



# Fine scale zooplankton distribution across the North Balearic Front during late spring

Maxime Duranson<sup>1,2</sup>, Léo Berline<sup>1</sup>, Loïc Guilloux<sup>1,3</sup>, Alice Della Penna<sup>4,5</sup>, Mark D. Ohman<sup>6</sup>, Sven Gastauer<sup>2</sup>, Cédric Cotte<sup>7</sup>, Daniela Banaru<sup>1</sup>, Théo Garcia<sup>8</sup>, Maristella Berta<sup>9</sup>, Andrea Doglioli<sup>1</sup>, Gérald Gregori<sup>1</sup>, Francesco D'Ovidio<sup>7</sup>, and François Carlotti<sup>1</sup>

<sup>1</sup>Université Aix Marseille, Université de Toulon, CNRS, IRD, MIO, Marseille, France

<sup>2</sup>Thünen Institute of Sea Fisheries, Bremerhaven, Germany

<sup>3</sup>CNRS, Univ Brest, IRD, IFREMER, LEMAR, IUEM, F-29280 Plouzané, France

<sup>4</sup>Institute of Marine Science, University of Auckland, New Zealand

<sup>5</sup>School of Biological Sciences, University of Auckland, New Zealand

<sup>6</sup>Integrative Oceanography Division, Scripps Institution of Oceanography, La Jolla, California, USA

<sup>7</sup>Sorbonne Université, CNRS, IRD, MNHN, Laboratoire d'Océanographie et du Climat: Expérimentations et Approches Numériques (LOCEAN-IPSL), Paris, France.

<sup>8</sup>Aix Marseille University CNRS, Centrale Marseille, I2M, Marseille, France

<sup>9</sup>Istituto di Scienze Marine- Consiglio Nazionale delle Ricerche (ISMAR-CNR), Sede Secondaria di Lerici, 19032 La Spezia, Italy

Correspondence to: Maxime Duranson (maxime.duranson@thuenen.de), François Carlotti (françois.carlotti@mio.osupytheas.fr)

# Keywords

BioSWOT-Med campaign, Northwestern Mediterranean Sea, Ocean Fronts, Vertical distributions, Copepod Community, Size,

0 Trophic categories

# Abstract

Observations, models and theory have suggested that ocean fronts are ecological hotspots, generally associated with higher diversity and biomass across many trophic levels. Nutrient injections are often associated with higher chlorophyll concentrations at fronts, but the response of the zooplankton community is largely unknown. The present study investigates mesozooplankton

- 5 stocks and composition during late spring, northeast of Menorca along two north-south transects that crossed the North Balearic Front (NBF) separating central water of the Northwestern Mediterranean Sea (NWMS) gyre from peripheral waters originating from the Algerian basin. During the BioSWOT-Med campaign, samples were collected using vertical triple net tows at three depths (100 m, 200 m, 400 m) with 200  $\mu$ m and 500  $\mu$ m mesh nets, processed with ZooScan, and the organisms assigned to eight taxonomic groups. Zooplankton distributions were analysed for the surface (0–100 m), intermediate (100–200 m),
- 10 and deeper (200–400 m) layers. The results showed no significant biomass increase at the front across all vertical layers. The NBF seems to act more like a boundary between communities rather than a pronounced area of active or passive zooplankton accumulation. Analysis of stratified vertical distributions of zooplankton highlighted distinct taxonomic compositions in the





surface, intermediate, and deeper layers, and a progressive homogenization of community structure with depth, reflecting a weaker impact of hydrological processes on deeper communities. The front's clearest impact was within the upper 100 meters,

- 15 where the taxonomic composition showed differences between the front and the adjacent water masses, with a decrease in Eumalacostraca and Foraminifera, while Cnidaria increased sharply. In the 100–200 m layer, the front also influenced community composition, although to a lesser extent, with a marked increase in Foraminifera and a strong decline in Cnidaria. Moreover, the northern water mass and the front were dominated by large copepods, while the southern water mass exhibited higher diversity of zooplankton and a community of smaller-sized copepods. The results of this study highlight the complexity of pro-
- 20 cesses shaping planktonic communities over time and space in the NBF zone and its adjacent waters. These processes include zooplankton stock reduction in the transitional post-bloom period, marked effect of diel variation linked to vertical migrations, and potentially the impact of storm-related mixing in the surface layer that can disrupt established ecological patterns.

# **1** Introduction

35

Fronts are key structural features of the ocean, affecting all trophic levels across a wide range of spatial and temporal scales
(Belkin et al. (2009)). Oceanic fronts are narrow regions of elevated physical gradients that separate water parcels with distinct properties, such as temperature, salinity, and thus density (Hoskins (1982); Joyce (1983); Pollard and Regier (1992); Belkin and Helber (2015)). These frontal zones act as dynamic boundaries between distinct water masses (Ohman et al. (2012); Mańko et al. (2022)), which play a crucial role in shaping marine ecosystems (Belkin et al. (2009)). Moreover, fronts display wide variations in spatial and temporal dimensions ranging from hundreds of meters to tens of kilometres, and from short-lived to permanent (Owen (1981); McWilliams (2016); Lévy et al. (2018)).

The relationships between fronts and plankton have received considerable attention in marine ecology due to the enhanced biological production and community changes that are sometimes observed in their vicinity (Le Fèvre (1987); Fernández et al. (1993); Pinca and Dallot (1995); Errhif et al. (1997); Pakhomov and Froneman (2000); Chiba et al. (2001); Munk et al. (2003)). As physical barriers or zones of mixing, fronts structure biomass and species distributions, generally leading to distinct ecological communities on either side (Ohman et al. (2012); Le Fèvre (1987); Prieur and Sournia (1994); Gastauer and Ohman (2024)). Fronts are also known to concentrate high phytoplankton concentration and, as a result, higher biomass levels and metabolism of zooplankton (Thibault et al. (1994); Ashjian et al. (2001); Ohman et al. (2012); Derisio et al. (2014); Powell and Ohman (2015a)). Frontal structures can enhance primary and secondary production essentially by promoting nutrient input through cross-frontal mixing and vertical circulation driven by horizontal density gradients (Durski and Allen (2005); Liu

40 et al. (2003); Derisio et al. (2014); Russell et al. (1999). These nutrient-rich conditions (bottom-up) sometimes fuel elevated chlorophyll concentrations, supporting the aggregation of zooplankton, fish larvae, and their higher trophic level predators such as tuna, sharks, seabirds and whales (Herron et al., 1989Herron et al. (1989); Olson et al. (1994); Royer et al. (2004); Queiroz et al. (2012); di Sciara et al. (2016); Druon et al. (2019)). Pronounced changes in zooplankton Diel Vertical Migration behaviour have also been observed across frontal gradients (Powell and Ohman (2015b), Gastauer and Ohman (2024)).





- 45 Recent studies (Mangolte et al. (2023); Panaïotis et al. (2024)) have highlighted the importance of investigating zooplankton at fine scales and their patchiness in the vicinity of fronts to understand their interactions with particles and the environment. Mangolte et al. (2023) revealed that the plankton community exhibits fine-scale variability across fronts, with biomass peaks of different taxa often occurring on opposite sides of the front, or with varying widths. This fine-scale cross-frontal patchiness suggests processes that spatially decouple plankton taxa, leading to the formation of multiple adjacent communities rather than
- 50 a single coherent frontal plankton community.

In the Northwestern Mediterranean Sea (NWMS), the role of mesoscale structures in the open ocean such as density fronts and eddies on the distribution and diversity of zooplankton has already been widely documented (Saiz et al. (2014)). These structures generally increase the patchiness and activity of plankton, and stimulate trophic transfers to large predators (Cotté et al. (2009), Cotté et al. (2011)). Due to its coastal proximity, the North Current's frontal zones have been widely studied

- 55 from physical and ecosystem perspectives, on both the Ligurian (Prieur et al. (1983); Stemmann et al. (2008)) and Catalan sides (Font et al. (1988); Sabatés et al. (2007)). Downstream of the Northern Current, the North Balearic current flows from northeast Menorca to southwest Corsica. This current is associated with the North Balearic Front (NBF), which marks the transition zone between the saltier surface waters from the Provencal basin at its northern edge, and the less salty surface waters from the Algerian basin at its southern edge. To the south is the Atlantic water that has recently penetrated the Strait of
- 60 Gibraltar, while to the north is the Atlantic water that has circulated in the western basin and has consequently become saltier due to evaporation typical of the Mediterranean Sea (Millot and Taupier-Letage (2005)). Recent contributions from glider data and satellite imagery have enabled us to better characterize the NBF (Barral (2022)). Its latitudinal position shows seasonal (from 40.2°N in spring to 41°N in autumn) and interannual variations, notably linked to the intensity and extension of the winter deep convection in the Provençal basin to the north, and to the south, where, mainly in summertime, mesoscale dynamics
- 65 spread lighter Atlantic water northward between Menorca and Sardinia (Millot (1999); Seyfried et al. (2019)). Among the most pronounced geostrophic frontal zones in the NWMS, the NBF and its ecological impacts are the least studied.

The BioSWOT-Med cruise, focused on the NBF, involved a wide range of sampling and data collection, including zooplankton sampling using nets. In this study, the following questions were addressed: how are zooplankton communities structured on each side of the front; is the zooplankton community at the front a mixture of communities from both sides, or does it form

70 a distinct community; does the front influence the vertical structure of zooplankton communities; and can weather events, such as storms, influence the structure of zooplankton communities within the water masses?

# 2 Materials and methods

### 2.1 Study area and sample strategy

The BioSWOT-Med cruise (https://doi.org/10.13155/100060; PIs: A. Doglioli and G. Grégori) was performed on board the 75 R/V L'Atalante (FOF-French Oceanographic Fleet) from 21 April to 14 May 2023 in an area about 100 km north-east of





Menorca Island (NWMS) (Fig. 1). Figure 1.a shows the zone as observed four days before the first transect due to cloud cover during this part of the survey (Fig. A1).



**Figure 1.** Maps of the sampling stations with surface chlorophyll concentration ( $\mu$ g L-1) from Sentinel 3. a) Map from April 21 showing conditions 4 days before the first transect. b) Map from May 5 showing conditions during the second transect. The colours representing the three water masses and the front will be maintained throughout the paper.

80

85

The strategy of the cruise was designed to take advantage of the novel SWOT satellite mission, in order to better resolve fine-scale oceanic features. During the "fast sampling phase," SWOT provided altimetry data characterized by high spatial resolution and a 1-day revisit period over 150 km-wide oceanic regions. With the support of the international SWOT AdAC (Adopt-A-Crossover, https://www.swot-adac.org/; PI F.d'Ovidio) Consortium, the BioSWOT cruise applied an adaptive multi-disciplinary approach by combining daily SWOT images and environmental bulletins provided daily by the SPASSO toolbox (https://spasso.mio.osupytheas.fr/, Rousselet et al., submitted), and in situ measurements using a suite of instruments to capture physical, biological, and chemical properties (Doglioli et al. (2024), cruise report). This strategy enabled the targeting of fine-scale features of the North-Balearic front.

This zone separates two water masses: the northern water mass, referred to as "A", which is more productive, colder, and characterised by higher salinity, and the southern mass, referred to as "B", which is less productive, warmer, and less saline (Fig. 1). This permitted us to sample two contrasting water masses separated by a sharp frontal region, designated "F". The

sampling strategy focused on two main transects, covering a total of six 24-hour drifting stations: A-F-B (to the west) and
B-F-A (to the east), in addition to three supplementary stations: *b3*, *m*, and *m2* (Fig. 1 and Table 1). The two f2 stations are relatively distant from each other due to a strong frontal current. Because of a storm (2nd May), a third area, "M", was sampled twice while the ship took shelter south of Menorca and conducted similar measurements as in zones A, B, and F. The M zone



95



is different from the three other sampled zones in terms of bathymetry, as it is located around 20 km from the continental shelf. On the way home, a final station was sampled in B (Table 1). At every station, physical properties were sampled using a CTD rosette, which was deployed four times daily at fixed intervals (06:00, 12:00, 18:00, and 00:00 local time). Hereafter, water masses will be designated by an uppercase letter (A, B, M, and F for the front), and stations by a lowercase letter (a, b, m, f).

Table 1. Stations details. In Abbreviation, 'D' stands for Day and 'N' stands for Night.

Campaign Stage	Station Name	Abbreviation	Date - Time	Latitude	Longitude
	a1_zoo1	a1_D	25/04 - 12:38	41.240	4.553
1st transect	a1_zoo2	a1_N	26/04 - 00:02	41.224	4.563
	f1_zoo1	f1_N	27/04 - 00:32	41.099	4.423
	f1_zoo2	f1_D	27/04 - 12:11	41.102	4.456
	b1_zoo1	b1_N	28/04 - 00:17	40.874	4.388
	b1_zoo2	b1_D	28/04 - 12:28	40.884	4.389
Storm	m_zoo1	m_N	02/05 - 00:37	39.555	4.101
	m_zoo2	m_D	02/05 - 12:22	39.493	4.087
	b2_zoo1	b2_D	04/05 - 12:16	40.795	4.933
	b2_zoo2	b2_N	05/05 - 00:13	40.849	4.936
2nd transact	f2_zoo1	f2_D	05/05 - 11:49	41.175	5.108
2nd transeet	f2_zoo2	f2_N	06/05 - 00:45	41.134	5.308
	a2_zoo1	a2_N	07/05 - 00:13	41.412	5.240
	a2_zoo2	a2_D	07/05 - 12:15	41.376	5.253
Return water mass M	m2_zoo1	m2_D	10/05 - 11:31	39.671	3.957
	m2_zoo2	m2_N	11/05 - 00:31	39.629	3.902
	m2_zoo3	m2_D'	11/05 - 11:53	39.603	3.885
Return water mass B	b3_zoo1	b3_D	12/05 - 12:15	40.782	5.152
	b3_zoo2	b3_N	12/05 - 23:58	40.746	5.112

# 2.2 Zooplankton collection

100

Zooplankton samples were collected using a triple net (Triple-WP2) equipped with three individual nets, each with a 60 cm mouth diameter but different mesh sizes (500  $\mu$ m, 200  $\mu$ m, 64  $\mu$ m). For this study, which focuses on mesozooplankton, only the samples collected by 200 and 500  $\mu$ m nets were used. The nets were deployed vertically to cover three integrated layers (400-0 m; 200-0 m; 100-0 m). Sampling occurred twice daily at local times: 12:00 and 00:00. Note that the net deployed to 400 m at station  $m_N$  (Table 1) could not be analysed because it was found folded up on itself upon retrieval. After collection, samples were preserved in 4% borate-buffered formaldehyde.

# 2.3 Zooplankton sample processing

105 In a shore-based laboratory, samples were digitized with the ZooScan digital imaging system (Gorsky et al. (2010)) to identify and determine the size structure of the zooplankton communities. Each sample was divided into one of two size fractions (<1000 and >1000  $\mu$ m) for better representation of rare large organisms in the scanned subsample (Vandromme et al. (2012)). Each fraction was split using a Motoda box (Motoda (1959)) until it contained an appropriate number of objects, approximately 1500. After scanning, each image was processed using ZooProcess (Gorsky et al. (2010)), which is written





- 110 in the ImageJ image analysis software (Rasband (1997–2011)). Only objects having an Equivalent Circular Diameter (ECD)  $> 300 \ \mu m$  were detected and processed (Gorsky et al. (2010)). Objects were first automatically classified using ECOTAXA (https://ecotaxa.obs-vlfr.fr/), then the results were validated by a taxonomist. With this approach 101 taxa were detected which were then grouped in 8 categories: Appendicularia, Chaetognatha, Copepoda, Cnidaria, Eumalacostraca, Foraminifera, Thaliacea and Other\_Organisms (Table 2). This last category includes all taxa which would not belong in any of the designated
- 115 classes and which were present in very low numbers across all samples. Zooplankton abundance (number of individuals  $m^{-3}$ ) was calculated from the number of validated vignettes in ZooScan samples, considering the scanned fraction and the sampled volume from the nets.

We also took into account trophic differences between the 40 taxonomic groups of copepods by assigning them into five major trophic groups based on their feeding behaviour, i.e., filter feeders, carnivorous cruise feeders, carnivorous ambush

120 feeders, non-carnivorous cruise feeders, and non-carnivorous ambush feeders, following the classification by Benedetti et al. (2015) (Table 2). Non-carnivorous copepods include herbivores and detritivores, and filter feeders include mixed feeders. The difference between cruise and ambush feeders is that cruise feeders have to swim actively to encounter their prey, while ambush feeders passively encounter them (Kiørboe et al. (2015)). The taxonomic categories other than copepods have well identified trophic behaviour, and were grouped in the different trophic groups (Table 2).

Table 2. Assignments of trophic groups to identified taxonomic classifications.	
---	--

Trophic groups	Categories	Abbreviations	Taxonomic Groups identified by ZooScan	
Carnivores	Chaetognatha	Cha	Chaetognatha	
	Copepods Carnivores Ambush-Feeders	Cop CAF	Candacia, Corycaeidae (Corycaeus, Urocorycaeus), Lubbockia	
	Copepods Carnivores Cruise-Feeders	Cop CCF	Aetideidae, Euchaeta, Haloptilus, Heterorhabdus, Sapphirinidae (Copilia, Sapphirina, Vettoria), Urocorycaeus	
	Cnidaria	Cni	Cnidaria (ephyra), Trachylinae (Aglaura, Solmundella), Obelia, Siphonophorae, Calycophorae (Diphyidae, Chelophyes, Chelophyes appendiculata, Eudoxoides, Eudoxoides spiralis, Hippopodius, Hippopodius hippopus, Lensia, Lensia subtilis, Muggiaea), Physonectae	
	Foraminifera	For	Foraminifera	
Filter-feeders	Appendicularians	Арр	Fritillariidae, Oikopleuridae	
	Copepods Filter-Feeders	Cop FF	Acartia, Calanus helgolandicus, Calocalanus, Centropages, Centropages bradyi, Centropages typicus, Eucalanus, Lucicutia, Lucicutia ovalis, Mormonilla, Nannocalanus minor, Pleuromamma, Pleuromamma abdominalis, Rhincalanus, Subeucalanus, Temora	
	Thaliacea	Tha	Doliolida, Pyrosomatida, Salpida ( <i>Salpa fusiformis</i> )	
Herbivores & detritivores	Copepods Non-Carnivores Ambush-Feeders	Cop NCAF	Distioculus minor, Oithona	
	Copepods Non-Carnivores Cruise-Feeders	Cop NCCF	Arietellus, Euchirella messinensis, Euchirella rostrata, Euterpina acutifrons, Mecynocera clausi, Microsetella, Oncaeidae, Peltidiidae, Scolecitrichidae	
Omnivores	Eumalacostraca	Eum	Amphipoda (Hyperiidae undetermined, <i>Phronima, Phrosina, Platyscelus, Primno, Pronoe, Scina, Vibilia</i> ), Decapoda (Dendrobranchiata, Penaeoidea), Euphausiacea larvae, Isopoda, Mysida	
All	Other_Organisms	Oth	Bivalvia larvae, Bryozoa larvae, Cirripedia larvae, Crustacea nauplii, Ctenophora, Cyphonaute, Echinodermata larvae, <i>Evadne</i> , Teleostei (larvae and eggs), Gastropoda (Atlantidae, Cavoliniidae, Cliidae, Creseidae, Cymbulioidea, Limacinidae), Halosphaera, Heteronemertea, Ostracoda, Polychaeta larva (Lopadorrhynchidae, Magelonidae, Phyllodocidae, Tomopteridae, Typhloscolecidae), Radiolaria, Tintinnida	





Enumerations from the 200 μm and 500 μm nets, which were deployed simultaneously, were combined. To avoid double counting of organisms large enough to be captured by both nets, a threshold value was established, based on the analysis of the Normalized Biomass Size Spectra (NBSS) (Sect. 2.7), considering all stations and depths (a specific value for each station would not have significantly altered the results). The threshold value (1148 μm ECD) identified the body size at which the 500 μm net samples more effectively (Fig. A2). Thus, organisms smaller than this size from the 200 μm net, and those larger from the 500 μm net, were combined to form a new count, called 'combined net' hereinafter.

### 2.4 Calculation of biovolumes

The biovolume of each organism can be calculated in different ways, including from ECD (Heath (1995)), or by determining the ellipsoid that best fits the object. However, mesozooplankton are generally not spherical in shape, meaning the ECD approach tends to overestimate biovolume because spheres have a higher volume-to-cross-sectional area ratio compared to other shapes
(Sprules et al. (2003); Beaulieu et al. (1999)). Consequently, a more accurate approach involves calculating biovolume based on the ellipsoid that best fits the organism (Alcaraz et al. (2003)). However, since the ellipsoid is a 3D object and the ZooScan image provides only 2D data, the following approach is used: (i) an ellipse is fitted to the 2D object, defining the major and minor axes; (ii) in 3D, the two minor axes are assumed to be equal; (iii) the biovolume is then calculated using formula (1) (Dubois et al. (2021)):

140 
$$B = \frac{4}{3} \times \pi \times \text{major} \times \text{minor}^2$$
 (1)

### 2.5 Definition of intermediate layers: 100-200 m and 200-400 m

Our nets sampled the layers: 0-100, 0-200, and 0-400 meters (Sect. 2.2). In order to study the community as a function of depth, the abundance of different taxonomic groups (Sect. 2.3) was calculated in each layer by differencing. For instance, subtracting the abundance measured at 0-100 m from that at 0-200 m provided the abundance of each taxonomic group, trophic group, or size class for the layer 100-200 m. A similar approach was applied to determine the abundance of taxonomic groups in the 200-400 m layer. This approach is assumed valid as the net's tows were carried out successively within a relatively short interval of time, typically 45 minutes. It is important to note that subtractions were performed on the eight major categories and not on each taxonomic group (see Table 2). In rare cases (12 %), especially for Eumalacostraca (particularly in the 100-200 m layer) and Cnidaria (particularly in the 200-400 m layer), resulting abundances were negative and thus set to zero.

### 150 **2.6** Analysis of variance and Post-Hoc comparisons

Using R version 4.4.1 (Team (2025)), one-way analyses of variance (ANOVA) was conducted to examine differences in absolute abundance across groups for each taxonomic category. Prior to performing the ANOVA, the normality of the data distribution was assessed using the Shapiro-Wilk test and the homogeneity of variances verified with Levene's test (car package, version 3.1-3; Fox and Weisberg (2019)). ANOVAs were then performed for five factors: Water masses, Layer, Period (day or





night), Transects (storm effect) and Copepod subgroups (mean abundances of the eight most abundant taxa within copepods).
 For each significant ANOVA result (p < 0.05), a Tukey's Honest Significant Difference test was applied to identify the groups that differ substantially from one another.</li>

### 2.7 Normalized Biomass Size Spectra (NBSS)

The size of organisms is considered a key indicator of community dynamics (Platt and Denman (1977)). NBSS (Platt and
Denman (1977)) are widely used to study this property. For constructing the NBSS, zooplankton organisms are grouped into logarithmically increasing size classes. The total biovolume of each class is then divided by the width of its size class (Platt and Denman (1977)). The x-axis [log2 zooplankton biovolume (mm<sup>3</sup>/ind)] is calculated as:

$$\log_2 \left[ \frac{\text{Zooplankton biovolume (mm^3.m^{-3})}}{\text{Abundance of each class size (ind.m^{-3})}} \right]$$
(2)

The y-axis [log2 normalized biovolume (m<sup>3</sup>)] is calculated as:

165 
$$\log_2\left[\frac{\text{Zooplankton biovolume (mm^3.m^{-3})}}{\text{Interval of each class size }(\Delta \text{volume.mm}^{-3})}\right]$$
 (3)

The NBSS thus represents the normalized biovolume as a function of the size of the organisms, both on a logarithmic scale. Biovolume data were estimated from ECD data provided by ZooProcess, using spherical shape.

A size-based analyses were conducted using PCA (Sect. 2.8) on copepod abundances per size at the different stations, using the size classes defined for the NBSS (Fig. A2). However, for clarity, the size classes were grouped from 15 size classes to just
170 five, and each class is defined by its ECD instead of its biovolume. These analyses were conducted on the copepod group only. Other taxonomic groups tend to have much larger size ranges, including large organisms such as Chaetognaths or Cnidaria,

### **2.8** Principal Component Analysis (PCA)

which introduce significant noise in the NBSS.

PCA was used to evaluate the similarities between the stations based on the abundance of the different taxonomic groups.
175 Distances between these stations were measured in the phase space of the PCA. Legendre and Gallagher (2001) showed that the Hellinger transformation, prior to PCA, is often preferable to Euclidean distance for calculating distances between samples. Hellinger distance (Rao (1995)) is obtained from:

$$D(x_1, x_2) = \sqrt{\sum_{j=1}^{p} \left(\sqrt{\frac{y_{1j}}{y_{1+}}} - \sqrt{\frac{y_{2j}}{y_{2+}}}\right)^2},\tag{4}$$

where p denotes the number of categories and  $y_{i+}$  is the sum of the frequencies of the ith object.

180 With this equation, the most abundant species contribute significantly to the sum of squares. The advantage of this approach is that it is asymmetric, meaning it is insensitive to double zeros, which is not the case for Euclidean distance (Prentice (1980); Legendre and Legendre (2012)).



190

195



The Hellinger transformation was performed with the labdsv package (Roberts (2023)). The abundance tables were centered and scaled, and the PCA was computed using FactoMineR (Lê et al. (2008)). Prior, to carrying out PCAs, the normality of the data and correlation of the variables were tested, using the Shapiro-Wilk test and Bartlett test of sphericity, respectively.

2.8.1 Fixed PCA axis for comparison across layers

To obtain comparable results across depth layers, the PCAs were always conducted in the same way with fixed axes. First, a PCA is performed using data from the 0–400 m layer. Then the datasets from all three layers were projected onto the 0-400 m layer's pair of PCA axes. This approach ensures that comparisons between communities in the three different layers are valid. This was done for community PCA, trophic PCA and size PCA, using matrices of station x groups of size 56 x 8, 56 x 11 and 56 x 5 respectively.

### 2.8.2 Pseudo-F calculation

To quantify the separation of the predefined groups (A, B, F, M) in PCA space, the pseudo-F (Caliński and Harabasz (1974)) is used. This calculation is based on a comparison between inter-group dispersion, measuring variability between groups, and intra-group dispersion, evaluating variability within each group. The pseudo-F calculation is:

$$Pseudo-F = \frac{Inter-group dispersion/(k-1)}{Intra-group dispersion/(n-k)},$$
(5)

where k is the number of groups and n the total number of individuals.

A high pseudo-F value suggests a clear separation between groups, indicating that inter-group variation is predominant compared to intra-group variation.

### 200 2.8.3 PCA with theoretical f stations

A fundamental question is: is the zooplankton community at the front a mixture of those from water masses A and B, or does it represent a specific community? A method was therefore developed to address this question by creating theoretical fT stations, which have a community composition that is a simple linear combination of those observed at stations a and b, as close as possible (i.e., minimal distance) to the observed f stations. The combination of a and b follows the formula:

205 
$$fT = \alpha \cdot a + (1 - \alpha) \cdot b,$$
 (6)

where  $\alpha$  is the proportional contribution from stations *a* and *b*, respectively. A total of 101 iterations was performed, with  $\alpha$  varying from 0 to 1 in increments of 0.01.





Therefore, four new theoretical stations are created per iteration:

$$flT\_D = \alpha_1 \cdot al\_D + (1 - \alpha_1) \cdot bl\_D$$
  
210 
$$flT\_N = \alpha_1 \cdot al\_N + (1 - \alpha_1) \cdot bl\_N$$
$$f2T\_D = \alpha_2 \cdot a2\_D + (1 - \alpha_2) \cdot b2\_D$$
$$f2T\_N = \alpha_2 \cdot a2\_N + (1 - \alpha_2) \cdot b2\_N$$

These fT stations were then added to the initial datasets. For each of these enriched datasets, a PCA is conducted. After obtaining the coordinates of the stations in the principal component space, the distances between the fT stations and their respective observed f stations are calculated for each iteration. For each iteration, the total distance is computed (sum of the distances between the theoretical stations and the original stations), for each transect. Finally, the fT station with the minimum total distance is selected, along with the corresponding  $\alpha$  value, for each transect. This process enables the generation of new intermediate observations that best reflect the theoretical composition of the front as a linear combination of a and b.

# **3** Results

225

### 220 3.1 Total abundance and biovolume across water masses and layers

The absolute values of abundances and biovolumes of zooplanktonic organisms across different depth layers and stations (Fig. 2) revealed distinct temporal and spatial patterns. In general, abundances and biovolumes of stations within the same water mass decreased over time (stations are presented in chronological order in Fig. 2), with the exception of abundances for the front (Fig. 2.a) and biovolumes for water mass B (Fig. 2.c). The same analysis conducted with the biovolume of organisms < 1 mm (Fig. 2.b) shows strong similarity with figure 2.a based on abundance. When individuals with an ECD greater than 1 mm are included (Fig. 2.c), large differences compared to the total abundances (Fig. 2.a) were observed for all stations.

Regarding the spatial differences during the two front crossings, abundance and biovolume tend to be lower at the front, than in water masses A and B. Indeed, for the first transect, abundances at the front were 67% lower compared to water mass A and 29% lower compared to water mass B. No such decrease was observed for the second transect (Fig. 2.a). Total biovolumes at

230 the front also showed a decline: 46% relative to water mass A for the first transect and 39% relative to water mass B for the second transect (Fig. 2.c). Interestingly, the second transect reveals greater homogeneity for organisms with a biovolume < 1 mm among water masses compared to the first transect (Fig. 2.b), reflecting the potential influence of post-storm dynamics.</p>

In figure 2.c, strong day-night differences were observed in the vertical distribution of organisms, probably influenced by diel vertical migration (DVM). During the night, between 49 and 62% of the total biovolume was concentrated in the upper

235 0-100 m layer, except for  $b3_N$  (due to cnidarians, see next Results section). During the day, between 51 and 83% of the biovolume was in the 200-400m layer, except for  $b1_D$  with 43%. On average, deep layer biovolume increases by 25% from night to day.







**Figure 2.** a)Total abundance of organisms by intermediate layers and across all sampled stations. b) Total biovolume of organisms, with ESD <1 mm only, by intermediate layers and across all sampled stations. c) Total biovolume of organisms by intermediate layers and across all sampled stations. Stations are in chronological order. The asterisk indicates that the net could not be analysed. Colours refers to the period of the day (blue for midday and black for midnight).





### 3.2 Taxonomic composition across nets and depth layers

240

The 200  $\mu$ m net more efficiently captures copepods, which constitute 45–95% of the relative abundances of taxa, while copepod abundances comprise only 5-55% in the 500  $\mu$ m net (Fig. 3). The larger mesh size is particularly effective for sampling larger taxa such as Eumalacostraca, Foraminifera, and Cnidaria. The combined samples, which include contributions from both mesh sizes, still heavily reflect the taxa distributions observed in the 200  $\mu$ m net, the abundances of larger organisms sampled with the 500  $\mu$ m net being low. This pattern was also observed in the layers 0-100 m and 0-400 m. Moreover, during the second transect (after the storm) the dominance of copepods is enhanced in water masses A and F.



**Figure 3.** Relative abundance of taxonomic groups for nets deployed from the surface to a depth of 200 m, for the two mesh sizes (200  $\mu$ m top, and 500  $\mu$ m, middle) across all sampled stations (chronological order). Bottom: Relative abundance combining the two mesh sizes. Colours refers to the period of the day (blue for midday and black for midnight).

In the 0–100 m layer, copepods consistently dominate, comprising at least 45% of the total abundance at nearly all stations except for *b2* and *b3* (Fig. 4). In the 100–200 m layer there is marked heterogeneity with many stations (8 out of 18) showing less than 60% copepods. This intermediate layer likely reflects a transitional zone where DVM results in taxonomic shifts. The 200–400 m layer returns to a dominance of copepods at most stations (15 out of 18), with the notable exceptions of station *b2\_N*, where Eumalacostraca account for an anomalously high 55% of the sampled taxa, and *b3\_N* where Cnidaria account for an anomalously high 55% of the sampled taxa.







Figure 4. Relative abundance of taxonomic groups for the combined nets for the three intermediate layers across all sampled stations (chronological order). The net from station  $m_N$  at a depth of 400 m could not be analysed. Colour refers to the period of the day (blue for midday and black for midnight).

#### 3.3 Diel variations in vertical structuring of zooplankton stocks

Biovolume values show similarities between the 0-100 m and 200-400 m layers which may correspond to a transfer of organisms between day and night layers (Fig. 2.c), whereas the 100–200 m layer (Fig. 4) presents a less structured composition as a potential transition zone for migratory organisms. The analysis of Hellinger distances for the eight taxonomic groups across 255 the three depth layers highlights that the communities in the 0-100 m and 200-400 m layers are more similar to each other, while the intermediate layer (100–200 m) harbours more distinct communities. This pattern was observed across all stations, as well as in separate analyses of the first and second transects. The lowest Hellinger distances were consistently found between the 0–100 m and 200–400 m layers for Copepoda (0.04 and 0.09, respectively), Eumalacostraca (0.03 and 0.08), and Other\_organisms (0.06 and 0.03). In contrast, distances involving the 100–200 m layer were approximately four times higher. 260 Furthermore, the vertical structuring of communities for the two migrant groups, Copepoda and Eumalacostraca, reveals a clear diel pattern: at night, the upper layers (0–100 m and 100–200 m) are more similar, while during the day, the similarity is

stronger between the deeper layers (100–200 m and 200–400 m). This pattern reflects a pronounced DVM for both groups with a significant day-night difference (post-hoc, p < 0.001 and 0.008 respectively). Indeed, distances between the 0–100 m and 200–400 m layers were notably higher both during the day (0.24 and 0.13 for Copepoda; 0.48 and 0.38 for Eumalacostraca) and at night (0.29 and 0.34 for Copepoda; 0.38 and 0.52 for Eumalacostraca, respectively for the two transects). In contrast,

265 Hellinger distances between the 0-100 m and 100-200 m layers at night were much lower (approximately 8 times lower for Copepoda and 3 times lower for Eumalacostraca), while during the day the lowest distances were observed between the 100-200 m and 200-400 m layers (approximately 21 times lower for Copepoda and 5 times lower for Eumalacostraca).





#### Community structure and water mass differentiation 3.4

#### 270 3.4.1 Community composition across depths and water masses

PCA\_Community summarizes the taxonomic composition of zooplankton communities across water masses and depths (Fig. 5). Axis 1 is inversely correlated to copepod abundance, which stems from the extreme dominance of this group. Axis 2 appears to be more to characteristics of other groups ranging from pure filter feeders (Appendicularians and Thaliacea) to omnivores (Eumalacostraca, Other\_organisms) and to carnivores (Chaetognatha and Cnidaria), Formanifera being at the extreme.

275

In the 0-100 m layer, water masses host distinct communities (Pseudo-F = 3.52). Copepods are more abundant in water masses A and the front, whereas other groups, particularly Foraminifera, Cnidaria, Chaetognatha and Eumalacostraca, dominate in water masses B and M. In this layer, for the first transect, stations f are closer to b than to a. For the second transect, there is a strong similarity between stations a and f. For the 100–200 m and 200–400 m layers, the taxonomic composition becomes progressively more homogeneous, with reduced differentiation among water masses (Pseudo-F = 1.66 and 1.06).

#### 3.4.2 Comparison of the front community composition with adjacent waters 280

The relative abundances of taxonomic groups across all stations in chronological order are used to compare the community compositions (Fig. 6). Results clearly reveal that the front appears very similar to water mass A in terms of the relative abundance of copepods which progressively decreases from A to F to B. To further investigate these observations, a PERMANOVA was conducted on the abundance of the entire community in water masses A, B, and the front. No significant difference was

- found between A and F (p = 0.312). However, significant differences were observed between B and F (p = 0.038) and between 285 A and B (p = 0.006). For copepods, significant differences were found between all pairs of water masses and for both transects, as determined by an ANOVA, except between F and A (p = 0.406 for the first transect and p = 0.459 for the second transect). For other groups, significant differences were only observed for Other\_organisms between B and A for both transects and between F and B for the second transect.
- 290 Figure 7 illustrates the theoretical community distribution at the front, derived from a combination of communities from water masses A and B (Sect. 2.8.3). The positioning of theoretical front stations (fT) is displayed within the PCA\_Community of figure 5 (Fig. 7.a). For the first transect (Fig. 6.b), the  $\alpha$  value (in Eq. 6) is low for the 0-100 m and 200-400 m layer (respectively 0.27 and 0.12) but high for the intermediate layer (0.73). This indicates that the front can be blurred by processes other than just the dynamics of water masses, for instance DVM through the 100-200 m layer. For the second transect, alpha is close to 1, even equal to 1 for the deeper layers, therefore the front is very similar to water mass A (Fig. 7.c).
- 295

A notable feature is the position of fT stations compared to observed f stations within the reduced PCA space. Focusing on the first transect (Fig. 7.b), observed f stations appear displaced relative to the fT stations. Indeed, in the 0–100 m, 100–200 m layers as well as during the night in the 200-400 m layer, observed f stations are positively shifted along axes 1 and/or 2. To examine these shifts, we can reconstitute the theoretical absolute abundance of these fT stations and then compare them to the

f stations. In the 0-100 m the different groups of fT are subjected to an average decrease of 35% (with a maximal decrease of 300 74% in Eumalacostraca). The observed shift is driven by a 72% increase in Cnidaria at the front compared to that is expected





(*fT*), while all other groups exhibit a decline. In the 100-200 m layer the important shift between *f* and *fT* is explained by the substantial increase of 90% in Foraminifera, this group is also the only one showing an increase in this layer (average decrease of 52% for other groups). During the night in the 200–400 m layer, this shift is explained by a pronounced increase of 605% in
Chaetognatha (average increase of 84% for other groups). In contrast, the second transect has much higher alpha values, which means a strong similarity between water mass A and F, with a strong domination of copepods in both water masses (Fig. 4). Thus, deviations between *fT* and *f* are very low and could not be analyzed.

### 3.4.3 Trophic composition and dynamics

PCA\_Trophic (Fig. 8) was conducted based on the abundance of each trophic category of Table 2. In the 0–100 m layer,
non-carnivorous groups (detritivores and herbivores) dominate in water masses A and F, while carnivorous groups are more abundant in water masses B and M. Indeed, the PCA\_Trophic reveals a clear trend in the distribution of trophic groups, with carnivores predominantly located in a quarter plan extending from the axes of Cop\_CCF to Cni, while herbivorous and non-carnivorous groups are situated on the rest of the plan (from Cop\_FF to App). Whereas the community in *a* stations is clearly dominated by non-carnivorous in 0-100 m and 200-400 m, a shift towards carnivory is observed in the 100–200 m layer for
these stations. The *b* and *m* stations are more dominated by filter-feeders and predators (Cni and Cha) in surface.

In the 200-400 m layer all the stations becoming less distinctly grouped and converging closer to the PCA center. In this layer, a positive shift along the x-axis is also observed. Upon closer examination, this shift is not driven by a high contribution of non-carnivorous copepods, but rather by a high contribution of Eumalacostraca and Foraminifera, and a low contribution of Chaetognatha. The discrimination of the stations among the water masses is stronger (approximately three times greater) for the 0-100 m layer compared to the other layers, indicating that, in terms of trophic functions, the water masses are less differentiated

0-100 m layer compared to the other layers, indicating that, in terms of trophic functions, the water masses are less differentiated for these deeper layers. Stations within each water mass appear relatively similar to one another, with the exception of station  $a1_N$  at 100-200 m, which stands out due to a high abundance of copepods carnivores ambush-feeders and Appendicularia.

### 3.4.4 Size composition of copepods

- In the 0–100 m layer, water masses exhibit the strongest differentiation in terms of copepod size structure, with *a* and *f* stations
  characterized by larger copepods (>1500 μm; Fig. 9). The *b* and *m* stations, on the other hand, are characterized by smaller copepods with *b* predominantly ranging from 300 to 450 μm and *m* from 450 to 950 μm. Meanwhile, the PCA\_Size shows a much more heterogeneous distribution for the *b* stations. As depth increases, size composition becomes more homogeneous, with all stations clustering near the PCA centre, but slightly shifted toward highest size. Indeed, there is a decrease in Pseudo-F with depth, respectively 4.47, 1.46 and 0.71. This concentration near the PCA centre and the decrease in Pseudo-F indicate a
  gradual decrease in variability among the deep stations, i.e., the differences between stations become less pronounced. This is
- also observed in the three previous PCAs (Fig. 5, 7, 8), but it is more pronounced here for the copepod size composition.







**Figure 5.** PCA\_Community illustrating the composition of communities, based on relative abundance data (Hellinger transformation) from all stations for each intermediate layers. The axis computed for 0 - 400 m were used for the three layers. Colour refers to the water mass (red for A, green for B, cyan for F and violet for M). In 0-100 m: stations  $a2_D$ ,  $a2_N$ ,  $f2_D$  and  $f2_N$  overlap at dim1 = -2.3 and dim2 = -0.3. In 100-200 m: stations  $a2_D$ ,  $a2_N$ ,  $f2_D$  and  $f2_N$  overlap at dim1 = -2.4 and  $f2_N$  and  $f2_N$  and  $f2_N$  overlap at dim1 = -1.7 and dim2 = -1.







**Figure 6.** Relative abundance of taxonomic groups across all stations, presented as a boxplot. Sorted by water mass and incorporating data from all three layers of each station. Colour refers to the period of the day (blue for midday and black for midnight).







**Figure 7.** a) PCA\_Community illustrating the composition of communities, based on relative abundance data (Hellinger transformation) from all stations for each of the intermediate layer (same as figure 5). The closest theorical *fT* of each observed *f* is plotted, with the corresponding  $\alpha_1$  and  $\alpha_2$  values of each *fT*'s couple for the 1st and 2nd transect, respectively. b) Zoom for the stations of the 1st transect. c) Zoom for the stations of the 2nd transect. In c), in 0-100 m stations *a2\_D*, *a2\_N*, *f2\_D*, *f2\_N*, *f2T\_D* and *f2T\_N* overlap at dim1 = -2.5 and dim2 = -0.2. In 100-200 m stations *a2\_D*, *a2\_N*, *f2\_D* and *f2T\_N* overlap at dim1 = -1.8 and dim2 = -0.6. In 200-400 m stations *f2T\_D* and *a2\_D* overlap at dim1 = -2 and dim2 = -0.2; *a2\_N*, *f2\_N* and *f2T\_N* overlap at dim1 = -1 and dim2 = -0.4.







**Figure 8.** PCA\_Trophic illustrating the trophic composition of communities, based on relative abundance data (Hellinger transformation) from all stations for each intermediate layer. The axis computed for 0 - 400 m were used for the three layers. Colour refers to the water mass (red for A, green for B, cyan for F and violet for M). In 0-100 m: stations  $a2_D$ ,  $a2_N$ ,  $f2_D$  and  $f2_N$  overlap at dim1 = -3.5 and dim2 = -2. In 200-400 m: stations  $a1_D$  and  $a1_N$  overlap at dim1 = -1.5 and dim2 = 0; stations  $b1_D$ ,  $m2_N$ ,  $m2_D$ ,  $b2_N$  and  $b3_D$  overlap at dim1 = -1 and dim2 = -0.5;  $b2_D$  and  $f1_N$  overlap at dim1 = -1.3 and dim2 = -1.3;  $b1_N$  and  $m_D$  overlap at dim1 = -0.1 and dim2 = -1.3.







**Figure 9.** PCA\_Size illustrating the body size composition of copepods, based on relative abundance data (Hellinger transformation) from all stations for each intermediate layer. The size classes (in  $\mu$ m) were defined according to those from NBSS. The axis computed for 0 – 400 m were used for the three layers. Colour refers to the water mass (red for A, green for B, cyan for F and violet for M). In 100-200 m: stations  $a2_N, b1_D, b2_D$  and  $m2_N$  overlap at dim1 = 0.8 and dim2 = 0.9. In 200-400 m: stations  $b2_N, m2_N, b3_D, m2_D'$  and  $f1_N$  overlap at dim1 = 1 and dim2 = 0.9;  $b1_N, b2_D$  and  $m_D$  overlap at dim1 = 0.6 and dim2 = -0.3.





# 4 Discussion

### 4.1 Zooplankton biovolumes, abundances and community structure across water masses

The spatial differences between water mass A and B in late spring can be linked to the regional hydrological and ecosystem functioning of the NWMS in the post-bloom period (D'Ortenzio and Ribera d'Alcalà (2009)). Water mass A has its origin in the Liguro-Provencal area (NWMS), characterized by intense convection and mixing (Barral et al. (2021)) and by high nutrient concentrations associated with strong gradients and shallower nutricline depths (Severin et al. (2017); Joël et al., 2025 in prep.). In the transitional post-bloom period (April-May), water mass A is still nutrient-rich (Severin et al. (2017); Joël et al., 2025 in prep.) and productive (Mayot et al. (2017); Hunt et al. (2017)) with the formation of a deep chlorophyll maximum around 50

- 340 m (Fig. A3; Lavigne et al. (2015); Doglioli et al. (2024)). Water mass A shows higher zooplankton stocks strongly dominated by copepods and larger forms (Fig. 4, 9). Water mass B in the southern part of the NBF comes from the epipelagic waters of the Algerian basin, which are warmer and fresher than waters from the NWMS, with virtually permanent stratification and a DCM deeper than 50 m (Fig. A3; Lavigne et al. (2015)). It is associated with lower nutrient concentration gradients and deeper nutricline depths (Joël et al., 2025 in prep.). In water mass B, zooplankton stocks fluctuate little over the year
- 345

350

the non-copepod organisms (Fig. 4, 9).

Mesozooplankton data from the two transects across the NBF during the BioSWOT-Med campaign can only be compared with a very limited number of other observations, particularly in the vicinity of the front. The two DEWEX campaigns (Conan et al. (2018)), in February and April 2013, studied dense water formation and the pelagic ecosystem's biogeochemical and ecological responses, including zooplankton mapping during the winter-spring transition (Donoso et al. (2017)). The stations

(Fernández de Puelles et al. (2004)), and community structure is dominated by small sizes and slightly more diversified within

- of water mass A were in the Deep Convection Zone (DCZ) near the LION Station (42°04′ N, 4°38′ E; Margirier et al. (2020)), whereas the nearest station of water mass B was situated north of Menorca Island (Donoso et al. (2017)). In winter, the DCZ exhibited its lowest zooplankton abundance and biomass in the top 200 m (200 ind/m<sup>3</sup> and 5 mg DW/m<sup>3</sup>), in contrast to the periphery of the DCZ and especially the Balearic stations (650 ind/m<sup>3</sup> and 10 mg DW/m<sup>3</sup>). In spring, this pattern reverses
- 355 with higher zooplankton stocks in the DCZ (4400 ind/m<sup>3</sup> and 100 mg DW/m<sup>3</sup>) and lower in the periphery, including the Balearic stations (2000 ind/m<sup>3</sup> and 30 mg DW/m<sup>3</sup>). During BioSWOT-Med, water mass A exhibited approximately twice the abundances and biovolumes compared to water mass B and the frontal zone F during the first transect (Fig. 2). However, during the second transect, abundances and biovolumes in water mass A were much lower compared to the first transect. This overall decrease during BioSWOT-Med might be explained by the 10 day interval between the two transects in a phase of
- 360

365

storm occurring between the two transects could have influenced these decreases.

Our biomasses in water mass A match the biomasses observed in the central part of the NWMS (converting our biovolumes to biomass using DW/WW of 10%, and 1 mg WW equal to 1 mm<sup>3</sup>). Biomasses in water mass B align with values observed around the Balearic Islands and in the Algerian Basin. A similar comparison also applies to abundances (Nival et al. (1975); Siokou-Frangou et al. (2010); Nowaczyk et al. (2011); Donoso et al. (2017); Fierro-González et al. (2023); Fernández de

seasonal decline in zooplankton stocks in the convection zone (Berline et al. (2011); Auger et al. (2014)). Additionally, the





Puelles et al. (2023)). Another interesting pattern is the highest percentage of copepods in mesozooplankton found in water mass A (around 85%) compared to water mass B (around 40%), which is consistent with other observations made in the same water masses (Nowaczyk et al. (2011); Feliú et al. (2020); Fierro-González et al. (2023); Fernández de Puelles et al. (2023); Siokou et al. (2019)). However, these previous observations in the NWMS were not dedicated to the NBF.

370

In the NWMS, studies dedicated to the impact of frontal dynamics on the stocks and structure of mesozooplankton have been carried out in the front associated with the northern current in the Ligurian (Prieur and Sournia (1994); Boucher et al. (1987); Panaïotis et al. (2024)) and the Catalan seas (Alcaraz et al. (2007); Saiz et al. (2007)). In both areas, the front does not appear to be an area of higher zooplankton biomass than adjacent waters (Boucher et al. (1987); Panaïotis et al. (2024); Alcaraz et al. (2007); Saiz et al. (2007)), but it is the site of higher physiological rates (spawning rates, larval growth) of the

375 organisms (Boucher et al. (1987); Alcaraz et al. (2007); Saiz et al. (2007)), enhanced by higher prey densities and turbulence levels (Alcaraz et al. (2007)). Studies of species distributions have shown that the front associated with the northern current can also represent a barrier for coastal species in their distribution to the central zone (Pedrotti and Fenaux (1992); Saiz et al. (2014)), or conversely, from the central zone to the coast (Berline et al. (2013)).

### 4.2 Complexity of concurrent processes impacting zooplankton biomass distribution at front

- 380 The decline in zooplankton abundances and biovolumes at the front during the first transect, and the weak increase in abundance but not in biovolume during the second transect (Fig. 2) could reflect specific hydrological and physical mixing characteristics of the northern Balearic front (Salat (1995); Alcaraz et al. (2007)), where dynamic turbulence and horizontal dynamics appeared less favourable for biomass accumulation. Actually, while turbulence at fronts is known to enhance nutrient diffusion to phytoplankton, promoting enriched food webs for zooplankton (Kiørboe (1993); Estrada and Berdalet (1997)), it can
- 385 also increase encounter rates between particles and consumers, influencing community interactions (Rothschild and Osborn (1988); Alcaraz et al. (1989); Saiz et al. (1992); Caparroy et al. (1998)). Indeed, the front in our area of interest, sampled by Lagrangian drifters at 1 and 15 m depth (Demol et al. (2023)), showed prevailing along-front deformation and patches of water mass convergence and divergence inducing variable vertical velocities up to approximately +/- 1 mm/s in the upper 15 m sea layer (Berta (2025)). Moreover, the core of the front, as identified by ADCP transects (Petrenko et al. (2024)), is found
- 390 within 100 m depth and 20 km width. Consequently, considering the frontal spatial scales, as well as the divergence and the vertical transport magnitude and variability, we expect that our results do not reveal significant effects beyond 100 m depth and that mixing has shorter time scales compared to zooplankton development times (several weeks to months). In a study of 154 glider-resolved fronts across the California Current System, Powell and Ohman (2015a) found that zooplankton biomass was often, but not always, enhanced, also indicating variations in matchup of frontal duration and zooplankton development
- 395 time. Finally, our campaign took place in late April to early May, corresponding to the post-bloom period, when phytoplankton biomass levels are already too low to sustain optimal growth of specific zooplankton groups.

The observed variability in zooplankton abundances and biovolumes over time and space (Fig. 2) underscores the complexity of concurrent processes acting at different scales, such as DVM or storm events that interact with the hydrological processes



405



creating the front. Furthermore, biomass peaks heavily depend on the taxa considered (Mangolte et al. (2023)) and our analyses 400 never consider a single taxon but always groups of organisms (Table 2) or even the whole sampled mesozooplankton.

#### Investigating the front: mixing zone or distinct community? 4.3

A fundamental question in this study was whether the front is a mixture of communities from water masses A and B, or if it hosts a distinct community with notably different abundances, or even the presence of new taxa. Our results reveal clearly that the front appears very similar to water mass A in terms of the relative abundance of copepods, with decreasing relative abundance from A to F to B (Fig. 6). Moreover, in the 0-100 m layer, the shifts between the projections of f and fT (Fig. 7) suggests a less disruptive influence of the front on Cnidaria and Foraminifera. In contrast, the pronounced decrease in Thaliacea may indicate active avoidance of the frontal region.

We note that the primary differences among higher taxa across the front concern not the most abundant groups, but secondary groups: Cnidaria, Foraminifera and Eumalacostraca for 0-100 m; Cnidaria and Foraminifera for 100-200m. This indicates that

- 410 the composition of zooplankton communities is shaped by species-specific responses, highlighting the importance of considering individual taxonomic groups rather than just overall abundance patterns when analysing community dynamics. In other frontal studies, some taxa have been found more abundant than in adjacent waters, showing that zooplankton abundance at fronts is species-specific (Molinero et al. (2008)). Gastauer and Ohman (2024) found that appendicularians, copepods, and rhizarians sometimes showed front-related increases in concentration. These patterns likely result from interactions between
- 415 species-specific behaviors and frontal dynamics. Indeed, certain zooplankton can maintain their depth through vertical swimming or buoyancy control, allowing them to counteract vertical currents and form aggregations at the front (Franks (1992)). As a result, fronts can enhance abundance or facilitate exchanges for particular species depending on their behavioral strategies (Wishner et al. (2006)).

To answer our initial question, the results suggest that for the first transect, the front is indeed a mix of A and B communities, 420 but it also shows higher abundances of organisms such as Cnidaria, Foraminifera and Chaetognatha, which are less represented in A and B. For the second transect, the storm of the previous days may have altered the community structure (a hypothesis that will be further discussed in Sect. 4.5), making it difficult to draw definitive conclusions. It can also be noted that no taxon was found exclusively at the front in either transect. We note that our more detailed analysis revealed the absence of a few low-abundance taxa at the front compared to water masses A or B (e.g., Magelonidae, cyphonautes, echinoderm larvae, radiolarians, and Heteronemertea).

425

430

#### 4.4 **Community structure with depth**

The zooplankton community reveals clear vertical stratification, with distinct compositions at different depths. Zooplankton taxonomic composition becomes more homogeneous with depth between water masses A, B, and the front, as the influence of the front weakens below 100 m and becomes negligible below 200 m. The deep zone below the front reflected reduced differentiation among water masses and weakening hydrological influence (Fig. 5, 8, 9), as generally observed (Scotto di Carlo





et al. (1984)). Thus, the higher discrimination between water masses in the upper 0-100 m layer emphasizes the stronger influence of hydrology and biological productivity at the surface. Besides physical processes (Sect. 4.2), differences in zooplankton taxonomic and size distributions between water masses A, B, and the front are influenced by phytoplankton size structure and trophic interactions with nano- and micro-grazers (Bănaru et al. (2014); Hunt et al. (2017); Tesán Onrubia et al. (2023)). Ongoing analyses will explore the food-web pathways to zooplankton in these water masses (Bănaru et al., in prep.; Joël et al.,

435

2025 in prep.). Our findings (Fig. 8, 9) already highlight the dominance of non-carnivorous, large copepods in water masses A and F.

The zooplankton structure in the surface layer was also clearly affected by the behaviour of migrating zooplankton. Our results clearly showed that the taxonomic and size distributions of migrant zooplankton were closer between layers of 0-100

m and 100-200 m at night and between layers of 100-200 m and 200-400 m during the day (Sect. 3.3, Fig. 4) which is a 440 direct consequence of the population movements that make up DVM. The copepod and eumalacostracan groups present in our samples are well documented as performing DVM (Andersen and Sardou (1992); Isla et al. (2015); Guerra et al. (2019)). This deep DVM observed in our samples is likely driven by species of Euchaeta, Pleuromamma, and Heterorhabdus, as Andersen et al. (2001b) demonstrated their DVM amplitude can reach up to 400-500 m for some taxa.

#### 4.5 Storm impact? 445

The Provençal basin in NWMS is an area highly exposed to strong regional winds (i.e., "mistral" and "tramontane") blowing southwards, with a marine origin in the Gulf of Lion that is particularly frequent in autumn and winter. Such winds can, nonetheless, occur more episodically but intensely in spring and even summer (Morales-Márquez et al. (2021)). During an intense two-day wind-storm in the Ligurian Sea, Barrillon et al. (2023) observed a significant hydrological effect on the upper

- 60 m, impacting nutrient supply and chlorophyll biomass over the following 24 hours. However, the influence on zooplankton 450 was not documented. Andersen et al. (2001a) sampled two one-day wind-storm events in the Ligurian Sea. The main effects also occurred in the upper surface layer. The main observed impact on the zooplankton was an increased production of nauplii linked to adult spawning, but a decrease of zooplankton biomass (mainly copepods). Analyzing the zooplankton vertical distributions from Niskin bottles sampling, Andersen et al. (2001b) noted upward aggregation within the upper 40 m of naupliar stages of copepods and euphausiids as well as small-sized copepods. They do not mention any effect on other groups than crustaceans,
- 455

460

notably gelatinous plankton, from either bottle samples or from plankton nets.

During BioSWOT-Med, a storm (NW winds) peaked on May 2nd, between the first and the second transect. The two transects were separated by nine days and : ~50 km. The observed changes in community structure and distribution between the two transects may have resulted from the mixing and dilution associated with the storm, but also from patchiness or temporal evolution. However, we observed a limited influence on mixed layer depth (deepening from ~15 m to 30 m) and also on chlorophyll-a fluorescence profiles, according to glider data, similar to the observations by Barrillon et al. (2023) and Andersen et al. (2001b). Comparison of abundances between the first to the second transect revealed significant differences for the 0–100

m layer, but not in deeper layer, therefore potentially linked to the storm (Table 3). In the surface layer, small and mid-sized





465

copepods were the most affected, but also chaetognaths and cnidarians. Large migrant copepods, such as *Pleuromamma* and *Euchaeta*, appeared weakly affected. A similar trend was observed for Calanoida, which includes both small and large, migrant and non-migrant species. Analyses of the whole planktonic community response to the storm (including phytoplankton) are required to better understand the observed zooplankton changes.

**Table 3.** Results of ANOVA tests (H0: no differences of averages between the first and the second transect) performed on the eight taxonomic groups and copepod subgroups (eight most abundant species). For each significant ANOVA result (p < 0.05), a Tukey's Honest Significant Difference test was applied to identify differences between the first and the second transect for each water mass (shown in the last four columns). For layers 100-200 m and 200-400 m, no significant differences were found.

Type of analysis	Depth	Taxonomic group / species	ANOVA p-value	p-value A1st vs A2nd	p-value B1st vs B2nd	p-value F1st vs F2nd	p-value M1st vs M2nd
ANOVA Depth	0-100 m	Appendicularians	0.124				
		Chaetognatha	0.039 *	<0.001 ***	0.0659	<0.001 ***	0.108
		Cnidaria	<0.001 ***	<0.001 ***	<0.001 ***	0.418	0.765
		Copepoda	<0.001 ***	0.189	< 0.001 ***	<0.001 ***	0.617
		Eumalacostraca	0.534				
		Foraminifera	0.429				
		Other_Organisms	0.375				
		Thaliacea	0.929				
ANOVA Copepod species	0-100 m	Calanoida	0.255				
		Centropages Typicus	0.014 *	0.002 **	0.104	< 0.001 ***	1
		Maxillopoda	0.052				
		Corycaeidae	0.0104 *	< 0.001 ***	0.797	< 0.001 ***	0.992
		Euchaeta	0.581				
		Oithona	0.0231 *	0.448	0.0197 *	0.923	0.876
		Oncaeidae	0.015 *	< 0.001 ***	0.031*	0.025 *	0.87
		Pleuromamma	0.928				

# 5 Conclusion

To our knowledge, this study represents the first detailed investigation of the fine-scale zooplankton distribution of the North Balearic Front in late spring, linking finescale dynamics to mesozooplankton distributions. Our findings reveal that the North Balearic Front exhibits characteristics more akin to a boundary between water masses than a zone of pronounced biological accumulation.

Key observations include the stratified vertical distribution of zooplankton communities, with distinct taxonomic compositions in the surface, intermediate, and deeper layers, and a progressive homogenization of community structure with depth.
DVM was particularly evident, underscoring the dynamic nature of zooplankton behaviour in relation to environmental gradients. Moreover, post-storm analyses highlighted the susceptibility of these communities to episodic weather events, which can disrupt established ecological patterns.

These results challenge generalized assumptions about the ecological role of oceanic fronts. They underscore the importance of high-resolution observations across horizontal and vertical spatial scales, temporal processes ranging from days to weeks, and precise taxonomic determination to fully understand the complexity of mesozooplanktonic communities in frontal zones.





Further trophic studies based on stable isotope ratios and the biochemical composition of zooplankton and phytoplankton size classes are still needed to decipher the complexity of trophic interactions in the two adjacent contrasting zones and in the frontal zone under the effect of nutrient input determined by physical processes. In addition, our net sampling approaches need to be complemented by continuous measurement techniques, such as autonomous gliders, bioacoustics and satellite data, with in-situ sampling to better capture the spatial and temporal variability of these systems. This approach would enable a more comprehensive assessment of how physical and biological processes interact to shape zooplankton communities at oceanic fronts.

# **Data Availability Statement**

The Sentinel-3 data used in this manuscript are available and freely accessible to the public (https://www.copernicus.eu/en), 490 accessed on 11 November 2024.

# **Author contribution**

Maxime Duranson was responsible for data curation, formal analysis, visualization, and conceptualization. François Carlotti and Léo Berline contributed to conceptualization and were responsible for supervision and validation. Loïc Guilloux performed the ZooScan processing and taxonomic identification of the samples. Maxime Duranson prepared the manuscript with review and editing of all co-authors.

```
495
```

485

# Acknowledgments

The authors thank the TOSCA program of the CNES (French Spatial Agency) which funds the BIOSWOT-AdAC project and the ANR – FRANCE (French National Research Agency) for its financial support to the BIOSWOT ANR-23-CE01-0027 project.

500 The FOF (French Oceanographic Fleet) and, in particular, the captain Gilles Ferrand and the crew of the R/V L'Atalante are acknowledged for their precious support during the BioSWOT-Med cruise.

M.D. Ohman and S. Gastauer were supported by U.S. National Science Foundation grant OCE-2243190.

Alice Della Penna was supported by the University of Auckland via the Faculty Research Development Fund (Grant 3724591)

505 Dr. Maristella Berta contribution was supported by the ITINERIS Project (IR0000032-Italian Integrated Environmental Research Infrastructures System-CUP B53C22002150006) and by CNR-ISMAR (Lerici, Italy) dedicated fundings.





# Appendix



Figure A1. Maps of the sampling stations with surface chlorophyll concentration for 3 different days (as complement of Figure 1).



**Figure A2.** NBSS incorporating all data (all stations and depths) for each mesh size. The threshold value represents the organism size above which the 500  $\mu$ m nets sample more efficiently than the 200  $\mu$ m nets.







Figure A3. Total particles abundance, temperature, salinity, and fluorescence profiles.





# References

Alcaraz, M., Estrada, M., and Marrasé, C.: Interaction between turbulence and zooplankton in laboratory microcosms, in: Proceedings of the

- 510 21st European Marine Biology Symposium, edited by Klekowski, R. Z., Styczynska-Jurewicz, E., and Falkowski, L., pp. 191–204, Polish Academy of Sciences-Institute of Oceanology, Ossolineum, Gdansk, 1989.
  - Alcaraz, M., Saiz, E., Calbet, A., et al.: Estimating zooplankton biomass through image analysis, Mar. Biol., 143, 307–315, https://doi.org/10.1007/s00227-003-1094-8, 2003.
- Alcaraz, M., Calbet, A., Estrada, M., Marrasé, C., Saiz, E., and Trepat, I.: Physical control of zooplankton communities in the Catalan Sea,
  Prog. Oceanogr., 74, 294–312, https://doi.org/10.1016/j.pocean.2007.04.003, 2007.
- Andersen, V. and Sardou, J.: The diel migrations and vertical distributions of zooplankton and micronekton in the Northwestern Mediterranean Sea. 1. Euphausiids, mysids, decapods and fishes, J. Plankton Res., 14, 1129–1154, https://doi.org/10.1093/plankt/14.8.1129, 1992.
  - Andersen, V., Nival, P., Caparroy, P., and Gubanova, A.: Zooplankton community during the transition from spring bloom to oligotrophy in the open NW Mediterranean and effects of wind events. 1. Abundance and specific composition, J. Plankton Res., 23, 227–242,
- 520 https://doi.org/https://doi.org/10.1093/plankt/23.3.227, 2001a.
  - Andersen, V., Gubanova, A., Nival, P., and Ruellet, T.: Zooplankton community during the transition from spring bloom to oligotrophy in the open NW Mediterranean and effects of wind events. 2. Vertical distributions and migrations, J. Plankton Res., 23, 243–261, https://doi.org/10.1093/plankt/23.3.243, 2001b.
  - Ashjian, C. J., Davis, C. S., Gallager, S. M., and Alatalo, P.: Distribution of plankton, particles, and hydrographic features across
- 525 Georges Bank described using the Video Plankton Recorder, Deep-Sea Res. Part II: Topical Stud. Oceanogr., 48, 245–282, https://doi.org/10.1016/S0967-0645(00)00121-1, 2001.
  - Auger, P. A., Ulses, C., Estournel, C., Stemmann, L., Somot, S., and Diaz, F.: Interannual control of plankton communities by deep winter mixing and prey/predator interactions in the NW Mediterranean: Results from a 30-year 3D modeling study, Progr. Oceanogr., 124, 12–27, https://doi.org/10.1016/j.pocean.2014.04.004, 2014.
- 530 Barral, Q.-B.: Caractérisation du front Nord-Baléares: Variabilité et rôle de la circulation des masses d'eau en Méditerranée Occidentale, Ph.D. thesis, Université de Toulon, 2022.
  - Barral, Q.-B., Zakardjian, B., Dumas, F., Garreau, P., Testor, P., and Beuvier, J.: Characterization of fronts in the Western Mediterranean with a special focus on the North Balearic Front, Prog. Oceanogr., p. 102636, https://doi.org/10.1016/j.pocean.2021.102636, 2021.
- Barrillon, S., Fuchs, R., Petrenko, A. A., Comby, C., Bosse, A., Yohia, C., Fuda, J. L., Bhairy, N., Cyr, F., Doglioli, A. M., Grégori, G.,
  Tzortzis, R., d'Ovidio, F., and Thyssen, M.: Phytoplankton reaction to an intense storm in the north-western Mediterranean Sea, Biogeo
  - sciences, 20, 141–161, https://doi.org/https://doi.org/10.5194/bg-20-141-2023, 2023. Beaulieu, S., Mullin, M., and Tang, V.: Using an optical plankton counter to determine the size distributions of preserved zooplankton

Belkin, I. M. and Helber, R. W.: Physical oceanography of fronts: an editorial, Deep-Sea Res. Part II, 119, 1–2, 2015.

samples, J. Plankton Res., 21, 1939-1956, https://doi.org/10.1093/plankt/21.10.1939, 1999.

- 540 Belkin, I. M., Cornillon, P. C., and Sherman, K.: Fronts in large marine ecosystems, Prog. Oceanogr., 81, 223–236, https://doi.org/10.1016/j.pocean.2009.04.015, 2009.
  - Benedetti, F., Gasparini, S., and Ayata, S.: Identifying copepod functional groups from species functional traits, J. Plankton Res., 38, 159–166, https://doi.org/10.1093/plankt/fbv096, 2015.





- Berline, L., Siokou-Frangou, I., Marasovic, I., Vidjak, O., Fernandes De Puelles, M. L., Mazzocchi, M. G., Assimakopoulou, G., Zervoudaki,
- 545 S., Fonda Umani, S., Conversi, A., Garcia-Comas, C., Ibanez, F., Gasparini, S., Stemmann, L., and Gorsky, G.: Intercomparison of Six Mediterranean Zooplankton Time Series, Prog. Oceanogr., 2011.
  - Berline, L., Zakardjian, B., Molcard, A., Ourmieres, Y., and Guihou, K.: Modeling jellyfish *Pelagia noctiluca* transport and stranding in the Ligurian Sea, Mar. Pollut. Bull., 70, 90–99, https://doi.org/10.1016/j.marpolbul.2013.02.016, 2013.
  - Berta, e. a.: Characterization of the fine scale dynamics during the BioSWOT experiment in the Balearic Sea, Manuscript in preparation,

550

565

2025.

- Boucher, J., Ibanez, F., and Prieur, L.: Daily and seasonal variations in the spatial distribution of zooplankton populations in relation to the physical structure in the Ligurian Sea Front, J. Mar. Res., 45, 133–173, 1987.
- Bănaru, D., Carlotti, F., Barani, A., Gregori, G., Neffati, N., and Harmelin-Vivien, M.: Seasonal variation of stable isotope ratios of size-fractionated zooplankton in the Bay of Marseille (NW Mediterranean Sea), J. Plankton Res., 36, 145–156, https://doi.org/10.1093/plankt/fbt083, 2014.

Caliński, T. and Harabasz, J.: A dendrite method for cluster analysis, Commun. Stat., 3, 1–27, https://doi.org/10.1029/98JC01537, 1974.
Caparroy, P., Pérez, M. T., and Carlotti, F.: Feeding behavior of Centropages typicus in calm and turbulent conditions, Mar. Ecol. Prog. Ser., 168, 109–118, https://doi.org/10.3354/meps168109, 1998.

Chiba, S., Ishimaru, T., Hosie, G. W., and Fukuchi, M.: Spatio-temporal variability of zooplankton community structure off east Antarctica
(90 to 160 E), Mar. Ecol. Prog. Ser., 216, 95–108, https://doi.org/10.3354/meps216095, 2001.

Conan, P., Testor, P., Estournel, C., D'Ortenzio, F., Pujo-Pay, M., and Durrieu de Madron, X.: Preface to the Special Section: Dense water formations in the northwestern Mediterranean: From the physical forcings to the biogeochemical consequences, J. Geophys. Res.: Oceans, 123, 6983–6995, https://doi.org/https://doi.org/10.1029/2018JC014301, 2018.

Cotté, C., Guinet, C., Taupier-Letage, I., Mate, B., and Petiau, E.: Scale-dependent habitat use by a large free-ranging predator, the Mediterranean fin whale, Deep-Sea Res. I, 56, 801–811, https://doi.org/10.1016/j.dsr.2008.12.008, 2009.

Cotté, C., d'Ovidio, F., Chaigneau, A., Lévy, M., Taupier-Letage, I., Mate, B., and Guinet, C.: Scale-dependent interactions of Mediterranean whales with marine dynamics, Limnol. Oceanogr., 56, 219–232, https://doi.org/10.4319/lo.2011.56.1.0219, 2011.

Demol, M., Berta, M., Gomez Navarro, L., Izard, L., Ardhuin, F., Bellacicco, M., Centurioni, L., d'Ovidio, F., Diaz-Barroso, L., Doglioli, A., Dumas, F., Garreau, P., Joël, A., Lizaran, I., Menna, M., Mironov, A., Mourre, B., Pacciaroni, M., Pascual, A., Ponte, A., Reyes,

- 570 E., Rousselet, L., Tarry, D. R., and Verger-Miralles, E.: A drifter dataset for the Western Mediterranean Sea collected during the SWOT mission calibration and validation phase, SEANOE, https://doi.org/10.17882/100828, 2023.
  - Derisio, C., Alemany, D., Acha, E. M., and Mianzan, H.: Influence of a tidal front on zooplankton abundance, assemblages and life histories in Península Valdés, Argentina, J. Mar. Syst., 139, 475–482, https://doi.org/10.1016/j.jmarsys.2014.08.019, 2014.

di Sciara, G. N., Castellote, M., Druon, J.-N., and Panigada, S.: Fin whales, Balaenoptera physalus: At home in a changing Mediterranean
Sea?, Adv. Mar. Biol., 75, 75–101, https://doi.org/10.1016/bs.amb.2016.08.002, 2016.

- Doglioli, A. M., Grégori, G., d'Ovidio, F., Bosse, A., Pulido, E., Carlotti, F., Lescot, M., Barani, A., Barrillon, S., Berline, L., Berta, M., Bouruet-Aubertot, P., Chirurgien, L., Comby, C., Cornet, V., Cotté, C., Della Penna, A., Didry, M., Duhamel, S., Fuda, J.-L., Gastauer, S., Guilloux, L., Lefèvre, D., Le Merle, E., Martin, A., McCann, D., Menna, M., Nunige, S., Oms, L., Pacciaroni, M., Petrenko, A., Rolland, A., Rousselet, L., and Waggonet, E. M.: BioSWOT Med. Biological applications of the satellite Surface Water and Ocean Topography in
- 580 the Mediterranean. Ref. Rapport de campagne, Université Aix-Marseille, https://doi.org/10.13155/100060, 2024.



585



Donoso, K., Carlotti, F., Pagano, M., Hunt, B. P. V., Escribano, R., and Berline, L.: Zooplankton community response to the winter 2013 deep convection process in the NW Mediterranean Sea, J. Geophys. Res. Oceans, 122, 2319–2338, https://doi.org/10.1002/2016JC012176, 2017.

D'Ortenzio, F. and Ribera d'Alcalà, M.: On the trophic regimes of the Mediterranean Sea: a satellite analysis, Biogeosciences, 6, 139–148, https://doi.org/10.5194/bg-6-139-2009, 2009.

- Druon, J. N., Hélaouët, P., Beaugrand, G., Fromentin, J. M., Palialexis, A., and Hoepffner, N.: Satellite-based indicator of zooplankton distribution for global monitoring, Sci. Rep., 9, 4732, https://doi.org/10.1038/s41598-019-41212-2, 2019.
  - Dubois, C., Irisson, J., and Debreuve, E.: Correcting estimations of copepod volume from 2-dimensional images, Limnol. Oceanogr.: Methods, 20, 361–371, https://doi.org/10.1002/lom3.10492, 2021.
- 590 Durski, S. M. and Allen, J. S.: Finite-amplitude evolution of instabilities associated with the coastal upwelling front, J. Phys. Oceanogr., 35, 1606–1628, https://doi.org/10.1175/JPO2762.1, 2005.

Errhif, A., Razouls, C., and Mayzaud, P.: Composition and community structure of pelagic copepods in the Indian sector of the Antarctic Ocean during the end of the austral summer, Polar Biol., 17, 418–430, https://doi.org/10.1007/s003000050136, 1997.
Estrada, M. and Berdalet, E.: Phytoplankton in a turbulent world, Scientia Marina, 61, 125–140, 1997.

- 595 Feliú, G., Pagano, M., Hidalgo, P., and Carlotti, F.: Structure and function of epipelagic mesozooplankton and their response to dust deposition events during the spring PEACETIME cruise in the Mediterranean Sea, Biogeosciences, 17, 5417–5441, https://doi.org/https://doi.org/10.5194/bg-17-5417-2020, 2020.
  - Fernández, E., Cabal, J., Acuña, J., Bode, A., Botas, A., and García-Soto, C.: Plankton distribution across a slope current-induced front in the southern Bay of Biscay, J. Plankton Res., 15, 619–641, https://doi.org/10.1093/plankt/15.6.619, 1993.
- 600 Fernández de Puelles, M. L., Valencia, J., and Vicente, L.: Zooplankton variability and climatic anomalies from 1994 to 2001 in the Balearic Sea (Western Mediterranean), ICES J. Mar. Sci., 61, 492–500, https://doi.org/https://doi.org/10.1016/j.icesjms.2004.03.026, 2004.
  - Fernández de Puelles, M. L., Gazá, M., Cabanellas-Reboredo, M., and O'Brien, T. D.: Decadal Trends in the Zooplankton Community of the Western Mediterranean, Water, 15, 4267, https://doi.org/10.3390/w15244267, 2023.
- Fierro-González, P., Pagano, M., Guilloux, L., Makhlouf, N., Tedetti, M., and Carlotti, F.: Zooplankton biomass, size structure, and associated
   metabolic fluxes with focus on its roles at the chlorophyll maximum layer during the plankton-contaminant MERITE-HIPPOCAMPE
   cruise, Marine Pollution Bulletin, 193, 115 056, https://doi.org/https://doi.org/10.1016/j.marpolbul.2023.115056, 2023.

Font, J., Salat, J., and Tintoré, J.: Permanent features of the circulation in the Catalan Sea, Oceanol. Acta, 9, 51–57, 1988.

Fox, J. and Weisberg, S.: An R Companion to Applied Regression, Sage, Thousand Oaks, CA, 3rd edn., 2019.

Franks, P. J.: Sink or swim: Accumulation of biomass at fronts, Mar. Ecol. Prog. Ser., 82, 1–12, 1992.

610 Gastauer, S. and Ohman, M. D.: Resolving abrupt frontal gradients in zooplankton community composition and marine snow fields with an autonomous Zooglider, Limnol. Oceanogr., https://doi.org/10.1002/lno.12642, 2024.

Gorsky, G., Ohman, M. D., Picheral, M., Gasparini, S., Stemmann, L., Romagnan, J. B., Cawood, A., Pesant, S., García-Comas, C., and Prejger, F.: Digital zooplankton image analysis using the ZooScan integrated system, J. Plankton Res., 32, 285–303, https://doi.org/10.1093/plankt/fbp124, 2010.

615 Guerra, D., Schroeder, K., Borghini, M., Camatti, E., Pansera, M., Schroeder, A., Sparnocchia, S., and Chiggiato, J.: Zooplankton diel vertical migration in the Corsica Channel (north-western Mediterranean Sea) detected by a moored acoustic Doppler current profiler, Ocean Sci., 15, 631–649, https://doi.org/https://doi.org/10.5194/os-15-631-2019, 2019.



640



- Heath, M.: Size spectrum dynamics and the planktonic ecosystem of Loch Linnhe, ICES J. Mar. Sci., 52, 627–642, https://doi.org/https://doi.org/10.1016/1054-3139(95)80077-8, 1995.
- 620 Herron, R. C., Leming, T. D., and Li, J.: Satellite-detected fronts and butterfish aggregations in the northeastern Gulf of Mexico, Cont. Shelf Res., 9, 569–588, https://doi.org/10.1016/0278-4343(89)90022-8, 1989.
  - Hoskins, B. J.: The Mathematical Theory of Frontogenesis, Annu. Rev. Fluid Mech., 14, 131–151, https://doi.org/10.1146/annurev.fl.14.010182.001023, 1982.
  - Hunt, B. P. V., Carlotti, F., Donoso, K., Pagano, M., D'Ortenzio, F., Taillandier, V., and Conan, P.: Trophic pathways of phytoplankton size
- 625 classes through the zooplankton food web over the spring transition period in the north-west Mediterranean Sea, J. Geophys. Res. Oceans, 122, 6309–6324, https://doi.org/10.1002/2016JC012658, 2017.
  - Isla, A., Scharek, R., and Latasa, M.: Zooplankton diel vertical migration and contribution to deep active carbon flux in the NW Mediterranean, J. Marine Syst., 143, 86–97, https://doi.org/https://doi.org/10.1016/j.jmarsys.2014.10.017, 2015.
  - Joyce, T. M.: Varieties of ocean fronts. Baroclinic instability and ocean fronts, Tech. Rep. 83-41, Technical Report, 1983.
- 630 Kiørboe, T.: Turbulence, phytoplankton cell size, and the structure of pelagic food webs, in: Adv. Mar. Biol., vol. 29, pp. 1–72, Academic Press, https://doi.org/10.1016/S0065-2881(08)60129-7, 1993.
  - Kiørboe, T., Ceballos, S., and Thygesen, U. H.: Interrelations between senescence, life history traits, and behaviour in planktonic copepods, Ecology, 96, 2225–2235, https://doi.org/10.1890/14-2205.1, 2015.
  - Lavigne, H., D'Ortenzio, F., Ribera D'Alcalà, M., Claustre, H., Sauzède, R., and Gacic, M.: On the vertical distribution of the
- 635 chlorophyll a concentration in the Mediterranean Sea: a basin-scale and seasonal approach, Biogeosciences, 12, 5021–5039, https://doi.org/https://doi.org/10.5194/bg-12-5021-2015, 2015.
  - Le Fèvre, J.: Aspects of the biology of frontal systems, in: Adv. Mar. Biol., vol. 23, pp. 163–299, Academic Press, https://doi.org/10.1016/S0065-2881(08)60109-1, 1987.

Legendre, P. and Gallagher, E.: Ecologically meaningful transformations for ordination of species data, Oecologia, 129, 271–280, https://doi.org/https://doi.org/10.1007/s004420100716, 2001.

- Legendre, P. and Legendre, P.: Numerical ecology ordination of plant communities, https://doi.org/https://doi.org/10.1007/978-94-009-7989-5, 2012.
- Liu, G. M., Sun, S., Wang, H., Zhang, Y., Yang, B., and Ji, P.: Abundance of \*Calanus sinicus\* across the tidal front in the Yellow Sea, China, Fish. Oceanogr., 12, 291–298, https://doi.org/10.1046/j.1365-2419.2003.00253.x, 2003.
- 645 Lévy, M., Franks, P. J., and Smith, K. S.: The role of submesoscale currents in structuring marine ecosystems, Nat. Commun., 9, 4758, https://doi.org/https://doi.org/10.1038/s41467-018-07059-3, 2018.
  - Lê, S., Josse, J., and Husson, F.: FactoMineR: an R package for multivariate analysis, J. Stat. Softw., 25, 1–18, https://doi.org/https://doi.org/10.18637/jss.v025.i01, 2008.

Mangolte, I., Lévy, M., Haëck, C., and Ohman, M. D.: Sub-frontal niches of plankton communities driven by transport and trophic interac-

- tions at ocean fronts, Biogeosciences, 20, 3273–3299, https://doi.org/https://doi.org/10.5194/bg-20-3273-2023, 2023.
- Margirier, F., Testor, P., Heslop, E., Mallil, K., Bosse, A., Houpert, L., Mortier, L., Bouin, M.-N., Coppola, L., D'Ortenzio, F., Durrieu de Madron, X., Mourre, B., Prieur, L., Raimbault, P., and Taillandier, V.: Abrupt warming and salinification of intermediate waters interplays with decline of deep convection in the Northwestern Mediterranean Sea, Sci Rep, 10, 20923, https://doi.org/https://doi.org/10.1038/s41598-020-77859-5, 2020.





655 Mayot, N., D'Ortenzio, F., Taillandier, V., Prieur, L., de Fommervault, O. P., Claustre, H., Bosse, A., Testor, P., and Conan, P.: Physical and biogeochemical controls of the phytoplankton blooms in North Western Mediterranean Sea: A multiplatform approach over a complete annual cycle (2012–2013 DEWEX experiment), J. Geophys. Res. Oceans, 122, 9999–10019, https://doi.org/10.1002/2016JC012052, 2017.

Mańko, M. K., Merchel, M., Kwasniewski, S., and Weydmann-Zwolicka, A.: Oceanic fronts shape biodiversity of gelatinous zooplankton in

the European Arctic, Frontiers in Marine Science, 9, 941025, https://doi.org/https://doi.org/10.3389/fmars.2022.941025, 2022.
 McWilliams, J. C.: Submesoscale currents in the ocean, P. R. Soc. A, 472, 20160117, https://doi.org/https://doi.org/10.1098/rspa.2016.0117, 2016.

- 665 Millot, C. and Taupier-Letage, I.: Circulation in the Mediterranean Sea, in: The Mediterranean Sea, edited by Saliot, A., vol. 5K of *Handbook* of *Environmental Chemistry*, pp. 29–66, Springer Berlin Heidelberg, https://doi.org/10.1007/b107143, 2005.
  - Molinero, J. C., Ibanez, F., Souissi, S., Bosc, E., and Nival, P.: Surface patterns of zooplankton spatial variability detected by high-frequency sampling in the NW Mediterranean. Role of density fronts, J. Mar. Syst., 69, 271–282, https://doi.org/https://doi.org/10.1016/j.jmarsys.2005.11.023, 2008.
- 670 Morales-Márquez, V., Hernández-Carrasco, I., Simarro, G., Rossi, V., and Orfila, A.: Regionalizing the impacts of wind-and wave-induced currents on surface ocean dynamics: A long-term variability analysis in the Mediterranean Sea, J. of Geophys. Res.: Oceans, 126, https://doi.org/https://doi.org/10.1029/2020JC017104, 2021.

Motoda, S.: Devices of simple plankton apparatus, Memoirs of the Faculty of Fisheries Hokkaido University, 7, 73–94, 1959.

Munk, P., Hansen, B. W., Nielsen, T. G., and Thomsen, H. A.: Changes in plankton and fish larvae communities across hydrographic fronts
 off West Greenland, J. Plankton Res., 25, 815–830, https://doi.org/10.1093/plankt/25.7.815, 2003.

- Nival, P., Nival, S., and Thiriot, A.: Influence des conditions hivernales sur les productions phyto-et zooplanctoniques en Méditerranée Nord-Occidentale. V. Biomasse et production zooplanctonique — relations phyto-zooplancton, Mar. Biol., 31, 249–270, https://doi.org/https://doi.org/10.1007/BF00387153, 1975.
- Nowaczyk, A., Carlotti, F., Thibault-Botha, D., and Pagano, M.: Distribution of epipelagic metazooplankton across the Mediterranean Sea
   during the summer BOUM cruise, Biogeosciences, 8, 2159–2177, https://doi.org/https://doi.org/10.5194/bg-8-2159-2011, 2011.
- Ohman, M. D., Powell, J. R., Picheral, M., and Jensen, D. W.: Mesozooplankton and particulate matter responses to a deep-water frontal system in the southern California Current System, J. Plankton Res., 34, 815–827, https://doi.org/https://doi.org/10.1093/plankt/fbs028, 2012.

Owen, R. W.: Fronts and eddies in the sea: mechanisms, interactions, and biological effects, in: Analysis of Marine Ecosystems, pp. 197–233, 1981.

Millot, C.: Circulation in the western Mediterranean Sea, J. Mar. Syst., 20, 423–442, https://doi.org/https://doi.org/10.1016/S0924-7963(98)00078-5, 1999.

Olson, D. B., Hitchcock, G. L., Mariano, A. J., Ashjan, C. J., Peng, G., Nero, R. W., and Podesta, G. P.: Life on the edge: marine life and fronts, Oceanography, 7, 52–60, 1994.

<sup>Pakhomov, E. A. and Froneman, P. W.: Composition and spatial variability of macroplankton and micronekton within the Antarctic Polar Frontal Zone of the Indian Ocean during austral autumn 1997, Polar Biol., 23, 410–419, https://doi.org/https://doi.org/10.1007/s003000050462, 2000.</sup> 



695



Panaïotis, T., Poteau, A., Riquier, E. D., Catalano, C., Courchet, L., Motreuil, S., Coppola, L., Picheral, M., and Irisson, J. O.: Temporal evolution of plankton and particles distribution across a mesoscale front during the spring bloom, Limnol. Oceanogr., 9999, 1–15, https://doi.org/https://doi.org/10.1002/lno.12566, 2024.

Pedrotti, M. L. and Fenaux, L.: Dispersal of echinoderm larvae in a geographical area marked by upwelling (Ligurian Sea, NW Mediterranean), Mar. Ecol. Prog. Ser., 86, 217–227, 1992.

- Petrenko, A. A., Arnaud, M., Barrillon, S., Comby, C., Fuda, J.-L., Berline, L., Bosse, A., Rousselet, L., Rolland, R., Bouruet-Aubertot, P., Oms, L., Demol, M., Didry, M., Gastauer, S., Pacciaroni, M., Berta, M., d'Ovidio, F., Gregori, G., and Doglioli, A.: Complex 3-D oceanic velocities at SWOT scales exhibited during the spring 2023 BioSWOTMed cruise, EGU General Assembly 2024, Vienna, Austria, 14–19 Apr 2024, EGU24-17343, https://doi.org/10.5194/egusphere-egu24-17343, 2024.
- 700 Pinca, S. and Dallot, S.: Meso- and macrozooplankton composition patterns related to hydrodynamic structures in the Ligurian Sea (Trophos-2 experiment, April-June 1986), Mar. Ecol. Prog. Ser., 126, 49–65, https://doi.org/10.3354/meps126049, 1995.
  - Platt, T. and Denman, K.: Organisation in the pelagic ecosystem, Helgol. Wiss. Meeresunters., 30, 575–581, https://doi.org/10.1007/BF02207862, 1977.

Pollard, R. T. and Regier, L. A.: Vorticity and vertical circulation at an ocean front, J. Phys. Oceanogr., 22, 609-625,

- 705 Powell, J. R. and Ohman, M. D.: Covariability of zooplankton gradients with glider-detected density fronts in the Southern California Current System, Deep-Sea Res. II, 112, 79–90, https://doi.org/https://doi.org/10.1016/j.dsr2.2014.04.002, 2015a.
  - Powell, J. R. and Ohman, M. D.: Changes in zooplankton habitat, behavior, and acoustic scattering characteristics across glider-resolved fronts in the Southern California Current System, Prog. Oceanogr., 134, 77–92, https://doi.org/http://dx.doi.org/10.1016/j.pocean.2014.12.011, 2015b.
- 710 Prentice, I. C.: Multidimensional scaling as a research tool in quaternary palynology: A review of theory and methods, Rev. Palaeobot. Palyno., 31, 71–104, https://doi.org/10.1016/0034-6667(80)90023-8, 1980.

Prieur, L. and Sournia, A.: "Almofront-1" (April-May 1991): an interdisciplinary study of the Almeria-Oran geostrophic front, SW Mediterranean, J. Mar. Syst., 5, 187–204, https://doi.org/https://doi.org/10.1016/0924-7963(94)90046-9, 1994.

Prieur, L., Bethoux, J., Bong, J., and Tailliez, D.: Particularités hydrologiques et formation d'eau profonde dans le bassin Liguro-Provencal
en 1981–1982, Rapp. Comm. Int. Mer Méditerr., 28, 51–53, 1983.

- Queiroz, N., Humphries, N., Noble, L., Santos, A., and Sims, D.: Spatial dynamics and expanded vertical niche of blue sharks in oceanographic fronts reveal habitat targets for conservation, PLoS ONE, 7, e32 374, https://doi.org/https://doi.org/10.1371/journal.pone.0032374, 2012.
  - Rao, C.: A review of canonical coordinates and an alternative to correspondence analysis using Hellinger distance, Qüestiió, 19, 1–3, 1995.
- 720 Rasband, W.: ImageJ, u.S. National Institutes of Health, Bethesda, Maryland, USA, 1997–2011. Roberts, D.: labdsv: Ordination and Multivariate Analysis for Ecology, https://CRAN.R-project.org/package=labdsv, r package version 2.1-0, 2023.
  - Rothschild, B. and Osborn, T.: Small-scale turbulence and plankton contact rates, J. Plankton Res., 10, 465–474, https://doi.org/10.1093/plankt/10.3.465, 1988.
- 725 Royer, F., Fromentin, J., and Gaspar, P.: Association between bluefin tuna schools and oceanic features in the western Mediterranean, Mar. Ecol. Prog. Ser., 269, 249–263, https://doi.org/https://doi.org/10.3354/meps269249, 2004.
  - Russell, R., Harrison, N., and Hunt, Jr, G.: Foraging at a front: hydrography, zooplankton, and avian planktivory in the northern Bering Sea, Mar. Ecol. Prog. Ser., 182, 77–93, https://doi.org/10.3354/meps182077, 1999.





- Sabatés, A., Olivar, M., Salat, J., Palomera, I., and Alemany, F.: Physical and biological processes controlling the distribution of fish larvae in the NW Mediterranean, Prog. Oceanogr., 74, 355–376, https://doi.org/10.1016/j.pocean.2007.04.017, 2007.
  - Saiz, E., Alcaraz, M., and Paffenhöfer, G.: Effects of small-scale turbulence on feeding rate and gross-growth efficiency of three Acartia species (Copepoda: Calanoida), J. Plankton Res., 14, 1085–1097, https://doi.org/https://doi.org/10.1093/plankt/14.8.1085, 1992.
    - Saiz, E., Calbet, A., Atienza, D., and Alcaraz, M.: Feeding and production of zooplankton in the Catalan Sea (NW Mediterranean), Progr. Oceanogr., 74, 313–328, https://doi.org/10.1016/j.pocean.2007.04.004, 2007.
- Saiz, E., Sabatés, A., and Gili, J.-M.: The Zooplankton, in: The Mediterranean Sea: Its History and Present Challenges, edited by Goffredo, S. and Dubinsky, Z., pp. 183–212, Springer Science, https://doi.org/https://doi.org/10.1007/978-94-007-6704-1\_11, 2014.
   Salat, J.: Interaction entre le courant Catalan et celui des Baléares dans le sud de la mer Catalane, Oceanol. Acta, 18, 227–234, 1995.
  - Scotto di Carlo, B., Ianora, A., Fresi, E., and Hure, J.: Vertical zonation patterns for Mediterranean copepods from the surface to 3000 m at a fixed station in the Tyrrhenian Sea, J. Plankton Res., 6, 1031–1056, https://doi.org/https://doi.org/10.1093/plankt/6.6.1031, 1984.
- 740 Severin, T., Kessouri, F., Rembauville, M., Sánchez-Pérez, E., Oriol, L., Caparros, J., Pujo-Pay, M., Ghiglione, J.-F., D'Ortenzio, F., Taillandier, V., Mayot, N., Durrieu De Madron, X., Ulses, C., Estournel, C., and Conan, P.: Open-ocean convection process: A driver of the winter nutrient supply and the spring phytoplankton distribution in the Northwestern Mediterranean Sea, J. Geophys. Res. Oceans, 122, 4587–4601, https://doi.org/10.1002/2016JC012664, 2017.

Seyfried, L., Estournel, C., Marsaleix, P., and Richard, E.: Dynamics of the North Balearic Front during an autumn tramontane and mistral

- 745 storm: air-sea coupling processes and stratification budget diagnostic, Ocean Sci., 15, 179–198, https://doi.org/https://doi.org/10.5194/os-15-179-2019, 2019.
  - Siokou, I., Zervoudaki, S., Velaoras, D., Theocharis, A., Christou, E., Protopapa, M., and Pantazi, M.: Mesozooplankton vertical patterns along an east-west transect in the oligotrophic Mediterranean Sea during early summer, Deep-Sea Res. II, 164, 170–189, https://doi.org/https://doi.org/10.1016/j.dsr2.2019.02.006, 2019.
- 750 Siokou-Frangou, I., Christaki, U., Mazzocchi, M., Montresor, M., Ribera d'Alcalá, M., Vaqué, D., and Zingone, A.: Plankton in the open Mediterranean Sea: a review, Biogeosciences, 7, 1543–1586, https://doi.org/10.5194/bg-7-1543-2010, 2010.
  - Sprules, W., Jin, E., Herman, A., and Stockwell, J.: Calibration of an optical plankton counter for use in fresh water, Limnol. Oceanogr., 43, 726–733, https://doi.org/10.4319/lo.1998.43.4.0726, 2003.
- Stemmann, L., Prieur, L., Legendre, L., Taupier-Letage, I., Picheral, M., Guidi, L., and Gorsky, G.: Effects of frontal processes on marine
   aggregate dynamics and fluxes: An interannual study in a permanent geostrophic front (NW Mediterranean), J. Mar. Syst., 70, 1–20, https://doi.org/https://doi.org/10.1016/j.jmarsys.2007.02.014, 2008.
  - Team, R. C.: R: A language and environment for statistical computing, https://www.R-project.org/, r Foundation for Statistical Computing, Vienna, Austria, 2025.
  - Tesán Onrubia, J., Tedetti, M., Carlotti, F., Tenaille, M., Guilloux, L., Pagano, M., Lebreton, B., Guillou, G., Fierro-González, P., Guigue, C.,
- 760 Chifflet, S., Garcia, T., Boudriga, I., Belhassen, M., Bellaaj Zouari, A., and Bănaru, D.: Spatial variations of stable isotope compositions and biochemical content of size-fractionated plankton in the Mediterranean Sea (MERITE-HIPPOCAMPE campaign). Special volume "Plankton and Contaminants in the Mediterranean Sea: Biological pump and interactions from regional to global approaches Plankton and Contaminants", Mar. Pollut. Bull., 189, 114787, https://doi.org/https://doi.org/10.1016/j.marpolbul.2023.114787, 2023.
  - Thibault, D., Gaudy, R., and Le Fèvre, J.: Zooplankton biomass, feeding and metabolism in a geostrophic frontal area (Almeria-Oran
- Front, western Mediterranean). Significance to pelagic food webs, J. Mar. Syst., 5, 297–311, https://doi.org/10.1016/0924-7963(94)90052-3, 1994.





- Vandromme, P., Stemmann, L., García-Comas, C., Berline, L., Sun, X., and Gorsky, G.: Assessing biases in computing size spectra of automatically classified zooplankton from imaging systems: A case study with the ZooScan integrated system, Methods Oceanogr., 1-2, 3–21, https://doi.org/https://doi.org/10.1016/j.mio.2012.06.001, 2012.
- 770 Wishner, K., Outram, D., and Ullman, D.: Zooplankton distributions and transport across the northeastern tidal front of Georges Bank, Deep-Sea Res. II, 53, 2570–2596, https://doi.org/10.1016/j.dsr2.2006.06.003, 2006.