



# Toward a typology of river functioning: a comprehensive study of POM composition at multi-

- <sup>3</sup> rivers scale
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- 18 study.





#### 19 Abstract

20 In riverine systems, particulate organic matter (POM) originates from various sources, each 21 having its proper dynamics related to production, decomposition, transport and burial. There is 22 a significant amount of spatiotemporal heterogeneity in the POM pool. The current study, based 23 on C and N elemental and isotopic ratios, applies Bayesian mixing models associated with 24 statistical multivariate analyses to 1) quantify and examine relationships between POM 25 composition and environmental forcings, and 2) draw a typology of river functioning based on 26 POM composition and its seasonal dynamics. Twenty-three rivers of temperate climate 27 accounting for a large diversity of environmental conditions were sampled fortnightly to 28 monthly for one to seven years at the River-Estuary Interface (REI). Phytoplankton and labile 29 terrestrial material were present in all rivers, contrary to sewage and refractory terrestrial 30 material that were present in only a few. At the twenty-three studied rivers scale, POM sources 31 are strongly related to watershed characteristics, phytoplankton being associated with 32 agricultural surfaces and labile terrestrial material to soil organic carbon content and erosion 33 rate. Overall, seasonal variations of phytoplankton, labile and refractory terrestrial material 34 were mainly related to drivers of phytoplankton growth, river flow, and sediment resuspension, 35 respectively. A statistical regionalization defined four river types: (1) systems whose POM is 36 dominated by labile terrestrial material all year long; (2) systems whose POM is composed of 37 labile and refractory terrestrial material, in addition to phytoplankton, with variable seasonality 38 according to rivers; systems whose POM is composed of phytoplankton and labile terrestrial 39 material (3) without and (4) with pronounced seasonality.

This work offers a comprehensive understanding of POM composition, dynamics and drivers
at the REI in temperate climates, complementing similar work dedicated to coastal systems.
Future work dedicated to estuaries is called to get a comprehensive understanding of POM
composition, dynamics and drivers along the Land-Ocean Aquatic Continuum.

This study examines particulate organic matter (POM) composition and dynamics in 23 temperate rivers. Carbon and nitrogen isotope analysis revealed four river types based on dominant POM sources (phytoplankton, terrestrial material). Watershed characteristics influence POM composition while seasonal variations in river flow and sediment resuspension drive POM dynamics. This study improves the understanding of river systems and calls for further studies exploring downstream estuarine functioning.





#### 50 1. Introduction

51 The River-Estuary Interface (REI) is a crucial biogeochemical interface for understanding the 52 transition between continental and coastal systems, beginning at estuaries, because of its key 53 location within the Land-Ocean Aquatic Continuum (LOAC) (Bate et al., 2002). Indeed, rivers 54 then estuaries are important filters for matters received from land, transporting and transforming 55 organic matter and nutrients along their courses (Bouwman et al., 2013; Dürr et al., 2011; 56 Middelburg and Herman, 2007). These processes are fundamental in understanding global 57 biogeochemical cycles (Regnier et al., 2013), as these matters directly fuel coastal ocean trophic 58 networks (Dagg et al., 2004). However, in a Human-impacted world, anthropogenic activities 59 and disturbances can modify natural matter fluxes. For example, damming rivers directly 60 impacts nutrient flows (Wang et al., 2022) and sediment transportation (Kang et al., 2021). 61 Indirectly, land use in river basins can lead to changes in the river matter quality (Lambert et 62 al., 2017).

63 In aquatic systems, particulate organic matter (POM), i.e., non-mineral particles, is composed 64 of different sources that originate from different compartments: phytoplankton, macrophytes 65 from the aquatic systems as well as soil particles and plant litter from terrestrial compartments 66 and even treated and untreated anthropogenic organic matter (Ke et al., 2019; Sun et al., 2021; Zhang et al., 2021). Depending on its composition, POM exhibits different levels of lability, 67 68 i.e., different levels of biogeochemical reactivity and bioavailability. For instance, phytoplankton is usually considered mainly labile and thus highly biogeochemically reactive 69 70 and bioavailable for primary consumers, while terrestrial POM is usually considered mainly 71 refractory and thus lightly biogeochemically reactive and poorly bioavailable for the food webs 72 (Brett et al., 2017; David et al., 2005; Etcheber et al., 2007). In other words, the determination 73 and quantification of POM composition (i.e., the relative proportion of each source composing 74 the POM) allow a better understanding of biogeochemical cycles and trophic ecology in aquatic 75 systems (e.g., Grunicke et al., 2023; Minaudo et al., 2016). Nevertheless, POM composition 76 and concentration are not only involved in biogeochemical and biological processes (e.g., 77 primary production, remineralization, feeding) but undergo other processes inside and at the 78 interface of the aquatic compartment (Canuel and Hardison, 2016). River hydrodynamics is one 79 of the main drivers of POM composition and concentration, leading to great variabilities in 80 terrestrial material quality and quantity (Dalzell et al., 2007; Lebreton et al., 2016; Marshall et 81 al., 2021), possibly leading to changes in source origins (Arellano et al., 2019; Barros et al., 82 2010). Also, changes in anthropic pressures can change POM composition and concentration 83 and their seasonal variations, like a decrease in nutrient load (Minaudo et al., 2015).

This dependency of POM composition and concentration on physical, biogeochemical and biological processes and their responses to environmental conditions and characteristics (Bonin et al., 2019; Falkowski et al., 1998; Field et al., 1998; Galeron et al., 2017; Goñi et al., 2009;





Lebreton et al., 2016) may lead to distinguishing different typology of rivers, i.e., the likeliness of rivers to carry preferential sources. For instance, highly turbid systems are more likely to carry refractory materials (Savoye et al., 2012), while eutrophicated rivers carry high biomass of phytoplankton (Hounshell et al., 2022; Minaudo et al., 2015) and contrasted processes can lead to a mixture between different detrital sources, as soil matter vs. fresh terrestrial plants (Ogrinc et al., 2008). However, to date, no study clearly determined typologies of rivers based on POM composition and its seasonal variability.

94 To distinguish POM sources and quantify their contribution to POM composition, different 95 tools such as elemental and isotopic ratios, pigments or specific compounds like fatty acids or 96 alkanes can be used (e.g., Chevalier et al., 2015; Liénart et al., 2020, 2017; Savoye et al., 2012). 97 Elemental and isotopic ratios are usually considered robust and allow the quantification of POM 98 composition in this kind of study (e.g., Liénart et al., 2016; Onstad et al., 2000; Wang et al., 99 2021). Indeed, they usually allow the discrimination of, e.g., riverine phytoplankton, terrestrial POM and wastewater POM (Ke et al., 2019) and they can be used for running mixing models 100 101 that quantify the proportion of the different sources into a POM mixture (Parnell et al., 2013). 102 However, studies using mixing models for quantifying POM composition in river systems are 103 still scarce (e.g., Ferchiche et al., 2025, 2024; Kelso and Baker, 2022, 2020; Zhang et al., 2021). 104 Within the scope of better understanding the role of the LOAC in modifying matter fluxes and 105 quality, the present study gathered published data and results from 23 rivers at the river-estuary 106 interface with the aim of 1) quantifying the POM composition each river, 2) describing the 107 seasonal variations of this composition, 3) determining the drivers of the seasonal variability

108 within each river and the spatial variability among the 23 rivers, and then 4) determining a 109 typology of rivers according to their POM composition and dynamics. This study is the first to 110 precisely quantify POM composition in numerous and various temperate river systems and 111 classify river types according to POM composition and dynamics.

#### 112 2. Material and methods

113 Twenty-three temperate rivers were studied at their river-estuary interface (i.e., right upstream of the tidal influence). All the data come from published studies or national open databases. To 114 115 minimize the heterogeneity of the datasets in terms of sampling strategy, we have considered for this study the datasets only when 1) C/N ratio along with isotopic ratio of carbon and/or 116 117 nitrogen were available, 2) particulate matter characteristics like, suspended particulate matter 118 (SPM), particulate organic carbon (POC), particulate nitrogen (PN), chlorophyll a (chl a) were 119 also available, 3) datasets exhibited at least a monthly temporal resolution for one full year. 120 When needed, published datasets were completed and harmonized thanks to national databases.

121 **2.1.Study sites** 





- 122 The studied rivers and associated watersheds are all located in France (except the upper basin 123 of the Rhône River) and distributed in all regions of the mainland. Three, fifteen and five of these rivers flow into the English Channel, the Atlantic Ocean and the Mediterranean Sea (Fig. 124 125 1). They encompass large gradients of environmental characteristics (Tab. 1). For instance, the 126 Loire River is one of the largest in Europe (length: 1006 km; watershed: 117,356 km<sup>2</sup>), while 127 the littlest studied river is a very small stream of the Arcachon lagoon (length: 3 km; watershed: 128 18 km<sup>2</sup>). They encompass large gradients of river flow (annual mean: 0.3 m<sup>3</sup>/s - 1572 m<sup>3</sup>/s), 129 turbidity (SPM annual mean: 2.7 mg/l - 40.9 mg/l) and trophic status (from oligotrophic to 130 eutrophic rivers; chl a annual mean:  $0.4 \mu g/l - 57.1 \mu g/l$ ). At last, they undergo a gradient of anthropic pressures as illustrated by the proportion of artificial surfaces (0.1 % - 5.6 %) and 131 agricultural areas (0 % - 86 %) in the watersheds (Fig. 1). 132
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135 <u>Figure 1</u> Studied rivers (thick blue lines), sampling locations (black stars) and watersheds (thin





- 137 Rance; 4: Elorn; 5: Aulne; 6: Loire; 7: Sèvre niortaise; 8: Charente; 9: Seudre; 10: Canal du
- 138 Porge; 11: Cirès; 12: Milieu; 13: Lanton; 14: Renet; 15: Tagon; 16: Leyre; 17: Canal des
- 139 Landes; 18: Adour; 19: Têt; 20: Aude; 21: Orb; 22: Hérault; 23: Rhône.





41	40
suspended particulate matter (SPM) and chlorophyll $a$ (chl.	<u>Table 1</u> Overview of river samplings and characteristics. V
1 a). Id: identification number; Number: number of sampling dates	Values are given as annual mean over the study period for river
. River types were	flow, temperature,

42 defined within the scope of the present study (see section 3.4).

143																										
	Rhône		Aude	Orb	Hérault	Têt	Adour	Landes	Leyre	•	Tagon	Milieu	Lanton	Renet	Cirès	Porge	Seudre	Charente	Sèvre	Loire	Aulne	Elorn	Rance	Orne	Seine	River
	23		22	21	20	19	18	17	10		15	14	13	12	Ξ	10	9	8	7	6	S	4	ω	2	-	Id
	Π		Π	Π	п	Π	Π	Ш	-	•	Ι	Ι	Ι	I	г	Ш	г	Π	N	N	N	N	N	Π	N	River type
	12/2003 to $01/2011$		01/2006 to 05/2010	01/2006 to 05/2010	01/2006 to 05/2010	01/2006 to 05/2010	04/2013 to 06/2014 and 05/2017 to	02/2008 to 02/2009	and $02/2014$ to $02/2015$	01/2008 to 03/2010	02/2008 to 02/2009	01/2008 to 02/2009	03/2014 to 09/2015	03/2014 to 03/2015	03/2014 to 03/2015	10/2009 to 07/2012	01/2014 to 06/2015	01/2014 to 06/2015	06/2014 to 05/2015	06/2014 to 06/2015	06/2014 to 06/2015	Sampled period				
	monthly		monthly	monthly	monthly	monthly	monthly	monthly	monthly	bi-monthly or	bi-monthly	monthly	monthly	bi-monthly	monthly	monthly	monthly	monthly	monthly	bi-monthly	monthly	monthly	monthly	monthly	monthly	Sampling Periodicity
	105		52	52	52	52	24	12	90	5	26	13	13	23	13	14	15	13	13	67	17	17	12	13	13	Num ber
	43.6787		43.2442	43.2850	43.3594	42.7137	43.4988	44.6169	44.0203		44.6590	44.6973	44.7002	44.7144	44.7598	44.7898	45.6740	45.8680	46.3153	47.3920	48.2127	48.4505	48.4916	49.1797	49.3067	Latitud e
	4.6212		3.1527	3.2813	3.4354	2.9935	-1.2949	-1.1091	-0.9961		-0.9891	-1.0225	-1.0244	-1.0441	-1.1107	-1.1612	-0.9331	-0.7131	-1.0039	-0.8604	-4.0944	-4.2483	-2.0014	-0.3491	1.2425	Longitu de
	812		223	136	148	115	308	14	011		10	7	15	ω	12	57	89	381	158	1006	144	56	103	169	774	River length (km)
	95590		5327	1585	2582	1369	16912	117	1/00	1000	30	21	36	18	45	222	855	9855	3650	117356	1875	385	1195	2932	79000	Catchment area (km²)
	1572		40	23	53	23	516	0.49	1/	i	0.64	0.58	0.26	0.56	0.58	3.48	1.81	89	3.72	630	30	6	1.37	16	496	River flow (m3/s)
	15.9		14.2	15.7	16.0	15.7	14.0	14.1	13.0	5	12.6	12.7	12.5	12.9	12.2	13.3	14.3	15.1	15.7	14.1	14.4	12.3	15.1	14.5	15.0	Water temperat ure (°c)
	41		31	8	7	8	48	з	Ξ	:	13	7	Ξ	10	S	12	17	13	13	19	7	16	21	Ξ	21	SPM (mg/l)
	1.9		NA	NA	NA	NA	2.4	1.1	0.9	5	1.3	0.4	1.2	0.6	0.4	5.0	0.5	1.3	3.8	18.7	3.3	3.0	57.1	1.8	2.8	Chl <i>a</i> (µg/l)
Higueras et al., 2014	/ Cathalot et al., 2013 /	Harmelin-Vivien et al., 2010	Higueras et al., 2014	Liénart et al., 2016 / Deborde, 2019	Polsenaere et al., 2013	Polsenaere et al., 2013 / Liénart et al., 2017, 2018	Dubois et al., 2012 /	Polsenaere et al., 2013	Liénart et al., 2017, 2018	Liénart et al., 2017, 2018	Liénart et al., 2017, 2018	Ferchiche et al., 2024	Liénart et al., 2017, 2018	References												





#### 144 **2.2.Data origin**

Regarding the core parameters (C/N ratio,  $\delta^{13}$ C,  $\delta^{15}$ N, water temperature, SPM, POC, PN, chl 145 146 a), most of the data sets come from published studies (Canton et al., 2012; Cathalot et al., 2013; 147 Dubois et al., 2012; Ferchiche et al., 2024; Harmelin-Vivien et al., 2010; Higueras et al., 2014; 148 Liénart et al., 2016, 2017, 2018; Polsenaere et al., 2013), while most of additional parameters come from national databases (Tab. A1). When not available in the cited studies, concentrations 149 150 of SPM, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup> and PO<sub>4</sub><sup>3-</sup>, pH and water temperature were retrieved from the Naïades 151 database (https://naiades.eaufrance.fr/, consulted the 07/10/2023). Note that these parameters 152 were not necessarily measured or sampled exactly at the same location or date for Naïades than 153 in the cited studies. In that case, the location was chosen as close as possible to the study 154 location and data values were time-interpolated to match the study date. Meteorological 155 variables (air temperature, zonal and meridional wind, irradiance) come from Météo France, 156 the French meteorological service. Wind data was received originally as direction and speed. 157 To remove the angular bias, they were combined using scalar products to get zonal and meridional wind speeds, which range between minus and plus infinity (see Lheureux et al., 158 159 2022, for more details). River flows were retrieved from the Banque Hydro database 160 (https://www.hydro.eaufrance.fr/, consulted the 07/10/2023) or from Polsenaere et al. (2013) 161 for the small streams. Land use proportions originate from the national Corine Land Cover 162 database (https://www.statistiques.developpement-durable.gouv.fr/corine-land-cover-0, 163 consulted the 10/01/2024). Soil organic carbon data originate from the SoilTrEC database 164 (https://esdac.jrc.ec.europa.eu/content/predicted-distribution-soc-content-europe-based-lucasbiosoil-and-czo-context-eu-funded-1, consulted the 10/01/2024). Net erosion soil data 165 originates from the WaTEM/SEDEM database (https://esdac.jrc.ec.europa.eu/content/estimate-166 167 net-erosion-and-sediment-transport-using-watemsedem-european-union, consulted the 10/01/2024). 168

169 It should be noted that a complete study was already dedicated to the Loire River and reported 170 as a companion article (Ferchiche et al., 2024). Consequently, the results are not reported in the 171 present study but are used for multi-system comparisons (Fig. 5 and 7, and corresponding text).

#### 172 **2.3.Determination of sources signatures**

To run mixing models for quantifying POM composition, it is previously needed to 1) determine sources of POM, and 2) associate elemental and isotopic signatures to these sources. In riverine systems, phytoplankton and terrestrial POM are the main sources that are usually considered as fueling the POM (e.g., Ferchiche et al., 2024; Pradhan et al., 2016; Sarma et al., 2014). Nevertheless, sewage POM may also contribute (Higueras et al., 2014). Consequently, phytoplankton, terrestrial POM and sewage POM were considered as potential sources in this study.





180 Phytoplankton cannot be easily picked up from bulk particles to measure its elemental and isotopic ratios. Therefore, the method developed and used by Savoye et al. (2012), Liénart et 181 al. (2017) and Ferchiche et al. (2024) was applied here. It consists of determining the elemental 182 and isotopic ratios from a subset of the bulk dataset. Briefly, phytoplankton-dominated POM is 183 184 characterized by a low POC/chl a ratio ( $\leq 200$  or even  $\leq 100$  g/g; Savoye et al., 2003 and 185 references therein). Thus, elemental and isotopic ratios of samples exhibiting a low POC/chl a 186 ratio can be considered as good estimates of phytoplankton elemental and isotopic ratios. When 187 the POC/chl a ratio is not available, samples exhibiting a high PN/SPM ratio can be used. 188 Additional constraints may be used to minimize potential overlap between phytoplankton and terrestrial elemental and isotopic signatures. Phytoplankton elemental and especially isotopic 189 190 ratios may deeply vary over time and space depending on primary production intensity and 191 potential limiting factors, nutrient origin, etc. (e.g., Miller et al., 2013; Savoye et al., 2003). 192 When existing, this variability has to be taken into account to avoid using elemental and isotopic 193 signatures that are not valid at the time or location of the sampling. This could be performed by 194 using regressions between elemental and/or isotopic ratios and environmental variables (see 195 Ferchiche et al., 2024; Liénart et al., 2017; Savoye et al., 2012). At last, when no samples exhibit 196 a low POC/chl a ratio, samples exhibiting the lowest (even if high) POC/chl a ratios can be 197 used but the data should be firstly corrected from the contribution of the terrestrial POM using 198 Equations 1-3.

$$199 \quad \delta^{13}C_{sample} = ([POC]_{phytoplankton} \times \delta^{13}C_{phytoplankton} + [POC]_{terrestrial} \times \delta^{13}C_{terrestrial}) / [POC]_{sample}$$

$$(eq. 1)$$

$$201 \quad [POC]_{phytoplankton} = [chl a]_{sample} \times (POC/chl a)_{mean}$$

$$(eq. 2)$$

$$202 \quad [POC]_{terrestrial} = [POC]_{sample} - [POC]_{phytoplankton}$$

$$(eq. 3)$$

203 where (POC/chl *a*)<sub>mean</sub> is the mean POC/chl *a* ratio of the samples used to determine 204 phytoplankton signatures. Similar equations are used for the N/C ratio,  $\delta^{15}$ N and C/N ratio but 205 using PN instead of POC for  $\delta^{15}$ N and C/N ratio.

206 Elemental and isotopic signatures of terrestrial POM can be estimated by directly measuring 207 elemental and isotopic ratios in terrestrial materials like soil and vascular plants (e.g., Sarma et 208 al., 2014). However, this does not take into account the reworking of this material within the 209 river system which can affect these signatures (Hou et al., 2021). Thus, similarly to phytoplankton, elemental and isotopic signatures of terrestrial POM can be estimated using 210 211 subsets of bulk data, following the approach of Savoye et al. (2012), Liénart et al. (2017) and 212 Ferchiche et al. (2025, 2024). Terrestrial POM is usually characterized by high POC/chl a and 213 C/N ratios and low POC/SPM ratios (Etcheber et al., 2007; Savoye et al., 2003 and references 214 therein). However, during its decay in aquatic systems, terrestrial POM is colonized by bacteria 215 (low C/N ratio) resulting in a consortium terrestrial POM + bacteria of lower C/N ratio than the 216 original terrestrial POM (Etcheber et al., 2007; Savoye et al., 2012). Finally, one can





217 discriminate two kinds of terrestrial POM: refractory terrestrial POM characterized by high POC/chl a and C/N ratios and very low POC/SPM ratio, and quite labile terrestrial POM 218 characterized by high POC/chl a ratio, intermediate C/N ratios and low POC/SPM ratio 219 220 (Etcheber et al., 2007; Savoye et al., 2012). Thus, subsets of high POC/chl a ratio can be selected to determine the elemental and isotopic signatures of terrestrial POM. The C/N ratio 221 222 can be used to discriminate labile from refractory terrestrial POM. When no samples exhibit a 223 high POC/chl a ratio, samples exhibiting the highest (even if quite low) POC/chl a ratio can be 224 used but the data should be firstly corrected from the contribution of the phytoplankton POM 225 using Equations 1-3.

Elemental and isotopic ratios of riverine POM can exhibit a departure from a simple phytoplankton-terrestrial POM mixing. In the present study, this was the case in only two rivers. For the Têt River, elemental and isotopic signature of anthropogenic POM was available in Higueras et al. (2014). It consisted of analyses of POM sampled in the wastewater treatment plant (WWTP) the closest to the sampling site. For the Orb River, the signatures were estimated using the sample exhibiting the lowest  $\delta^{15}$ N, typical of anthropogenic POM (Ke et al., 2019).

The estimation of POM-source signatures was performed independently for each river, except for some of the tributaries of the Arcachon Lagoon (rivers 11 to 15), where data sets were gathered, thanks to very similar characteristics (same  $\delta$ 13C of dissolved inorganic carbon; Polsenaere et al., 2013), to get a larger subset of data for estimating elemental and isotopic signatures more accurately. All criteria used for defining the above-described subsets are reported in Table 2.

238 Table 2 Elemental and isotopic signatures of POM sources and criteria used to choose the data 239 subset to determine them. When the signature did not vary over time, average  $\pm$  standard 240 deviation are reported. When the signature did vary over time, minimum and maximum values, 241 standard deviations as well as equations are reported. The types of mixing models performed for each river are also indicated (carbon mixing models were performed using  $\delta^{13}C$  and N/C 242 243 ratio or only  $\delta^{13}$ C; nitrogen mixing models were performed using  $\delta^{15}$ N and C/N ratio; mixed 244 mixing models were performed using  $\delta^{13}$ C,  $\delta^{15}$ N and N/C ratio). POC% (or PN%) = Particulate 245 Organic Carbon (or Particulate Nitrogen) to Suspended Particulate Matter ratio (%); C/N =246 POC/PN ratio (mol/mol); chla = chlorophyll a (µg/l); phaeo = phaeopigments (µg/l); conduc = conductivity ( $\mu$ S); temp = water temperature (°C); Q7 = mean of past seven days river flow; 247

248  $NO_3^- = nitrate (mg(NO_3^-)/l).$ 





		Source	discriminants		Mo	del perforn	ned	L	abile terre	strial mat	ter	Refi	actory ter	restrial n	atter
River	Labile terrestrial matter	Refractory terrestrial matter	Phytoplankton	WWTP's POM	Carbon	Nitrogen	Mixed	$\delta^{13}C$	$\delta^{15}\!N$	C/N	N/C	$\delta^{13}C$	$\delta^{15}N$	C/N	N/C
Seine	C/N > 10		POC/chla < 200		х	х		-28.5 ± 0.3	6.6 ± 0.9	$\begin{array}{c} 10.6 \\ \pm \ 0.3 \end{array}$	$\begin{array}{c} 0.093 \\ \pm \ 0.002 \end{array}$				
Orne	C/N > 11		POC/chla < 500		х	х		-28.4 ± 0.3	5.8 ± 1.0	12.4 ± 0.4	$\begin{array}{c} 0.082 \\ \pm \ 0.003 \end{array}$				
Rance	POC/chla > 200 and chla < 10		POC/chla < 150		х	х		-26.8 ± 0.2	6.1 ± 0.7	$\begin{array}{c} 8.8 \\ \pm \ 0.4 \end{array}$	$\begin{array}{c} 0.113 \\ \pm \ 0.007 \end{array}$				
Elorn	C/N > 12		POC/chla < 200		х	х		-28.4 ± 0.7	5.8 ± 0.9	$13.0 \pm 0.8$	0.077 ± 0.005				
Aulne	C/N > 11		POC/chla < 200 and C/N < 9		х	х		-28.9 ± 0.8	$\begin{array}{c} 5.8 \\ \pm \ 0.8 \end{array}$	12.1 ± 1.1	$\begin{array}{c} 0.08 \\ \pm \ 0.008 \end{array}$				
Loire	POC/chla > 500		POC/chla < 200		х	х		-28.1 ± 0.1	$\begin{array}{c} 5.9 \\ \pm \ 0.3 \end{array}$	$\begin{array}{c} 10.3 \\ \pm \ 0.2 \end{array}$	$\begin{array}{c} 0.097 \\ \pm \ 0.002 \end{array}$				
Sèvre Niortaise	C/N > 14		POC/chla < 300		х			-28.0 ± 0.4			$0.057 \pm 0.040$				
Charente	C/N > 12		POC/chla < 300		х	х		-29.0 ± 0.4	4.7 ± 0.2	14.5 ± 0.5	0.069 ± 0.002				
Seudre	POC/chla > 2000 and C/N > 12		POC/chla < 1000		х			-28.5 ± 0.1							
Porge	C/N > 15		δ <sup>13</sup> C : POC/chla < 100 ; N/C : mean of Cirès to Landes		х			-26.5 ± 1.1			$\begin{array}{c} 0.050 \\ \pm \ 0.007 \end{array}$				
Cirès / Renet / Milieu / Lanton / Tagon	C/N > 15 and chla < 1		POC/chla < 1000 and POC% > 10		х			-28.5 ± 0.5			$\begin{array}{c} 0.053 \\ \pm \ 0.013 \end{array}$				
Leyre	C/N > 15 and chla < 1		$\begin{array}{l} POC/Chla < 1000.\\ \delta^{13}C < 28.59 \text{ and}\\ POC\% > 10 \end{array}$		х			-28.3 ± 0.5			$\begin{array}{c} 0.06 \\ \pm \ 0.005 \end{array}$				
Landes	C/N > 12		$POC/Chla \le 600.$ $\delta^{13}C \le -29.1$		х			-29.1 ± 0.4			$\begin{array}{c} 0.075 \\ \pm \ 0.002 \end{array}$				
Adour	POC/chla > 3000		POC/chla < 200		х			-26.0 ± 0.9			$\begin{array}{c} 0.099 \\ \pm \ 0.008 \end{array}$				
Têt	C/N > 11.5	POC% < 4.25	$\begin{array}{l} PN\% > 2. \ \delta^{13}C < \\ 26 \ and \ \delta 15N > 5 \end{array}$	Measured			х	-26.0 ± 0.2	$\begin{array}{c} 3.7 \\ \pm \ 0.6 \end{array}$	$\begin{array}{c} 12.2 \\ \pm \ 0.5 \end{array}$	$\begin{array}{c} 0.082 \\ \pm \ 0.002 \end{array}$	-26.0 ± 0.6	6.7 ± 1.4	5.8 ± 1.4	$\begin{array}{c} 0.180 \\ \pm \ 0.045 \end{array}$
Aude	C/N > 12	Q7 > 70	PN% > 1 or 2 and C/N < 6				х	-28.1 ± 0.6	6.3 ± 0.1	15.3 ±1.6	$\begin{array}{c} 0.066 \\ \pm \ 0.007 \end{array}$	-28.0 ± 0.7	4.7 ± 0.4	7.3 ± 1.0	$\begin{array}{c} 0.139 \\ \pm \ 0.018 \end{array}$
Orb	C/N > 10		$PN\% > 2. \delta^{15}N > 4.06$	Lower δ <sup>15</sup> N			х	-27.1 ± 0.4	3.7 ± 0.4	$\begin{array}{c} 10.5 \\ \pm \ 0.3 \end{array}$	$\begin{array}{c} 0.095 \\ \pm \ 0.350 \end{array}$				
Hérault	C/N > 12	Q > 45	PN% > 2				Х	-27.7 ± 0.2	6.1 ± 0.7	13.7 ±1.2	$\begin{array}{c} 0.073 \\ \pm \ 0.007 \end{array}$	-27.8 ± 0.4	4.7 ± 0.6	8.2 ± 1.5	$\begin{array}{c} 0.124 \\ \pm \ 0.019 \end{array}$
Rhône	C/N > 12	POC% < 1.25	$C/N \le 6.68$ and $\delta^{15}N \ge 3.92$				х	-26.4 ± 1.3	5.2 ± 1.0	17.0 ± 3.2	$\begin{array}{c} 0.061 \\ \pm \ 0.012 \end{array}$	-25.9 ± 0.4	$\begin{array}{c} 3.1 \\ \pm \ 0.8 \end{array}$	8.8 ± 3.1	$\begin{array}{c} 0.119 \\ \pm \ 0.032 \end{array}$





#### 251 <u>Table 2</u> (continued)

			Pl	nytoplankton					WWTI	P's POM	
River	δ <sup>13</sup> C	C + equations	$\delta^{15}N$	+ equations	C/N	N/C +	equations	$\delta^{13}C$	$\delta^{15}N$	C/N	N/C
Seine	-32.8 ± 1.1		8.4 ± 1.7		7.4 ± 0.7	0.136 ± 0.012					
Orne	-31.4 ± 0.8		$\begin{array}{c} 4.3 \\ \pm \ 0.8 \end{array}$		6.6 ± 1.3	$\begin{array}{c} 0.141 \\ \pm \ 0.010 \end{array}$					
Rance	[-31.4;-25;6] ± 1.7	5.7x10 <sup>-4</sup> ×[chla+phaeo] <sup>2</sup> - 0.04×[chla+phaeo]-30.6	$\substack{[4.7;11.4]\\\pm 0.7}$	-0.28×[NO3 <sup>-</sup> ]+12.7	$\begin{array}{c} 6.2 \\ \pm \ 0.4 \end{array}$	$\begin{array}{c} 0.161 \\ \pm \ 0.010 \end{array}$					
Elorn	-27.4 ± 0.3		6.9 ± 0.5		$\begin{array}{c} 10.0 \\ \pm \ 0.9 \end{array}$	0.101 ± 0.007					
Aulne	-28.1 ± 0.2		8.6 ± 0.2		$\begin{array}{c} 8.2 \\ \pm \ 0.2 \end{array}$	$0.122 \pm 0.003$					
Loire	[-30.6;-25.0] ± 0.9	5x10 <sup>-4</sup> ×[chla+phaeo] <sup>2</sup> - 0.02[chla+phaeo]- 0.39[chla/phaeo]-27.9	[3.0;10.4] ± 1.2	4.2x10 <sup>-4</sup> [chla] <sup>2</sup> - 0.08[chla]+8.2	$\begin{array}{c} 7.2 \\ \pm \ 0.6 \end{array}$	$\begin{array}{c} 0.140 \\ \pm \ 0.011 \end{array}$					
Sèvre Niortaise	[-35.7;-29.2] ± 1.0	-258×exp([chla+phaeo] <sup>2/</sup> 16055)-0.15×[temp]+229				[0.106;0.145] ± 0.006	2.9x10 <sup>-</sup> <sup>3</sup> ×[chla+phaeo] +				
Charente	-30.8 ± 0.03		7.5 ± 1.6		$\begin{array}{c} 6.6 \\ \pm \ 0.3 \end{array}$	$0.152 \pm 0.006$					
Seudre	-33.3 ± 0.1										
Porge	-33.6 ± 0.4					$\begin{array}{c} 0.128 \\ \pm \ 0.008 \end{array}$					
Cirès / Renet / Milieu / Lanton / Tagon	-34.9 ± 0.4					$\begin{array}{c} 0.133 \\ \pm \ 0.006 \end{array}$					
Leyre	-30.1 ± 0.3					$\begin{array}{c} 0.140 \\ \pm \ 0.016 \end{array}$					
Landes	-29.9 ± 0.3					0.112 ± 0.010					
Adour	-28.2 ± 0.6					$\begin{array}{c} 0.111 \\ \pm \ 0.010 \end{array}$					
Têt	$\substack{[-29.7;-27.8]\\\pm0.6}$	-5.2×10 <sup>-3</sup> [temp] <sup>2</sup> +0.08×[temp]-27.5	${[5.3;13.3]\atop \pm 1.8}$	5.53×[temp]-5.5	$\begin{array}{c} 5.6 \\ \pm \ 0.7 \end{array}$	$\begin{array}{c} 0.181 \\ \pm \ 0.021 \end{array}$		-26.3 ± 0.1	-0.7 ± 0.1	$\begin{array}{c} 6.3 \\ \pm \ 0.3 \end{array}$	$\begin{array}{c} 0.160 \\ \pm \ 0.017 \end{array}$
Aude	[-32.6;-27.8] ± 0.6	-0.21×[temp]-26.5	[5.2;10.6] ± 1.6	-1.13×δ <sup>13</sup> C-26.2	$\begin{array}{c} 5.0 \\ \pm \ 0.8 \end{array}$	0.205 ± 0.033					
Orb	[-30.7;-23.4] ± 0.6	-0.19×[temp]-26.0	[4.9;8.4] ± 0.6	8.44-(3.63×(conduc- 505))/(conduc-111)	$\begin{array}{c} 4.8 \\ \pm \ 0.9 \end{array}$	0.213 ± 0.039		-27.1 ± 0.4	1.9 ± 1.9	3.7 ± 3.7	$\begin{array}{c} 0.270 \\ \pm \ 0.270 \end{array}$
Hérault	[-31.5;-27.5] ± 1.0	-0.19×[temp]-26.0	[6.3;10;9] ± 1.3	3.6x10 <sup>-2</sup> ×[temp] <sup>2</sup> - 1.15×[temp]+14.6	$\begin{array}{c} 5.0 \\ \pm \ 0.7 \end{array}$	$\begin{array}{c} 0.203 \\ \pm \ 0.031 \end{array}$					
Rhône	-27.8 ± 1.2		5.6 ± 0.8		$\begin{array}{c} 5.5 \\ \pm \ 0.8 \end{array}$	$\begin{array}{c} 0.180 \\ \pm \ 0.030 \end{array}$					

#### 252

#### 253 **2.4. Quantification of POM composition**

254 POM composition was quantified using a Bayesian mixing model ('simmr' R package version 255 0.4.5, Govan and Parnell, 2023) which solves the equations system based on bulk and source POM elemental and isotopic signatures. Mixing models were computed for each sampling date 256 of each river (Tab. 1), using carbon ( $\delta^{13}$ C and N/C ratio, Eq. 4, 7, 8), nitrogen ( $\delta^{15}$ N and C/N 257 ratio, Eq. 5, 6, 8), and/or a combination of three ( $\delta^{13}$ C,  $\delta^{15}$ N and N/C ratio, Eq. 4, 5, 7, 8) tracers. 258 259 From the three mixing models performed for each sampling date and river (carbon, nitrogen or 260 mixed), one model was selected as the best estimation of bulk POM data. It should be noted that N/C and C/N ratios give information on the mixing of C and N, respectively (Perdue and 261 262 Koprivnjak, 2007). We used at least the same number of equations as unknowns (sources) to 263 avoid running underdetermined models that result in large uncertainty in model outputs(Phillips et al., 2014). Equations of the models were: 264





265	$\delta^{13}C_{\text{mixture}} = x_1 \ \delta^{13}C_{\text{source }1} + x_2 \ \delta^{13}C_{\text{source }2} + x_3 \ \delta^{13}C_{\text{source }3} + x_4 \ \delta^{13}C_{\text{source }4}$	(Eq. 4)
266	$\delta^{15}N_{mixture} = x_1 \ \delta^{15}N_{source \ 1} + x_2 \ \delta^{15}N_{source \ 2} + x_3 \ \delta^{15}N_{source \ 3} + x_4 \ \delta^{15}N_{source \ 4}$	(Eq. 5)
267	$C/N_{mixture} = x_1 C/N_{source 1} + x_2 C/N_{source 2} + x_3 C/N_{source 3} + x_4 C/N_{source 4}$	(Eq. 6)
268	$N/C_{mixture} = x_1 N/C_{source 1} + x_2 N/C_{source 2} + x_3 N/C_{source 3} + x_4 N/C_{source 4}$	(Eq. 7)
269	$x_1 + x_2 + x_3 + x_4 = 1$	(Eq. 8)

270 As there was no *a priori* knowledge of sources contributions to the POM mixture, the models 271 were set with an uninformative prior (1, 1, 1, 1) following a Dirichlet distribution (all sources 272 have an equal probability to contribute to the mix; Phillips et al., 2014). Model runs were set 273 following the recommendations of Phillips et al. (2014). Models outputs were evaluated with 274 Gelman-Rubin diagnostic (verification of chain convergence) and predictive distributions to 275 ensure the good fit of the models to the observed data. Models outputs are given as medians. 276 Absolute uncertainties for the models varied from 1 to 18 % (range of average for each river) 277 with an overall average of 8 % (all models).

#### 278

#### 2.5. Forcings at local and multi-systems scales

Environmental forcings driving POM composition were determined using redundancy analysis
(RDA; 'dudi.pca' and 'pcaiv' functions; R package {ade4} version 1.7-19). RDA summarizes
multiple linear regressions between the response variable (POM composition: mixing model
outputs) and a set of explanatory variables (environmental forcings) to assess causality links
(Legendre et al., 2011). RDAs were performed at single-river and multi-river scales. Regarding
the multi-rivers scale, the annual mean POM composition of each river was used to determine
the drivers of spatial (i.e., between-rivers) variations of POM composition.

286 The proxies of the environmental forcings were chosen to directly or indirectly reflect processes 287 that occur in the river and the adjacent ecosystems (e.g., primary production, soil leaching or 288 WWTP's discharge) influencing POM source inputs and isotopic values. To homogenize the 289 data sets for running the single-river RDAs, the same combination of twelve proxies for 290 environmental forcings was used for each river: SPM, chlorophyll a, phaeopigments, 291 temperature, daily river flow, pH, ammonium, nitrates, phosphates, irradiance, zonal and 292 meridional wind. For the multi-river RDA, environmental proxies were selected to reflect 293 processes occurring at large spatial scales and in the river basin. Hence, a new combination of 294 sixteen proxies was used: river flow, conductivity, air temperature, precipitations, zonal and 295 meridional wind, artificial, agricultural, forest and natural and wetlands surface areas, net soil erosion, organic carbon content in the soil, river length, basin surface area, latitude, longitude. 296 297 From this initial list of proxies, some were removed to limit the auto-correlation (use of the 298 Variance Inflation Factor, Borcard et al., 2011) and to improve the adjusted R<sup>2</sup> of each RDA 299 analysis.





#### **2.6. Typology of systems**

301 Rivers were classified based on POM composition and its temporal dynamics by performing a 302 regionalization analysis as in Liénart et al. (2018) (Fig. A1). This method based on multivariate 303 cluster analysis allows to consider the temporal (seasonal) variations specific to each river in 304 addition to the spatial (between-rivers) component. The regionalization analysis was based on POM composition data (i.e., proportions of sources) computed for each river and each month. 305 306 When more than twelve months were available (bi-monthly sampling or more years), a standard 307 year of twelve months was chosen (averaged by month if fortnight dates). (Souissi et al., 2000) 308 A contingency matrix (rivers, sources, months) was created from monthly values of source 309 contributions (i.e., mixing model outputs). For each month, a dendrogram was performed and 310 ten cut-off levels were considered. Then, for each cut-off level, similarities between stations were identified within the twelve-monthly dendrograms. Ultimately, global similarities 311 312 between rivers were computed using a fuzzy cluster that returns probabilities of membership of 313 each river to each cluster type. The best number of river types, i.e., river typology, was 314 determined considering the best Dunn coefficient (Dunn, 1974) and Silhouette score (Rousseeuw, 1987). 315

#### 316 **3. Results**

Hereafter, four rivers were selected and considered as representative of each type of studied
river (see section 3.4). Thus, most of the results are illustrated using these four rivers. Graphs
of all the other rivers are reported in the supplementary material.

#### 320 **3.1.Contrasted seasonalities in river characteristics**

As stated in section 2.1, the 23 studied rivers encompassed large gradients of environmental 321 322 characteristics, as illustrated by the lowest and highest annual means of river flow (0.3 and 1572 323 m<sup>3</sup>/s; Lanton and Rhône Rivers), water temperature (12.3 to 17.1 °C; Cirès and Têt Rivers), 324 SPM (2.7 and 40.9 mg/l; Cirès and Rhône River), POC (0.3 and 5.1 mg/l; Hérault and Loire 325 Rivers) and chlorophyll a (0.4 to 57.1 µg/l; Cirès and Rance Rivers) concentrations as well as POC/chl a (199 and 6444 g/g; Loire and Leyre Rivers) and C/N (5.9 and 20.3 mol/mol; Têt and 326 327 Lanton Rivers) ratios; this was less contrasted among rivers for  $\delta^{13}$ C (-30.2 and -26.2 ‰; Sèvre and Têt Rivers) and especially  $\delta^{15}N$  (4.0 and 8.0 %; Leyre and Rance Rivers) (Fig. 2, A2). 328

As generally observed in rivers from mid-latitude, studied rivers exhibited clear seasonal patterns in water temperature with lower and higher values in winter and summer, respectively. However, such clear seasonal patterns were not always recorded for all the parameters, as there were contrasted patterns of seasonal variability among rivers. Indeed, the seasonal variability of river flow was quite smooth (e.g., the Rance and Charente Rivers) with a higher flow in winter/spring and lower flow in summer/fall for some rivers, whereas it was highly pulsed for some others with constant low levels marked by short and strong floods (e.g., 53m<sup>3</sup>/s in mean





336 but 1169m<sup>3</sup>/s in flood time for the Hérault River) (Fig. 2). Overall, one can distinguish rivers 337 that are characterized by high concentrations of chlorophyll a and clear seasonal patterns of most parameters (e.g., 53  $\mu$ g/l of chlorophyll *a* in mean ranging from 3 to 135  $\mu$ g/l in the Rance 338 339 River) from rivers characterized by low concentrations of chlorophyll a, high POC/chl a and low seasonal variability for most of the parameters (e.g., 1.1 µg/l of chlorophyll a in mean 340 341 ranging from 0.7 to 1.7  $\mu$ g/l in the Milieu River) and from rivers that are characterized by high 342 seasonal variability of most parameters but without a clear seasonal pattern (e.g., Hérault 343 River). Other rivers exhibited intermediate behavior (e.g., Charente River) (Fig. 2, A2). 344 Usually, Rance-like rivers exhibited high concentrations of chlorophyll a in spring/summer associated with POC/chl a ratio lower than 200 g/g, C/N ratio lower than 8 mol/mol and low 345 δ<sup>13</sup>C (down to -31 ‰ or -33 ‰; e.g., Seine River, Fig. A2). In contrast, Milieu-like rivers 346 exhibited high POC/chl a (> ~ 700 g/g) and C/N ratio (> 15 mol/mol) and quite constant  $\delta^{13}$ C 347 348 (~-29 - -28 ‰) all year round (e.g., Cirès and Renet Rivers). These rivers are tributaries of the 349 Arcachon Lagoon. Hérault-like rivers flowing into the Mediterranean Sea exhibited highly and suddenly variable C/N ratios (4 – 17 mol/mol),  $\delta^{13}$ C (~-33 – -26 ‰) and  $\delta^{15}$ N (~2 – 12 ‰) (e.g., 350 351 Aude and Orb Rivers; Fig. A2).











353 <u>Figure 2</u> Temporal variations of matter characteristics for representative rivers along the studied 354 periods for  $\delta^{13}$ C (left axis; black dotted line) and  $\delta^{15}$ N (right axis; blue line) (first column); C/N 355 (left axis; black dotted line) and POC/chl a(right axis; blue line) ratios (second column); 356 SPM(left axis; black dotted line), POC (right axis; blue line) and chl *a* (right axis; blue dotted 357 line) concentrations (third column) and river flow (left axis; black dotted line) and temperature 358 (right axis; blue line) (fourth column).

#### 359 **3.2. Elemental and isotopic signatures of POM sources**

360 Elemental and isotopic signatures of phytoplankton were estimated for each of the twenty-three 361 rivers (Tab. 2, Fig. 3 and A3). Most of them (all of them for the C/N ratio) were found constant 362 over time. Their annual mean values varied between -34.9 ‰ (some tributaries of the Arcachon Lagoon) and -27.4 ‰ (Elorn River) for  $\delta^{13}$ C, between 4.3 ‰ (Elorn River) and 8.6 ‰ (Aulne 363 River) for  $\delta^{15}$ N and between 4.8 mol/mol (Orb River) and 10.0 mol/mol (Elorn River) for the 364 365 C/N ratio. Some of them varied over time along with pigment concentration and ratio or with temperature for  $\delta^{13}$ C, and with pigment concentration, nitrate concentration, temperature,  $\delta^{13}$ C 366 or conductivity for  $\delta^{15}$ N (Tab. 2). The range of temporal variability was usually 4-6 % for  $\delta^{13}$ C 367 and  $\delta^{15}$ N. Overall, phytoplankton signatures are comprised between -35.6 and -23.8 % for the 368  $\delta^{13}$ C and between 3.0 and 13.2 ‰ for the  $\delta^{15}$ N. 369

All other signatures were found constant over time (Tab. 2 and A2, Fig. 3 and A3) but may 370 371 differ between rivers. Signatures mean annual values of labile terrestrial POM were comprised between -29.1 and -26.0 % for the  $\delta^{13}$ C, between 3.7 and 6.6 % for the  $\delta^{15}$ N and between 8.8 372 and 17.0 mol/mol for the C/N ratio. Signatures mean annual values of refractory terrestrial POM 373 374 were comprised between -28.0 and -25.9 % for the  $\delta^{13}$ C, between 3.1 and 6.7 % for the  $\delta^{15}$ N and between 5.8 and 8.8 mol/mol for the C/N ratio. Signatures mean annual values of sewage 375 POM were -27.1 and -26.3 ‰ for  $\delta^{13}$ C, 1.9 and -0.7 ‰ for  $\delta^{15}$ N and 3.7 and 6.3 mol/mol for 376 C/N ratio for Orb and Têt Rivers, respectively. 377





378







379 <u>Figure 3</u>  $\delta^{13}$ C,  $\delta^{15}$ N, N/C or C/N values of bulk POM (black crosses) and sources. The latter are 380 presented as closed circles (average) and bars (standard deviation) when the signatures were 381 constant over time and by colored area when at least one of the proxies was variable over time 382 (see Table 2). This colored area corresponds to the dispersion of the values including their 383 uncertainties.

#### **384 3.3.Dynamics of particulate organic matter composition**

385 Particulate organic matter composition resulting from mixing models outputs is presented hereafter, for each river, as the relative contribution of each source to the POM pool (Fig. 4). 386 387 Among rivers whose POM is composed of only two sources (terrestrial POM and phytoplankton), one can distinguish rivers with terrestrial-dominated POM (e.g., Milieu River: 388 389 terrestrial POM accounted for  $94 \pm 3$  % of the mixture) to rivers of intermediate POM composition (e.g., Charente and Rance Rivers where phytoplankton accounted for  $34 \pm 10$  % 390 391 and  $62 \pm 10$  % of the mixture, respectively). All these rivers flow in the English Channel and 392 the Atlantic Ocean. The rivers whose POM is composed of three or four sources flow in the 393 Mediterranean Sea. In these rivers, terrestrial POM is present as refractory and labile materials. 394 The contribution of labile terrestrial POM ranged between  $16 \pm 15$  % (Têt River) and  $46 \pm 21$ % (Orb River), and of refractory terrestrial POM between  $21 \pm 9$  % (Rhône River) and  $39 \pm 15$ 395 396 % (Aude River). The contribution of phytoplankton ranged between  $34 \pm 15$  % (Aude River) 397 and  $51 \pm 30$  % (Hérault River) for the Mediterranean rivers. The fourth source of POM was 398 WWTP's POM. It was identified as a source in the Orb and Têt Rivers and accounted for  $15 \pm$ 399 6 % and 10  $\pm$  7 % in these two rivers respectively. Regarding temporal variations of POM 400 composition, some rivers exhibited clear seasonal patterns whereas others revealed a 401 homogeneous composition over the annual cycle (Fig. 4). The rivers where POM was highly 402 dominated by terrestrial POM (Seudre, Cirès, Renet, Lanton, Milieu, Tagon, Leyre Rivers) 403 showed almost no seasonal variability. In contrast, some rivers like the Rance, the Elorn or the 404 Aulne River showed a clear seasonal pattern with the dominance of terrestrial material in winter 405 and phytoplankton in summer. At last, other rivers exhibited less clear (e.g., Landes, Porge, Charente Rivers) or even no clear seasonal pattern but a quite stochastic variability over the 406 annual cycle (e.g., Sèvre, Adour, Aude, Orb). 407

It should be noted that the above is valid for carbon and mixed as well as nitrogen models (cf.Tab. 2; Fig. 4 and A4).













411

412 <u>Figure 4</u> Temporal dynamic (rectangle graphs) and (inter-)annual mean (pie charts) of POC
413 source proportions. Sources are phytoplankton (green), labile terrestrial material (brown),
414 refractory terrestrial material (yellow) and anthropogenic POM (orange).





#### 416 **3.4. Typology of rivers**

417	Four types of rivers were determined by the regionalization analysis based on river POM
418	composition and its temporal dynamics (Fig. 5). The seven rivers (Renet, Cirès, Lanton, Milieu,
419	Seudre, Tagon and Leyre River) mainly belonging to Type I are characterized by terrestrial-
420	dominated POM and no/low seasonality. Six of them are small streams/rivers flowing to the
421	Arcachon Lagoon. The five rivers (Aude, Hérault, Têt, Rhône and Orb River) mainly belonging
422	to Type II are characterized by the co-occurrence of labile and refractory terrestrial POM and
423	large temporal variability but, except for the Hérault River, without a clear seasonal pattern.
424	They all flow to the Mediterranean Sea. The five rivers (Porge, Adour, Charente, Orne and
425	Landes River) mainly belonging to Type III are composed of phytoplankton and terrestrial
426	POM, and exhibit moderate seasonality. Type III is clearly an intermediary between Type I and
427	Type IV. These five rivers flow to the Atlantic Ocean or the English Channel. Among the seven
428	rivers flowing to the Arcachon Lagoon, the two that mainly belong to Type III are man-
429	managed streams and flow through lakes, contrary to the six other ones, which mainly belong
430	to Type I and are natural streams that do not flow through lakes. Finally, the six rivers (Rance,
431	Elorn, Aulne, Loire, Seine and Sèvre River) mainly belonging to Type IV are composed of
432	phytoplankton and terrestrial POM, and exhibit high seasonality. These six rivers flow to the
433	Atlantic Ocean or the English Channel.



434

435 <u>Figure 5</u> Typology of rivers following a hierarchical cluster analysis on POM source
 436 proportions. The percentages of membership for each type attributed to each river are shown.

#### 437 **3.5.Environmental forcings driving POM composition**



457



438 One redundancy analysis was performed for each river to relate environmental parameters, 439 considered as proxies of drivers, to the POM composition, i.e., to assess the drivers of the temporal variability of POM composition for each river (Fig. 6 and A5). It should be kept in 440 mind that the POC or PN concentration of each source was used for these analyses and not the 441 relative proportion of the sources. In type-I rivers, i.e., rivers characterized by terrestrial-442 443 dominated POM and no/low seasonality, terrestrial POM is usually linked to river flow and/or 444 SPM concentration (e.g., Milieu River on Fig. 6, Leyre and Tagon Rivers in Fig. A5). However, 445 this feature is not always clear since the POM of these rivers is always dominated by terrestrial 446 material, almost whatever the environmental conditions are. In type-II rivers, i.e., rivers characterized by the co-occurrence of labile and refractory terrestrial POM and large temporal 447 448 variability, phytoplankton POM is usually positively linked to temperature and negatively 449 linked to river flow, whereas labile and refractory terrestrial POM is positively linked to SPM 450 and/or river flow. Interestingly, labile terrestrial POM is usually better linked to river flow and 451 refractory terrestrial POM to SPM (e.g., Hérault River in Fig. 6 and Rhône River in Fig. A5). 452 In the Têt River, anthropogenic POM was linked to nitrate concentration (Fig. A5). In rivers 453 characterized by phytoplankton and terrestrial-POM composition with moderate (Type III) or high (Type IV) seasonality, terrestrial POM was almost always positively linked to river flow 454 455 and/or SPM concentration while phytoplankton was usually linked with chl a concentration (e.g., Charente and Rance River on Fig. 6, Landes and seine River on Fig. A5). 456







458 <u>Figure 6</u> Redundancy analyses (correlation circles) of rivers standing for each type of river. 459 Black arrows represent explained variables (concentration of POC sources) and red arrows 460 represent explaining variables (environmental variables). River types are recalled (Roman 461 numerals). POC<sub>lt</sub> = Labile terrestrial POC; POC<sub>rt</sub> = Refractory terrestrial POC; POC<sub> $\phi$ </sub> = 462 Phytoplankton POC; Chl a = chlorophyll a; Phaeo. = phaeopigments; M. wind = meridional 463 wind; Z. wind = zonal wind; R. flow = river flow; Temp. = temperature; Irrad. = Irradiance; 464 NH<sub>4</sub><sup>+</sup> = ammonium; NO<sub>3</sub><sup>-</sup> = nitrate; PO<sub>4</sub><sup>3-</sup> = phosphates ; Adj. R<sup>2</sup> = adjusted R<sup>2</sup>.

465 At last, another RDA was performed gathering all rivers to relate environmental parameters to 466 the mean annual POM composition at the multi-rivers scale (Fig. 7). As anthropogenic POM 467 was only detected in two rivers (Orb, Têt), it was not included in the multi-rivers analysis to 468 avoid analysis bias. At this scale, phytoplankton is strongly positively linked to agricultural 469 surfaces, labile terrestrial material to soil erosion rate and soil organic carbon and refractory 470 terrestrial material to river flow and water temperature.



471

472Figure 7Multi-rivers redundancy analysis. Black arrows represent explained variables (relative473proportions), red arrows represent explaining variables (environmental variables) and numbers474are river identifiers (cf. Fig.1). R. flow = river flow; OC soil = percentages of organic carbon475in soil; Soil erosion = soil erosion rate; Rain. = precipitations; NO<sub>3</sub><sup>-</sup> = nitrate concentration ;476Adj. R<sup>2</sup> = adjusted R<sup>2</sup>.

#### 477 4. Discussion

#### 478 **4.1.Bulk data and source signatures in temperate rivers**

479 Over the 23 studied rivers,  $\delta^{13}$ C,  $\delta^{15}$ N and C/N ratios of bulk POM ranged between -35.2 and -480 24.5 ‰, -0.3 and 12.6 ‰, and 3 and 23.4 mol/mol, respectively. This corresponds to usual 481 values recorded for riverine POM over temperate systems, except for the lowest C/N ratios

482 (Ferchiche et al., 2024; Kendall et al., 2001; Ogrinc et al., 2008).





483 In the present study, isotopic and elemental signatures of terrestrial POM and phytoplankton 484 were determined from subsets of the bulk data sets following the approaches of Savoye et al. (2012), Liénart et al. (2017) and Ferchiche et al. (2025, 2024). It has the double advantage of 485 1) taking into account the reworking of terrestrial POM within the river and thus discriminating 486 487 labile from refractory terrestrial POM, and 2) taking into account the variability of phytoplankton signature over time, due to differences in growth conditions (see below). Labile 488 489 terrestrial POM mainly appears during high river flow (Fig. 6 and A5; Savoye et al., 2012) and 490 is usually composed of riparian litter (e.g., Veyssy et al., 1998). In the studied rivers,  $\delta^{13}C$ ,  $\delta^{15}N$ 491 and C/N ratio of labile terrestrial POM ranged between  $-28.9 \pm 0.8$  ‰ and  $-26 \pm 0.9$  ‰,  $3.7 \pm$ 492 0.6 % and 6.6  $\pm$  0.9 %, and 8.8  $\pm$  0.4 and 17  $\pm$  3.2 mol/mol, respectively. These values are very 493 similar to values found in other temperate systems like the Gironde Estuary ( $\delta^{13}C = -28.7 \pm 0.9$ %; Savoye et al., 2012), the Sava River ( $\delta^{13}C = -28 \pm 5$  %;  $\delta^{15}N = 5 \pm 2$  %; C/N = 33 ± 15 494 mol/mol Ogrinc et al., 2008) or Taiwanese rivers ( $\delta^{13}C = -26.6 \pm 1.8$  ‰; C/N = 31.1 ± 23.4 495 mol/mol Hilton et al., 2010) and very similar to direct measurement of C3 plants ( $\delta^{13}C = -28.1$ 496  $\pm 2.5$  ‰; O'Leary, 1981 and references therein;  $\delta^{13}C = -28 \pm 1.3$  ‰;  $\delta^{15}N = 0.8 \pm 2.9$  ‰; C/N 497 = 39.6 ± 25.7 mol/mol; Dubois et al., 2012;  $\delta^{13}C = -27.9 \pm 0.1$  ‰; Fernandez et al., 2003). 498 Refractory terrestrial POM is terrestrial POM that has undergone large reworking within river 499 500 water, river sediment or even estuarine maximum turbidity zone (e.g., Etcheber et al., 2007; Veyssy et al., 1998). In the studied rivers where it was found,  $\delta^{13}$ C,  $\delta^{15}$ N and C/N ratios of 501 502 refractory terrestrial POM ranged between  $-28 \pm 0.7$  ‰ and  $-25.9 \pm 0.4$  ‰,  $3.2 \pm 0.8$  ‰ and 6.7503  $\pm$  1.4 ‰, and 5.8  $\pm$  1.4 and 8.8  $\pm$  3.1 mol/mol, respectively. These values are very similar to those found in other systems like the Gironde Estuary (France) ( $\delta^{13}C = -25.2 \pm 0.3 \%$ ;  $\delta^{15}N =$ 504  $5.5 \pm 0.4$  %; C/N =  $8.5 \pm 0.8$  mol/mol; Savoye et al., 2012), Taiwanese rivers ( $\delta^{13}C = -23.6 \pm$ 505 1.1 ‰; C/N = 6.5 ± 1.6 mol/mol; Hilton et al., 2010) and in the Pearl River (China) ( $\delta^{13}$ C: 506 between  $-28.3 \pm 0.8$  ‰ and  $-21.7 \pm 0.7$  ‰; C/N: between  $8.9 \pm 1.1$  and  $17.9 \pm 3.6$  mol/mol; Yu 507 508 et al., 2010).

509 Isotopic signatures of phytoplankton vary depending on biogeochemical conditions and processes like nutrient availability and utilization, growth rate and limitation (e.g., Fry, 1996; 510 511 Liénart et al., 2017; Lowe et al., 2014; Miller et al., 2013; Savoye et al., 2003; Sigman et al., 512 2009; Yan et al., 2022) and can be estimated using measured environmental parameters 513 (Ferchiche et al., 2024, 2025; Liénart et al., 2017; Savoye et al., 2012). For the seven rivers where phytoplankton isotopic signatures were found variable over time, phytoplankton  $\delta^{13}$ C or 514  $\delta^{15}$ N were correlated to: concentrations and ratio of chlorophyll a and phaeopigments, water 515 516 temperature, nitrate concentration and/or conductivity (Tab. 2). Chlorophyll a and 517 phaeopigments concentrations are direct proxies of phytoplankton fresh and degraded 518 biomasses and are related to phytoplankton growth and decay, two processes that increase 519 phytoplankton  $\delta^{13}$ C (Golubkov et al., 2020; Michener and Kaufman, 2007 and references





520 therein). Similar processes may explain phytoplankton- $\delta^{15}$ N increase with chlorophyll a 521 increase. An increase in water temperature accelerates bio-mediated carbon remineralization processes bringing a lower  $\delta^{13}$ C value than CO<sub>2</sub> coming from water-atmosphere equilibration 522 523 and rock-leaching CO<sub>2</sub> (Polsenaere et al., 2013 and references therein). Consequently, 524 phytoplankton  $\delta^{13}$ C decreases as phytoplankton uses remineralized CO<sub>2</sub> and thus as water temperature increases. Phytoplankton  $\delta^{15}N$  depends on N-nutrient origin and availability 525 526 (Savoye et al., 2003 and references therein). Especially, it increases with nutrient concentration 527 decrease (Sigman et al., 2009) as reported for the Rance River (Tab. 2). Water conductivity could be considered as a proxy of water mass and thus of nitrate origin. This may explain the 528 relationship between phytoplankton  $\delta^{15}$ N and water conductivity in the Orb River (Tab. 2). 529

530 In the studied rivers, phytoplankton  $\delta^{13}$ C,  $\delta^{15}$ N and C/N ratio ranged between -34.9 ± 0.4 and -531  $23.8 \pm 0.6$  ‰,  $4.3 \pm 0.8$  and  $13.2 \pm 1.8$  ‰, and  $4.8 \pm 0.9$  and  $10 \pm 0.9$  mol/mol, respectively. This is similar to values reported for the Loire River, another French river (-30.6  $\leq \delta^{13}C \leq -25.0$ 532 533  $3.0 ≤ \delta^{15}$ N ≤ 10.4 ‰; C/N = 7.2 ± 0.6 mol/mol: Ferchiche et al, 2024) but narrower ranges 534 can be found in the literature. In the Sava River (Eastern Europe), phytoplankton signature was  $-30.4 \pm 2.1$  ‰,  $5.0 \pm 1.5$  ‰ and  $6.5 \pm 1.5$  mol/mol for  $\delta^{13}$ C,  $\delta^{15}$ N and C/N ratio, respectively 535 (Ogrine et al., 2008), similar to that of Indian ( $\delta^{13}C = -30.6 \pm 1.7 \%$ ,  $\delta^{15}N = 7.0 \pm 2.3 \%$ ; 536 Gawade et al., 2018) and Texan ( $\delta^{13}C = -31.4$  %; Lebreton et al., 2016) rivers. Lower  $\delta^{13}C$ 537 538 values ( $\leq$  -32 ‰) were also found (Finlay et al., 2010; Hellings et al., 1999; Sato et al., 2006; Savoye et al., 2012). However, values of elemental and isotopic ratios for riverine 539 540 phytoplankton are scarce in the literature. Indeed, it is not easy to estimate phytoplankton 541 signature since it cannot be separated from other particles. Thus, literature estimates may not 542 be perfectly representative of the variability of phytoplankton isotopic signatures.

## 543 4.2.Watershed characteristic drive on spatial dynamics of POM 544 composition

545 At the annual scale, we observed deep variations between studied rivers concerning the mean 546 POC proportion of the different sources ( $5 \le phytoplankton \le 80\%$ ;  $17 \le labile$  terrestrial POC  $\leq$  95 %; 0  $\leq$  refractory terrestrial POC  $\leq$  39 %). Interestingly, phytoplankton proportion was 547 highly correlated to the proportion of agriculture surface area (Fig. 7;  $R^2 = 0.59$  or even 0.72 548 549 when Seudre River is not included in the statistics) and nitrate concentration (Fig. 7;  $R^2 = 0.41$ 550 when Seudre River is not included in the statistics). Such relationship between agriculture 551 surface area and phytoplankton is well-known, as agricultural activities increase nutrient input 552 to river bodies (Khan and Mohammad, 2014), leading to better conditions for phytoplankton 553 growth (Dodds and Smith, 2016; Minaudo et al., 2015). Interestingly, the proportion of labile 554 terrestrial matter was strongly linked to soil erosion and soil organic carbon content (Fig. 7) 555 indicating a strong relationship between terrestrial matter in rivers and soil nature with 556 undecomposed and fresh detrital matter (McCorkle et al., 2016).





#### 557 **4.3. Temporal dynamics of POM composition and river typology**

558 In aquatic systems, phytoplankton likely appears during spring and summer in favorable 559 conditions, related to low discharge, high-temperature conditions and enough nutrients to support its growth, while in winter, high turbidity and low-temperature conditions limit its 560 presence (Turner et al., 2022). Quantitative differences between rivers can be due to differences 561 in nutrient availability, either because of anthropic mitigation (Minaudo et al., 2015), 562 563 competition with other nutrient users (Descy et al., 2012; Minaudo et al., 2016) or sewage inputs (Codiga et al., 2022) in addition to agricultural inputs (see section 4.2). Terrestrial material 564 likely appears during winter conditions, related to floods that transport great amounts of 565 terrestrial material (Dalzell et al., 2007). Such a seasonal dichotomy between phytoplankton 566 and terrestrial POM was clearly visible for most of the studied rivers (Fig. 4), especially for 567 type-IV and type-III rivers, but even for some of those highly dominated by the labile terrestrial 568 569 POM (e.g., Milieu and Tagon Rivers; Type-I rivers). This was illustrated by the relationships 570 between phytoplankton POM and chlorophyll a concentration and/or temperature (as proxies 571 of favorable conditions for phytoplankton production) on the one hand, and between labile terrestrial POM and river flow and/or SPM concentration on the other hand (Fig. 6 and A5). 572 573 This dichotomy in POM composition was also reported in other similar studies (e.g., Kelso and Baker, 2020; Lu et al., 2016). In rivers where refractory terrestrial POM was present in addition 574 575 to the labile one (type-II rivers), it was interesting to see that the refractory terrestrial POM was 576 more related to SPM concentration than river flow and inversely for the labile terrestrial POM. This indicates that labile and refractory terrestrial POM were preferentially associated with 577 578 direct river input and sediment resuspension, respectively. The origin of the refractory terrestrial 579 POM may be fossil/bedrock/petrogenic OM (e.g., Copard et al., 2022; Hilton et al., 2010; Sun et al., 2021) brought by river flow (in quantity undetectable in the bulk POM using our tools) 580 and then accumulated in the sediment, and/or labile terrestrial POM brought by the river flow 581 582 and then accumulated and reworked/decayed in the sediment (e.g., Etcheber et al., 2007; Savoye et al., 2012). 583

584 Sewage POM was detected in two of the studied rivers but with different associated temporal dynamics. In the Têt River, because the former WWTP was dysfunctional, a new one replaced 585 586 it in late 2007 (https://www.assainissement.developpement-durable.gouv.fr/pages/data/fiche-587 060966136002, last visit 10/09/24). This explains the shift in sewage POM between the two 588 studied periods (2006-2007 versus 2008-2010 without anthropogenic POM). In the Orb River, sewage POM was detected throughout the studied periods. The WWTP is located only a few 589 590 kilometers upstream of the sampling site and is large enough (220 000 inhabitant equivalent) compared to the river flow (annual mean: 23m<sup>3</sup>/s) to make the sewage POM detectable in the 591 bulk POM using  $\delta^{15}$ N. Such a result is quite common for urban rivers (e.g., Kelso and Baker, 592 593 2020).





#### 594 **4.4.Originality of the study**

595 The originality of the present study firstly lies in its approach. Even if C and N stable isotopes 596 have been used for decades to investigate POM origins within river waters, the quantification 597 of POM composition (i.e., the relative proportion of each source composing the POM) using 598 mixing models, and especially Bayesian mixing models, is not so common. In addition, most 599 of the previous studies either use literature data for phytoplankton isotopic signature (e.g., 600 Zhang et al., 2021) or use lake or autochthonous POM as a proxy of phytoplankton (e.g., Kelso 601 and Baker, 2020). Also, most of these studies use direct measurements of soil or plants to assess 602 the isotopic signature of terrestrial POM whereas this material is able to rework within the water 603 column or the sediment, which changes its elemental and isotopic values (e.g., Savoye et al., 604 2012). These approaches do not consider that isotopic signatures of phytoplankton and 605 terrestrial material may change over time. In the present study, we used the approach developed 606 by Savoye et al. (2012) in an estuary, Liénart et al. (2017) in coastal systems and Ferchiche et 607 al. (2025, 2024) in a river to assess the elemental and isotopic signatures from subsets of bulk 608 POM and when needed, empirical equations. This approach has the great advantage of 1) using 609 signatures dedicated to the sampling area and 2) taking into account the potential variability of 610 these signatures over time, i.e., depending on the environmental conditions for phytoplankton 611 growth and taking into account its decay for phytoplankton and terrestrial POM. Especially, we 612 were able to discriminate labile from refractory terrestrial POM in some rivers, as Savoye et al. 613 (2012) in an estuary. Another great originality of the present study lies in the multi-systems 614 approach: studying 23 rivers in a single study allowed the detection of four types of river 615 functioning regarding the POM composition and its temporal dynamics. It also highlights the great influence of land use (agriculture) and characteristics (erosion, organic carbon content) 616 on the POM composition of rivers. At last, using  $\delta^{13}$ C,  $\delta^{15}$ N and C/N ratio all together allowed 617 618 either to perform mixing models with up to four end-members or to study separately POC and 619 PN composition. It showed that POC and PN display very similar compositions and dynamics 620 in rivers.

#### 621 5. Synthesis and perspectives

622 The present study proposes a comprehensive estimation of POM composition and its spatial 623 and seasonal variability in temperate rivers. Thanks to the inclusion of twenty-three rivers, 624 encompassing large gradients of environmental conditions under a temperate climate, a river 625 typology is proposed based on the POM composition and its temporal dynamics. In type-I 626 rivers, POM is dominated by labile terrestrial material all year long. This material is mainly 627 associated with suspended particulate matter. Phytoplankton slightly contributes, especially during summertime. Type-II rivers are mainly characterized by the presence of both labile and 628 629 refractory terrestrial material, in addition to phytoplankton. The temporal variability between 630 these sources is high but the seasonality is not always pronounced even if phytoplankton and





631 terrestrial POM can dominate the POM composition during summer and winter, respectively. 632 Nevertheless, labile terrestrial POM is mainly related to river flow and refractory terrestrial POM to SPM, indicating the sedimentary origin of the latter. In type-III rivers, POM is 633 composed of phytoplankton and labile terrestrial material. The seasonality of POM composition 634 is not very pronounced even if the contribution of labile terrestrial POM is deeply related to 635 river flows. Type III is an intermediary between type I and type IV. In type-IV rivers, POM is 636 637 also composed of phytoplankton and labile terrestrial material but the seasonality of POM 638 composition is very pronounced with a clear balance between high phytoplankton contribution 639 in summer and high terrestrial contribution in winter. Labile terrestrial POM is deeply related to river flow. Beyond this typology, the main difference in POM composition between the 640 641 studied rivers is that the phytoplankton contribution to the POM composition is related to the 642 proportion of agricultural surface in the watershed and the contribution of labile terrestrial POM 643 is related to soil erosion and organic carbon content in the watershed.

644 The originality of this study mainly lies in 1) the approach used to determine the elemental and 645 isotopic signatures of POM sources, which allowed to discriminate labile from refractory 646 terrestrial POM and to take into account, when any, the variability of the signatures over time, 647 and 2) determining a typology of temperate rivers based on the POM composition and its 648 temporal dynamics.

649 Overall, this study, which focuses on the River-Estuary Interface, brings meaningful 650 information for the comprehension of C and N cycles along the LOAC and especially the 651 behavior, dynamics and drivers of POM that leaves the river and enters the estuary.

652 From a methodological perspective, such a study could be strengthened by the use of non-653 exchangeable  $\delta^2 H$  as an additional tool to even better distinguish and quantify more sources in 654 mixing models. This tool has been recently shown to be powerful for such purposes (Ferchiche 655 et al., 2025). From a fundamental perspective, aggregating more datasets from other temperate rivers would allow testing the robustness of this typology and probably detecting additional 656 657 types, but also datasets from polar and tropical rivers to perform an even more comprehensive 658 study at a global climate scale. In addition, a similar study dedicated to the estuarine systems 659 would even increase our comprehensive understanding of the origin and fate of POM along the 660 Land-Ocean Aquatic Continuum by complementing the present study dedicated to the River-661 Estuary Interface and those of Liénart et al. (2017, 2018) dedicated to the coastal systems.





#### Appendix a typology of river functioning: Toward a comprehensive study of POM composition at multi-rivers scale Ferchiche F.<sup>1</sup>, Liénart C.<sup>1</sup>, Charlier K.<sup>1</sup>, Deborde J.<sup>2,3</sup>, Giraud M.<sup>4</sup>, Kerhervé P.<sup>5</sup>, Polsenaere P.<sup>1,3</sup>, Savoye N.1\* <sup>1</sup> Univ. Bordeaux, CNRS, EPHE, Bordeaux INP, UMR 5805 EPOC, F-33600 Pessac, France <sup>2</sup> Univ. Pau & Pays Adour, CNRS, E2S UPPA - MIRA, UMR 5254 IPREM, F-64000 Pau, F-64600 Anglet, France <sup>3</sup> Ifremer, COAST, F-17390 La Tremblade, France <sup>4</sup> MNHN, CRESCO, Station Marine de Dinard, F-35800 Dinard, France <sup>5</sup> Univ. Perpignan, CNRS, UMR 5110 CEFREM, F-66860 Perpignan, France \*Corresponding author nicolas.savoye@u-bordeaux.fr Station marine - 2 rue du Professeur Jolyet 33120 ARCACHON Keywords: River-Estuary Interface; particulate organic matter; isotopes; multi-ecosystems study

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726 727 725 728 black dotted line) and temperature (right axis; blue line) (fourth column). δ<sup>15</sup>N (right axis; blue line) (first column); C/N (left axis; black dotted line) and POC/chl a(right axis; blue line) ratios (second column); SPM(left axis; black dotted line), POC (right axis; blue line) and chl a (right axis; blue dotted line) concentrations (third column) and river flow (left axis; Figure A2 Temporal variations of matter characteristics for representative rivers along the studied periods for  $\delta^{13}C$  (left axis; black dotted line) and



















731







- 732 733 bars (standard deviation) when the signatures were constant over time and by colored area when at least one of the proxies was variable over time Figure A3 8<sup>13</sup>C, 8<sup>15</sup>N, N/C and/or C/N values of bulk POM (black crosses) and sources. The latter are presented as closed circles (average) and
- 734 (see Table 2). This colored area corresponds to the dispersion of the values including their uncertainties.







Figure A4 Temporal dynamic (rectangle graphs) and (inter-)annual mean (pie charts) of PN
 source proportions. Sources are phytoplankton (green) and labile terrestrial material (brown).

- 743
- 744





745







·PN

NH4<sup>+</sup>

748 749 750 751 Seine C (IV) Aulne C (IV) Elorn C (IV) Adj. R<sup>2</sup> 0.95 Adj. R<sup>2</sup> 0.82 Adj. R<sup>2</sup> 0.23 Pheo Гетр. Temp. Chl. a 🖕 Chl. a R. flow Chl. a R. flow Pheo. SPM POC PO43-NH POC, Pheo. POC POC<sub>\u03cb</sub> NH4<sup>+</sup> pН SPM Seine N (IV) Aulne N (IV) Elorn N (IV) Adj. R<sup>2</sup> 0.80 Adj. R<sup>2</sup> 0.74 Adj. R<sup>2</sup> 0.04 Temp. R. flov









765 Figure A5 Redundancy analyses (correlation circles) of rivers standing for each type of river. 766 Black arrows represent explained variables (concentration of POC or PN sources) and red arrows represent explaining variables (environmental variables). River types are recalled 767 (Roman numerals). POC or PNlt = Labile terrestrial POC or PN; POCrt = Refractory terrestrial 768 POC; POC or PN\u03c6 = Phytoplankton POC or PN; POCanth. = Anthropic POC; SPM = 769 Suspended Particulate Matter; POC or PN= Particulate Organic Carbon or nitrogen; Chl. a = 770 Chlorophyll a; Pheo. = Phaeopigments; R. flow = River flow; Temp. = Temperature; M. wind 771 772 = Meridional wind; Z. wind = Zonal wind; Irrad. = Irradiance;  $NH_4^+$  = Ammonium;  $NO_3^-$  = Nitrate;  $PO_4^{3-}$  = Phosphate; Adj.  $R^2$  = adjusted  $R^2$ . 773





#### 774

#### 775 Author contributions

- 776 FF: Formal analysis, Investigation, Visualization, Writing original draft. CL: Formal analysis,
- 777 Conceptualization, Supervision, Writing original draft, Writing review & editing. KC:
- 778 Investigation. JD: Investigation, Visualization, Writing review & editing. MG: Investigation.
- 779 PK: Investigation. PP: Investigation, Visualization, Writing review & editing.
- 780 NS: Conceptualization, Formal analysis, Investigation, Methodology, Supervision, Writing -
- 781 original draft, Writing review & editing.

782

#### 783 **Competing interest**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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#### 787 **Data availability**

- All POM and environmental data used in this article are stored in Figshare, accessible for the
- review process through this private link : <u>https://figshare.com/s/a7101028e6ab5452c4db</u>.

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