Micromesistius poutassou</i> )	MICR-POU	(15, 20]	2004 - 2017 (11)
		(20, 25]	2004 - 2019 (15)
		(25, 30]	2005 - 2019 (14)
European sardine ( <i>Sardina pilchardus</i> )	SARD_PIL	(10, 15]	2005 - 2019 (13)
		(15, 20]	2004 - 2019 (16)
		(20, 25]	
Atlantic chub mackerel ( <i>Scomber colias</i> )	SCOM-COL	(15, 20]	2004 - 2018 (14)
		(20, 25]	2004 - 2019 (16)
		(25, 30]	2004 - 2019 (15)
		(30, 35]	2004 - 2019 (14)
		(35, 40]	2005 - 2019 (9)
Atlantic mackerel ( <i>Scomber scombrus</i> )	SCOM-SCO	(15, 20]	2005 - 2019 (11)
		(20, 25]	2004 - 2019 (16)
		(25, 30]	
		(30, 35]	
		(35, 40]	
European sprat ( <i>Sprattus sprattus</i> )	SPRA-SPR	(5, 10]	2004 - 2019 (13)
		(10, 15]	2004 - 2019 (16)
Atlantic horse mackerel ( <i>Trachurus trachurus</i> )	TRAC-TRU	(10, 15]	2004 - 2019 (16)
		(15, 20]	
		(20, 25]	
		(25, 30]	
Mediterranean horse mackerel ( <i>Trachurus mediterraneus</i> )	TRAC_MED	(30, 35]	2004 - 2019 (13)
		(15, 20]	2005 - 2018 (10)
		(20, 25]	2006 - 2019 (10)
		(40, 45]	2006 - 2019 (9)

287 A summary of the data used in this study can be found in Fig. 2 (top panel).

288

## 2.6 Data gridding and missing data imputation

289 Although the biological and physico-chemical sampling are rather regular and  
290 homogeneous on the PELGAS cruise (Doray et al., 2018c), the spatial resolution and  
291 number of samples may vary across years and between variables. Therefore, a block  
292 averaging procedure (Petitgas et al., 2009, 2014) was consistently applied on each of  
293 the datasets before conducting any further analysis. All variable used were gridded  
294 over a common spatial grid. The grid mesh size was set at  $0.3^\circ$  in both latitude and  
295 longitude, a compromise between the grid mesh size commonly set at  $0.25^\circ$  in previous  
296 studies based on spatially resolved datasets (Doray et al., 2018a; Masse et al., 2018;  
297 Petitgas et al., 2018) and the distance between the mesozooplankton sampling  
298 stations, to limit the number of grid cells without data. The grid origin  $x_0$  was initially  
299 positioned at  $43^\circ\text{N}$  and  $6^\circ\text{W}$  and then drawn randomly within a two cells radius, 300  
300 times. Data were averaged in each grid cell for every origin position, in order to  
301 minimize the influence of the origin position on gridded values. Finally, 300 mean  
302 values were averaged to calculate a spatially smoothed estimate in each grid cell (Fig.  
303 2, step 1). Grid cells having their center point inside the polygon defining the survey  
304 area were kept for the analysis ( $n = 121$ ) (Fig. 1). The block averaging step was  
305 realized with the *EchoR R* library (Doray et al., 2013).

306 After the block averaging step, some cells were still empty (missing data). These  
307 missing data are the consequence of missing sampling stations, especially in the  
308 northern part of the BoB, mostly at the beginning of the time series (Appendix A, Table  
309 A.1). To fill data gaps in particular grid cells and years, we applied a missing data  
310 imputation procedure (Josse and Husson, 2016). Each variable was organized into  
311 matrix form, grid cells x years. An algorithm based on iterative Principal Component  
312 Analysis (PCA) was applied to each data matrix, to impute predicted data point values  
313 to the empty cells (Fig. 2, step 2). The iterations were run until the difference between  
314 two successive estimated values was smaller than a threshold (set at  $1e-06$ ).  
315 Sometimes, the final predicted values were negative. In this case, we imputed the  
316 mean of the adjacent grid cells to the empty cell. A comparison of annual spatial  
317 patterns and annual means, calculated with and without the imputed values, was used  
318 as a quality check of the imputed values. This missing data imputation method was  
319 implemented using the *MissMDA R* library (Josse and Husson, 2016).

320 Eventually, the complete dataset is composed of sets of gridmaps for parameters in  
321 the hydrology, primary producers, mesozooplankton and SPF components. All the  
322 gridmaps have the same spatial extent and resolution and span from 2004 to 2019,  
323 except for primary producers dataset which span from 2009 to 2019.

## 324 **2.7 Data analysis**

325 All the data analyses were performed using the *R* statistical language version 4.0.3  
326 (R Core Team, 2020). The analytic pipeline is organized as follow. First, we applied a  
327 Multiple Factor Analysis (MFA, see below for details) on each ecosystem component  
328 (Fig. 2, step 3). Then the mesozooplankton MFA results were used to input a  
329 hierarchical clustering (Fig. 2, step 4). Finally, results from the hydrology, primary  
330 producers and fish MFA were used as explanatory variables in Generalized Linear  
331 Models (GLMs) to explain the observed mesozooplankton space-time patterns (Fig. 2,  
332 step 5).

### 333 *2.7.1 Multiple Factor Analysis*

334 Multiple Factor Analysis (MFA, Escofier and Pagès, 1994; Pagès, 2014) has  
335 recently been applied to characterise space-time patterns in series of multivariate  
336 gridmaps of ecological variables (Abdi et al., 2013; Doray et al., 2018a; Petitgas et al.,  
337 2018). MFA is a multi-table statistical method based on PCA and designed to analyse  
338 3D structured datasets, in which the variables are organized in tables of the same size,  
339 over time. Here, the grid cells correspond to the tables' rows (1<sup>st</sup> dimension:  $n$   
340 individuals), the variables the columns (2<sup>nd</sup> dimension:  $m$  variables), and the tables are  
341 stacked (3<sup>rd</sup> dimension:  $p$  groups), to obtain a multivariate time series of yearly tables.  
342 It is worth noticing that MFA applies even when the number of available variables differ  
343 over time.

344 Similarly as with PCA and a 2D dataset, MFA results in summarizing a 3D dataset  
345 into a factorial space, in which all individuals (*i.e.* grid cells), variables and tables  
346 (years) can be represented (Pagès, 2014). Eventually, each grid cell can be located in  
347 each year in the factorial space around its (time) average position and similarly for  
348 each variable. MFA enables the estimation of a common multivariate correlation  
349 structure over all years, and its associated time variability. The individuals being grid  
350 cells here, the target structure is a spatial pattern with its time variability. For a detailed

351 description of MFA implementation, see Petitgas et al., 2018, § 3.3 - 3.5, pp 191-193,  
352 and Doray et al., 2018 § 2.5-2.7, p 91.

### 353 2.7.2 *MFA implementation and principal components selection*

354 Specific MFAs were applied separately on the hydrology, primary producers,  
355 mesozooplankton and SPF datasets, to characterize their space-time patterns. Data  
356 from biotic components (primary producers, mesozooplankton and SPF) were  
357 log-transformed to reduce the skewness in their distributions. Hydrology and primary  
358 producers variables were centred and normalized in each year to account for the  
359 differences in units and ranges. Mesozooplankton and SPF variables were centred  
360 only (all variables in these two datasets have the same units), thus leaving the  
361 differences in variance between years affect time variability. The MFAs were  
362 implemented using the function *MFA* from the FactoMineR *R* library (Lê et al., 2008).

363 Similarly to a classic PCA, MFA can be used to reduce the dimensionality, filter  
364 and synthesize complex datasets by selecting only a few first principal components  
365 (*PCs*) for subsequent analyses. When using MFA on time resolved datasets, it is  
366 important to consider not only the percentage of total variance explained by the MFA  
367 *PCs* (Fig. 4), but also the number of significant correlations between each *PC* and the  
368 years (*i.e.* tables) to assess the representativity of each *PC* over time (Pagès, 2014).  
369 We selected the *PCs* that showed correlations with the years higher than a threshold  
370 for at least half of the total number of years. Such threshold was estimated following  
371 Pagès (2014), using the maximum value of correlation between years and high orders  
372 *PCs*, here *PC.4* and *PC.5* (see the results of this procedure in Fig. 5).

373 Then, each selected *PC* was interpreted considering its correlation with the  
374 variables. The variables retained to interpret the *PCs* showed a correlation coefficient  
375 greater than |0.5| with the *PCs* for at least half of the total number of years. Further,  
376 mean individuals' (grid cells') coordinates on the *PCs* were mapped at their  
377 geographical positions. The MFA selected *PCs* were used as synthetic descriptors for  
378 each of the ecosystem component datasets.

### 379 2.7.3 *Mesozooplankton space-time patterns: hierarchical clustering*

380 The average spatial structure of the mesozooplankton community was identified  
381 by applying a Hierarchical Agglomerative Clustering (function *hclust* in *R* language,

382 Ward's method, Euclidean distance, without spatial constrain) on the mean individuals'  
383 (grid cells') coordinates in the mesozooplankton MFA factorial space. The function  
384 NbClust from the R library NbClust (Charrad et al., 2014) was used to determine the  
385 optimal number of clusters through the calculation of thirty partitioning indices. The  
386 best number of clusters is defined as the one suggested by the highest number of  
387 indices. Then, grid cell clusters were mapped, to characterise regions with  
388 time-consistent specific mesozooplankton assemblages. The taxonomic composition  
389 of each cluster was assessed by calculating the annual mean of the variables (taxa) in  
390 each cluster. Only the taxa showing a correlation coefficient greater than  $|0.5|$  with the  
391 mesozooplankton MFA *PCs* for at least half of the total number of years were  
392 considered. Variables being centered, the annual mean biomass of a taxa within a  
393 cluster represents its residual variation relative to the global mean taxon biomass over  
394 all the grid cells and years. Finally, each grid cell in each year was associated to the  
395 nearest cluster centroid in the MFA space to estimate the annual spatial distribution of  
396 the clusters and to characterize the interannual variability in the clusters. The frequency  
397 of grid cells affiliation within each cluster was also calculated over the time series, and  
398 mapped.

#### 399 *2.7.4 Correlates of the observed mesozooplankton space-time patterns*

400 The MFA *PCs* for the hydrology, primary producers and SPF ecosystem  
401 components (referred to as predictors hereafter) were used as explanatory variables  
402 in GLMs, where the MFA *PCs* for the mesozooplankton were the dependent variables.  
403 For each *PC* selected in the mesozooplankton MFA, we first built a model using all the  
404 predictors. We checked for collinearity in the predictors by using the Variance Inflation  
405 Factor (VIF, Fox and Monette, 1992). A VIF value higher than five was used to identify  
406 problematic multicollinearity among predictors. VIF was calculated using the *car* R  
407 library (Fox and Weisberg, 2019). Predictors explaining less variance and displaying  
408 high multicollinearity (VIF criterion) were removed. Then, a stepwise backward model  
409 selection procedure was applied (function *stepAIC* in R language) to select the most  
410 significant predictors based on Akaike's Information Criterion (AIC) (Burnham et al.,  
411 2002). Models were considered significantly different when their AIC difference was  
412 higher than two. Finally, an ANOVA was used on the model selected to rank the  
413 predictors by their explicative power.



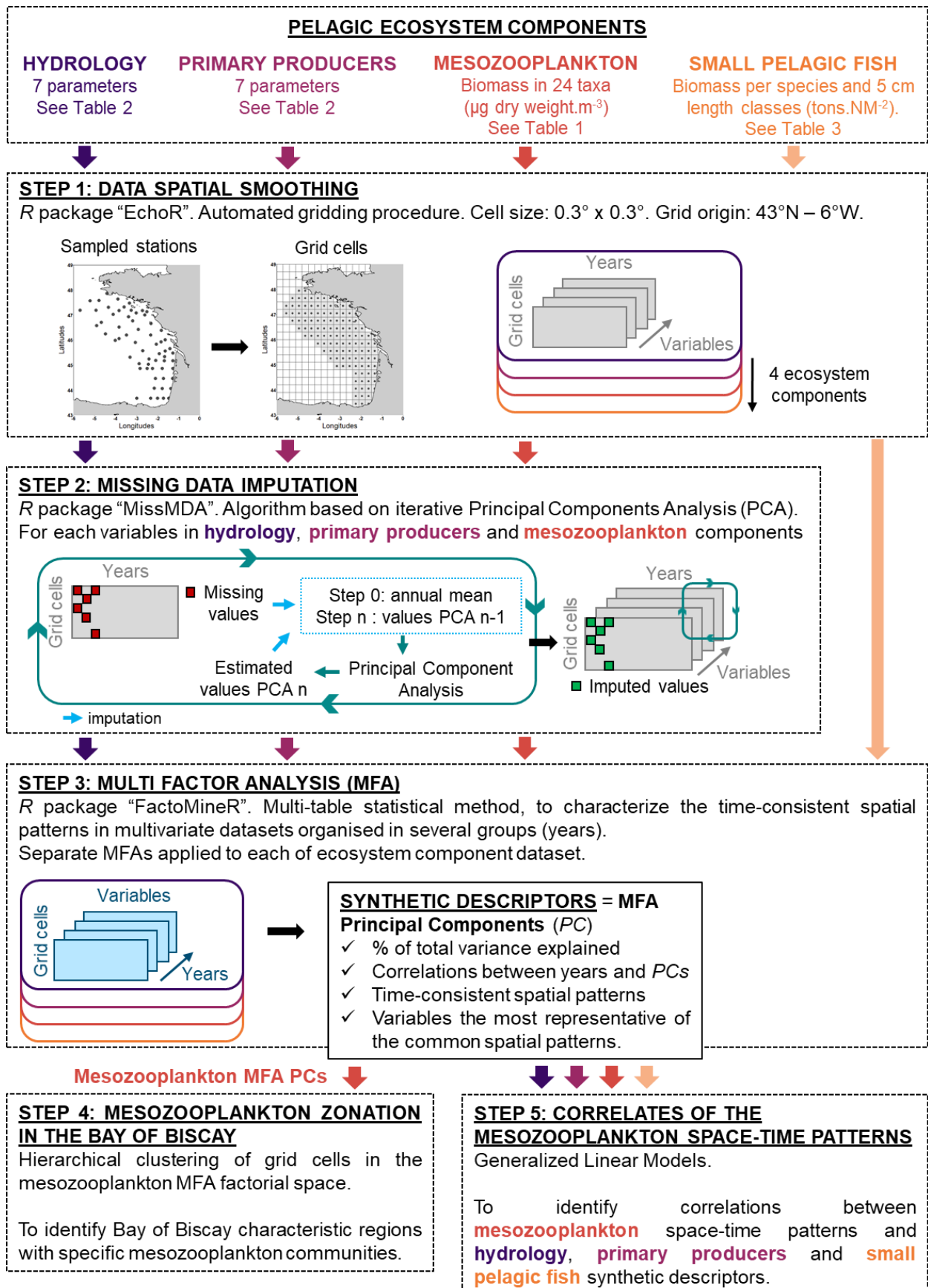


Fig. 2: Summary of the steps followed for the datasets construction and analysis.

415 **3 Results**

416 **3.1 Taxonomic composition and total biomass of the**  
417 **mesozooplankton community**

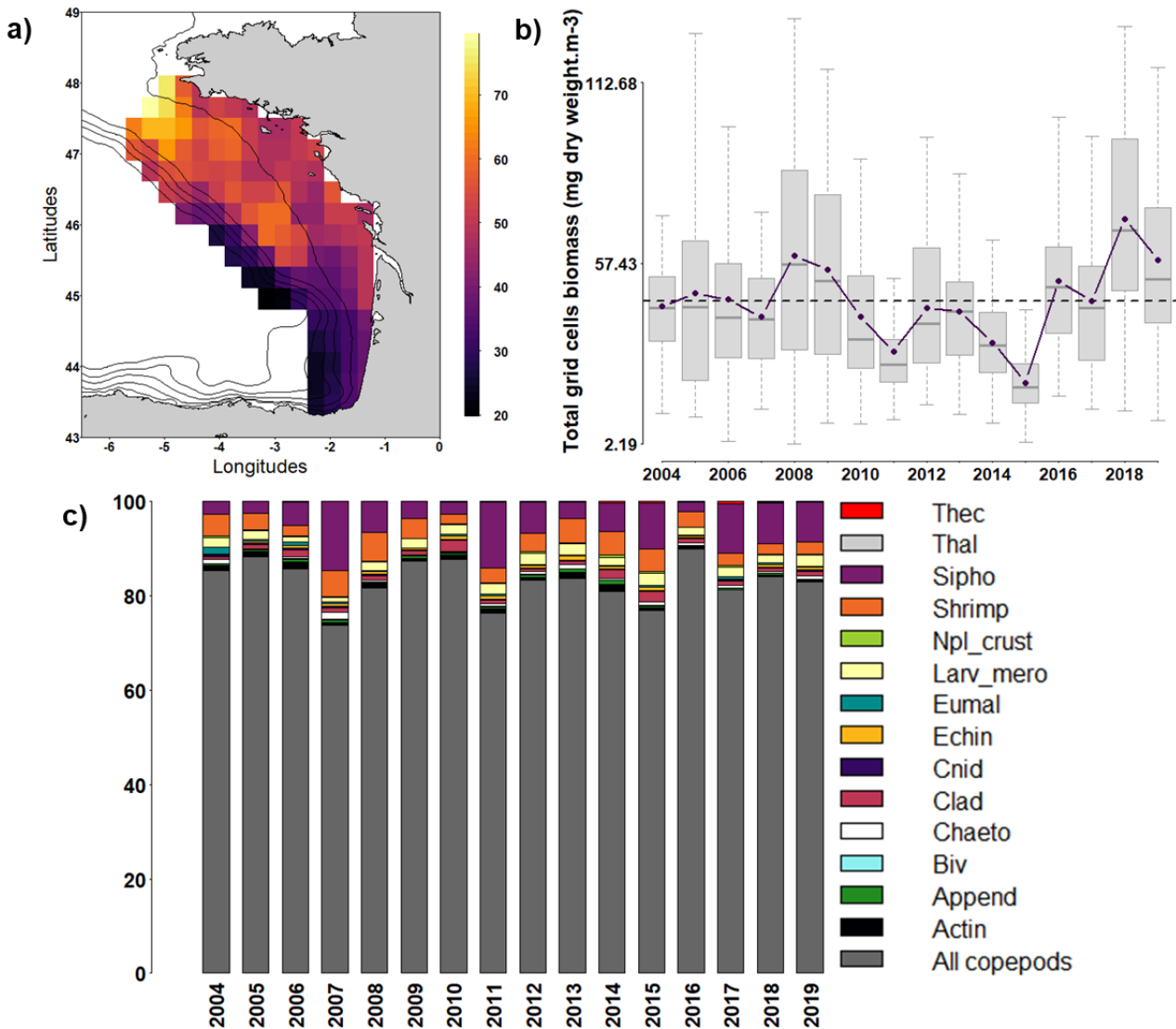


Fig. 3: (a) Mesozooplankton community total biomass (mg dry weight.m<sup>-3</sup>) mean map, averaged between 2004 and 2019. (b) Grid cells total biomass yearly distributions from 2004 to 2019. Boxplots: total biomass by grid cells (n = 121). The dashed line is the mean total biomass over the time series. Solid line: annual means. (c) Taxonomic groups relative contributions to the mesozooplankton community total biomass (percentage). The ten copepod groups were summed and plotted as one group for more clarity.

419 The mean map of total biomass per grid cell, from 2004 to 2019, revealed a  
420 North-South gradient. Highest total biomass was observed north to 45.5°N between  
421 the 100 m and 200 m isobaths, where the biomass ranged between 60 and 80 mg dry  
422 weight.m<sup>-3</sup>, whereas the lowest values ranged between 20 and 30 mg dry weight.m<sup>-3</sup>  
423 south of 45.5°N, on the continental shelf slope (Fig. 3a). Total biomass yearly  
424 distributions did not show a clear temporal trend, but the lowest values occurred from  
425 2010 to 2015. The lowest annual means occurred in 2011 and 2015 (30.5 ± 16 mg dry  
426 weight.m<sup>-3</sup> and 21 ± 8.6 mg dry weight.m<sup>-3</sup>, respectively), while the highest ones  
427 occurred in 2008 and 2018 (59.8 ± 33.3 mg dry weight.m<sup>-3</sup> and 71 ± 30.5 mg dry  
428 weight.m<sup>-3</sup>, respectively) (Fig. 3b). The copepods dominated the mesozooplankton  
429 community every year, contributing from 73% (2007) to 90% (2016) to the community  
430 total biomass. Siphonophores was the second most important taxonomic group,  
431 representing up to 14% of the total biomass in 2007 and 2011. These two groups had  
432 opposite trends in their contributions, the years of minimum contribution of one  
433 corresponding to the years of the maximum contribution of the other (Fig. 3c).

### 434 **3.2 MFA PCs selection for each ecosystem component.**

435 The first three *PCs* of the hydrology MFA (referred to as hydroMFA1, hydroMFA2  
436 and hydroMFA3 hereafter) explained 72.8 % of the space-time variance in the  
437 hydrology dataset (Fig. 4). In this case, the correlation threshold for retaining MFA *PCs*  
438 as meaningful was set to 0.73 (see Methods section 2.7.2). All the years of the  
439 hydrology time series showed correlations higher than 0.73 to the hydroMFA1 (n = 16)  
440 and the hydroMFA2, except 2007 (hydroMFA2, n = 15). Ten years were correlated with  
441 a coefficient > 0.73 to hydroMFA3 (Fig. 5a). The first *PC* of the primary producers MFA  
442 (referred to as phytoMFA1 hereafter) explained 36.7 % of the space-time total variance  
443 in the primary producers dataset (Fig. 4). It was the only *PC* showing correlation  
444 coefficients with years (n = 11) higher than the threshold (here set to 0.67, see Methods  
445 section 2.7.2) (Fig. 5b). The first two *PCs* of the mesozooplankton MFA (referred to as  
446 zooMFA1 and zooMFA2 hereafter) explained 50.5 % of the space-time variance in the  
447 mesozooplankton dataset (Fig. 4). All the years (n = 16) of the mesozooplankton time  
448 series were correlated to the zooMFA1 with a coefficient higher than the threshold  
449 (here set to 0.83, see Methods section 2.7.2). Ten years were correlated with a  
450 coefficient > 0.83 to zooMFA2 (Fig. 5c). Finally, the first two *PCs* of the SPF MFA

451 (referred to as spfMFA1 and spfMFA2 hereafter) explained 45.5 % of the space-time  
452 variance in the SPF biomass dataset (Fig. 4). Eleven years were correlated with a  
453 coefficient higher than the threshold (here set to 0.9, see Methods section 2.7.2) to the  
454 spfMFA1, and half the years ( $n = 8$ ) were correlated to the spfMFA2 (Fig. 5d).  
455 Eventually, eight MFA PCs were retained as synthetic descriptors for the next steps of  
456 the analysis (*i.e.* hydroMFA1, hydroMFA2, hydroMFA3; phytoMFA1; zooMFA1,  
457 zooMFA2; spfMFA1, spfMFA2).

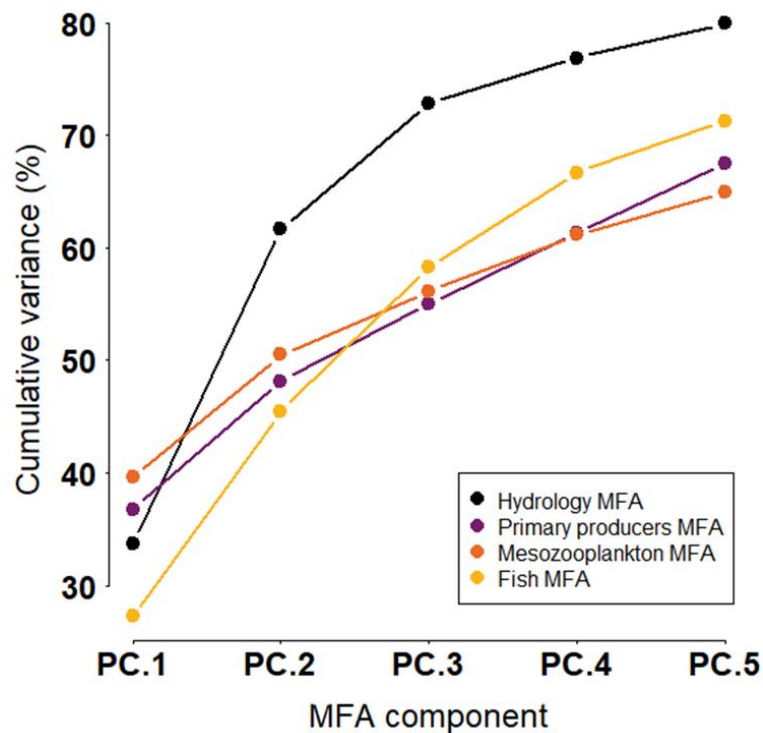


Fig. 4: Cumulative percentage of variance explained by the first five principal components of the Multiple Factor Analyses applied to the series of multivariate maps characterising hydrology, mesozooplankton and small pelagic fish (2004 - 2019) and primary producers data (2009 - 2019).

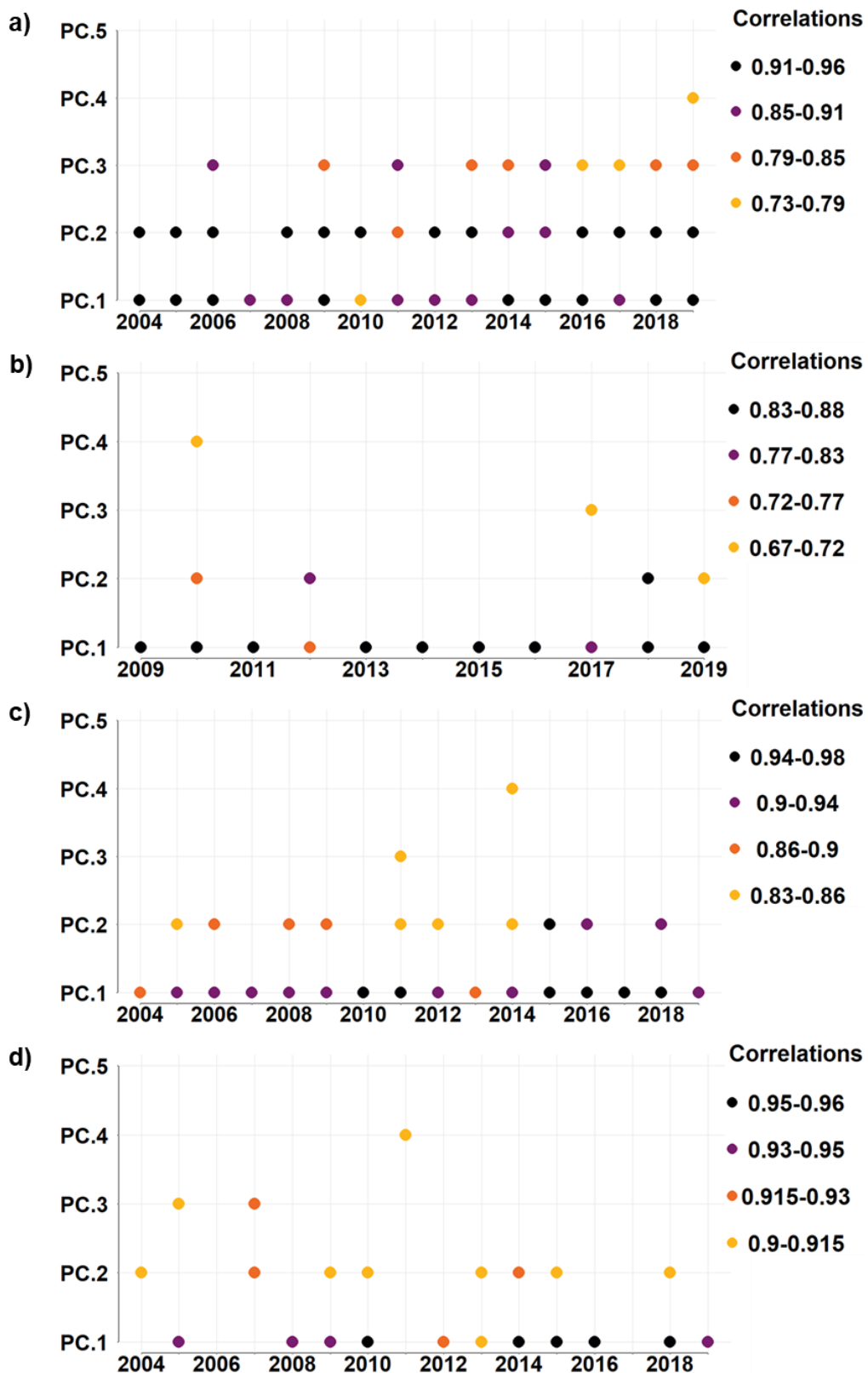


Fig. 5: Correlations between the years and the first five principal components of the Multiple Factor Analyses applied to the series of multivariate maps characterising (a) hydrology (2004 – 2019), (b) primary producers (2009 - 2019), (c) mesozooplankton (2004 – 2019) and (d) small pelagic fish (2004 - 2019). Only the correlations higher than the thresholds are shown (see Methods section 2.7.2).

### 459 **3.3 Mesozooplankton ecosystem component**

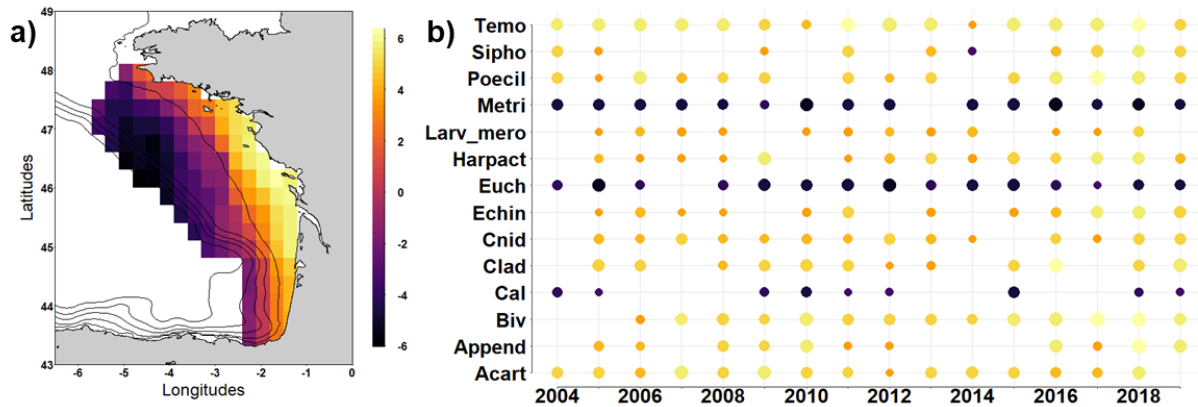
#### 460 *3.3.1 Mesozooplankton MFA selected PCs*

461 Here we present the spatial patterns (MFA *PCs*) in the mesozooplankton  
462 ecosystem component, based on the mapping of the mean individuals' (grid cells')  
463 coordinates in the mesozooplankton MFA factorial space (Fig. 6a and c). The variables  
464 defining the MFA *PCs* are shown in Fig. 6b and d.

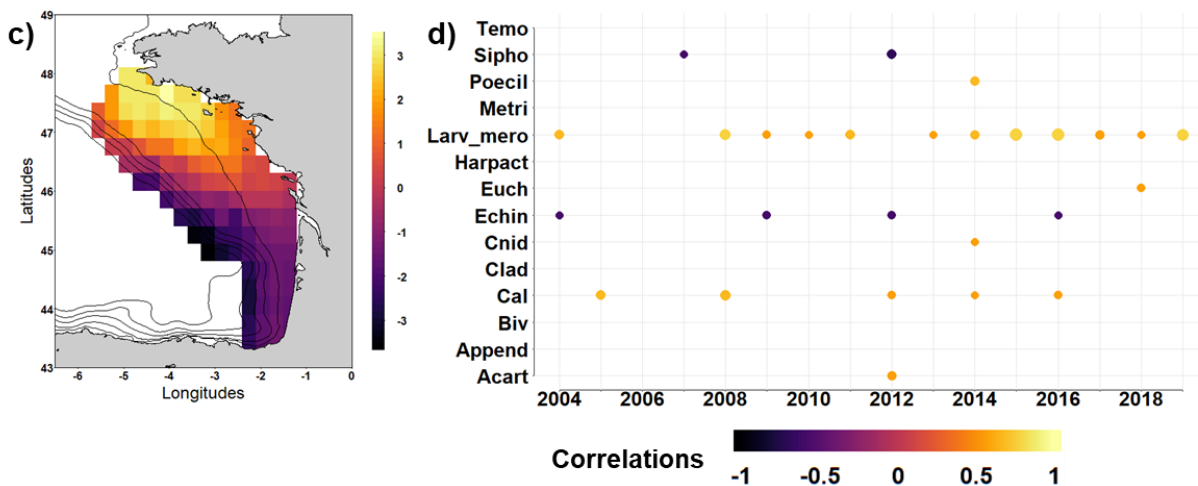
465 The zooMFA1 showed a dominant coastal – offshore gradient in the  
466 mesozooplankton community. This structure was highlighted by the highest values in  
467 coastal areas, between the Loire estuary and the Arcachon Bay (Fig. 6a, lightest cells),  
468 and the lowest along the shelf break, north to 46°N (Fig. 6a, darkest cells). Small  
469 copepods, such as Acartiidae, Temoridae, Poecilostomatoida and Harpacticoida, as  
470 well as Cnidarians, Cladocerans, Bivalvia, Echinoderms and meroplanktonic  
471 crustaceans larvae groups displayed consistent positive correlations with zooMFA1  
472 over the time series, revealing higher biomass in coastal areas. Significant positive  
473 correlations with zooMFA1, and consequently higher biomass in coastal areas, were  
474 also observed for appendicularians between 2005 and 2012 and between 2016 and  
475 2019, indicating a higher variability across years for this taxon. On the contrary,  
476 Metridinidae, Euchaetidae, and to a lesser extent Calanidae showed consistent  
477 negative correlations with zooMFA1, and consequently high biomass offshore,  
478 especially north to 46°N (Fig. 6b).

479 The zooMFA2 revealed a North – South gradient, highlighted by the highest  
480 values in the northern part of the BoB along the Brittany southern coast (Fig. 6c, lightest  
481 cells), and the lowest ones in the south, especially over the shelf break, south to 45.5°N  
482 (Fig. 6c, darkest cells). The only taxa displaying a consistent positive correlation with  
483 this component were the meroplanktonic crustaceans larvae, after 2008 (Fig. 6d),  
484 showing higher biomass in the northern part of the BoB.

### Principal component 1, mesozooplankton MFA



### Principal component 2, mesozooplankton MFA



**Fig. 6:** Maps of mean individuals' (grid cells') coordinates on mesozooplankton Multiple Factor Analysis (MFA) principal component *PC.1* (a) and *PC.2* (c), and time series of significant correlations ( $> |0.5|$ ) between the 24 taxonomic groups and MFA *PC.1* (b) and *PC.2* (d). The disk radii are proportional to the absolute value of the correlation coefficient. The mesozooplankton MFA was performed on the maps time series of 24 taxonomic groups biomass calculated from the PELGAS data, from 2004 to 2019 (Table 1).

#### 485 3.3.2 Mesozooplankton zonation in the Bay of Biscay and its temporal evolution

486 The hierarchical clustering of the mean individuals' (grid cells') coordinates in the  
 487 MFA factorial space made by the two selected *PCs* (*i.e.* zooMFA1 and zooMFA2,  
 488 explaining 50.5 % of the mesozooplankton dataset total variance, Fig. 4), resulted in  
 489 identifying three clusters, on the basis of the classification tree presented in Fig. 7a  
 490 and the Nbclust method to identify significant clusters (Fig. 7b). The clusters' spatial  
 491 distribution and their variability in time are presented in Fig. 8.

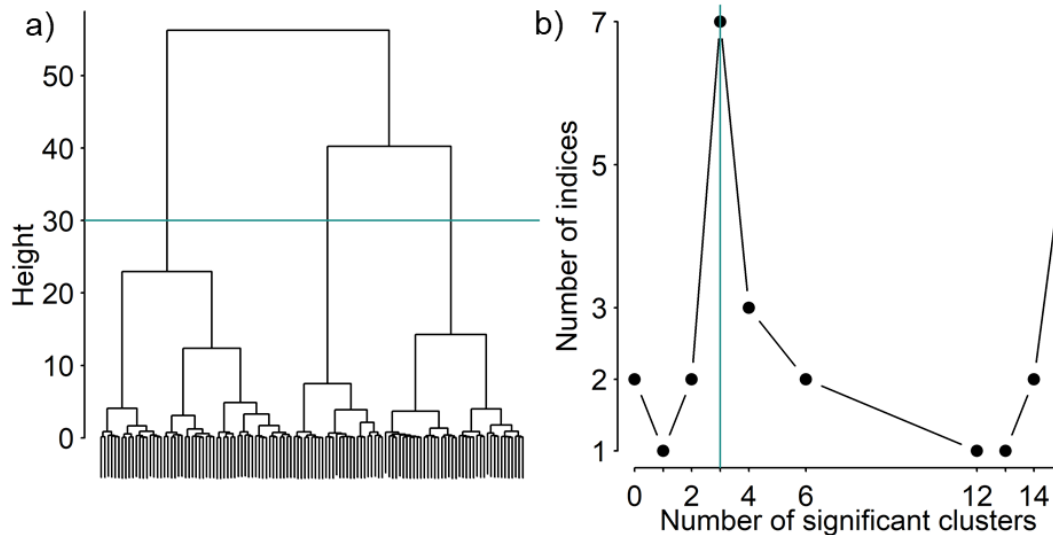


Fig. 7: (a) Classification tree of the hierarchical clustering of mean individuals (grid cells) coordinates in the mesozooplankton Multi Factor Analysis factorial space, made by the two selected mesozooplankton new descriptors (*i.e.* zooMFA1 and zooMFA2). (b) Number of clusters suggested by the Nbclust method (see section 2.7.3). Three clusters were retained (cut line in blue).

492

493 The time-consistent mesozooplankton spatial pattern combined coastal –  
 494 offshore and north – south gradients (Fig. 8a). A coastal cluster (G2, blue / medium  
 495 coloured cluster) extended from the Loire estuary to the south coast of the BoB, where  
 496 all taxa displayed higher-than-average biomass except large copepods (*i.e.* Calanidae,  
 497 Metridinidae and Euchaetidae) (Fig. 8c). A northern cluster (G3, yellow / light coloured  
 498 cluster) was located North of 45.5°N, from offshore waters along the shelf break to the  
 499 south Brittany coast. It was characterised mostly by Euchaetidae, and to a lesser extent  
 500 Metridinidae, Calanidae and meroplanktonic crustaceans' larvae, all of them showing  
 501 higher-than-average biomasses in this area (Fig. 8b). Finally, a southern cluster (G1,  
 502 purple / dark coloured cluster) extended from the central – shelf areas, south of 46°N,  
 503 mostly along the shelf break. Metridinidae dominated this southern mesozooplankton  
 504 community, and Appendicularians, Cladocerans, Echinodermata, Euchaetidae and  
 505 Siphonophorae also exhibited higher-than-average biomasses. On the contrary,  
 506 Cnidarians, meroplanktonic crustaceans' larvae and Temoridae displayed the lowest  
 507 overall biomass in this cluster (Fig. 8d). The spatial clusters corresponded to habitats  
 508 of particular communities that were consistent in time. The largest temporal variability  
 509 was located offshore in the northern cluster. The coastal and northern clusters covered



510 large spatial areas, whereas the southern cluster was less spatially extended and  
511 confined to the small southern outer-shelf and shelf break (Fig. 8a).

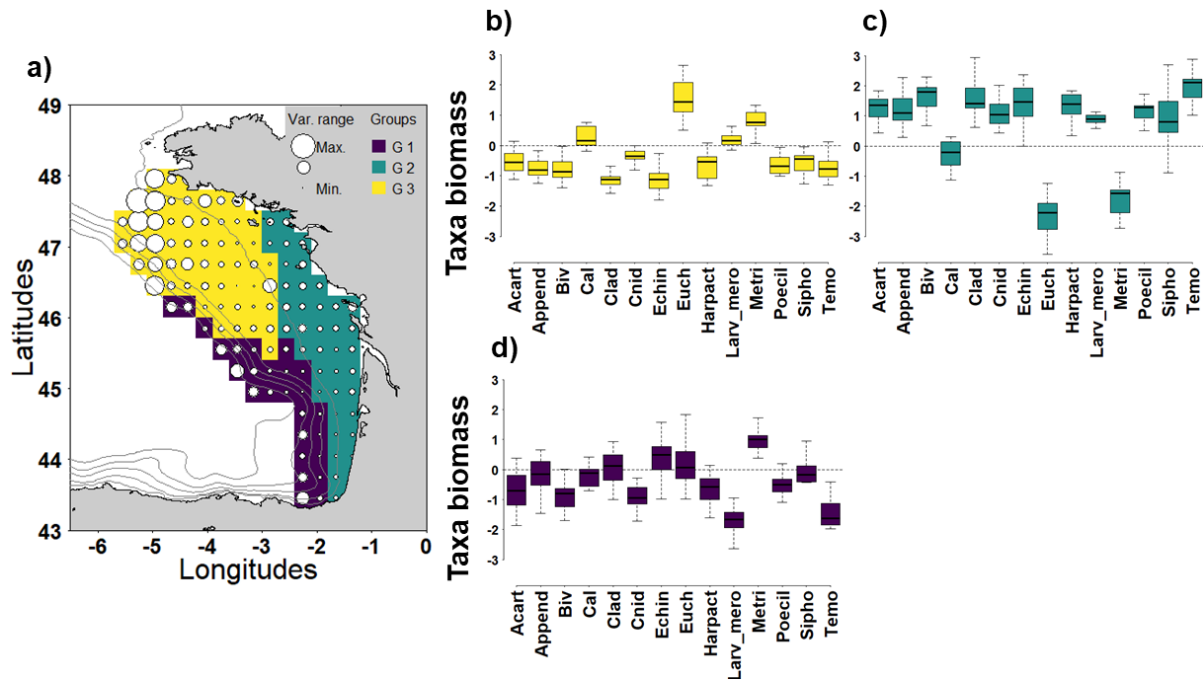


Fig. 8: (a) Time-consistent spatial patterns of the mesozooplankton community, derived from the hierarchical clustering of grid cells in the mesozooplankton Multiple Factor Analysis (MFA) space (two first principal components). White dots diameters are proportional to the inter-annual variability over the 16 years of the study. (b – d) Boxplots showing the inter-annual variability of the taxa biomasses ( $\mu\text{g dry weight.m}^{-3}$ ) in each cluster. The taxa shown are significantly correlated to the mesozooplankton MFA *PCs* with a frequency in time higher than 0.5. The horizontal dashlines mark the overall mean biomass (see Methods section 2.7.3).

512

513 Most of the grid cells within the coastal cluster G2 showed high occurrence  
514 frequencies in time ( $> 0.8$ ), meaning that they consistently belonged to this cluster over  
515 time (Fig. 9c). The occurrence frequency of the grid cells within the southern and  
516 northern offshore clusters were smaller and more variable over years. Some grid cells  
517 along the shelf break between  $46^\circ\text{N}$  and  $47^\circ\text{N}$  switched between both clusters,  
518 depending on the year (Fig. 9a and b). The annual distributions of clusters can be  
519 found in the Appendix B, Fig. B1.

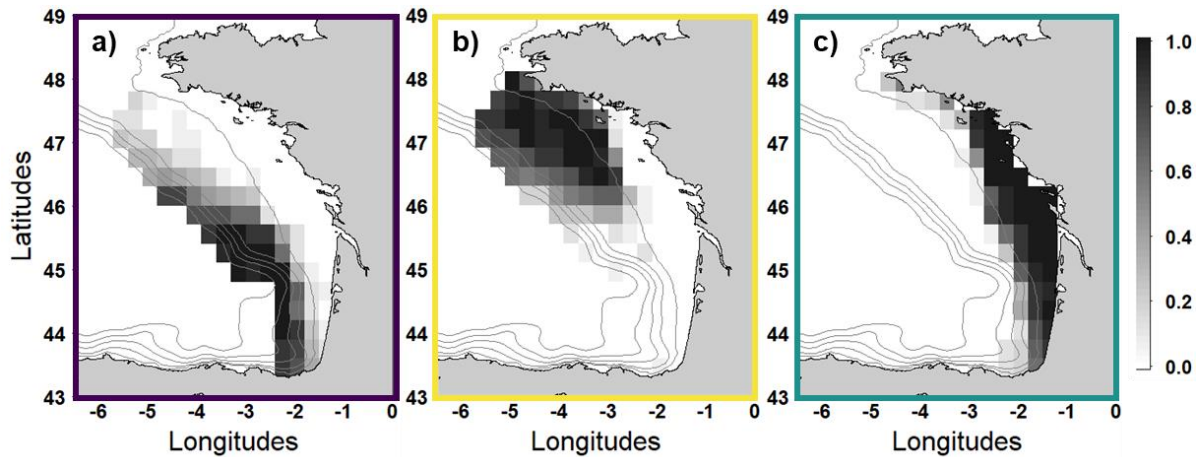


Fig. 9: Frequencies in time of grid cells' occurrences in each of the clusters forming the time-consistent spatial pattern: southern cluster G1 (a), northern cluster G3 (b), coastal cluster G2 (c).

### 521 3.4 Other ecosystem components MFA selected *PCs*

#### 522 3.4.1 Hydrology ecosystem component

523 Here we present the spatial patterns (MFA *PCs*) in the hydrology ecosystem  
 524 component, based on the mapping of the mean individuals' (grid cells') coordinates in  
 525 the hydrology MFA factorial space (Fig. 10a, c, e). The variables best correlated to the  
 526 MFA *PCs* are shown in Fig. 10b, d, f.

527 The hydroMFA1 underpinned a dominant coastal – offshore gradient. This structure  
 528 was highlighted by the lowest values off the Loire, Gironde and Adour estuaries (Fig.  
 529 10a, darkest cells) and the highest along the shelf break in the northern part of the BoB  
 530 (Fig. 10a, lightest cells). The surface salinity and the mixed layer maximum depth had  
 531 consistently positive correlations with hydroMFA1, meaning that these variables drove  
 532 the coastal-offshore gradient of hydroMFA1 with the highest values located offshore in  
 533 the northern part of the BoB between 2004 and 2019. On the contrary, the equivalent  
 534 freshwater height was constantly negatively correlated with hydroMFA1, revealing the  
 535 influence of the river plumes in coastal waters. Finally, surface temperature displayed  
 536 negative correlations with hydroMFA1 at the beginning (from 2004 to 2008) and at the  
 537 end (from 2014 to 2019) of the time series, suggesting that it contributed significantly  
 538 to the coastal – offshore gradient those years, with higher values in coastal areas (Fig.  
 539 10b).

540 The hydroMFA2 displayed a north – south gradient, with the highest values in the  
541 southern part of the BoB (Fig. 10c, lightest cells) and the lowest values in the northern  
542 part of the BoB, along the Brittany southern coast, north to the Loire estuary (Fig. 10c,  
543 darkest cells). The constant positive correlations of both surface and bottom  
544 temperature with the hydroMFA2 throughout the series highlighted higher water  
545 temperatures, in the south of the BoB, and colder water temperatures in the north (Fig.  
546 10d).

547 The potential energy deficit was the only variable significantly correlated to the  
548 hydroMFA3, with positive values (Fig. 10f), explaining the spatial pattern in hydroMFA3  
549 with highest values in the middle of the shelf centered on the 100 m isobaths (Fig. 10e,  
550 lightest cells).

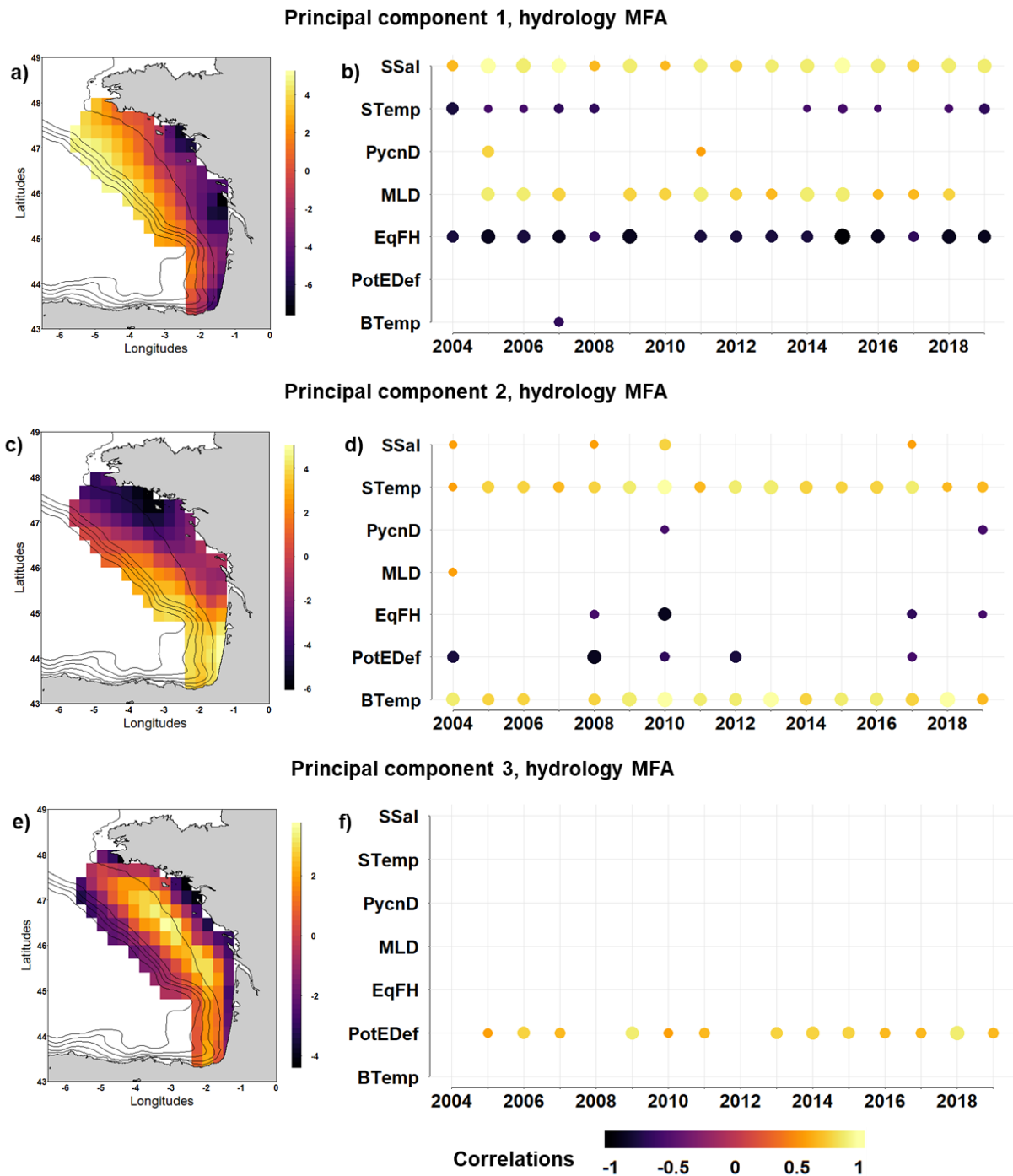


Fig. 10: Maps of mean individuals' (grid cells') coordinates on hydrology Multiple Factor Analysis (MFA) principal component (*PC*) 1 (a), 2 (c) and 3 (e), and time series of significant correlations (> |0.5|) between the seven hydrological parameters and the MFA *PC.1* (b), *PC.2* (d) and *PC.3* (f). The disk radii are proportional to the absolute value of the correlation coefficient. The hydrology MFA was performed on the maps time series of seven hydrological parameters derived from PELGAS and satellite data, from 2004 to 2019 (Table 2).

552 3.4.2 Primary producers ecosystem component

553 Here we present the spatial patterns of the primary producers ecosystem  
 554 component, based on the mapping of grid cells coordinates in the primary producers  
 555 MFA factorial space (Fig. 11a). The variables best correlated with the MFA PCs are  
 556 shown in Fig. 11b.

557 The phytoMFA1 exhibited a dominant coastal – offshore gradient. This structure  
 558 was highlighted by the low values offshore, along the shelf break in the south of the  
 559 BoB (Fig. 11a, darkest cells). High value were located in coastal areas, with highest  
 560 values off the Loire estuary and to a lesser extent, off the Gironde and Adour estuaries  
 561 (Fig. 11a, lightest cells). Surface chlorophyll-a concentrations of organisms larger than  
 562 20  $\mu\text{m}$ , total surface chlorophyll-a concentrations and satellite-derived chlorophyll-a  
 563 concentrations were consistently positively correlated to phytoMFA1, showing highest  
 564 values in coastal areas, especially in the Loire estuary, between 2009 and 2019 (Fig.  
 565 11b). On the contrary, the depth of the chlorophyll maximum displayed negative  
 566 correlations with phytoMFA1 approximately every other year, revealing higher values  
 567 along the shelf break in the south and contributing significantly to the coastal-offshore  
 568 gradient in 2009, 2011, 2013, 2015, 2016 and 2019 (Fig. 11b).

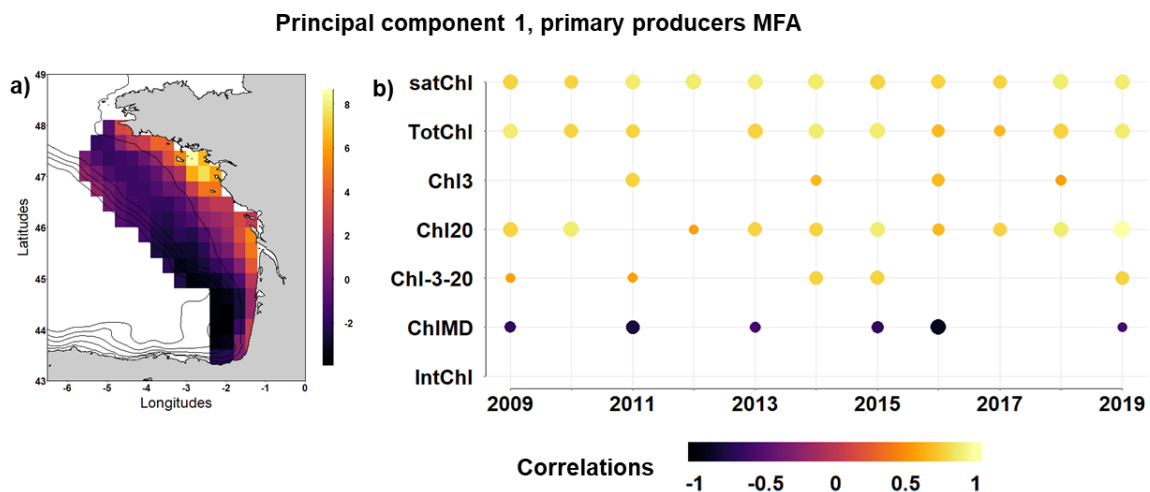


Fig. 11: Maps of mean individuals' (grid cells) coordinates on the primary producers Multiple Factor Analysis (MFA) principal component *PC.1* (a) and time series of significant correlations ( $> |0.5|$ ) between the seven phytoplankton variables and *PC.1* (b). The disk radii are proportional to the absolute value of the correlation coefficient. The primary producers MFA was performed on the maps time series of chlorophyll-a parameters derived from PELGAS and satellite data, from 2009 to 2019 (Table 2).

569           3.4.3 *Small pelagic fish ecosystem component*

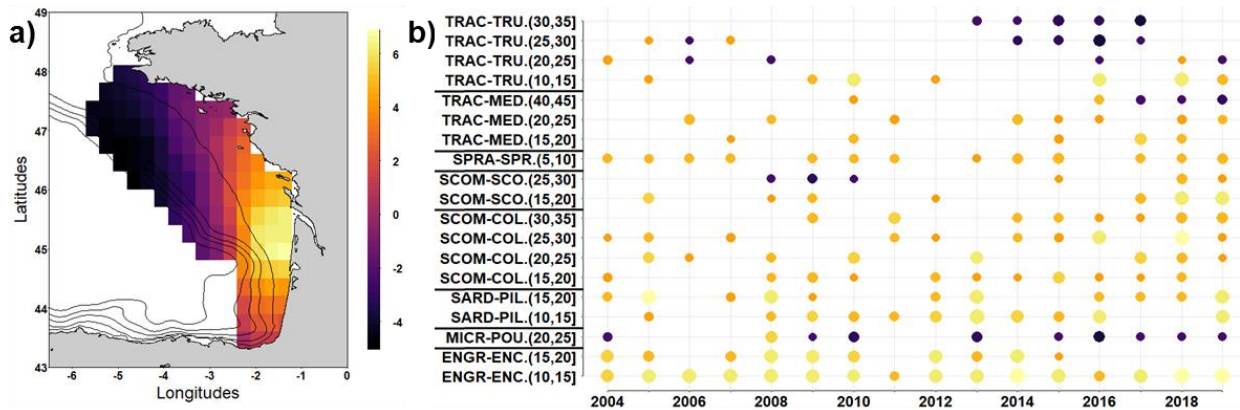
570           Here we present the spatial patterns of the SPF ecosystem component, based on  
571 the mapping of grid cells coordinates in the SPF MFA factorial space (Fig. 12a). The  
572 variables best correlated with the MFA PCs are shown in Fig. 12b.

573           The spfMFA1 map showed a south-eastern – north-western gradient. This structure  
574 was underpinned by the lowest values offshore north to 45.5°N, along the shelf break  
575 (Fig. 12a, darkest cells), and the highest in the south-eastern part, from the Gironde  
576 estuary to 44.5°N (Fig. 12a, lightest cells). Small anchovy (*i.e.* ENGR-ENC. (10-15)),  
577 sardine (*i.e.* SARD-PIL.), sprat (*i.e.* SPRA-SPR.) and Atlantic chub mackerel (*i.e.*  
578 SCOM-COL.) of all sizes were positively correlated to spfMFA1 (Fig. 12b), displaying  
579 higher biomass in the south-eastern part of the BoB. Note this pattern is particularly  
580 consistent for large Atlantic chub mackerel after 2013 (six consecutive years for > 30  
581 cm Atlantic chub mackerel). On the other hand, high biomass of blue whiting (*i.e.*  
582 MICR-POU) and large Atlantic horse mackerel (*i.e.* TRAC-TRU > 25 cm) were found  
583 offshore in the northern part of the BoB, showing negative correlations with spfMFA1,  
584 at the end of the time series (after 2014 and between 2013 and 2017, respectively)  
585 (Fig. 12b).

586           The spfMFA2 map showed an East – West gradient. This structure was underpinned  
587 by the low values in the coastal area, from 46°N to 47.5°N, including the Brittany  
588 southern coast (Fig. 12c, darkest cells), and high values in the southern part of the  
589 BoB and along the shelf break, south to 46°N (Fig. 12c, lightest cells). Sprat between  
590 10 and 15 cm (*i.e.* SPRA-SPR. (10-15)) had consistent negative correlations with the  
591 spfMFA2, meaning that they had higher biomass in coastal areas. On the contrary,  
592 large Atlantic mackerel (> 30 cm) (*i.e.* SCOM-COL.) had positive correlations with the  
593 spfMFA2, showing higher biomass in the south. The same pattern was observed for  
594 large Atlantic chub mackerel before 2013 (Fig. 12d).

595

### Principal component 1, small pelagic fish MFA



### Principal component 2, small pelagic fish MFA

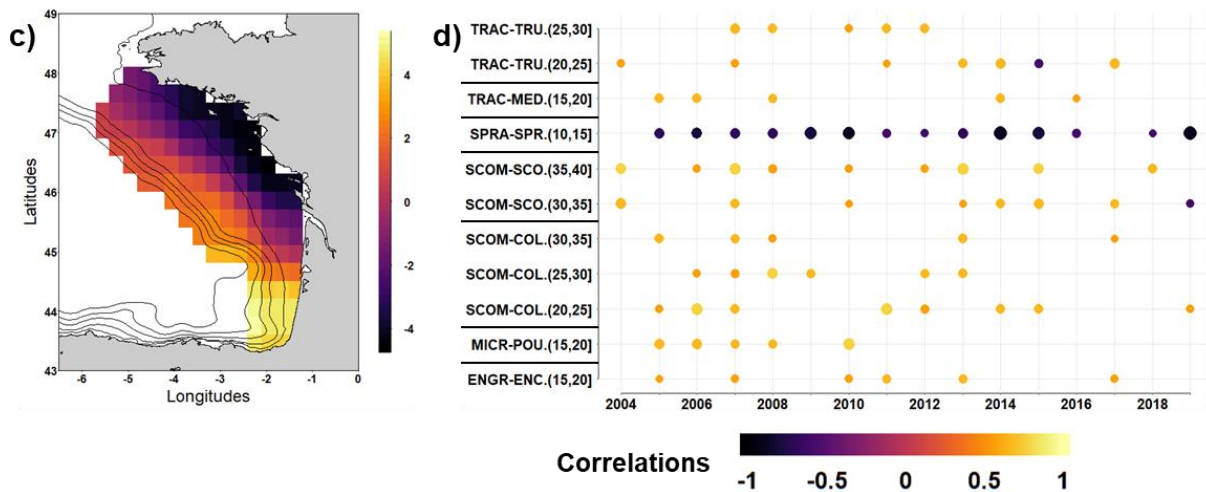


Fig. 12: Maps of mean individuals (grid cells) coordinates on small pelagic fish (SPF) Multiple Factor Analysis (MFA) principal component PC.1 (a) and PC.2 (c), and time series of significant correlations (> |0.5|) between the SPF species considered by 5 cm size classes and MFA PC.1 (b) and PC.2 (d). The disk radii are proportional to the absolute value of the correlation coefficient. The small pelagic fish MFA was performed on the maps time series of SPF species biomass considered by 5 cm size classes, calculated from the PELGAS data, from 2004 to 2019 (Table 3).

## 596 3.5 Correlates of the observed mesozooplankton space-time 597 patterns

598 We searched for correlates of the mesozooplankton main space-time patterns  
599 (zooMFA1 and zooMFA2) using GLMs. We used the six selected *PCs* in the hydrology,  
600 primary producers and SPF MFA as predictors.

601 *3.5.1 First PC in the mesozooplankton space-time pattern: zooMFA1*

602 For the initial full model built with the six predictors, all the VIF values were higher  
 603 than five (except that of hydroMFA3, VIF value = 3.7), showing collinearity between all  
 604 the predictors except hydroMFA3 (Appendix C, Table C.1). The predictors phytoMFA1  
 605 and spfMFA2 had highest linear correlation coefficient with the other predictors  
 606 (Appendix C, Fig. C.1) and were thus removed from the initial full model. The model  
 607 without these two predictors showed acceptable VIF values and thus the predictors  
 608 hydroMFA1, hydroMFA2, hydroMFA3 and spfMFA1 were kept for the subsequent  
 609 model selection procedure.

610 The stepwise backward model selection procedure showed that the four  
 611 predictors together constituted the best model explaining the zooMFA1 (lowest AIC,  
 612 Appendix C, Table C.2). The salinity and water column stratification gradient  
 613 (hydroMFA1, Fig. 5a and b) explained most of the variance in the model. To a lesser  
 614 extent, the North - South gradient in the water temperature (hydroMFA2, Fig. 5c and  
 615 d) and SPF biomass spatial patterns (spfMFA1, Fig. 12a and b) were also significant  
 616 contributors explaining the mesozooplankton coastal-offshore gradient (zooMFA1, Fig.  
 617 8a and b) (Table 4). On the other hand, the potential energy deficit spatial pattern  
 618 (hydroMFA3, Fig. 5e and f) was not significant in the model (Table 4), but a model  
 619 without this predictor was not significantly different from the best model, as their AIC  
 620 difference was smaller than two (Appendix C, Table C.2).

621 Table 4: Predictors included in the model explaining the first principal component of the  
 622 mesozooplankton MFA, with the estimated coefficients (“Estimate” and “Std. Error”) and their  
 623 significance in the model (“P value”).

Parameters	Estimate	Std. Error	P value
(Intercept)	2.49	0.17	< 2e-16 ***
hydroMFA1	-0.093	0.015	4.88e-09 ***
hydroMFA2	-0.047	0.009	2.55e-06 ***
hydroMFA3	-0.023	0.012	0.07
spfMFA1	0.088	0.015	8.79e-08 ***

624



625 3.5.2 Second PC in the mesozooplankton space-time pattern: zooMFA2

626 Using the same procedure as previously (VIF value of the predictors in the initial  
627 full model and correlation coefficients among predictors, see Appendix C, Table C.1  
628 and Fig. C.1), the predictors hydroMFA1, phytoMFA1 and spfMFA2 were removed  
629 from the analysis.

630 The stepwise backward model selection procedure was applied with the three  
631 selected predictors (*i.e.* hydroMFA2, hydroMFA3 and spfMFA1). The best model  
632 explaining zooMFA2 included hydroMFA2 and spfMFA1 (lowest AIC, Appendix C,  
633 Table C.3) but hydroMFA2 as the only significant predictor (Table 5). The map of  
634 zooMFA2 was underpinned by the North – South gradient in the biomass of  
635 meroplanktonic crustaceans larvae (zooMFA2, Fig. 8c and d). Thus, this structuration  
636 followed the North – South gradient in water temperature (hydroMFA2, Fig. 5c and d).  
637 Both predictors hydroMFA3 and spfMFA1 were not significantly explicative in the  
638 model (Table 5), but models without those predictors were not significantly different,  
639 as their AIC difference was smaller than two (Appendix C, Table C.3). This suggested  
640 that the potential energy deficit (hydroMFA3) and the SPF biomass spatial patterns  
641 (spfMFA1) were less influential in structuring in the mesozooplankton North – South  
642 gradient (zooMFA2).

643 Table 5: Predictors included in the model explaining the second principal component of the  
644 mesozooplankton MFA, with the estimated coefficients (“Estimate” and “Std. Error”) and their  
645 significance in the model (“P value”).

Parameters	Estimate	Std. Error	P value
(Intercept)	2.35	0.063	< 2e-16 ***
hydroMFA2	-0.13	0.006	< 2e-16 ***
hydroMFA3	0.007	0.009	0.434
spfMFA1	-0.008	0.005	0.113

646

## 647 **4 Discussion**

648 This study is the first space-time characterisation of mesozooplankton springtime  
649 assemblages, over the BoB continental shelf and almost two decades. Results show  
650 that the spatial patterns in the mesozooplankton assemblage was coherent with that  
651 observed in other pelagic ecosystem components (hydrology, primary producers and  
652 small pelagic fish). Coastal-offshore and North-South gradients were the main  
653 mesoscale spatial patterns. Small copepods, gelatinous and meroplanktonic  
654 organisms characterised coastal areas. Euchaetidae and meroplanktonic crustaceans'  
655 larvae showed higher biomass in the northern part of the BoB while Metridinidae,  
656 Cladocera, Appendicularia and Echinodermata had higher biomass in the southern  
657 part. Moreover, the high correlations between the years and the mesozooplankton  
658 MFA *PCs* underlined the stability in time of the above-mentioned spatial patterns. The  
659 high occurrence frequencies of grid cells to the clusters defining specific regions in the  
660 BoB evidenced the temporal consistency of the mesozooplankton assemblage spatial  
661 structure. In the other pelagic ecosystem components taken into account, salinity,  
662 water column stratification and primary producers supported a coastal-offshore  
663 gradient, while surface and bottom water temperature drew a North-South gradient. In  
664 addition, the spatial structure in the SPF biomass by size classes obtained here was  
665 consistent with that obtained by Doray et al. (2018) with a similar methodology. Finally,  
666 the observed mesozooplankton patterns were correlated with those of hydrology and  
667 SPF ecosystem components.

### 668 **4.1 Mesozooplankton assemblages temporal stability**

669 The highest inter-annual variability occurred offshore in the northern part of the BoB,  
670 where some grid cells had missing data at the beginning of the time series. Therefore,  
671 this result can be due to the missing data imputation in this area. Yet, a true biological  
672 variability in the mesozooplankton assemblage could also be real close to the shelf  
673 break. The consistency in time of the mesozooplankton assemblage was a conclusion  
674 drawn by Irigoien et al. (2008), studying spring zooplankton distribution between 1998  
675 and 2006 in the southern part of the BoB. Feuilloley et al. (2022) also showed the high  
676 stability through time of the zooplankton density, body size and taxonomic composition  
677 in the North Western Mediterranean Sea, during the period 1995-2019. Yet, abrupt

678 shifts in community structure could be expected when analysing long-term dynamics  
679 in the context of environmental changes. Such regime shifts were detected in the late  
680 1980s and in the 1990s in several marine regions in the Northern Hemisphere  
681 (Beaugrand et al., 2015; Morse et al., 2017; Bode et al., 2020; Chust et al., 2022).  
682 Recently, Dessier et al. (2018) found a change in the large copepod species  
683 dominance in the South of the BoB between the period 2003 – 2006 and the period  
684 2007 – 2009, preceding a return toward the initial situation over the period 2010 –  
685 2013. Also, Iriarte et al. (2022) distinguished three periods in the relative density of  
686 copepod species in the Southeastern BoB: 1998 – 2007, 2008 – 2013 and 2014 –  
687 2015. Here, the mesozooplankton community was studied using 24 taxonomic groups,  
688 a less detailed level than the species level used in these previous studies. Fine  
689 taxonomic resolution could reveal shifts in species dominance that could compensate  
690 each other when analysing the community at a broader taxonomic resolution and,  
691 consequently, the overall structure could appear consistent in time. This point  
692 constitutes a potential limit of our study. Shifts in species dominance could have  
693 occurred in the zooplankton community between 2004 and 2019 but the broad  
694 taxonomic resolution used here may not have enable the observation of such possible  
695 changes, resulting in an apparent temporal stability over the time series. On the other  
696 hand, these studies were done at a smaller spatial scale (southern part of the BoB in  
697 Dessier et al. (2018); two neritic stations in the Cantabrian Sea in Iriarte et al. (2022)),  
698 and the shifts occurring in restricted areas could be undetectable at a broader spatial  
699 scale or not affect the large scale distribution pattern.

## 700 **4.2 Coherent spatial patterns across pelagic ecosystem** 701 **components**

702 Our results agreed with the spatial organisation in mesozooplankton assemblages  
703 along a coastal-offshore gradient previously showed in the BoB (Albaina and Irigoien,  
704 2004, 2007a; Sourisseau and Carlotti, 2006; Irigoien et al., 2008; Vandromme et al.,  
705 2014; Dessier et al., 2018) and in other shelf ecosystems (Marcolin et al., 2013; Pepin  
706 et al., 2015; Noyon et al., 2022). Our study also confirms that surface and bottom water  
707 temperature, salinity-related parameters and water column stratification indices were  
708 key hydrographic variables that correlated with this mesozooplankton springtime  
709 spatial structure (Albaina and Irigoien, 2004; Zarauz et al., 2007; Irigoien et al., 2011;

710 Dessier et al., 2018; Iriarte et al., 2022). Hydrological parameters integrating the  
711 historical seasonal conditions, such as the equivalent freshwater height, mixed layer  
712 depth and the deficit of potential energy, were key variables to study the  
713 mesozooplankton spring habitat, taking into account the onset of spring conditions.  
714 The correlations between the hydrology and primary producers space-time patterns  
715 observed here confirm the important role of nutrients enrichment at the rivers' mouths  
716 in the primary producers development in spring in the BoB (Guillaud et al., 2008).  
717 Primary producers are known to drive the mesozooplankton dynamics in the BoB  
718 (Zarauz et al., 2008; Dessier et al., 2018; Iriarte et al., 2022) and elsewhere (Pepin et  
719 al., 2015; Capuzzo et al., 2018). Here, because of the method used, the primary  
720 producers' space-time pattern was not included in the models to explain the  
721 mesozooplankton space-time structure. Yet, the correlations between the hydrology  
722 and primary producers ecosystem components (Appendix C, Fig. C.1) combined with  
723 the hydrology space-time patterns significance in the linear models (Tables 4 and 5)  
724 suggest that the hydrological landscapes may influenced the mesozooplankton  
725 community structure through the primary production and trophic relationships between  
726 phytoplankton and mesozooplankton in the BoB.

727 The North-South gradient in the mesozooplankton assemblage in the BoB was  
728 never described before, as previous studies mostly focused on the southern part of the  
729 BoB (Albaina and Irigoien, 2004, 2007a; Irigoien et al., 2008; Dessier et al., 2018). At  
730 the scale of the Northeast Atlantic, a strong latitudinal effect potentially driven by  
731 temperature regimes has been shown on the seasonal and interannual variability of  
732 two copepod species (Valdes et al., 2022). Moreover, Fanjul et al. (2019) demonstrated  
733 that some zooplankton groups' abundances displayed temperature-mediated  
734 latitudinal differences, and that meroplankton contributed more than holoplankton to  
735 the main between-site differences. In the BoB, the spring meroplanktonic crustaceans'  
736 larvae distribution could be driven by the temperature-induced latitudinal gradient,  
737 resulting in higher biomass in the northern coastal area. Nevertheless, this result could  
738 also highlight the timing of the surveys, lasting one month and starting in the south of  
739 the BoB, potentially before the development of the meroplanktonic crustaceans' larvae  
740 in this area.

741 Furthermore, the correlation between mesozooplankton and SPF spatial patterns  
742 confirms the consistent spatial structure across pelagic ecosystem components that

743 have been reported in the BoB (Doray et al., 2018a; Petitgas et al., 2018), possibly  
744 mediated via predation and/or forced by hydrological structuring. For example, taking  
745 into account the eight major SPF species inhabiting the BoB, Bachiller and Irigoien  
746 (2015) described an overall high diet overlap and concluded to a top – down control by  
747 these planktivorous fishes on zooplankton. Elsewhere, such links between ecosystem  
748 components were highlighted in the Mediterranean Sea (Feuilloley et al., 2020), on the  
749 Portuguese continental coast (Fonseca et al., 2022), in the North Sea (Capuzzo et al.,  
750 2018) and in Barents Sea (Stige et al., 2014). Here, the spatial structures in the  
751 hydrological parameters, primary producers and SPF were similar to that in  
752 mesozooplankton assemblages and were also consistent in time. Spatial structures  
753 had been described for different ecosystem components in the BoB (Special Issue No.  
754 166 in Progress in Oceanography, 2018), including river plumes, primary producers,  
755 fish and top-predators but this is the first time they are described for the  
756 mesozooplankton and shown to be similar to the other ecosystem components. Such  
757 consistency in the spatial structure across ecosystem components and over time is  
758 remarkable. Donohue et al. (2013) discussed the multiple and related facets of  
759 ecosystem stability. In the case of the BoB continental shelf at springtime, hydrographic  
760 landscapes resulting from river plumes' extents on the shelf and temperature gradients  
761 (Koutsikopoulos and Le Cann, 1996; Koutsikopoulos et al., 1998; Castaing et al., 1999;  
762 Planque et al., 2004; Guillaud et al., 2008) certainly impede strong structuring meso-  
763 scale forcings on the BoB regional pelagic ecology. The structure could also be  
764 maintained and reinforced by top-down and bottom-up controls through trophic levels  
765 as generally reported for the North Atlantic (Frank et al., 2007) although not directly  
766 studied here.

### 767 **4.3 Data and methods used**

768 Such regional study linking several ecological components was made possible by  
769 the unique long-term and spatially resolved datasets constructed from integrated  
770 surveys. Thus, maintaining such surveys is of primary importance for ecosystem  
771 assessment research, especially in the context of climate change, as they provide a  
772 platform dedicated to gather data from several biological and abiotic components at  
773 the same temporal and spatial resolution (Kupschus et al., 2016; Doray et al., 2018c).  
774 Nevertheless, it should be noted that these *in situ* data sets provide a temporal

775 snapshot of the ecosystem state at the time of the surveys. Any apparent interannual  
776 variability could be caused by interannual differences in the seasonal evolution of the  
777 system (Huret et al., 2018). The combination of imaging and deep learning tools  
778 provided an efficient method to analyse a high quantity of mesozooplankton samples  
779 quickly and in a standardized way. This analytical pipeline, including also automated  
780 procedures for spatial gridding and missing data imputation, allowed to construct a  
781 mesozooplankton spatially resolved time series over 16 years. Finally, the MFA was  
782 well suited to estimate time-average spatial patterns in a multivariate dataset and  
783 perhaps less designed to evidence temporal trends in the mean distribution, at least in  
784 the way we implemented the method. Moreover, the results showing consistency in  
785 time of the spatial patterns were based on a subset of MFA principal components,  
786 accounting only for part of the total variance. The residual variance, not fully explained  
787 here, showed exceptional events, which we did not fully explore. Yet, the major time  
788 variations happened in particular years (for example, lowest mesozooplankton  
789 biomass in 2011 and 2015, Fig. 3) and were located in particular areas (offshore in the  
790 northern part of the BoB, Fig. 8), suggesting no obvious trend in time.

## 791 **5 Conclusion**

792 Thanks to long-term spatially resolved time series, we provided the first space-time  
793 characterisation of mesozooplankton springtime assemblage, correlated with other  
794 pelagic ecosystem components (hydrology, primary producers and SPF), over 16  
795 years and the whole BoB continental shelf. We demonstrated coherent spatial patterns  
796 across pelagic ecosystem components supported by coastal-offshore and North-South  
797 gradients. Moreover, we highlighted the remarkable stability in time of these spatial  
798 patterns. The springtime spatial structure of biotic components seems to be based on  
799 spring hydrological landscapes and is potentially related to bottom-up and top-down  
800 trophic controls, although a more detailed study on the ecosystem components'  
801 correlations is needed to fully confirm this hypothesis. Our regional study relies on  
802 long-term spatially resolved datasets originating from integrated surveys and provide  
803 key understanding of the ecosystem structure and dynamics. Therefore, our work  
804 emphasize the importance of such surveys for ecosystem assessment research  
805 especially in the context of climate change.

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## 820 **7 Authors contributions**

821 NG: conceptualized the study, generated and scrutinized the imaging data, conducted  
822 the numerical analyses, drafted the paper. J-BR: conceptualized the study, supervised  
823 NG work, scrutinized the imaging data, drafted the paper. CD: conceptualized the  
824 study, supervised NG work, drafted the paper. MD: supervised numerical analyses and  
825 provided fish data, participated in the drafting. MH: provided hydrology data,  
826 participated in the drafting. PP: conceptualized the study, supervised NG work, drafted  
827 the paper. All co-authors reviewed and revised the paper before submission.

828

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1136

## 1137 **9 Supplementary Materials**

### 1138 **A. Missing data imputation**

1139 Following the spatial gridding procedure, some data were missing, due to missing  
1140 sampling stations in the northern part of the Bay, mostly at the beginning of the time  
1141 series (Table A.1). Following Josse & Husson (2016), an algorithm based on iterative  
1142 Principal Components Analysis (PCA) from the *MissMDA R* library was used to impute  
1143 estimated values in place of the missing ones in the hydrology, phytoplankton and  
1144 mesozooplankton datasets. For each variable, the data were organized in matrix grid  
1145 cells x years. The first step of the algorithm consisted in the imputation of missing  
1146 values with the annual means (for example, if the grid cell numbered 33 had not  
1147 Acartiidae's biomass in 2004, it was filled with the 2004 mean Acartiidae's biomass).  
1148 Then, a first PCA was performed on the imputed matrix and the estimated values  
1149 replaced the annual mean values previously imputed to fill the missing data. Then, a  
1150 second PCA was performed and the estimated values from the first PCA were replaced  
1151 by the estimated value from the second PCA. The algorithm kept running iteratively  
1152 until the difference between two successive estimated values was smaller than a  
1153 threshold (set at 1e-06 by default ; Josse & Husson, 2016).

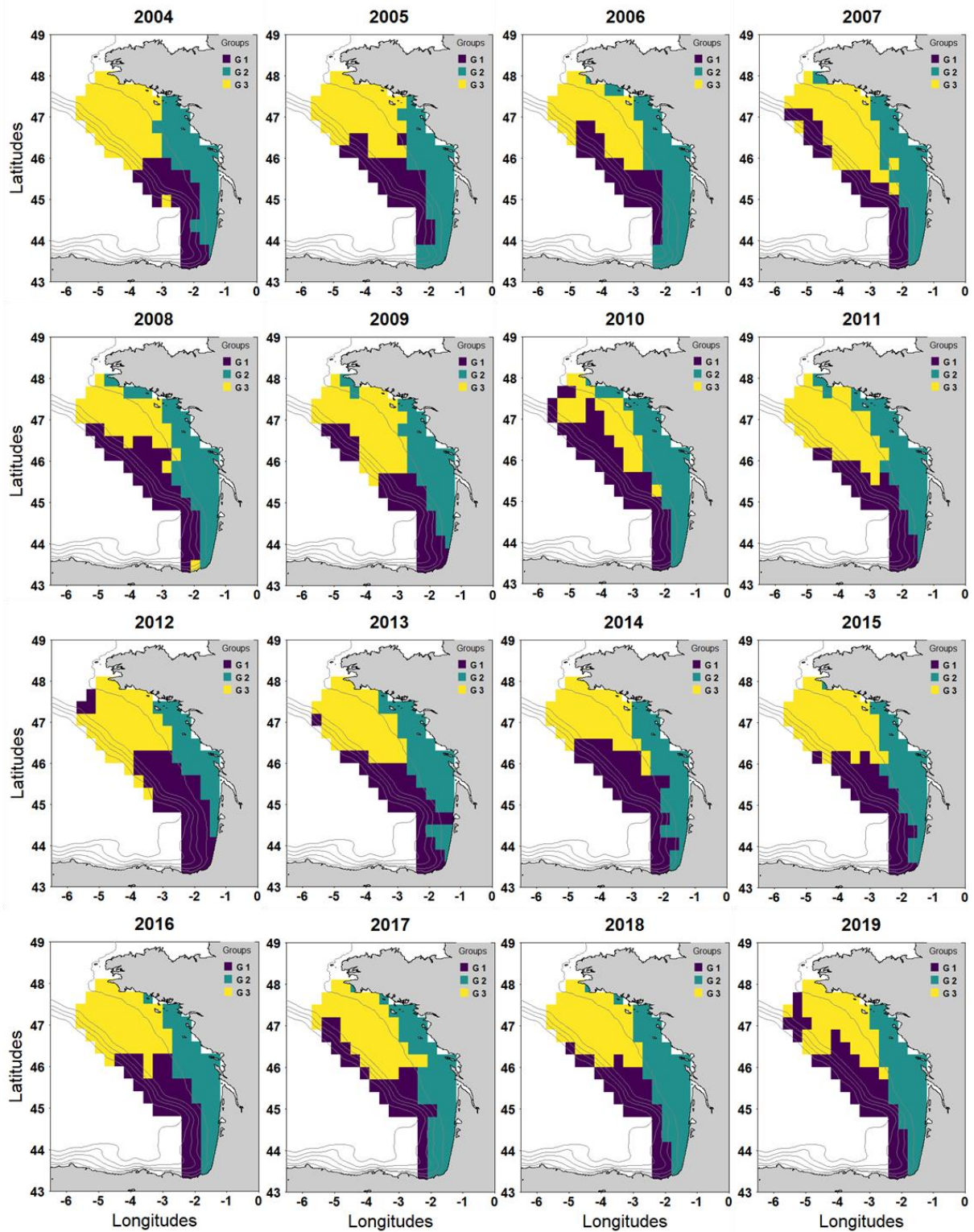
Table A1: Number and percentage of grid cells with missing data each year.

Years	Number of missing cells	Percentage of missing cells
2004	6	5
2005	17	14.1
2006	22	18.2
2007	12	9.9
2008	4	3.3
2009	0	0
2010	0	0
2011	2	1.7
2012	5	4
2013	5	4
2014	2	1.7
2015	0	0
2016	0	0
2017	0	0
2018	0	0
2019	1	0.8

1154

1155

## B. Annual spatial clustering



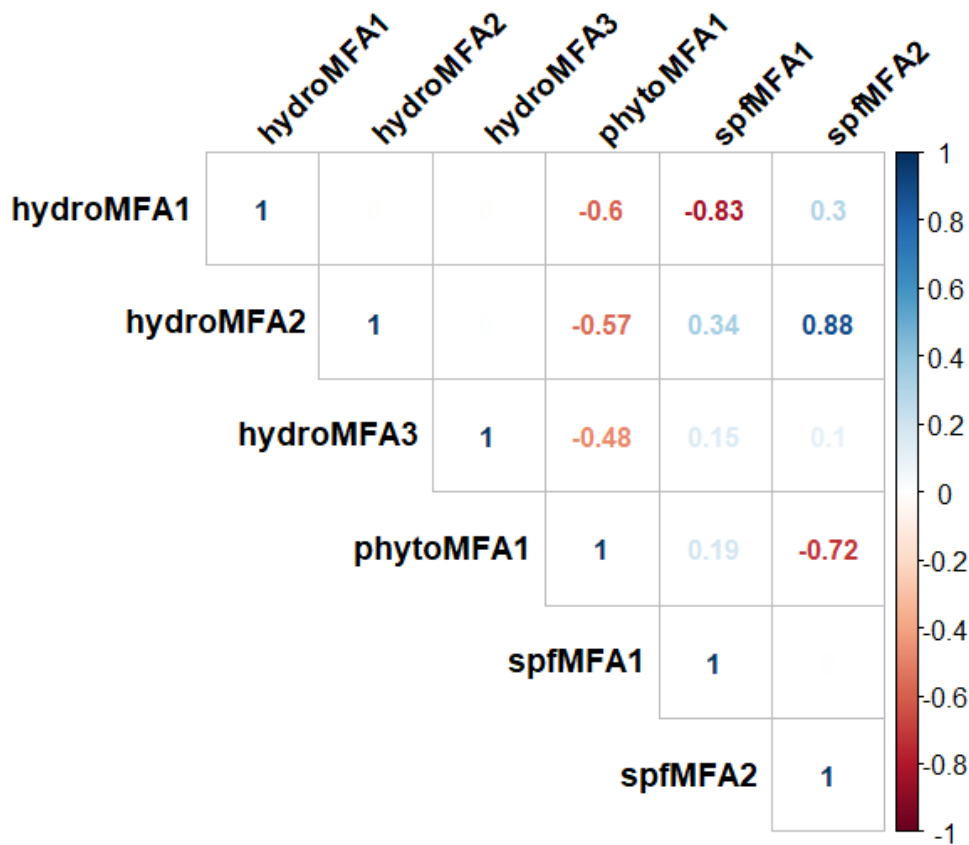
1157 Fig. B1: Annual spatial extent of the clusters identified with the hierarchical clustering of grid  
 1158 cells in the mesozooplankton Multi Factor Analysis factorial space.

1160 **C. Correlates of the observed mesozooplankton space-time**  
 1161 **patterns**

1162 Table C.1: Variance Inflation Factor (VIF) calculated for each Multi Factor Analysis selected  
 1163 principal component in the hydrology, primary producers and small pelagic fish datasets, which  
 1164 are used as predictors in Generalized Linear Models. A VIF value higher than five identifies  
 1165 problematic multicollinearity among predictors.

Principal components	hydroMFA1	hydroMFA2	hydroMFA3	phytoMFA1	spfMFA1	spfMFA2
VIF values	12.6	13.9	3.7	12.3	7.4	10.1

1166



1167 Fig. C.1: Correlation coefficients among all the predictors.

1168

1169 Table C.2: Results of the stepwise backward model selection procedure showing the models  
 1170 explaining the first principal component in the mesozooplankton Multi Factor Analysis  
 1171 zooMFA1 and their Akaike's Information Criterion (AIC).

Predictors included in the model	AIC
hydroMFA1 + hydroMFA2 + hydroMFA3 + spfMFA1	450.05
hydroMFA1 + hydroMFA2 + spfMFA1	451.31
hydroMFA1 + hydroMFA3 + spfMFA1	473.49
hydroMFA1 + hydroMFA2 + hydroMFA3	479.17
hydroMFA2 + hydroMFA3 + spfMFA1	491.39

1172

1173 Table C.3 Results of the stepwise backward model selection procedure showing the models  
 1174 explaining the second principal component in the mesozooplankton Multi Factor Analysis  
 1175 zooMFA2 and their Akaike's Information Criterion (AIC).

Predictors included in the model	AIC
hydroMFA2 + hydroMFA3 + spfMFA1	311.93
hydroMFA2 + spfMFA1	310.57
hydroMFA2	310.73
spfMFA1	771.64

1176

1177