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RESEARCH ARTICLE

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Key Points:

- We use observationally trained statistical models to obtain decadal predictions of ocean carbon flux from initialized GCM-based predictors
- The hybrid GCM-statistical ocean carbon flux predictions show improved skill over hindcast predictions from the GCM's biogeochemical models
- The hybrid models are used to make decadal predictions for the ocean-atmosphere carbon flux over the decade ending in 2029

Supporting Information:

Supporting Information may be found in the online version of this article.

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Improving GCM-Based Decadal Ocean Carbon Flux Predictions Using Observationally-Constrained Statistical Models

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Abstract An essential step toward meeting agreed climate targets and policies is the ability to understand and predict near-term changes in global carbon cycle, and importantly, ocean carbon uptake. Initialized climate model simulations have proven skillful for near-term predictability of the key physical climate variables, for example, temperature, precipitation, etc. By comparison, predictions of biogeochemical fields like ocean carbon flux, are still emerging. Initial studies indicate skillful predictions are possible for lead-times up to 6 years at global scale for some CMIP6 models. However, unlike core physical variables, biogeochemical variables are not directly initialized in existing decadal prediction systems, and extensive empirical parametrization of ocean-biogeochemistry in Earth System Models introduces a significant source of uncertainty. Here we propose a new approach for improving the skill of decadal ocean carbon flux predictions using observationally-constrained statistical models, as alternatives to the ocean-biogeochemistry models. We use observations to train multi-linear and neural-network models to predict the ocean carbon flux. To account for observational uncertainties, we train using six different observational estimates of the flux. We then apply these trained statistical models using input predictors from the Canadian Earth System Model (CanESM5) decadal prediction system to produce new decadal predictions. Our hybrid GCM-statistical approach significantly improves prediction skill, relative to the raw CanESM5 hindcast predictions over 1990–2019. Our hybrid-model skill is also larger than that obtained by any available CMIP6 model. Using bias-corrected CanESM5 predictors, we make forecasts for ocean carbon flux over 2020–2029. Both statistical models predict increases in the ocean carbon flux larger than the changes predicted from CanESM5 forecasts. Our work highlights the ability to improve decadal ocean carbon flux predictions by using observationally-trained statistical models together with robust input predictors from GCM-based decadal predictions.

Plain Language Summary Using initialized Earth system model simulations for near term predictions of ocean biogeochemical variables is an emerging field of research. In particular, near term predictability of ocean carbon flux is central to efforts for planing and limiting climate change. Unlike physical variables whose predictability have been established, these simulations are only indirectly initialized and rely on heavily parameterized ocean biogeochemistry models. Here, we propose a new approach to acquire decadal predictions of air-sea carbon flux as alternatives to those based on ocean biogeochemistry models. Our methodology combines the explanatory power of statistical models that have widely been used for gap filling purposes for informing full coverage ocean carbon flux data products, and well established predictability skill of key physical predictors. We provide hybrid GCM-statistical ocean carbon flux hindcasts using predictors from CanESM5 and doing so, show that we can beat all CMIP6 decadal prediction system hindcast skills. We use our models to provide near future hybrid model forecast for ocean carbon flux. Our results shows the potential for improving predictability skill of ocean carbon sink by combining GCMs and observationally trained statistical models.

1. Introduction

The ocean accounts for sequestering nearly 25% percent of human CO₂ emissions annually (Friedlingstein et al., 2020, 2022; Hauck et al., 2020), playing a key role in mitigating climate change. Future changes in the ocean carbon flux are of direct relevance to climate change science (Friedlingstein et al., 2022) and policy making related to climate and emissions targets. Ocean carbon uptake has increased substantially over the past several decades in response to human induced increases in atmospheric CO₂ concentrations (Gooya et al., 2023;

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Lovenduski et al., 2016; McKinley et al., 2016; Rodgers et al., 2020; L. Wang et al., 2016). However, there is also substantial internal variability in the magnitude of the flux on seasonal to decadal time scales both regionally and globally (Gruber et al., 2019; Landschützer et al., 2016; McKinley et al., 2017, 2020). Decadal scale variability of ocean carbon flux is believed to be mainly driven by variability in external forcing (McKinley et al., 2020), but also changes in circulation (DeVries et al., 2019; Keppler & Landschützer, 2019). Specifically, the deviations of atmospheric growth of CO₂ from the long term trend in the 90s left a smooth weakening (multi)decadal footprint on ocean carbon flux. Higher frequency inter-annual variability is largely attributable to modes of climate variability such as ENSO on global scale and other modes of high latitude variability on regional scales (McKinley et al., 2017). Predicting future variations in the ocean carbon sink on inter-annual to (multi)decadal time scales in the face of these multiple drivers is therefore challenging.

Decadal predictions, such as those made under the Decadal Climate Prediction Project (DCCP) are produced by Global Climate Models (GCMs) that are initialized with observations and also driven by external forcing (Kirtman et al., 2013). Predictive skill of key physical climate variables from such simulations have been well established in the literature (Boer et al., 2016). However, near term predictability of the ocean carbon flux and other biogeochemical variables have only become possible with the recent advent of Earth System Models (ESMs) (Meehl et al., 2021) and are still at their infancy. For ocean carbon flux in particular, previous studies have shown potential predictability for up to 7 years (Li et al., 2019; Lovenduski et al., 2019; Séférian et al., 2018) and actual skill versus observation-based estimates for 2–6 years based on different ESMs (Ilyina et al., 2021; Li et al., 2019). However, ESM simulations are subject to biases, drifts (Kharin et al., 2012) and exhibit a wide range of prediction skill globally and regionally (Ilyina et al., 2021). Predictions of ocean carbon flux using ESMs are especially challenging given that ocean biogeochemical variables are not directly initialized in current decadal prediction systems (Sospedra-Alfonso et al., 2021), and that the ocean biogeochemical models themselves are heavily parameterized using empirical parameterizations (Christian et al., 2022). While perfect-model studies show direct carbon cycle reconstruction adds little improvement to the global carbon cycle predictions because imperfect reconstruction of the physical climate state impedes better biogeochemical reconstruction (Fransner et al., 2020; Spring et al., 2021), the question remains whether an alternative model can leverage initialization skill in physical predictors for better predictions of ocean carbon flux?

Here we propose using observationally-trained statistical models forced by predictors from GCM/ESM-based decadal predictions, as an alternative to using the raw predictions of ocean carbon flux obtained from the ESMs ocean biogeochemistry models. It is well established that the surface ocean partial pressure of CO₂, and by extension the surface carbon flux, is closely related to physical predictors, such as sea-surface temperature and salinity, atmospheric CO₂ concentration and wind speed. These empirical relationships are widely exploited in the observational community to infill sparse direct observations of the ocean carbonate system (e.g., Surface Ocean CO₂ Atlas, SOCAT), using indirect but more widely sampled physical variables (Landschützer et al., 2016). It is also common to post-process raw GCM results to produce more skillful predictions, for example, through bias correction (Kharin et al., 2012). Our proposal is a logical extension of these two established practices that combines the explanatory power that statistical models learn from the relationships between observational predictors, and the established prediction skill of the process based physical models. Our principal goal is to establish a methodology that allows us to improve near-term predictions of the ocean carbon sink over and above the skill obtained from raw ESM predictions.

We begin by introducing the methodology and our statistical models of choice in Section 2. In Section 3 we evaluate observational uncertainties and the performance of our statistical models when forced by observation based predictors. In Section 4, we apply the observationally trained statistical models to physical predictors from CanESM5 simulations, and evaluate the skill of this hybrid approach relative to the raw CanESM5 predictions over the hindcast period of 1990–2019. We go on to provide forecasts for ocean carbon flux over the decade 2019 to 2029 in Section 5. We conclude by reflecting on how our approach could be improved and expanded on in future work.

2. Materials and Methods

2.1. Surface CO₂ Flux Data

For observations of the atmosphere-ocean CO₂ flux we use the SeaFlux Ocean carbon sink ensemble product (Fay et al., 2021). SeaFlux contains an ensemble of flux estimates, based on six global observation-based mapping

products for surface ocean partial pressure of CO₂ (pCO_2), and wind speeds from ERA5. The six products include three neural-network-derived products (CMEMS-FFNN, MPI-SOMFFN, NIES-FNN), a mixed layer scheme product (JENA-MLS), a multiple linear regression (JMA-MLR), and a machine learning ensemble (CSIR-ML6) (Fay et al., 2021). We also use the mean across the products, which we refer to as SF-MEAN. Given the sparseness of actual pCO_2 measurements, using the ensemble of products allows us to quantify uncertainties associated with the data infilling and mapping techniques, and avoids overfitting to a single product.

All six SeaFlux products show strong agreement in the long term (trended) changes in ocean carbon flux (not shown here). Comparing linearly detrended versions of the SeaFlux products shows cross correlation coefficients between them ranging from 0.47 to 0.95 (Figure S1 in Supporting Information S1). The MPI-SOM-FFN and JENA-MLS are least correlated with others. The lower correlation skills for the two show that there are variabilities specific to these products that are not common to other data sets, and known biases linked to data sparsity (Gloege et al., 2021; Hauck et al., 2023). The averaged SF-MEAN contains signals common to all of the products, and we use this as the most reliable estimate moving forward.

2.2. Statistical Models and Observed Predictors

For each individual SeaFlux input data set and SF-MEAN, we train a multi-linear regression model and a neural network (NN) model to predict the surface atmosphere ocean carbon flux, using three observation-based physical predictors - sea surface temperature (SST), sea surface salinity (SSS), surface wind speed (sfcWind), one biological predictor -surface chlorophyll concentrations (CHL), as well as atmospheric CO₂ concentrations (xCO_2) (Table S1 in Supporting Information S1). These are mainly physical predictors for which full coverage observational products are available and are believed to drive the variability in ocean carbon flux (Landschützer et al., 2016) on different time scales. Linear models are trained for each grid cell on a standard one degree grid, while the NN models are trained over 16 biomes (Landschützer et al., 2016), as explained further in SI (Sect. S1.1). By combining these biomes, we can produce spatially resolved maps of the surface CO₂ flux, given the set of five input predictors at any point. In total that gives us 14 sets of models (7 set of linear models, and 7 NN models, one for each SeaFlux target predictand) that are later used to make hindcasts and forecasts using modeled predictors from CanESM5. We have chosen to illustrate our approach using the linear and NN models, which have different structures and levels of complexity, as illustrative examples. However, alternative models and predictor variables could be used.

We did not perform any tests on the optimization of predictors, but rather follow some of the most commonly applied predictors used for global assessment studies (see Friedlingstein et al., 2023; Table S3 in Supporting Information S1). Predictors are usually selected to represent physical, chemical and biological controls of the carbon cycle and often limited to data availability (Landschützer et al., 2013, 2014). Particularly the use of chlorophyll-a as predictor has been challenged in the recent literature (Ford et al., 2022; Rödenbeck et al., 2022), however, it still represents one of the few globally available proxies for primary production, that is, an essential process describing the local draw-down of CO₂ in surface waters, and has been shown to regionally improve reconstructions of the sea surface pCO_2 variability (Zhong et al., 2022).

2.3. Decadal Predictions Using GCM Base Predictors

To make predictions, the five predictors from Table S1 in Supporting Information S1 are obtained from CanESM5 simulations (Sospedra-Alfonso et al., 2021; Swart et al., 2019). We use a range of simulations, including standard free running CMIP6 historical simulations (Eyring et al., 2016), as well as assimilation (dcppA-assim), hindcast (retrospective forecasts as in dcppA-hindcast) and forecast runs (Boer et al., 2016). In assimilation runs, CanESM5 is nudged toward observations for key physical variables (Sospedra-Alfonso et al., 2021). For historical, hindcast and forecast simulations, the five predictors are bias corrected to the same observational predictors used for training the models following the approach of (Kharin et al., 2012). This bias correction adjusts the mean and trend of the predictors to be consistent with observations. These CanESM5 predictors are fed to the each of the 14 statistical models mentioned above to produce hybrid predictions of surface ocean CO₂ flux. For hindcasts and forecasts, predictions are made for lead years 1–10. To test significance of prediction skill differences, we use a nonparametric bootstrap test (Goddard et al., 2013) to generate the probability distributions. We resampled the initialized/uninitialized simulations and the observations in time using the same sets of randomly selected indexes with replacement. The correlations are calculated using the resampled time-series with

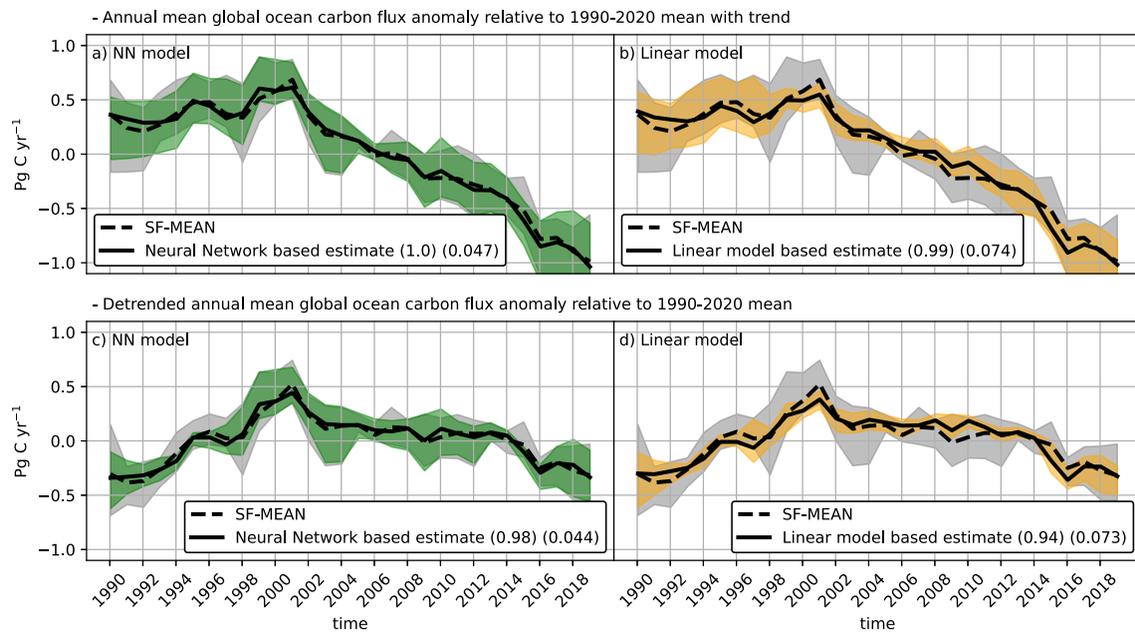


Figure 1. Time series of the global ocean CO₂ flux anomalies for the (a) NN model (left panel) and (b) linear model (right panel) reconstruction using observational predictors. The black lines shows reconstruction using models that are trained on mean of SeaFlux products (SF-MEAN; solid) as well the mean product itself (dashed). The shadings represent the range estimates from the six different SeaFlux products (gray) and from NN and linear models reconstructions (green and orange). The numbers in the legends are correlation coefficients between the solid black lines and dashed black lines (first number) and root mean square error of the two time series (second number). (c) and (d) are same as panels (a) and (b) but are linearly detrended.

the resampled observation. We repeated this process 1000 times to generate a large sample for the distribution of the predictive skill improvements. The randomly selected time steps are separated for at least 5 years to account for auto-correlation and scores are considered significant at 95% confidence limit.

3. Evaluation of Statistical Models

In this section, we consider the performance of the statistical models trained on the SeaFlux ensemble and using observed predictors, for predicting the global surface carbon flux as defined by SF-MEAN (Figure 1). When trained on SF-MEAN, both the NN and linear models can accurately reconstruct the changes of the SF-MEAN ($r > 0.9$), indicating that the statistical models are able to capture the majority of the variance in the global surface flux. The NN model shows higher skill in reconstructing SF-MEAN relative to the linear model, reflected in higher correlations and lower root mean square error (Figure 1). Similarly, both linear and NN models are able to successfully reproduce individual SeaFlux products on which they are trained (Figure S2 in Supporting Information S1), with the NN models again achieving tighter fits than the linear models. The orange and green shading in Figure 1 represents the spread across models trained on individual SeaFlux products. These models are still able to successfully reproduce SF-MEAN, which gives an indication of their generalizability. The smaller spread for the linear models (Figure 1b, orange shading), suggests they may be more generalizable (i.e., successful in predicting data they were not trained on) than the NN models. We further explore the idea of generalizability when using model-derived predictors in the following section.

4. Applying Statistical Models to Physical Predictors From the ESM

4.1. Assimilation Run

The CanESM5 assimilation run is relaxed toward the observed physical state of the system, which forces physical variables, including our input predictors, to be close to observations. However, the detrended CO₂ flux from the CanESM5 biogeochemical component is not in good agreement with observations (Figure 2 bottom row). We have identified issue in the model derived CO₂ flux, including seasonality that is out of phase with observations (not shown here), and it appears that the data ingestion in the assimilation run degrades the biogeochemical models performance. Indeed, previous results have shown that atmosphere-ocean CO₂ flux predictability is low in

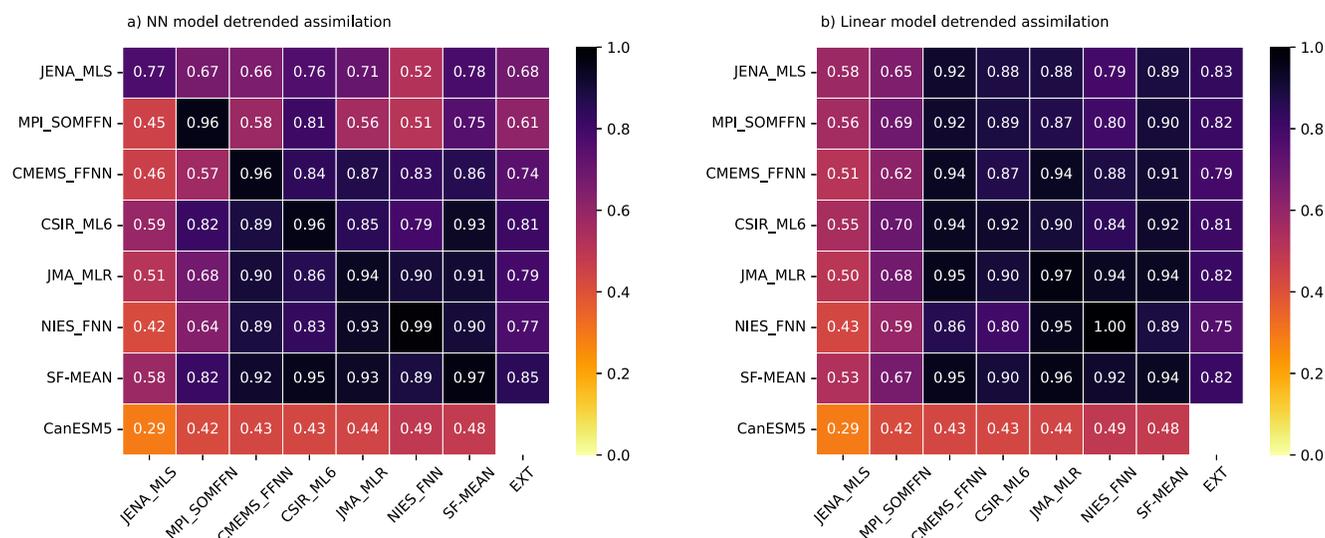


Figure 2. (a) Correlation matrix for the detrended global ocean carbon flux anomaly. The y axis indicate the product on which the NN model is trained and the x axis shows the data products against which the skill is evaluated. Predictors are chosen from assimilation runs for this test. The EXT column measures the mean of skills excluding the diagonal element for each row. (b) Same as panel (a) but for the Linear model.

CanESM5, and particularly poor in the early lead years immediately following the assimilation run (Ilyina et al., 2021). A major goal of our effort is to see whether by replacing the CanESM5 biogeochemical model derived flux with one computed based on the statistical models leads to improvement.

We use the linear and NN models previously trained using observed predictors, and for each of the six individual SeaFlux products and SF-MEAN as predictands (for a total of 14 model sets). We then extract the five input predictors from the (ensemble mean of 10) CanESM5 DCPD assimilation runs, apply the statistical models on these GCM-based predictors, and compare their skill against the original SeaFlux observational products (Figure 2).

The statistical models forced by CanESM5 assimilation predictors obtain similar skills in reproducing the individual SeaFlux products to the skills of the reconstructions that used predictors from observations (compare Figure 2 and supplementary Figure S2 in Supporting Information S1). This is a somewhat expected result given that assimilation runs assimilate physical predictors and are very close to the observations, but nonetheless it is first step in applying the models on data on which they were not directly trained. For both the linear and NN statistical models, the skill in all cases is significantly higher than skill of the raw CanESM5 CO₂ flux. These results indicate that statistical models trained on observations can usefully be applied to GCM-derived predictors. By using this approach we are able to avoid biases in the CanESM5 biogeochemical model by combining the observationally constrained statistical models with the directly initialized physical predictors from CanESM.

We compute the cross-correlation matrix for statistical models trained on one SeaFlux product in reproducing all the other five product and SF-MEAN (Figure 2). This allows us to assess the impacts of observational uncertainty, and the potential consequences of overfitting statistical models to a single observational product. As expected, the statistical models are most skillful in reproducing the product on which they were trained (diagonal in Figure 2). Correlation in reproducing other products can be lower than 0.5. The extent to which a model trained on one observational product can be generalized to others is measured with the mean of scores versus all other observational data products (mean of rows excluding the diagonal values as indicated in Figure 2 EXT column). Overall, the linear models have larger extendibility scores, while the NN models produce better fits for the products on which they were trained. Our results illustrate that care should be taken in tightly fitting statistical models to a single observation based CO₂ flux product, as uncertainties exist. Moving forward, we will use statistical model trained on the SF-MEAN product as the best estimate. Based on the encouraging success so far, in the next section we will apply our approach to decadal predictions.

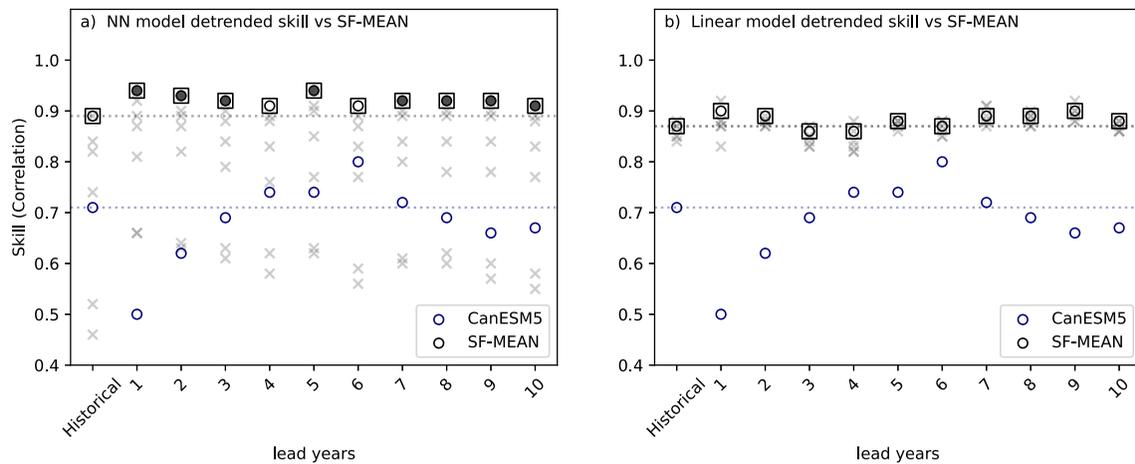


Figure 3. (a) Detrended global ocean carbon flux skills versus SF-MEAN for raw CanESM5 model (blue dots) and NN model trained on the SF-MEAN using bias corrected historical/hindcast predictors from CanESM5 (black dots). The scores that are statistically better than the raw CanESM5 score based on 1000 iteration bootstrap tests are shown with black boxes and the lead years where scores are significantly better than the corresponding historical score are filled. The gray marks in the background show scores from models trained on individual SeaFlux products versus the SF-MEAN. (b) Same as panel (a) but for the linear model.

4.2. Prediction Skill of CO₂ Flux Over the Hindcast Period

Hindcasts are ESM simulations that use the observationally constrained assimilation simulation as initial conditions, and which are then run freely under standard CMIP6 external forcings for 10 years (Boer et al., 2016). Generally, as lead years increase (i.e., number of year since initialization) the hindcasts simulations lose memory of initialization and drift toward the preferred state of the model (historical simulations). However, raw CanESM5 ocean carbon flux DCPD scores show a decrease in the skill after initialization in hindcast compared to the historical free runs (Ilyina et al., 2021). This is not the expected result of initialization and indicates possible discrepancies with interactions between initialization and the CanESM5 biogeochemical decadal prediction system (initialization “shocks”).

As an alternative to the biogeochemical model flux, we apply our SF-MEAN trained statistical models on predictors from the CanESM5 hindcast simulations over the period 1990 to 2019. The hindcast skill from both the linear and NN model when trained and evaluated against SF-MEAN are significantly larger than raw CanESM5 skills, with NN yielding slightly better scores (Figure 3). Both statistical models show increase in skill after initialization and a gradual drop with lead time, as expected and seen in physical predictors (Boer et al., 2013). Next, we compare the skill of the statistical models driven by CanESM5 predictors against the skill from all other available CMIP6 models that participated in DCPD. The NN model skill is higher than that shown by any raw CMIP6 model, when evaluated against SF-MEAN (Figure S3 in Supporting Information S1) over 1990–2017 that is the period common to all models. Linear model scores are higher than all CMIP6 models on all lead years except lead year 3 where CESM1 (Danabasoglu, 2019a, 2019b) yields slightly larger score (Figure S3 in Supporting Information S1). Analyzing individual CMIP6 ESMs skills is beyond the scope of this study. However, these results clearly show the potential of our approach for improving the decadal CO₂ flux prediction skills relative to that achieved when using result directly from the biogeochemical component of ESMs. For a comprehensive analysis of prediction skills of different ESMs refer to (Ilyina et al., 2021).

To this point we have considered the absolute skill in predicting global mean surface CO₂ flux. An important concept in decadal prediction is the relative contribution to the absolute skill that is provided by the initialization which itself can be decomposed into a forced and an unforced component (Sospedra-Alfonso et al., 2021). To assess whether initialization has added additional value to the predictions, the hindcast simulation skill can be compared to that found in standard, non-initialized CMIP6 historical simulations (Figure 3). For the linear statistical models, hindcast skills are close to the corresponding historical skill, and do not show statistically significant improvement. That is, the linear model scores do not show significant added skill due to initialization. For the NN model, the hindcast skills are significantly larger than the historical skills at least for the first 3 years, based on a bootstrapping test (Figure 3). This is the range where temperature variations largely control short term predictability of ocean carbon sink (Li et al., 2019). The NN hindcast scores are not significantly better than

historical for lead years 4–6, but show re-emergence of significance afterward. NN models consistently show better fits to the data set used for training them (Figure S2 in Supporting Information S1), but are also more subject to overfitting than the linear models (Figure 2). While more work is needed to understand difference in model structure, our results show that initialization does add value to predictions made with the NN models (see also Figure S6 in Supporting Information S1).

While bootstrapping shows improved skill compared to the historical simulations for the statistical models, high correlation skills on detrended statistical model-based historical timeseries indicate the presence of a residual signal common to historical and hindcast simulations. With atmospheric forcing being the only common predictor, this can be mainly attributed to the (multi)decadal signal imprinted by the slow down and re-acceleration in the growth rate of atmospheric CO₂ over the 90s and early 2000s (McKinley et al., 2020). To assess whether there is any skill added due to initialization for the unforced interannual variabilities, we remove this (multi)decadal signal using a locally weighted regression (Gloege et al., 2021) (Figure S4 in Supporting Information S1). Given that the (multi)decadal signal is mainly driven by external forcing, we approximate the interannual residuals from this signal as unforced. However, this approach has its own discrepancies as it could conflate low frequency internal variability with the forced signal. The results show the NN skills are higher than that of the linear models, with linear skills being lower than raw CanESM5 skills on longer lead years (Figure S5 in Supporting Information S1). This suggests that the increase in skills observed on the detrended time series could largely be attributed to the (multi)decadal signal. NN models show statistically significant improvement relative to CanESM5 on nearly all lead times as opposed to only 1 year in the Linear models. NN and linear models show statistically significant improvement in skill due to initialization on at least the first 2 years for the former and 1 year for the later. Finally, the degree to which the (multi)decadal signal in ocean carbon flux is accurately estimated in the observation based products is itself a question of ongoing research (Gloege et al., 2021; Hauck et al., 2023) which will affect these results. However, within the current accuracy on the observational side for ocean carbon flux, our results show clear improvements from statistical models as alternatives (emulators) to biogeochemistry models.

Both the hindcasts and historical runs used observed atmospheric CO₂ concentrations (as do our statistical models, as an input predictor). We expect that skills estimated from the hindcast are higher than those achievable in true forecasts, because in true forecasts the atmospheric CO₂ concentration will not be known. It is not just the background rate of increase that is relevant, but deviations in the growth rate of atmospheric CO₂ are also known to be a key driver of (multi)decadal scale variability in the ocean CO₂ sink (McKinley et al., 2020). This is an issue common to any DCP-style hindcast. Regardless, the improved skill that the statistical models driven by CanESM5 based predictors show over and above CanESM5 or other raw CMIP6 DCP model hindcast skills encourages us to apply our methods to making future predictions in the following section. First however, we turn to considering the spatial pattern of skill over the hindcast period.

We compare spatially resolved temporal correlations between SF-MEAN, the CanESM5 raw biogeochemical model, and the two statistical models for the historical, assimilation and lead years 1–10 of the hindcast experiments. Both the NN and linear models show large correlations for the detrended flux over the majority of global ocean, when driven by predictors from the CanESM5 assimilation run (Figure 4). Compared to the raw flux from the CanESM5 assimilation run, the statistical models significantly improve skill over more than 55% of the global ocean (56% for NN and 65% for linear). The linear model shows better average grid scale correlation compared to the NN model for assimilation and lead year one hindcast. This is most likely due to the high grid scale training resolution of the linear model as opposed to biome scale resolution of the NN model (see supplements). Notably, the linear models has improved skill regionally, while the skill of the globally integrated sink is better from the NN model. On longer hindcasts lead years, the mean grid scale skill for the linear models drop faster than NN model and NN model beats the linear model with small offsets and more percentage of grid cells (not shown here) with significantly improved skills.

The regions that show significant improvements relative to raw CanESM5 model include but are not limited to the highly active regions for the sink (Gooya et al., 2023) which makes them important for both the flux magnitude and uncertainty. These are regions where the largest sink is concentrated in smallest ocean surface area and where internal and model uncertainty tend to be largest. Specifically, significant improvements over the Southern Ocean is the common feature to all simulations. The Southern Ocean is of key importance for ocean carbon sink (Gruber et al., 2019) where the models disagree most (Frölicher et al., 2015; Gooya et al., 2023). The added skill in the

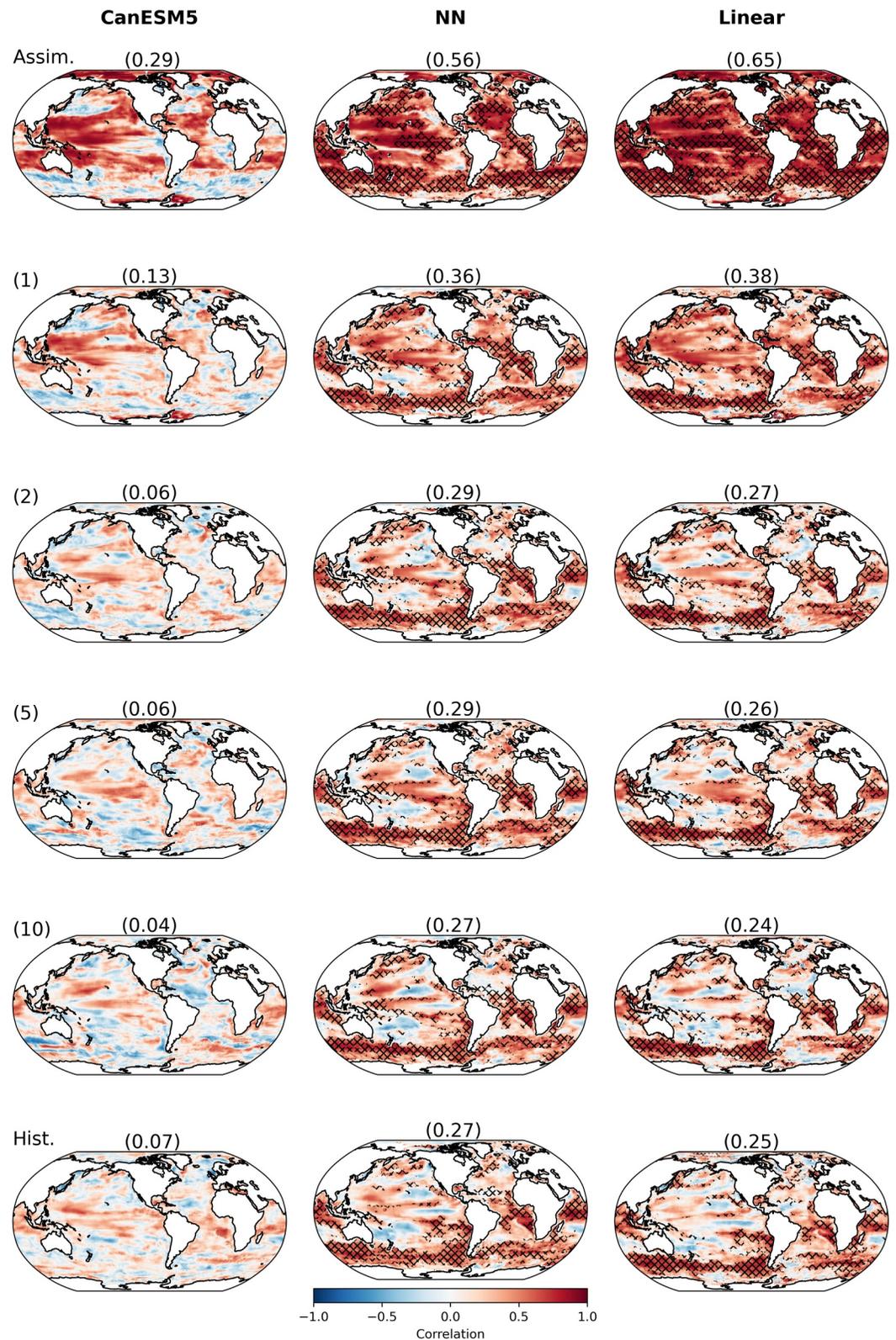


Figure 4. Grid wise correlation for detrended ocean carbon flux anomalies versus SF-MEAN. Rows show predictions using assimilation, historical as well as lead years 1, 2, 5, and 10 predictors from CanESM5. The first column shows raw CanESM5 model skills, while the second and third columns show the NN and linear model based simulations. Hatches show regions

Southern Ocean is seen in hindcast as well as historical simulations. As discussed above, the improvement thus likely is mainly driven by the improvement in representing the (multi)decadal signal, specially on longer lead years. In the hindcast simulations, skills decrease with lead year, approaching the corresponding historical simulation skill on longer lead times (>7), as expected. For all lead years there is significant improvement beyond the raw CanESM5 results regionally over more than 30% of the global ocean (hatched areas in Figure 4). Our results offer a potential pathway to better quantification of ocean carbon sink predictions both regionally and globally.

5. Hybrid Forecast of the 2020–2029 Ocean Carbon Sink

The ultimate purpose of decadal prediction systems is to provide forecasts of the short term future evolution of the climate system, including the ocean carbon flux. In this section, we use the statistical models trained on the SF-MEAN, and evaluated over the hindcast period, to make predictions for the near term evolution of ocean carbon flux. We extract ensemble means of our five predictors from CanESM5 DCPD forecasts for the period 2019–2029, and bias correct them according to lead time following (Kharin et al., 2012). We apply the statistical models on these predictors, and include the atmospheric concentration of CO_2 from SSP245 (Eyring et al., 2016), which is the same procedure applied to the hindcasts in the previous section.

Both NN model and linear model based forecasts predict that ocean carbon sink is going to grow with a faster than linear rate over the next decade under the SSP245 scenario (Figure 5). The linear model predicts slower rate of increase until 2022 compared to the NN model, and an accelerated increase after to nearly 1.29 pgC yr^{-1} relative to 2019 by 2029. The rate of change in the linear model is consistent with the rate of change of the atmospheric CO_2 concentrations under the SSP245 scenario. The NN model predicts a more steady yet faster than linear increase of approximately 1.09 pgC yr^{-1} in global ocean carbon sink relative to 2019. Both models are in close agreement regarding decadal scale changes in the flux and predict larger changes compared to the bias corrected flux from the CanESM5 biogeochemical component. The fact that the results are largely consistent between the two statistical models over 1990–2019 as well as the future forecast globally and regionally (Figure S7 in Supporting Information S1), increases our confidence in the results. Based on the skill demonstrated in the hindcasts, we assert that our hybrid statistical-GCM predictions represent a more reliable estimate of future changes in the ocean carbon flux than the raw model predictions.

6. Discussion and Conclusions

We have proposed a methodology to improve the decadal predictability of the ocean carbon flux by using statistical models as alternatives to the ocean biogeochemistry components of decadal prediction systems. Through their training, the statistical models encode the relationships between physical predictors and the surface carbon flux found in observations. Predictions are made by applying these observationally trained statistical models on (largely) physical predictors obtained from the GCM-based decadal prediction systems. Unlike biogeochemical variables, the physical variables are directly initialized in current prediction systems, have a more established track record of skill, and are based on less heavily parameterized processes than ocean biogeochemistry. In principal, our approach can be thought of as an extension of traditional bias correction (Kharin et al., 2012). Statistical bias correction schemes using linear/NN algorithms have previously been used for physical parameters in prediction system for example, (Hess et al., 2023; Sospedra-Alfonso et al., 2022; F. Wang & Tian, 2022). Unlike those, our approach uses statistical models as emulators and relies primarily on key physical predictors whose predictability have been well evaluated.

We have demonstrated that in hindcasts, our hybrid statistical-GCM system improves prediction skill for the surface ocean carbon flux relative to the ocean biogeochemical model, both in the global flux, and regionally over broad areas of the ocean. The added skill can be largely attributed to the low frequency (multi)decadal signal present on observation-based products which is believed to be mainly driven by external forcing from atmospheric CO_2 concentrations. Indeed, for the global flux, our hybrid skills based on CanESM5 predictors beat all

where there is an statistically significant improvement in skill using a 1,000 iteration bootstrap test compared to the raw CanESM5 results. The numbers on top of each panel are global mean of correlations.

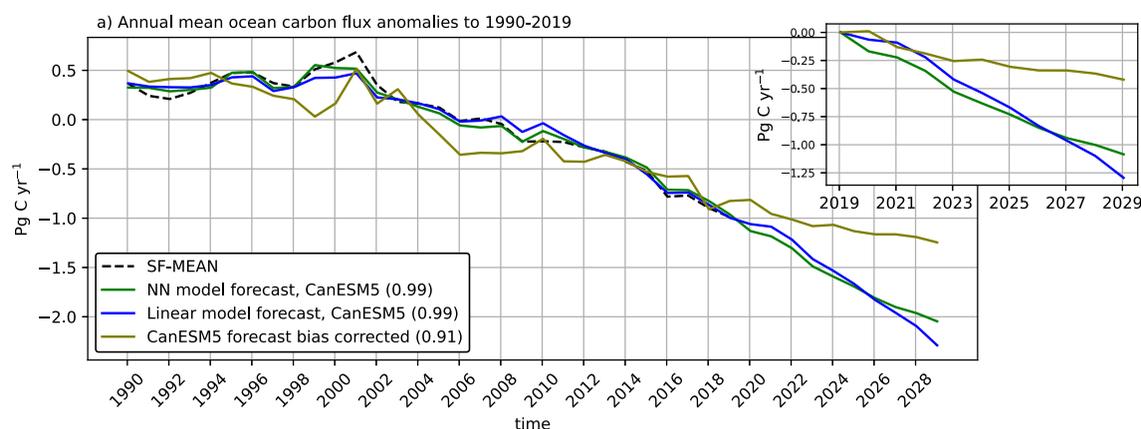


Figure 5. Global ocean carbon flux decadal forecast based on bias corrected CanESM5 (olive), NN model (green), and linear model (blue) trained on SF-MEAN. The dashed black line shows SF-MEAN over the period of 1990–2019. The forecasts show assimilation runs over this period and forecast initialized in 2019 after. The subplot shows anomalies relative to the 2019 ocean carbon flux on each product and shows the predicted changes until 2029 from different estimates. All global timeseries are scaled based on the spatial coverage of the SF-MEAN to account for differences in coverage.

available CMIP6 DCPM models. Globally, the NN model can retain the memory of initialization of the predictors at least up to lead year three after initialization.

We have demonstrated our approach using two examples of observationally constrained statistical models of different complexities; a linear and a neural network model. The two statistical models used here have different structures and use different combinations of predictors. Both statistical models are able to reconstruct observed CO₂ fluxes when forced by observed predictors, and both perform well in hindcast evaluations driven by CanESM5-based predictors (i.e., beating the skill of the raw CanESM5 flux). In general, the NN model was able to achieve higher correlations when trained and evaluated against a given surface flux product, but the linear model showed more “generalizability” across products. In addition, while the linear model was quite robust to changes in structure (predictors), the NN model was quite sensitive to changes in the number of predictors or neurons used. This shows the need for carefully adjusting such complex models and validation against other such models to avoid possible overfitting and to make reliable estimates.

We emphasize that the two statistical models we have used are just examples of our more general approach of applying observationally trained statistical models to GCM predictors. Our method is not limited to the choice of ESM, observation based product, or to the choice of the alternative model. Future work should test the ability of different types of statistical models to improve upon our results, and could draw upon the large body of work in developing empirical relationships for the purposes of infilling sparse pCO₂ observations (Fay et al., 2021). Currently, CanESM5 is the only model with sufficient number of simulations publicly available for 10-year hindcasts and forecast for all of the required predictors. More robust estimates of the future changes of ocean carbon sink would be possible with multimodel averages of predictors, since such multi-model predictions are generally more skillful (Tebaldi & Knutti, 2007). Furthermore, we would like to note that by testing and optimizing predictor data, there is a potential to further improve the prediction skill. Finally, we also note that our approach is not limited to surface ocean carbon flux, but could also be applied to other biogeochemical predictors, or even less certain physical variables that could benefit from exploiting empirical relationships based on well predicted quantities such as SST.

Based on the demonstrated skill of our hybrid approach in hindcasts, we have made forecasts of the near term evolution of ocean carbon flux using both the linear and NN models under SSP245 scenario. Both hybrid statistical models show consistent changes over the period of 2019–2029 with faster than linear increase in the sink that are larger than bias corrected CanESM5 forecasts. This information about predicted future changes in the ocean carbon sink might be useful to climate science and policy effort, for example, the assessment of the global carbon budget (Friedlingstein et al., 2022). Moving forward we encourage further research into improving decadal predictions by optimally exploiting all available observational information, and data science techniques, in conjunction with traditional GCM based predictions.

Data Availability Statement

The SeaFlux observation based ensemble is available publicly (Gregor & Fay, 2021). All model data used in this study are part of the World Climate Research Programme's (WCRP) 6th Coupled Model Intercomparison Project (CMIP6) and open-access through Earth System Grid Federation (ESGF) repositories. Observational predictors used for training the statistical models were obtained from (Lan et al., 2023) for atmospheric CO₂ concentrations, (Copernicus Climate Change Service (C3S), 2017) for surface wind speed, <https://www.ncei.noaa.gov/data/sea-surface-temperature-optimum-interpolation/v2.1/access/avhrr/> for sea surface temperature, (Good et al., 2013a, 2013b) for sea surface salinity, and (Copernicus Marine Service, 2023) for surface Chlorophyll concentrations. All other inquiries should be directed to P. Gooya.

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References

- Boer, G. J., Kharin, V. V., & Merryfield, W. J. (2013). Decadal predictability and forecast skill. *Climate Dynamics*, 41(7), 1817–1833. <https://doi.org/10.1007/s00382-013-1705-0>
- Boer, G. J., Smith, D. M., Cassou, C., Doblas-Reyes, F., Danabasoglu, G., Kirtman, B., et al. (2016). The decadal climate prediction project (DCPP) contribution to CMIP6. *Geoscientific Model Development*, 9(10), 3751–3777. <https://doi.org/10.5194/gmd-9-3751-2016>
- Christian, J. R., Denman, K. L., Hayashida, H., Holdsworth, A. M., Lee, W. G., Riche, O. G. J., et al. (2022). Ocean biogeochemistry in the Canadian earth system model version 5.0.3: CANESM5 and CANESM5-canoe. *Geoscientific Model Development*, 15(11), 4393–4424. <https://doi.org/10.5194/gmd-15-4393-2022>
- Copernicus Climate Change Service (C3S). (2017). ERA5: Fifth generation of ECMWF atmospheric reanalyses of the global climate [Dataset]. *Copernicus Climate Change Service Climate Data Store (CDS)*. Retrieved from <https://cds.climate.copernicus.eu/cdsapp#!/home>
- Copernicus Marine Service. (2023). Global ocean colour (copernicus-globcolour), bio-geo-chemical, 14 (monthly and interpolated) from satellite observations (1997–ongoing) [Dataset]. *Copernicus Marine Service*. <https://doi.org/10.48670/moi-00281>
- Danabasoglu, G. (2019a). NCAR CESM1-1-CAM5-CMIP5 model output prepared for CMIP6 DCPP [Dataset]. *Earth System Grid Federation*. <https://doi.org/10.22033/ESGF/CMIP6.11542>
- Danabasoglu, G. (2019b). NCAR CESM1-1-CAM5-CMIP5 model output prepared for CMIP6 DCPP dcppA-hindcast [Dataset]. *Earth System Grid Federation*. <https://doi.org/10.22033/ESGF/CMIP6.11552X>
- DeVries, T., Quéré, C. L., Andrews, O., Berthet, S., Hauck, J., Ilyina, T., et al. (2019). Decadal trends in the ocean carbon sink. *Proceedings of the National Academy of Sciences*, 116(24), 11646–11651. <https://doi.org/10.1073/pnas.1900371116>
- Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., & Taylor, K. E. (2016). Overview of the coupled model Intercomparison project phase 6 (CMIP6) experimental design and organization. *Geoscientific Model Development*, 9(5), 1937–1958. <https://doi.org/10.5194/gmd-9-1937-2016>
- Fay, A. R., Gregor, L., Landschützer, P., McKinley, G. A., Gruber, N., Gehlen, M., et al. (2021). Seaflux: Harmonization of air–sea CO₂ fluxes from surface pCO₂ data products using a standardized approach. *Earth System Science Data*, 13(10), 4693–4710. <https://doi.org/10.5194/essd-13-4693-2021>
- Ford, D. J., Tilstone, G. H., Shutler, J. D., & Kitidis, V. (2022). Identifying the biological control of the annual and multi-year variations in South Atlantic air–sea CO₂ flux. *Biogeosciences*, 19, 4287–4304. <https://doi.org/10.5194/bg-19-4287-2022>
- Fransner, F., Counillon, F., Bethke, I., Tjiputra, J., Samuelsen, A., Nummelin, A., & Olsen, A. (2020). Ocean biogeochemical predictions—Initialization and limits of predictability. *Frontiers in Marine Science*, 7, 386. <https://doi.org/10.3389/fmars.2020.00386>
- Friedlingstein, P., O'Sullivan, M., Jones, M. W., Andrew, R. M., Bakker, D. C. E., Hauck, J., et al. (2023). Global Carbon Budget 2023. *Earth System Science Data*, 15, 5301–5369. <https://doi.org/10.5194/essd-15-5301-2023>
- Friedlingstein, P., O'Sullivan, M., Jones, M. W., Andrew, R. M., Gregor, L., Hauck, J., et al. (2022). Global carbon budget 2022. *Earth System Science Data*, 14(11), 4811–4900. <https://doi.org/10.5194/essd-14-4811-2022>
- Friedlingstein, P., O'Sullivan, M., Jones, M. W., Andrew, R. M., Hauck, J., Olsen, A., et al. (2020). Global carbon budget 2020. *Earth System Science Data*, 12(4), 3269–3340. <https://doi.org/10.5194/essd-12-3269-2020>
- Frölicher, T. L., Sarmiento, J. L., Paynter, D. J., Dunne, J. P., Krasting, J. P., & Winton, M. (2015). Dominance of the Southern Ocean in anthropogenic carbon and heat uptake in CMIP5 models. *Journal of Climate*, 28(2), 862–886. <https://doi.org/10.1175/JCLI-D-14-00117.1>
- Gloege, L., McKinley, G. A., Landschützer, P., Fay, A. R., Frölicher, T. L., Fyfe, J. C., et al. (2021). Quantifying errors in observationally based estimates of Ocean Carbon sink variability. *Global Biogeochemical Cycles*, 35(4), e2020GB006788. <https://doi.org/10.1029/2020GB006788>
- Goddard, L., Kumar, A., Solomon, A., Smith, D., Boer, G., Gonzalez, P., et al. (2013). A verification framework for interannual-to-decadal predictions experiments. *Climate Dynamics*, 40(1), 245–272. <https://doi.org/10.1007/s00382-012-1481-2>
- Good, S. A., Martin, M. J., & Rayner, N. A. (2013a). EN4.2.1: Quality controlled ocean temperature and salinity profiles and monthly objective analyses with uncertainty estimates [Dataset]. *Journal of Geophysical Research: Oceans*. <https://doi.org/10.1002/2013JC009067>
- Good, S. A., Martin, M. J., & Rayner, N. A. (2013b). EN4: Quality controlled ocean temperature and salinity profiles and monthly objective analyses with uncertainty estimates. *Journal of Geophysical Research: Oceans*, 118(12), 6704–6716. <https://doi.org/10.1002/2013jc009067>
- Gooya, P., Swart, N. C., & Hamme, R. C. (2023). Time-varying changes and uncertainties in the CMIP6 ocean carbon sink from global to local scale. *Earth System Dynamics*, 14(2), 383–398. <https://doi.org/10.5194/esd-14-383-2023>
- Gregor, L., & Fay, A. (2021). SeaFlux: Harmonised sea-air CO₂ fluxes from surface pCO₂ data products using a standardised approach (2021.04.03) [Dataset]. *Zenodo*. <https://doi.org/10.5281/zenodo.5482547>
- Gruber, N., Clement, D., Carter, B. R., Feely, R. A., Heuven, S. v., Hoppema, M., et al. (2019). The oceanic sink for anthropogenic CO₂ from 1994 to 2007. *Science*, 363(6432), 1193–1199. <https://doi.org/10.1126/science.aau5153>
- Hauck, J., Nissen, C., Landschützer, P., Rödenbeck, C., Bushinsky, S., & Olsen, A. (2023). Sparse observations induce large biases in estimates of the global ocean CO₂ sink: An ocean model subsampling experiment. *Philosophical Transactions of the Royal Society A: Mathematical, Physical & Engineering Sciences*, 381(2249), 20220063. <https://doi.org/10.1098/rsta.2022.0063>
- Hauck, J., Zeising, M., Le Quéré, C., Gruber, N., Bakker, D. C. E., Bopp, L., et al. (2020). Consistency and challenges in the Ocean Carbon sink estimate for the global carbon budget. *Frontiers in Marine Science*, 7, 571720. <https://doi.org/10.3389/fmars.2020.571720>

- Hess, P., Lange, S., Schotz, C., & Boers, N. (2023). Deep learning for bias-correcting CMIP6-class earth system models. *Earth's Future*, 11(10), e2023EF004002. <https://doi.org/10.1029/2023EF004002>
- Ilyina, T., Li, H., Spring, A., Müller, W. A., Bopp, L., Chikamoto, M. O., et al. (2021). Predictable variations of the carbon sinks and atmospheric CO₂ growth in a multi-model framework. *Geophysical Research Letters*, 48(6), e2020GL090695. <https://doi.org/10.1029/2020GL090695>
- Keppeler, L., & Landschützer, P. (2019). Regional wind variability modulates the southern ocean carbon sink. *Scientific Reports*, 9(1), 7384. <https://doi.org/10.1038/s41598-019-43826-y>
- Kharin, V. V., Boer, G. J., Merryfield, W. J., Scinocca, J. F., & Lee, W.-S. (2012). Statistical adjustment of decadal predictions in a changing climate. *Geophysical Research Letters*, 39(19), L19705. <https://doi.org/10.1029/2012GL052647>
- Kirtman, B., Power, S. B., Adedoyin, A. J., Boer, G. J., Bojariu, R., Camilloni, I., et al., (2013). Near-term climate change: Projections and predictability.
- Lan, X., Tans, P., & Thoning, K., & NOAA Global Monitoring Laboratory. (2023). NOAA greenhouse gas marine boundary layer reference—CO₂ [Dataset]. *NOAA GML*. <https://doi.org/10.15138/DVNP-F961>
- Landschützer, P., Gruber, N., & Bakker, D. C. E. (2016). Decadal variations and trends of the global ocean carbon sink. *Global Biogeochemical Cycles*, 30(10), 1396–1417. <https://doi.org/10.1002/2015GB005359>
- Landschützer, P., Gruber, N., Bakker, D. C. E., & Schuster, U. (2014). Recent variability of the global ocean carbon sink. *Global Biogeochemical Cycles*, 28, 927–949. <https://doi.org/10.1002/2014GB004853>
- Landschützer, P., Gruber, N., Bakker, D. C. E., Schuster, U., Nakaoka, S., Payne, M. R., et al. (2013). A neural network-based estimate of the seasonal to inter-annual variability of the Atlantic Ocean carbon sink. *Biogeosciences*, 10, 7793–7815. <https://doi.org/10.5194/bg-10-7793-2013>
- Li, H., Ilyina, T., Müller, W. A., & Landschützer, P. (2019). Predicting the variable ocean carbon sink. *Science Advances*, 5(4), eaav6471. <https://doi.org/10.1126/sciadv.aav6471>
- Lovenduski, N. S., McKinley, G. A., Fay, A. R., Lindsay, K., & Long, M. C. (2016). Partitioning uncertainty in ocean carbon uptake projections: Internal variability, emission scenario, and model structure. *Global Biogeochemical Cycles*, 30(9), 1276–1287. <https://doi.org/10.1002/2016GB005426>
- Lovenduski, N. S., Yeager, S. G., Lindsay, K., & Long, M. C. (2019). Predicting near-term variability in ocean carbon uptake. *Earth System Dynamics*, 10(1), 45–57. <https://doi.org/10.5194/esd-10-45-2019>
- McKinley, G. A., Fay, A. R., Eddebar, Y. A., Gloege, L., & Lovenduski, N. S. (2020). External forcing explains recent decadal variability of the ocean carbon sink. *AGU Advances*, 1(2), e2019AV000149. <https://doi.org/10.1029/2019AV000149>
- McKinley, G. A., Fay, A. R., Lovenduski, N. S., & Pilcher, D. J. (2017). Natural variability and anthropogenic trends in the ocean carbon sink. *Annual Review of Marine Science*, 9(1), 125–150. <https://doi.org/10.1146/annurev-marine-010816-060529>
- McKinley, G. A., Pilcher, D. J., Fay, A. R., Lindsay, K., Long, M. C., & Lovenduski, N. S. (2016). Timescales for detection of trends in the ocean carbon sink. *Nature*, 530(7591), 469–472. <https://doi.org/10.1038/nature16958>
- Meehl, G. A., Richter, J. H., Teng, H., Capotondi, A., Cobb, K., Doblus-Reyes, F., et al. (2021). Initialized Earth System prediction from sub-seasonal to decadal timescales. *Nature Reviews Earth & Environment*, 2(5), 340–357. <https://doi.org/10.1038/s43017-021-00155-x>
- Rödenbeck, C., DeVries, T., Hauck, J., Le Quéré, C., & Keeling, R. F. (2022). Data-based estimates of interannual sea–air CO₂ flux variations 1957–2020 and their relation to environmental drivers. *Biogeosciences*, 19, 2627–2652. <https://doi.org/10.5194/bg-19-2627-2022>
- Rodgers, K. B., Schlunegger, S., Slater, R. D., Ishii, M., Frölicher, T. L., Toyama, K., et al. (2020). Reemergence of anthropogenic carbon into the ocean's mixed layer strongly amplifies transient climate sensitivity. *Geophysical Research Letters*, 47(18), e2020GL089275. <https://doi.org/10.1029/2020GL089275>
- Séférian, R., Berthet, S., & Chevallier, M. (2018). Assessing the decadal predictability of land and ocean carbon uptake. *Geophysical Research Letters*, 45(5), 2455–2466. <https://doi.org/10.1002/2017GL076092>
- Sospedra-Alfonso, R., Exenberger, J., McGraw, M. C., & Dang, T. K. (2022). Deep learning-based bias adjustment of decadal climate predictions. In *Neurips 2022 workshop on tackling climate change with machine learning*. Retrieved from <https://www.climatechange.ai/papers/neurips2022/102>
- Sospedra-Alfonso, R., Merryfield, W. J., Boer, G. J., Kharin, V. V., Lee, W.-S., Seiler, C., & Christian, J. R. (2021). Decadal climate predictions with the Canadian earth system model version 5 (CanESM5). *Geoscientific Model Development*, 14(11), 6863–6891. <https://doi.org/10.5194/gmd-14-6863-2021>
- Spring, A., Dunkl, I., Li, H., Brovkin, V., & Ilyina, T. (2021). Trivial improvements in predictive skill due to direct reconstruction of the global carbon cycle. *Earth System Dynamics*, 12(4), 1139–1167. <https://doi.org/10.5194/esd-12-1139-2021>
- Swart, N. C., Cole, J. N. S., Kharin, V. V., Lazare, M., Scinocca, J. F., Gillett, N. P., et al. (2019). The Canadian earth system model version 5 (CanESM5.0.3). *Geoscientific Model Development*, 12(11), 4823–4873. <https://doi.org/10.5194/gmd-12-4823-2019>
- Tebaldi, C., & Knutti, R. (2007). The use of the multi-model ensemble in probabilistic climate projections. *Philosophical Transactions of the Royal Society A: Mathematical, Physical & Engineering Sciences*, 365(1857), 2053–2075. <https://doi.org/10.1098/rsta.2007.2076>
- Wang, F., & Tian, D. (2022). On deep learning-based bias correction and downscaling of multiple climate models simulations. *Climate Dynamics*, 59(11–12), 3451–3468. <https://doi.org/10.1007/s00382-022-06277-2>
- Wang, L., Huang, J., Luo, Y., & Zhao, Z. (2016). Narrowing the spread in cmip5 model projections of air-sea CO₂ fluxes. *Scientific Reports*, 6(1), 37548. <https://doi.org/10.1038/srep37548>
- Zhong, G., Li, X., Song, J., Qu, B., Wang, F., Wang, Y., et al. (2022). Reconstruction of global surface ocean pCO₂ using region-specific predictors based on a stepwise FFNN regression algorithm. *Biogeosciences*, 19, 845–859. <https://doi.org/10.5194/bg-19-845-2022>

References From the Supporting Information

- Bethke, I., Wang, Y., Counillon, F., Kimrutz, M., Fransner, F., Samuelsen, A., et al. (2019). NCC NorCPM1 model output prepared for CMIP6 DCPD dcppA-hindcast [Dataset]. *Earth System Grid Federation*. <https://doi.org/10.22033/ESGF/CMIP6.10865>
- Boucher, O., Denvil, S., Levvasseur, G., Cozic, A., Caubel, A., Foujols, M.-A., et al. (2020). IPSL IPSL-CM6A-LR model output prepared for CMIP6 DCPD dcppA-hindcast [Dataset]. *Earth System Grid Federation*. <https://doi.org/10.22033/ESGF/CMIP6.5137>
- Fay, A. R., & McKinley, G. A. (2014). Global open-ocean biomes: Mean and temporal variability. *Earth System Science Data*, 6(2), 273–284. <https://doi.org/10.5194/essd-6-273-2014>
- Landschützer, P., Gruber, N., & Bakker, D. (2020). An observation-based global monthly gridded sea surface PCO₂ and air-sea CO₂ flux product from 1982 onward and its monthly climatology [Dataset]. *NOAA National Centers for Environmental Information*. <https://doi.org/10.7289/V5Z899N6>

- Pohlmann, H., Müller, W., Modali, K., Pankatz, K., Bittner, M., Früh, B., et al. (2019). MPI-M MPI-ESM1.2-HR model output prepared for CMIP6 DCPD dcppA-hindcast [Dataset]. *Earth System Grid Federation*. <https://doi.org/10.22033/ESGF/CMIP6.6490>
- Reynolds, R. W., Rayner, N. A., Smith, T. M., Stokes, D. C., & Wang, W. (2002). An improved in situ and satellite SST analysis for climate. *Journal of Climate*, *15*(13), 1609–1625. [https://doi.org/10.1175/1520-0442\(2002\)015<1609:aissas>2.0.co;2](https://doi.org/10.1175/1520-0442(2002)015<1609:aissas>2.0.co;2)
- Sospedra-Alfonso, R., Lee, W., Merryfield, W. J., Swart, N. C., Cole, J. N., Kharin, V. V., et al. (2019). CCCma CanESM5 model output prepared for CMIP6 DCPD dcppA-hindcast [Dataset]. *Earth System Grid Federation*. <https://doi.org/10.22033/ESGF/CMIP6.3557>