

Supplementary Information for

**Contributions of biological and physical dynamics to deglacial CO₂ release from
the polar Southern Ocean**

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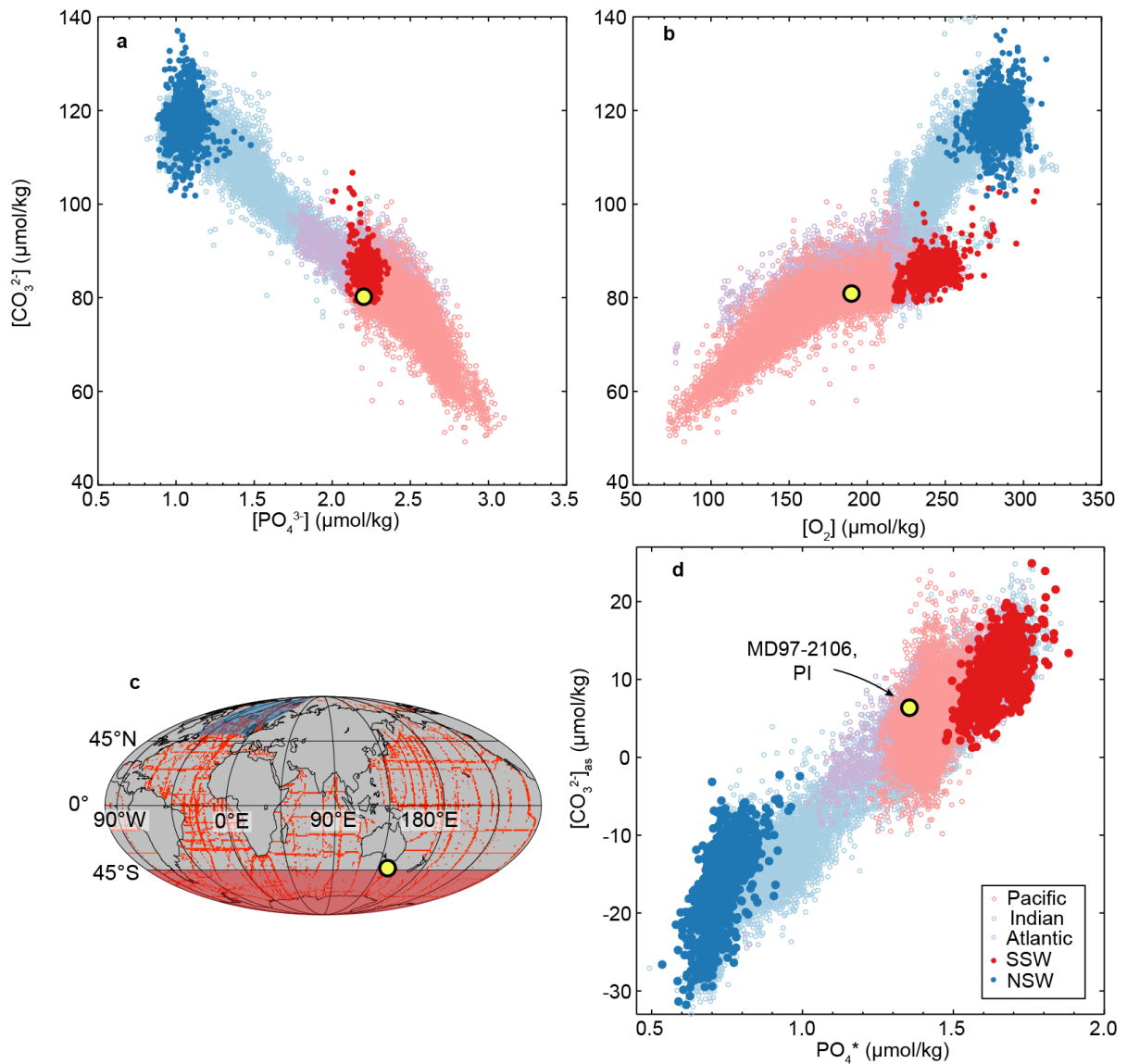
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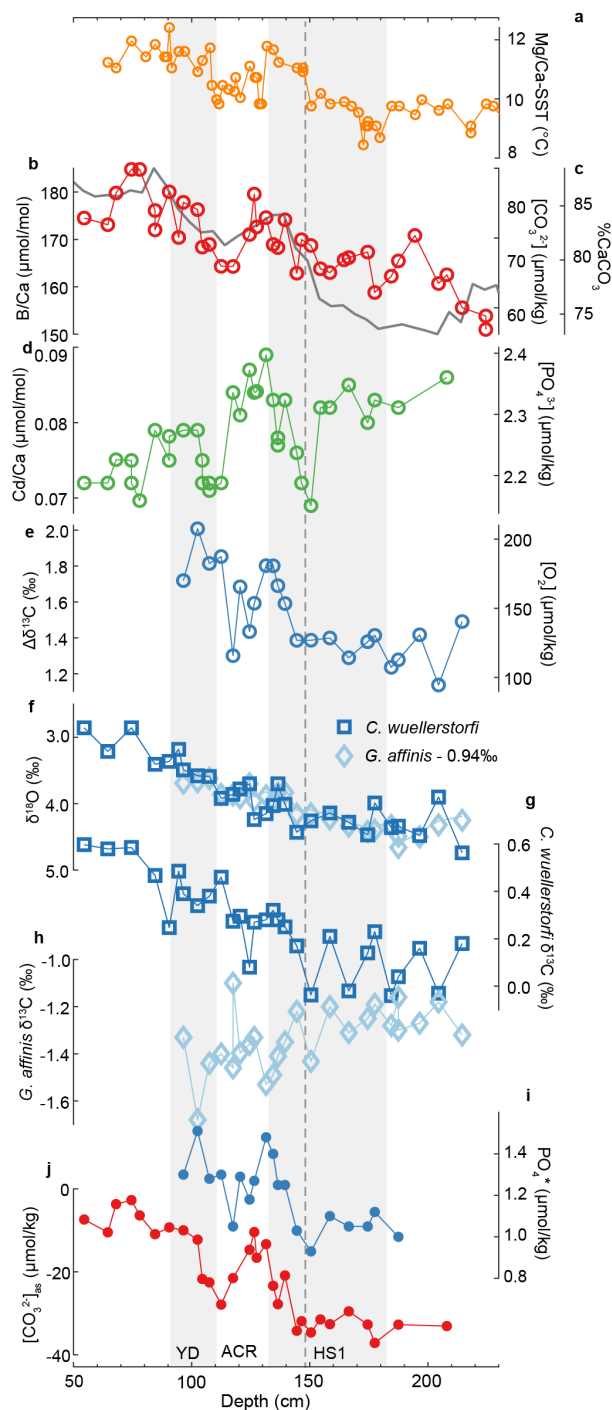
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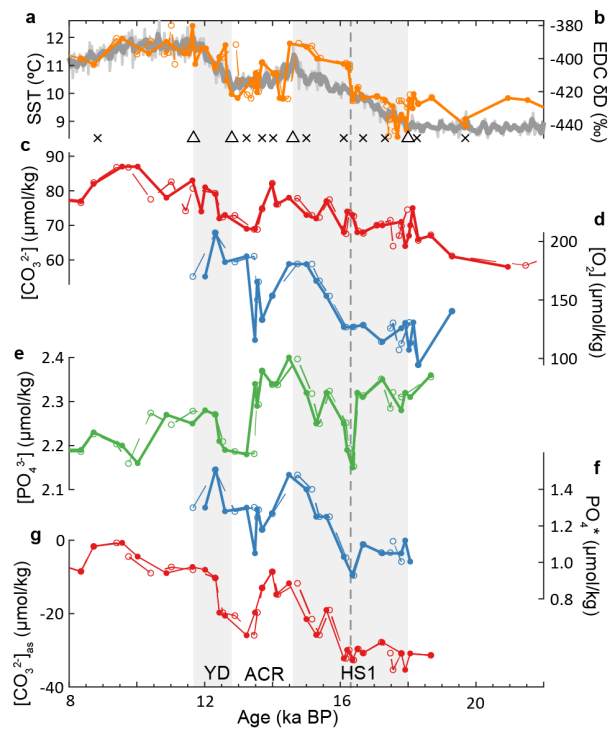
Supplementary Figures 1 to 10



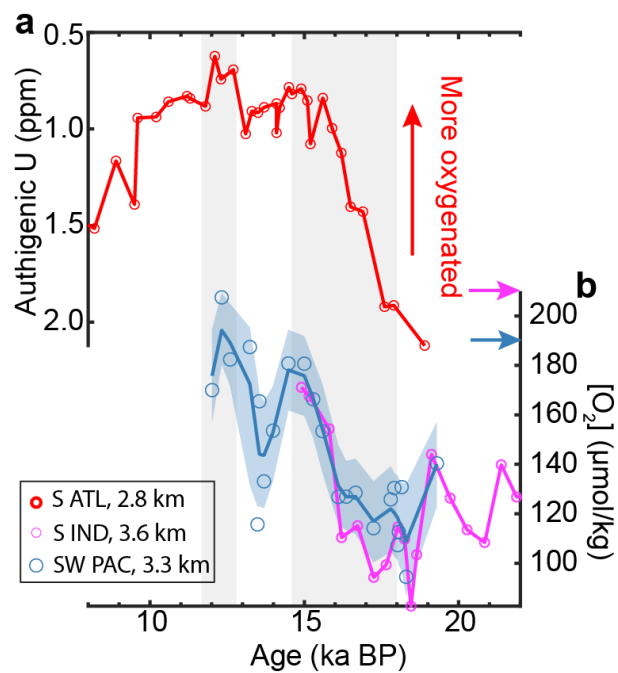
Supplementary Fig. 1. Preindustrial deep ocean biogeochemistry. **a**, $[\text{PO}_4^{3-}]$ versus $[\text{CO}_3^{2-}]$, **b**, $[\text{O}_2]$ versus $[\text{CO}_3^{2-}]$, **c**, Map showing hydrographic locations, **d**, PO_4^* versus $[\text{CO}_3^{2-}]_{\text{as}}$. Data are from GLODAPv2^{1,2}. Data in the Pacific, Indian, and Atlantic oceans are from >2500 m water depths. The Northern Sourced Water (NSW) and Southern Sourced Water (SSW) are defined following Ref. ³. The yellow circles in **a**, **b**, and **d** show the preindustrial deep-water biogeochemistry at site MD97-2106. The yellow circle in **c** indicates the location of Site MD97-2106. Red and blue shading regions in **c** highlight regions where SSWs and NSWs, respectively, are defined.



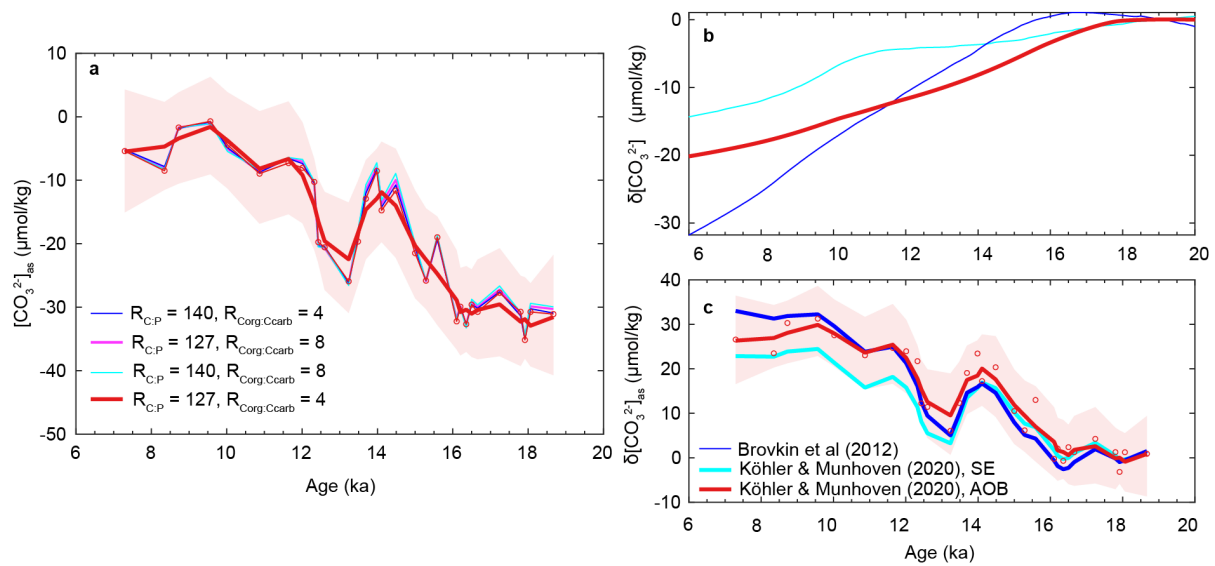
Supplementary Fig. 2. Proxy data against core depth at MD97-2106. **a**, Sea surface temperature (SST) estimated from Mg/Ca in *G. bulloides*⁴. **b**, *C. wuellerstorfi* B/Ca and the estimated $[\text{CO}_3^{2-}]$. **c**, Sedimentary carbonate content (grey curve)⁵. **d**, *H. elegans* Cd/Ca and the estimated $[\text{PO}_4^{3-}]$. **e**, $\delta^{13}\text{C}$ gradient between *C. wuellerstorfi* and *G. affinis* and estimated $[\text{O}_2]$. **f**, $\delta^{18}\text{O}$ in *C. wuellerstorfi* (dark blue) and *G. affinis* (light blue), where $\delta^{18}\text{O}$ in *G. affinis* is subtracted by 0.94‰ to account for the interspecies offset. **g**, *C. wuellerstorfi* $\delta^{13}\text{C}$. **h**, *G. affinis* $\delta^{13}\text{C}$. **i**, PO_4^{3-} . **j**, $[\text{CO}_3^{2-}]_{\text{as}}$. In **b** and **d**, replicated measurements of subsamples are shown as different data points. Grey vertical bars indicate the depth ranges corresponding to the Younger Dryas (YD) and Heinrich Stadial 1 (HS1). Between these two periods is the Antarctic Cold Reversal (ACR). The vertical dashed line shows the depth corresponding to the boundary between early and late HS1.



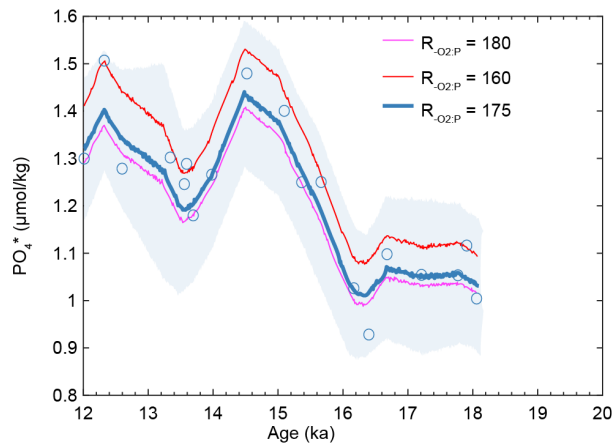
Supplementary Fig. 3. Influences of age control points on records at site MD97-2106. **a**, Sea surface temperature (SST)⁴. **b**, EPICA Dome C (EDC) δD , a proxy for Antarctic temperature)⁶. Triangles at the bottom are age tie points for SST at site MD97-2106 and EDC δD , and crosses represent ¹⁴C dates^{4,7}. **c**, $[CO_3^{2-}]$; **d**, $[O_2]$; **e**, $[PO_4^{3-}]$; **f**, PO_4^* ; **g**, $[CO_3^{2-}]_{as}$. In **a**, **c-g**, solid curves with solid dots are the same as shown in the main text. **a**, **c-g**, dashed curves with empty dots are records based on an alternative age model relying only on radiocarbon dates.



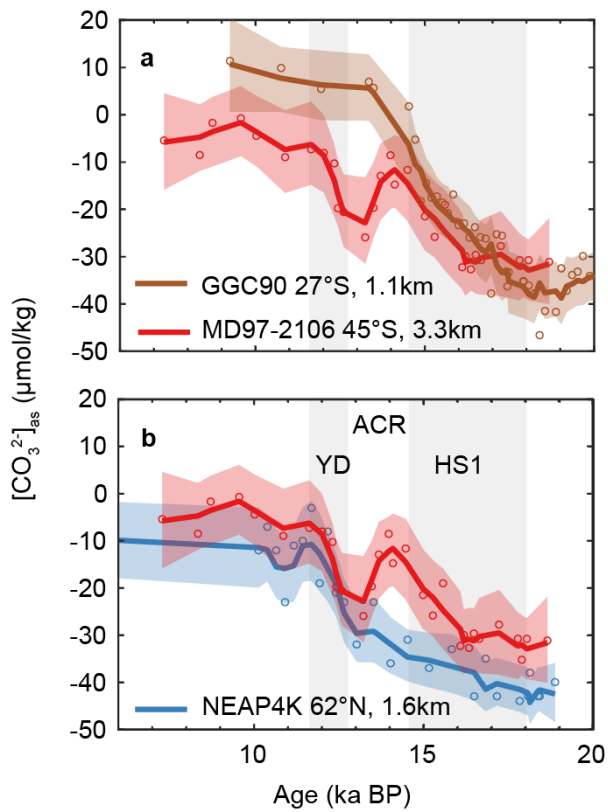
Supplementary Fig. 4. Comparison of deep-water [O₂] at site MD97-2106 to other deep Southern Ocean oxygenation reconstructions. a, South Atlantic Ocean oxygenation based on authigenic uranium⁸ (red circles). **b**, South Indian Ocean oxygenation⁹ based on benthic carbon isotope gradients (magenta circles), compared to [O₂] at site MD97-2106 (blue circles with error envelope; this study).



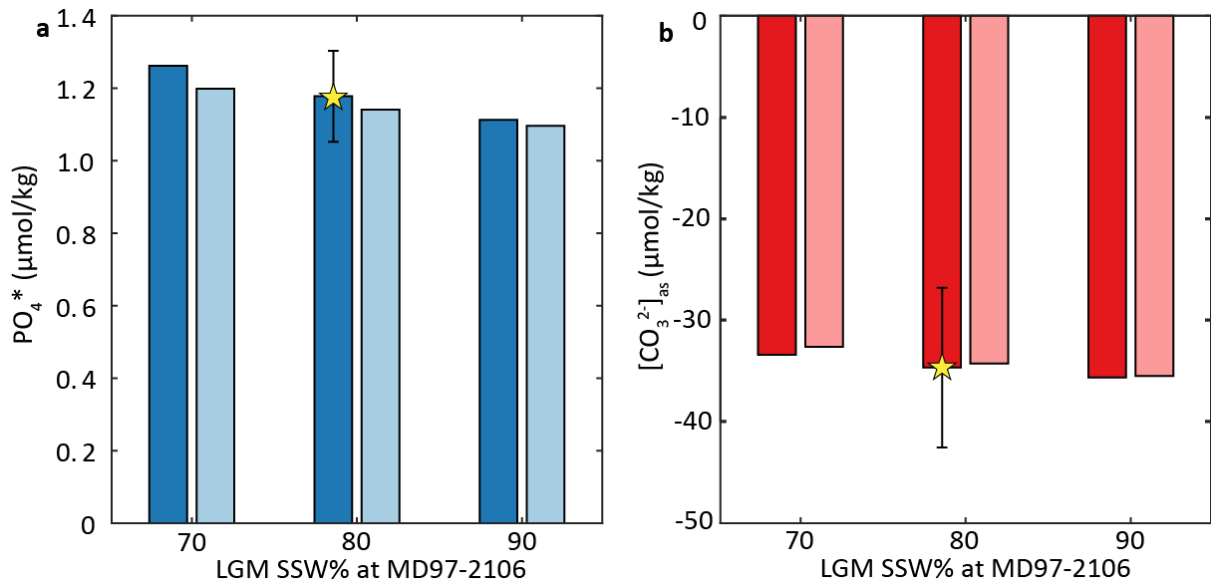
Supplementary Fig. 5. Sensitivity tests for deglacial $[\text{CO}_3^{2-}]_{\text{as}}$ at site MD97-2106. a, $[\text{CO}_3^{2-}]_{\text{as}}$ calculated using various sensitivities to evaluate effects from changes in organic matter composition ($R_{\text{C:P}}$) and rain ratio ($R_{\text{Corg:Ccarb}}$)¹⁰. Red curve with envelope is calculated using $R_{\text{C:P}} = 127$ and $R_{\text{Corg:Ccarb}} = 4$ shown in the main text. **b,** Effects of the global alkalinity change on deep Southern Ocean $[\text{CO}_3^{2-}]$, assuming no air-sea CO_2 exchange from three different simulations^{11,12}. **c,** changes in $[\text{CO}_3^{2-}]_{\text{as}}$ relative to the LGM calculated using three global alkalinity histories^{11,12}.



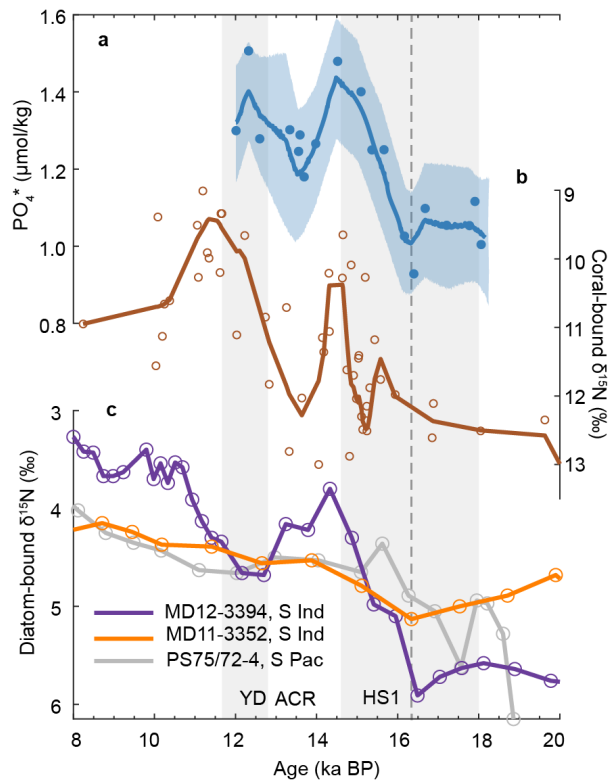
Supplementary Fig. 6. Deglacial PO_4^* derived from different Redfield ratios ($R_{\text{O}_2:\text{P}}$) at site MD97-2106. The blue curve with circles and error envelope is based on a $R_{\text{O}_2:\text{P}}$ of 175 following the definition in Ref. ^{13,14}. Red and magenta curves are calculated using $R_{\text{O}_2:\text{P}}$ of 160 and 180, respectively, which encompasses the entire deep-water $R_{\text{O}_2:\text{P}}$ variability in the modern oceans¹⁵.



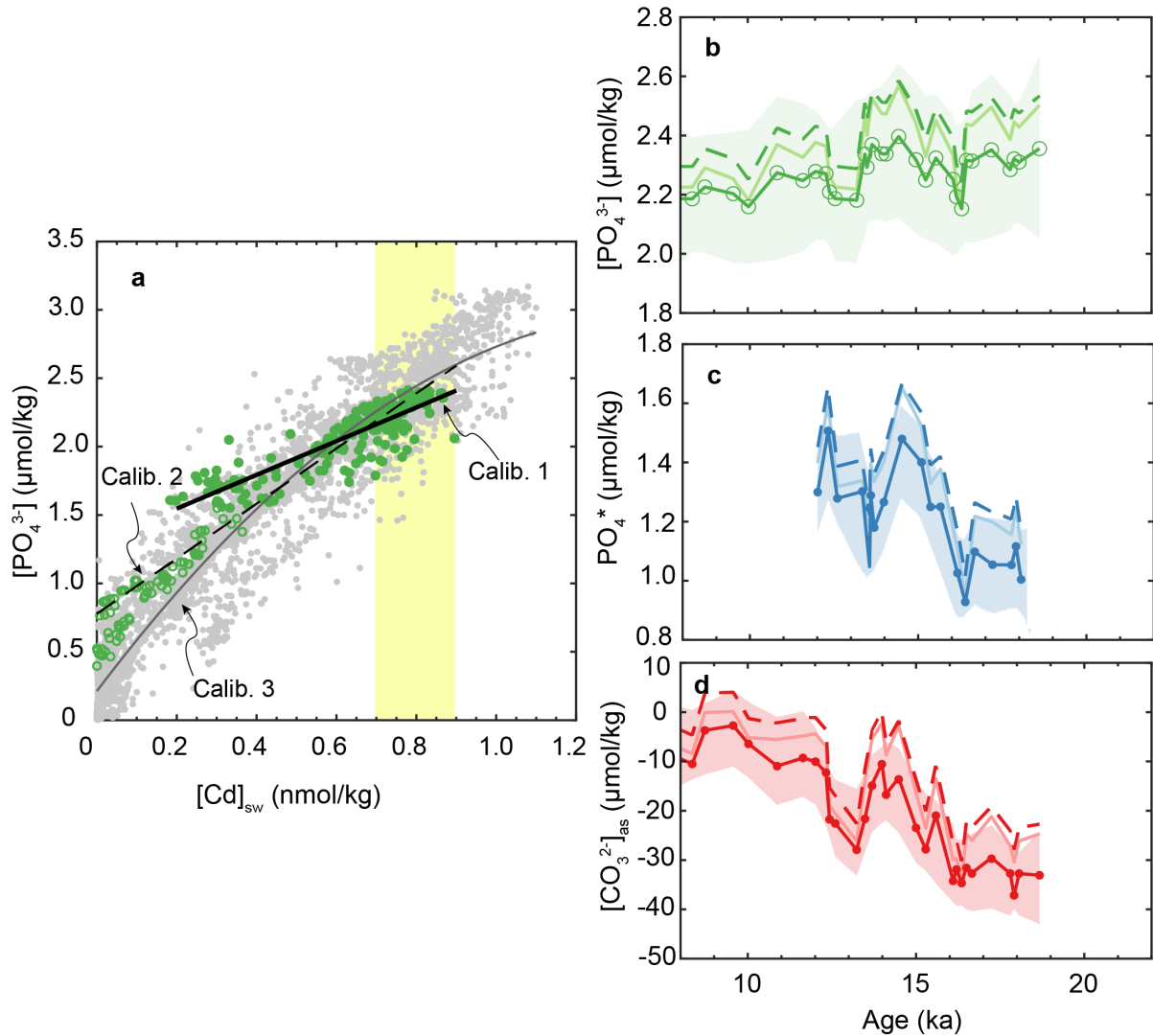
Supplementary Fig. 7. Comparisons of MD97-2106 $[\text{CO}_3^{2-}]_{\text{as}}$ with published records¹⁶. **a**, Results at MD97-2106 (this study) compared with those at core GGC90 (South Atlantic) bathed in the downstream of AAIW¹⁶. **b**, Results at core NEAP 4K bathed in the North Atlantic Deep Water¹⁶ and at MD97-2106 (this study).



Supplementary Fig. 8. Sensitivity tests for LGM SSW PO_4^* (a) and $[CO_3^{2-}]_{as}$ (b). LGM SSW PO_4^* and $[CO_3^{2-}]_{as}$ are calculated based on varying SSW% (70%-90%) at site MD97-2106. In **a**, bars in dark and light blue, respectively, show LGM SSW PO_4^* calculated assuming 20% decline and no decline in the NSW PO_4^* during the LGM compared to the late Holocene. In **b**, bars in dark and light red, respectively, show LGM SSW $[CO_3^{2-}]_{as}$ calculated assuming 0% and 10% SSW at the NSW-dominated site NEAK-4K during the LGM compared to the late Holocene. Yellow stars highlight the preferred calculation discussed in the main text. Error bars show the $\pm 1\sigma$ uncertainties associated with the reconstruction of the LGM SSW endmember values. These sensitivity tests show that LGM SSW PO_4^* and $[CO_3^{2-}]_{as}$ calculated here are insensitive to LGM SSW% at site MD97-2106 and NSW endmember values.



Supplementary Fig. 9. Comparison of deep-water PO_4^* at MD97-2106 with fossil-bound $\delta^{15}N$ records from the Southern Ocean. a, Deep-water PO_4^* at site MD97-2106. b, Coral-bound $\delta^{15}N$ from the Drake Passage^{17,18}. c, Diatom-bound $\delta^{15}N$ from core located near the modern Antarctic Polar Front^{19,20}. Despite weaker age constraints than coral-bound $\delta^{15}N$ records, diatom-bound $\delta^{15}N$ data show little change during early HS1 as revealed by the coral-bound $\delta^{15}N$ data. $\delta^{15}N$ data suggest reduced nutrient utilization and thus a weaker biological pump in the PAZ during late HS1, consistent with PO_4^* at site MD97-2106.



Supplementary Fig. 10. Sensitivity test of [Cd]_{sw}-[PO₄³⁻] calibrations on our reconstructions. **a**, modern seawater [Cd]_{sw} and [PO₄³⁻] data from the GEOTRACES dataset²¹. Empty green dots are from GIPY06 Section south of Tasmania and filled green dots are sites with [PO₄³⁻] > 1.5 μmol/kg. Calibrations 1 and 2 are linear calibrations based on GIPY06 data with [PO₄³⁻] > 1.5 μmol/kg and all GIPY06 data, respectively. Calibration 3 is a quadratic calibration based on all GEOTRACES data. The yellow vertical bar highlights the [Cd]_{sw} range reconstructed at site MD97-2106. **b**, Reconstructed [PO₄³⁻]; **c**, PO₄*; and **d**, [CO₃²⁻]_{as}. In **b-d**, bold lines with dots and shadings are reconstructions based on calibration 1 shown in the main text, while light-coloured lines and dashed lines are reconstructions based on calibrations 3 and 2, respectively. Different calibrations yield similar patterns of [PO₄³⁻], PO₄*, and [CO₃²⁻]_{as}, despite their varying absolute values.

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