

Supplementary Information for

Ice and ocean constraints on early human migrations into North America along the Pacific coast

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Datasets S1 to S5

Supplemental Text

Climate change impacts on marine ecosystems

High latitude Northern Hemisphere climate was volatile during the last glacial and deglacial periods, with abrupt warming and cooling events occurring on timescales of years to decades (1). These events would have transformed landscapes, weather systems, and terrestrial and marine ecosystems, requiring relatively rapid adaptation adjustments for ecological and human communities. Here we address how changing climate conditions may have affected marine ecosystems along the Pacific Rim, and how that may have in turn, affected human migration or coastal habitability at various times.

Kelp ecosystems and sea ice environments

Kelp forests have been hypothesized to have been key for supporting human dispersal to the Americas (2,3). Expanded convoluted shorelines and shallow rocky intertidal zones in the Beringian Transitory Archipelago during the LGM would likely have been conducive to kelp ecosystems (4). The persistence of the Steller's sea cow, a ~10-ton sirenian that relied exclusively on kelp and ranged along the Pacific Coast from Japan to California and across the Bering Sea shelf, suggests that kelp habitats remained at least regionally intact throughout the Pleistocene and early Holocene. This marine mammal may have been an important species in maintaining healthy kelp ecosystem prior to its extinction (5), and an important marine resource for ancient coastal people, as it was docile, confined to shallow waters, and provided enormous quantities of meat and oil (6). Neither Steller's sea cow nor kelp forests can survive in persistent multiyear ice; ice scouring can dislodge kelp hold fasts, and excessive freshwater input and sediment loads can negatively impact growth (7). In modern Arctic environments, however, kelp can be highly productive and well adapted to seasonal sea ice and light limitation (7). Thus, kelp environments may have thrived during the glacial period throughout coastal Beringia, despite seasonal sea ice. However, local conditions were likely highly variable in space and time, and conditions in the Gulf of Alaska during the Siku ice surge events may have been particularly antagonistic for kelp growth.

Expansions of the oxygen minimum zone

The abrupt ocean warming that occurred during the Bølling-Allerød and early Holocene periods caused major changes to marine ecosystems across the North Pacific. These events were accompanied by an expansion of the oxygen minimum zone (OMZ) across the North Pacific and rapid transitions to deep-water hypoxia, with a contraction of benthic faunal diversity at intermediate depths along the Alaskan and California margins (8,9). However, it remains unclear how this event impacted upper ocean ecosystems, such as kelp forests. On the one hand, increased primary productivity during the Bølling-Allerød may have supported thriving marine ecosystems, providing energy to fuel extensive trophic webs (Fig. S8a). As is seen in the modern Bering Sea 'Green Belt' along the shelf edge where both macro and micronutrients are available from shelf mixing, high levels of primary productivity are associated with some of the most productive marine waters, supporting large stocks of fish, squid,

waterfowl, and marine mammals (10). On the other hand, high *export* productivity doesn't necessarily indicate a healthy marine ecosystem, as export efficiency can be decoupled from primary productivity and instead reflect rapid removal of organic carbon from the upper ocean, limiting its utilization for heterotrophic grazing (11). For example, diatomaceous algal blooms can have high efficiency of carbon export that can lead to a self-sedimentation effect which promotes laminated opal-rich sediments (12), like those observed during the Bølling-Allerød. In events like this, algal blooms may induce poisoning of the food chain and/or subsurface hypoxia through rapid deoxygenation of the water column as sinking organic material is respired (Fig. S8b). It remains unclear whether ocean deoxygenation at intermediate depths impacted upper ocean ecosystems during the Bølling-Allerød and early Holocene hypoxic events.

While it is unlikely that sustained hypoxia extended into shallower depths, it is possible that the vertical expansion of the OMZ led to more frequent hypoxic events, especially in upwelling areas, such as the California Current system. Even episodic events that caused detrimental effects to reliable food sources may have been an impetus to seek additional resources inland, possibly promoting movement into the hinterlands. This may in part account for inland archaeological sites that date to the Bølling period, but no coastal evidence for occupation that has been uncovered during this period. Submergence of coastal sites by postglacial sea level rise makes locating earlier deglacial coastal sites challenging, however, a few coastal areas have remained above sea level throughout the deglaciation (13,14), so additional factors affecting the timing of habitability must be assessed. For example, the Daisy Cave site on San Miguel Island (one of the Channel Islands off southern California) contains a few artifacts and charcoal from a possible hearth feature that has been dated to ~18.6 ka (15,16). Although situated several kilometers from the coast at this time, the site was available for occupation throughout the deglaciation and remains above sea level.

With Pacific Coast archaeological sites occupied prior to the Holocene, the early Holocene hypoxic event presents a possible opportunity to assess whether OMZ expansion had discernable impacts on nearshore coastal ecosystems. A possible multi-centennial hiatus in the Channel Island occupation between ~11.1-10.7 ka (17) is similar in timing to the early Holocene hypoxic event (8,18,19), suggesting a possible connection between hypoxia and a temporary abandonment of the Channel Islands. Confirming such connections will require greater precision in linking terrestrial and marine age models, as well as additional proxy evidence for changes in species composition in response to changing climate conditions.



Figure S1. Simulations of ocean currents in the North Pacific during glacial and interglacial periods. Mean annual surface ocean velocity (m/s) for the modern climate state (top) and the last glacial maximum (bottom), showing a strengthening of most currents during the LGM, including the Kuroshio, Oyashio, Kamchatka, Bering Slope, and Alaska Current systems. Boundary currents flow in a cyclonic (anticlockwise) direction. The California Current is the only major current that shows a weakening during glacial conditions relative to the modern period.



Figure S2. Ocean current vectors for the Northeast Pacific for various climate states and sea levels.



Figure S3. Mean current velocity (m/s) in glacial simulations with varying freshwater fluxes (averaged from July-Dec). Note the difference in scale shown here (0-0.7 m/s) relative to the scale shown in Fig. 2 (0-0.3 m/s) to accommodate the larger velocities associated with the high freshwater fluxes.



Figure S4. Maximum and mean current velocity for various meltwater fluxes shown in Figure S3.



Figure S5. Changes in Winter (DJF) and Summer (JJA) surface salinity, sea surface temperature, surface air temperature, and current strength for the 1m sea level equivalent (SLE) / 500 yr freshwater flux (FWF) experiment minus control. Change in current strength (lower panel) is shown as percent change with the directional vectors overplotted from the control simulation.









1

FWF changes in Sea Surface Temperature (degC)





2

FWF changes in Near-Surface Air Temperature (K)





Figure S6. Same as Figure S5 for the 2m sea level equivalent / 500 yr experiment.



Fig. S7. Reconstructed SST records from the Northeast Pacific margin (A-G) that contribute to the stacked average records (H-J). A) $U_{37}^{K'}$ record from U1419 (20); **B**) $U_{37}^{K'}$ record from EW0408-85JC (8), **C**) $U_{37}^{K'}$ record from EW0408-87JC (21); **D**) $U_{37}^{K'}$ record from EW0408-66JC & EW0408-26JC (22); **E**) Mg/Ca-based SST reconstruction from MD02-2496 on *Globigerina bulloides* (Gbull; light green) and *Neogloboquadrina pachyderma* (Npl; dark green) (23); **F**) $U_{37}^{K'}$ record from JT96-09 (24); **G**) $U_{37}^{K'}$ record from ODP 1019 (25,26, age/depth model modified by ref. 21); **H**) Average stack using the high-resolution records (~100-yr average or greater); **I**) Average stack using all records (on a 200 yr timestep); **J**) Normalized average using all cores; **K**) Number of records contributing to the high-resolution stack (blue) and the all-core and normalized stacks (pink). Error bars on the stacked records reflect the standard error of the mean.



b) High productivity algal-dominated 'unhealthy' ecosystem



Figure S8: Contrasting scenarios for coastal ecosystems during periods of OMZ expansion, such as occurred during the Bølling-Allerød and early Holocene. Paleoceanographic reconstructions indicate ocean warming and an expansion/intensification in sedimentary hypoxia at depths between 400-1600 meters (indicated by dark blue shading), concurrent with increased burial of biosilica (diatoms) and organic carbon; however little information on the composition of coastal marine ecosystems is currently available. In the 'healthy' scenario (a), favorable conditions—such as high nutrient availability, mixing, and a stable but shallow warm surface layer—promote high primary productivity which supports a thriving ecosystem; adequate mixing keeps the upper ocean well ventilated while excess carbon export drives subsurface deoxygenation. In the 'unhealthy' scenario (b) abrupt ocean warming leads to heat stress, stratification, and a system dominated by diatom algal blooms, which deplete subsurface oxygen as sinking organic material is respired. Figure adapted and modified from (8).

Table S1. Ages for earliest archaeological sites in North America and Beringia										
Site	General Location	Latitude (N)	Longitude (W)	Earliest 14C age (kyr BP)	Calibrated minimum age (ka)	Calibrated maximum age (ka)	Calibrated mean age (ka)	Primary reference for site/most recent radiocarbon age	Compilation reference	
North of Ice Sheets					-8- ()	-8- ()	()			
Bluefish Cave	Yukon Canada	67.2	140.8	19.65 ± 0.13	23.31	24.04	23.67	Bourgeon et al., 2017	Beccera-Valdivia & Higham 2020	
Swan Point	Alaska	63.3	146.0	12.50 ± 0.15	14.15	14.65	14.40	Holmes et al., 2011	Beccera-Valdivia & Higham 2020	
Little John (KdVo-6)	SW/ Yukon	70.5	144.0	12.35 ± 0.06 12.02 ± 0.07	13.84	14.32	14.08	Faston et al., 2014	Beccera-Valdivia & Higham 2020	
Walker Road	Alaska	64.0	140.5	11.82 + 0.20	13.08	14.95	14.05	Goebel et al., 1996	Beccera-Valdivia & Higham 2020	
Broken Mammoth	Alaska	64.3	146.1	11.77 ± 0.22	13.26	14.13	13.69	Holmes et al., 2001	Beccera-Valdivia & Higham 2020	
Dry Creek	Alaska	63.9	149.0	11.58 ± 0.04	13.35	13.70	13.53	Graf et al., 2015	Beccera-Valdivia & Higham 2020	
Moose Creek	Alaska	64.1	149.1	11.19 ± 0.06	12.98	13.74	13.36	Powers et al., 1989	Beccera-Valdivia & Higham 2020	
Owl Ridge	Alaska	64.0	149.6	11.06 ± 0.06	12.75	14.32	13.53	Graf et al., 2019	Beccera-Valdivia & Higham 2020	
Mead	Alaska	64.3	146.1	11.46 ± 0.05	13.01	13.52	13.27	Potter et al., 2011	Beccera-Valdivia & Higham 2020	
Mesa	Alaska	68.4	155.8	11.82 ± 0.10 11.66 ± 0.08	9.36	13.44	11.63	Kunz et al., 1994	Beccera-Valdivia & Higham 2020 Beccera-Valdivia & Higham 2020	
South of Ice Sheets										
Chiquihuite Cave	Mexico	24.6	101.1	27.93 ± 0.08	31.41	33.15	32.28	Ardelean et al, 2020	Beccera-Valdivia & Higham 2020	
Goult	New Mexico	32.8	106.3	22.80 ± 32 21.70 ± 1.4	20.53	23.62	22.08	Williams et al. 2018	Roccora Valdivia & Higham 2020	
Meadowcroft Bockshelter	Pennsylvania	40.3	80.5	16 18 + 0.98	18.62	20.44	21.52	Adovasio et al. 1998	Beccera-Valdivia & Higham 2020	
Cactus Hill	Virginia	37.0	77.3	18.30 ± 1.4	18.97	20.59	19.78	Feathers et al., 2006/8	Beccera-Valdivia & Higham 2020	
Cooper's Ferry	Idaho	45.9	116.4	13.17 ± 0.07	15.28	16.56	15.92	Davis et al., 2019	Beccera-Valdivia & Higham 2020	
Debra Friedkin	Texas	30.9	97.7	16.65 ± 1.075	14.69	16.33	15.51	Waters et al., 2018	Beccera-Valdivia & Higham 2020	
Hebior	Wisconsin	42.6	88.0	12.59 ± 0.05	13.98	15.62	14.80	Overstreet et al., 2003	Beccera-Valdivia & Higham 2020	
Page-Ladson	Florida	30.2	84.0	13.945 ± 0.05	14.45	14.71	14.58	Halligan et al., 2016	Beccera-Valdivia & Higham 2020	
Lindsay	Montana	47.2	105.1	12.395 ± 0.055	13.95	14.63	14.29	Hill et al., 1998	Beccera-Valdivia & Higham 2020	
Paisley Caves	Oregon	42.8	120.7	12.40 ± 0.06	13.78	14.76	14.27	Jenkins et al., 2014	Beccera-Valdivia & Higham 2020	
wally's Beach	Alberta Canada	49.4	114.1	11.53 ± 0.05	13.22	13.34	13.28	Waters et al., 2015	Beccera-Valdivia & Higham 2020	
Ronneville Estates Rockshelter	Nevada	33.0 40.3	101.9	12.05 ± 0.29	12.77	12.79	13.20	Goebel et al. 2007	Beccera-Valdivia & Higham 2020	
Buhl	Idaho	42.7	114.8	10.675 ± 0.095	12.42	12.74	12.58	Green et al., 1998	Beccera-Valdivia & Higham 2020	
Clovis Sites										
Sheriden Cave	Ohio	41.0	83.3	12.84 ± 0.10	13.40	14.85	14.12	Waters et al., 2009	Beccera-Valdivia & Higham 2020	
Aubrey	Texas	33.4	97.8	11.59 ± 0.09	13.25	14.75	14.00	Ferring 2001	Beccera-Valdivia & Higham 2020	
El Fin del Mundo	Mexico	29.7	111.8	11.55 ± 0.06	13.47	14.14	13.80	Sanchez et al., 2014	Beccera-Valdivia & Higham 2020	
Blackwater Draw	New Mexico	34.3	103.3	12.79 ± 0.16	12.68	13.44	13.06	Haynes et al., 1995	Beccera-Valdivia & Higham 2020	
Dent	Colorado	40.3	104.8	11.155 ± 0.05	12.42	13.55	12.98	Deviese et al., 2018	Beccera-Valdivia & Higham 2020	
Murray Springs	Arizona	32.2	110.3	11.19 ± 0.18	12.62	13.02	12.82	Haynes et al., 2007	Beccera-Valdivia & Higham 2020	
Anzick	Montana	46.0	110.7	10.915 ± 0.05	12.49	13.18	12.83	Beccera-Valdivia et al., 2018	Beccera-Valdivia & Higham 2020	
Shawnee-Minisink	Pennsylvania	41.0	75.1	11.02 + 0.03	12.71	12.91	12.81	McNett et al. 1977	Beccera-Valdivia & Higham 2020	
Colby	Wyoming	44.0	107.9	10.95 + 0.03	12.35	13.17	12.76	Frison et al., 1986	Beccera-Valdivia & Higham 2020	
Jake Bluff	Oklahoma	36.6	99.5	10.885 ± 0.035	12.65	12.78	12.71	Bement & Carter 2003	Beccera-Valdivia & Higham 2020	
Lange/Ferguson	South Dakota	43.3	102.1	11.11 ± 0.04	11.11	14.29	12.70	Hannus 2018	Beccera-Valdivia & Higham 2020	
Lehner	Arizona	31.4	110.1	12.0 ± 0.45	9.39	13.58	11.48	Haynes 1992	Beccera-Valdivia & Higham 2020	
Coastal Sites										
CA-SMI-261 Daisy Cave	Channel Islands	34.0	120.2	15.78 ± 0.12	18.53	18.81	18.67	Erlandson et al., 1996	McLaren et al., 2020	
EkTb-9	Triquet Island	51.8	128.3		13.70	14.00	13.85	Gavreau & McLaren 2017	McLaren et al., 2020	
Manis Matodon	Olympic Peninsula	48.0	123.1	11.96 ± 0.02	13.76	13.86	13.81	Waters et al., 2011	McLaren et al., 2020	
EJTa-4	Calvert Island	51.7	128.1	11.44 ± 0.025	13.26	13.32	13.29	McLaren et al., 2018	McLaren et al., 2020	
Each 91	Ouedra Island	43.0	124.3	11.60 ± 0.04	12.63	13.44	13.04	Endio et al. 2018a	McLaren et al., 2020	
CA-SRI-173 Arlington Springs	Channel Islands	34.0	120.1	11.58 + 0.05*	12.70	13.02	12.86	Johnson et al., 2002	McLaren et al., 2020	
EbSh-98	Quadra Island	50.2	125.2	10.94 ± 0.06	12.72	12.97	12.85	Fedje et al., 2018a	McLaren et al., 2020	
K1 cave	Haida Gwaii	52.9	132.5	10.96 ± 0.035	12.73	12.90	12.81	Fedje et al., 2011a	McLaren et al., 2020	
EbSh-1	Quadra Island	50.2	125.2	10.74 ± 0.07	12.57	12.74	12.66	Fedje et al., 2018a	McLaren et al., 2020	
Gaadu Din 1	Haida Gwaii	52.3	131.5	10.615 ± 0.03	12.55	12.60	12.58	Fedje et al., 2011a	McLaren et al., 2020	
Bear Creek	Puget Sound	47.6	120.1	10.49 ± 0.03	12.37	12.67	12.52	Kopperl 2016	McLaren et al., 2020	
Gaadu Din 2	Haida Gwaii	52.3	131.5	10.53 ± 0.02	12.43	12.55	12.49	Fedje et al., 2011	McLaren et al., 2020	
Indian Sands	Oregon Coast	42.1	124.2	10.43 ± 0.015	11.69	12.93	12.31	Davis 2008	McLaren et al., 2020	
PAIC-49 Richard's Ridge	Cedros Island Baia CA	49.4 28.0	122.5	10 75 + 0 03*	12.07	12.50	12.25	Des lauriers et al. 2017	McLaren et al. 2020	
DhRn-18	Stave Watershed	49.4	122.3	10.35 + 0.03	12.04	12.45	12.25	McLaren 2017	McLaren et al. 2020	
DhRo-11	Stave Watershed	49.4	122.3	10.00 2 0.00	12.00	12.38	12.19	McLaren et al., 2020	McLaren et al., 2020	
DhRo-16	Stave Watershed	49.4	122.3	10.29 ± 0.05	11.83	12.38	12.11	McLaren 2017	McLaren et al., 2020	
Shuka Kaa/On Your Knees Cave	Southeast Alaska	56.3	133.6	10.30 ± 0.05	11.83	12.46	12.08	Dixon et al., 2014		
CA-SRI-723	Channel Islands	33.6	120.1	10.94 ± 0.05*	11.93	12.17	12.05	Rick et al., 2013	McLaren et al., 2020	
CA-SMI-679	Channel Islands	34.0	120.2	10.8 ± 0.05	11.71	12.20	11.96	Erlandson et al., 2011	McLaren et al., 2020	
DhRn-29	Stave Watershed	49.4	122.3	10.37 ± 0.04	11.29	12.46	11.88	McLaren 2017	McLaren et al., 2020	
CA-SMI-678	Channel Islands	34.0	120.2	10.95 ± 0.05	11.40	12.20	11.80	Erlandson et al., 2011	McLaren et al., 2020	
CA-SKI-512	Channel Islands	34.0	120.1	10.2 ± 0.05 10.7 ± 0.04	11.41	12.00	11./1	Erlandson, Rick, et al., 2011 Erlandson, Rick, et al., 2011	McLaren et al., 2020	
CA-SMI-701	Channel Islands	34.0	120.1	10.7 ± 0.04	11.41	11.30	11.70	Friandson et al 2013	McLaren et al. 2020	
CA-SRI-997	Channel Islands	34.0	120.0	10 ± 0.03	11.77	11.74	11.51	Gill et al. 2021	metaren et al., 2020	
CA-SRI-706	Channel Islands	33.6	120.1	10.6 ± 0.07*	11.24	11.62	11.43	Rick et al., 2013	McLaren et al., 2020	
CA-SRI-725	Channel Islands	33.6	120.1	10.59 ± 0.05*	11.22	11.36	11.29	Rick et al., 2013	McLaren et al., 2020	
CA-SRI-708	Channel Islands	33.6	120.0	10.4 ± 0.05*	10.79	11.25	11.02	Rick et al., 2013	McLaren et al., 2020	

= requires marine reservoir correction
 Latitude/longitude values have been rounded to protect archaeological sites from vandalism or looting.

Table S1. Summary of archaeological sites discussed and plotted in Fig. 5. For additional site and reference information, see attached Dataset S1.

#	Core	Location	Lat (°)	Lon (°)	Elev/Depth (m)	Reference	Proxy	Deglacial resolution (yr)
1	SO202-18-6	Bering Sea	60.1	179.4	1105	Meheust et al., 2018	U ^{K'} 37	80
2	SO201-2-114KL	Western Bering Sea	59.3	167.0	-1376	Meyer et al., 2016	TEX ₈₆	150
3	SO201-2-114KL	Western Bering Sea	59.3	167.0	-1376	Max et al., 2012	U ^{K'} 37	170
4	EW0408-85JC	Gulf of Alaska	59.6	-144.2	-682	Praetorius et al., 2015	U ^{K'} 37	100
5	U1419	Gulf of Alaska	59.5	-144.1	-698	Romero et al., 2022	U ^{K'} 37	155
6	EW0408-87JC	Gulf of Alaska	58.8	-144.5	-3680	Praetorius et al., 2020	U ^{K'} 27	200
7	SO201-2-101KL	Bering Sea	58.8	170.7	-630	Riethdorf et al., 2013	Mg/Ca	160
8	EW0408-66JC	Gulf of Alaska	58.5	-137.2	-426	Praetorius et al., 2016	U ^{K'} 37	50
9	EW0408-26JC	Gulf of Alaska	57.6	-136.7	-1623	Praetorius et al., 2016	U ^{K'} 27	50
10	SO201-2-85KL	Bering Sea	57.5	170.4	-968	Riethdorf et al., 2013	Mg/Ca	225
11	SO201-2-85KL	Bering Sea	57.5	170.4	-968	Max et al., 2012	U ^{K'} 27	170
12	SO201-2-77KL	Bering Sea	56.3	170.7	-2135	Max et al., 2012	U ^{K'} 37	190
13	SO201-2-77KL	Bering Sea	56.3	170.7	-2135	Riethdorf et al., 2013	Mg/Ca	310
14	SO202-27-6	Gulf of Alaska	54.2	-149.6	-2919	Meheust et al., 2018	U ^{K'} 37	880*
15	SO201-2-12KL	Northwestern Pacific	54.0	162.4	-2145	Meyer et al., 2016	TEX	140
16	SO201-2-12KL	Northwestern Pacific	54.0	162.4	-2145	Riethdorf et al., 2013	Mg/Ca	80
17	SO201-2-12KL	Northwestern Pacific	54.0	162.4	-2145	Max et al., 2012	U ^{K'} 37	70
18	U1340	Bering Sea	53.4	179.5	-1294	Schlung et al., 2013	U ^{K'} 37	410*
19	XP07-C9	Okhotsk Sea	52.3	146.0	-1431	Harada et al., 2012	U ^{K'} 27	340
20	MR06-04-PC7	Okhotsk Sea	51.3	149.2	-1247	Seki et al., 2009	TEX.	1500*
21	SO202-07-6	Northwestern Pacific	51.3	167 7	-2340	Meheust et al. 2018	LI ^{K'}	870*
22	XP98-PC-2	Okhotsk Sea	50.4	148 3	-1258	Seki et al 2004	U ^K .	670*
23	XP08-PC-4	Okhotsk Sea	49.5	146.0	-664	Seki et al. 2004	U 37	640*
20	MR00K03-PC-04	Okhotsk Sea	49.5	153.0	-1821	Harada et al. 2004	U 37	310
25	1//20-11/-3	Okhotsk Sea	40.4	152.0	-1765	Max et al. 2012	U 37	230
20	1/20 114 3	Okhotsk Sea	40.4	152.0	-1765	Righthorf et al. 2012	0 37 Ma/Ca	110
20	MD02-2496	Vancouver margin	49.4	-127.0	-1703	Taylor et al. 2013	Mg/Ca	110
28	MD02-2496	Vancouver margin	49.0	-127.0	-1243	Taylor et al., 2014	Mg/Ca	80
29	JT96-09PC	Vancouver margin	48.9	-126.9	-920	Kienast & McKay 2001	U ^{K'}	90
30	MR00K03-PC-01	Northwestern Pacific	46.3	152.5	-2793	Harada et al., 2004	U ^{K'} 27	200
31	MR9702-St-8s	Okhotsk Sea	44.8	170.2	-1780	Harada et al., 2004	U ^{K'} 27	750*
32	MD01-2412	Okhotsk Sea	44.5	145.0	-1225	Harada et al., 2006	U ^{K'} 27	210
33	MR06-04-PC04	Okhotsk Sea	44.5	145.0	-1225	Harada et al. 2012	U ^{K'}	110
34	W8709A-8TC	Northeastern Pacific	42.2	-127.7	-3111	Prahl et al., 1995	U ^{K'}	1010*
35	GH02-1030	Northwestern Pacific	42.2	144.2	-1212	Inagaki et al., 2009	U ^{K'}	390
						Barron et al., 2003, Herbert et al.	0 3/	
36	ODP 1019	California margin	41.7	-124.9	-980	2003	U ^{K'} 37	70
37	ODP 1020	California margin	41.0	-126.4	-3042	Herbert et al., 2001	U ^{K'} 37	640*
38	PC-6	Japan margin	40.4	143.5	-2215	Minoshima et al., 2007	U ^{K'} 37	250
39	MR98-05-St5	Central Pacific	40.0	165.0	-5498	Harada et al., 2004	U ^{K'} 27	1300*
40	MR98-05-St6	Central Pacific	37.5	162.7	-3130	Harada et al., 2004	U ^{K'}	2800*
41	MD01-2421	Northwestern Pacific	36.0	141.8	-2224	Yamamoto et al., 2005	U ^{K'}	150
42	ODP 1017	California margin	34.5	-121 1	-955	Pak et al 2012	Mg/Ca	160
43	ODP 1017	California margin	34.5	-121.1	-955	Seki et al., 2002	U ^{K'}	250
44	ODP 1016	California margin	34.5	-122.3	-3834	Yamamoto et al., 2007	U ^{K'}	1120*
45	ODP 893	Santa Barbara Basin	34.3	-120.0	-575	Hendy 2010	foram assemblages	60
46	ODP 1014	California margin	32.8	-118.9	-1165	Yamamoto et al., 2007	U ^{K'} 37	1470*
47	KT92-17 St. 14	Japan margin	32.6	138.6	-3252	Sawada and Handa, 1998	U ^{K'} 37	440*
48	ODP 1012	California margin	32.3	-118.4	-1783	Herbert et al., 2001	U ^{K'} 37	610*
49	KY07-04-01	Japan margin	31.6	128.9	-758	Kubota et al., 2010	Mg/Ca	100
50	MD98-2195	East China Sea	31.6	129.0	-746	ljiri et al., 2005	U ^{K'} 37	160
51	MD02-2515	Guaymas Basin	27.5	-112.1	-881	McClymont et al., 2012	U ^{K'} 37	80
52	MD02-2515	Guaymas Basin	27.5	-112.1	-881	McClymont et al., 2012	TEX ₈₆	85
53	LaPaz 21P	Baja Penisula	23.0	-109.5	-624	Herbert et al., 2001	U ^{K'} 37	1320*
54	17940	South China Sea	20.1	117.4	-1968	Pelejero et al., 1999	U ^{K'} 37	150
55	ODP 1144	South China Sea	20.1	117.6	-2037	Wei et al, 2007	Mg/Ca	300
56	GIK17927-2	South China Sea	17.3	119.5	-2804	Sadatzki et al., 2016	U ^K '37	150
57	GIK17954-2	South China Sea	14.8	111.5	-1520	Pelejero et al., 1999	U ^{K'} 37	1070*
58	MD97-2141	Sulu Sea	8.8	121.3	-3633	Rosenthal et al., 2003	Mg/Ča	70
59	MD02-2529	Eastern equatorial Pacific	8.2	-84.1	-1619	Leduc et al., 2007	U ^{K'} 37	260
60	ME0005A-43JC	Eastern equatorial Pacific	7.9	-83.6	-1368	Benway et al., 2006	Mg/Ca	220
61	MD01-2390	South China Sea	6.6	113.4	-1545	Steinke et al., 2008	U ^{K'} 37	200
62	MD01-2390	South China Sea	6.6	113.4	-1545	Steinke et al., 2008	Mg/Ca	200
63	MD98-2181	West Pacific warm pool	6.3	125.8	-2114	Stott et al., 2007	Mg/Ca	60
64	TR163-22	Eastern equatorial Pacific	0.5	-92.4	-2830	Lea et al., 2006	Mg/Ca	250
65	ME0005A-24JC		0.0	-86.5	-2941	Kienast et al., 2006	U ^r 37	150

* denotes datasets with deglacial resolution that exceeds the 400-yr cutoff; used these datasets for LGM-Holocene estimates only

Table S2. Summary of marine sediment cores used in the North Pacific SST compilation (Fig. 4). For additional information on SST data and age models, see Dataset S4.

Legends for Data Files:

Dataset S1 (Microsoft Excel format): Metadata, references, and radiocarbon data for archaeological sites listed in Table S1.

Dataset S2 (Microsoft Excel format): Age models recalibrated using the Marine20 calibration for the following cores: EW0408-66JC, EW0408-87JC, EW0408-26JC, JT96-09JPC, ODP1019, with variable marine reservoir corrections that generally follow those applied in (37).

Dataset S3 (Microsoft Excel format): New alkenone-derived %C_{37:4} records from the following marine sediment cores in the Gulf of Alaska: EW0408-85JC, EW0408-66JC, EW0408-87JC, EW0408-26JC.

Dataset S4 (Microsoft Excel format): Metadata, references, and age model information for North Pacific marine sediment cores used in the SST compilation (Fig. 4, Table S2).

Dataset S5 (Microsoft Excel format): Average Northeast Pacific SST stacks.

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