

Tectonics

December 2008; Volume 27 (TC6006) : Pages 1-20

<http://dx.doi.org/10.1029/2008TC002263>

© 2008 American Geophysical Union

Archimer <http://www.ifremer.fr/docelec/>

Archive Institutionnelle de l'Ifremer

An edited version of this paper was published by [AGU](#).

Tectonic history of northern New Caledonia Basin from deep offshore seismic reflection: Relation to late Eocene obduction in New Caledonia, southwest Pacific

Julien Collot^{1,2,*}, Louis Geli¹, Yves Lafoy², Roland Vially³, Dominique Cluzel⁴,
Frauke Klingelhoefer¹, and Hervé Nouzé¹

¹ Department of Geodynamics and Geophysics, Ifremer, Centre de Brest, Plouzané, France.

² Service de la Géologie de Nouvelle Calédonie, Direction de l'Industrie, des Mines et de l'Energie de Nouvelle Calédonie, Nouméa, New Caledonia.

³ Institut Français du Pétrole, Rueil-Malmaison, France.

⁴ Institut des Sciences de la Terre d'Orléans, University of Orleans, CNRS UMR 6113, Orléans, France.

*: Corresponding author : Collot J., email address : julien.collot@ifremer.fr

Abstract:

New, high-quality multichannel seismic reflection data from the western New Caledonia offshore domain allow for the first time the direct, continuous connection of seismic reflectors between the Deep Sea Drilling Project 208 drill hole on the Lord Howe Rise and the New Caledonia Basin. A novel seismic interpretation is hence proposed for the northern New Caledonia Basin stratigraphy, which places the Eocene/Oligocene unconformity deeper than previously thought and revisits the actual thickness of the pre-Oligocene sequences. A causal link is proposed between the obduction of the South Loyalty Basin over New Caledonia (NC) and the tectonic history of the northern New Caledonia Basin. Here it is suggested that as the South Loyalty Basin was being obducted during early Oligocene times, the NC Basin subsided under the effect of the overloading and underthrust to accommodate the compressional deformation, which resulted in (1) the uplift of the northern Fairway Ridge and (2) the sinking of the western flank of New Caledonia. This event also had repercussions farther west with the incipient subsidence of the Lord Howe Rise.

Keywords: obduction; seismics; New Caledonia.

I. Introduction

The South West Pacific region east of Australia consists of a succession of ridges and basins, that result from the fragmentation of Gondwanaland since Cretaceous time [Symonds et al., 1996; Gaina et al., 1998; Auzende et al., 2000b; Crawford et al., 2002]. From west to east: the Tasman sea, the Dampier Ridge (DR), the Middleton Basin (MB), the Lord Howe Rise (LHR), the Fairway Basin (FB), the Fairway Ridge (FR), the New Caledonia Basin (NCB), the Norfolk Ridge (NR), the South Loyalty Basin (SLB) and the Loyalty Ridge (LR). Since late Miocene, this region is the foreland area of the Vanuatu Subduction Zone where the Australian Plate subducts beneath the Pacific Plate. A crustal flexure associated to this process is observed with the uplift of the Loyalty Islands [Dubois et al., 1974; Dubois et al., 1977; Guyomard, 1996] and the southern tip of New Caledonia and the Island of Pines [Launay and Recy, 1972; Launay, 1985; Cabioch, 1988; Cabioch et al., 1996]. Vertical motions from the 125 Ka reef indicate different tectonic blocks and a general slow subsidence (0.03 to 0.16 mm/year) of New Caledonia west coast [Lagabrielle et al., 2005]. This paper more particularly focuses on the tectonics and the sedimentary history of the New Caledonia Basin. The New Caledonia Basin is generally thought to have formed during the Cretaceous, based on rifting evidence from the Taranaki Basin, or during the Late-Cretaceous-early Paleocene [Burns et al., 1973a; Ravenne et al., 1976; Mignot, 1984; Collot et al., 1987; Uruski and Wood, 1991; Willcox et al., 2001; Lafoy et al., 2005a; Lafoy et al., 2005b], as mid-Paleocene fossil oozes were found at DSDP 206 [Burns et al., 1973b] drill site (see Figure 1 for location

and Figure 2 for stratigraphy). The basin can be divided into two parts, having different structural characteristics, separated by a NE-trending morpho-tectonic feature at 23°S (Figure 3) [Lafoy *et al.*, 2005a] : based on recent, wide-angle seismic data [Klingelhoefer *et al.*, 2007], the NW-trending northern New Caledonia Basin (NNCB) is characterized by a thick sedimentary cover and a 15 km thick crust of undetermined type with velocities of 6.4-7.4 km/s, whereas the N-trending Southern New Caledonia Basin has a thinner sedimentary cover with a back-arc basin type oceanic crust (8-9 km thick) with velocities of 6 km/s. Although the nature of the crust underlying the New Caledonia and Fairway basins is now documented, the history of the sedimentary infill remains unclear, mainly because of the lack of age determinations.

Rock outcrops on land provide well dated milestones that help us understand the tectonic history of New Caledonia [Avias, 1958; Aitchison, 1995; Cluzel *et al.*, 1997; Aitchison *et al.*, 1998; Baldwin *et al.*, 2007]. By contrast, in the offshore domain, the ages of the sedimentary sequences are known only at a limited number of widely spaced, isolated drill hole sites (Figure 1). The base of the Eocene/Oligocene unconformity that is documented at some sites has been used for a long time as a major, regional milestone for seismic stratigraphy. However, due to the lack of seismic data, no direct link was possible until now between the borehole sites and the New Caledonia Basin.

In this paper, we use all available seismic data, including new high-quality, 3.5 km long 360-channel, seismic reflection data collected during the Noucaplac-2 [Loubrieu *et al.*, 2004] (August 2004) and Zoneco-11 [Lafoy *et al.*, 2004] (September 2004) cruises of R/V L'Atalante to construct NE-trending transects between the Lord Howe Rise and the western margin of the New Caledonia / Norfolk Ridge system, which, for the first time ever, extend the Eocene/Oligocene unconformity documented at the DSDP 208 Drill hole site [Burns *et al.*, 1973d], without interruption into the NNCB. The previous attempts that have been made

by correlating the seismic reflectors within this basin [Van de Beuque, 1999; 2003; Vially *et al.*, 2003] are hence revisited.

II. Marine Stratigraphy from regional borehole data: the Eocene/Oligocene unconformity

Few boreholes were drilled in the geological structures linked to the New Caledonia Basin. The eight closest holes were drilled during the early seventies, during leg 21 and leg 90 of the DSDP program. These sites are DSDP 587, DSDP 591, DSDP 590, DSDP 593, DSDP 206, DSDP 207, DSDP 208 and DSDP 284 (Figure 1). All sites but one (DSDP 206) are located in relatively shallow areas, on the crest or on the flanks of the Lord Howe Rise and the penetration beneath seafloor varies from one site to another:

- five sites (small blue circles on Figure 1), DSDP 587, DSDP 591, DSDP590, DSDP593 and DSDP 284, do not reach the Eocene/Oligocene section [Kennett *et al.*, 1975; Kennett James *et al.*, 1986a; c; b; Kennett James *et al.*, 1986d]
- on the other hand, DSDP 206, 207 and 208 [Burns *et al.*, 1973d; c; b] (big blue circles on Figure 1) penetrated Eocene and older rocks, and are important to this study. Late Eocene to mid Oligocene sediments are absent in the regional unconformity (U1) at all three sites. Late Paleocene and early Eocene sediments are also missing (U2) at all three sites [Burns *et al.*, 1973d; c; b]. U1 has a common base over all three sites (mid-Eocene) whereas its top varies from mid Oligocene to mid Miocene (Figure 2-D).

DSDP208 being the nearest hole of the studied area and the only one that intersects our available seismic line FAUST1-S206-2 [Lafey *et al.*, 1998], it will be taken as the reference for this interpretation.

Taking into account U1 at site DSDP208, we suggest that erosion of the Lord Howe Rise started after the mid Eocene (ca 45 Ma) [Burns *et al.*, 1973d], and that the northern part of the

LHR returned to bathyal water depth in the early Oligocene, ca 34-28.5 Ma [*Burns et al.*, 1973d]

U1 has a particular property of being the boundary between calcareous sediments (overlying) and calcareous sediments with an important radiolarian siliceous component (underlying) [*Burns et al.*, 1973d]. The presence of radiolarians in the underlying layer induces a contrast of porosity between the two layers and therefore a contrast of density with a less dense underlying layer. The reflection coefficient associated with U1 is therefore negative and a local negative jump of seismic and sonic velocities is observed. The bottom of hole DSDP 208 reaches Campanian sedimentary strata.

III. Seismic data description

The seismic lines that we use here to connect DSDP 208 site to the NNCB, were all collected after 1998 (in the digital era), using modern multi-channel seismic facilities (see details in Table 1). These lines (underlined in yellow in Figure 3) are the following:

- Lines FAUST1-S206 from the French-Australian-Seismic-Transect survey [*Lafoy et al.*, 1998] (also known as LHRNC - Lord Howe Rise / New Caledonia) -2 and -1 collected in 1998 with R/V Rig Seismic using a 3.3 km-long, 264-channel streamer and a 3000 cu-inch, tuned, seismic array producing signal in the 50-60 Hz frequency band. Line FAUST1-S206-2 runs over DSDP 208 drill site, heading northward along the western, upper slope of the Lord Howe Rise. Line FAUST1-S206-1 bears northeast, from the crest of the Lord Howe Rise to the Fairway Basin.
- High Resolution seismic lines from the Fairway Basin collected during the ZoNeCo-11 cruise [*Lafoy et al.*, 2004] (September 2004) to characterize the reflectors that were once interpreted as gas hydrate-related bottom simulating reflectors (BSR) [*Exon et al.*, 1998; *Auzende et al.*, 2000a; *Pecher*, 2004].

These lines are used to connect the Fairway Basin and regional seismic transects across the New Caledonia Basin.

- Deep frequency, regional seismic transects collected in 2004 during the Noucaplac-2 [Loubrieu *et al.*, 2004] and Zoneco-11 [Lafoy *et al.*, 2004] cruises of R/V L'Atalante, using a 4.5 km long, 360-channel streamer, together with a large (8000 cu-inch) air gun array tuned in the single bubble mode [Avedik *et al.*, 1993]. These lines fill the data gap that used to exist between the Fairway Basin and the New Caledonia Basin.

In addition we use the ZoNeCo-5 lines [Auzende *et al.*, 1999] collected in 1999 (and reprocessed in 2006) with R/V L'Atalante, using a 300-m long, 6-channel streamer and two 250 cu-inch GI-Guns. Although the penetration of the seismic energy is restricted to the upper part of the sediment cover (the maximum penetration depth is ~ 1.5 to 2 s twt), these closely spaced lines (spacing is less than 12 km) are of particular interest because they cover the entire Fairway Basin and the eastern flank of the Lord Howe Rise.

IV. Linking DSDP208 stratigraphy to the New Caledonia Basin

IV.1 Correlation between seismic and drill hole data at DSDP 208

The starting point is the DSDP 208 borehole, where the Eocene/Oligocene unconformity (U1) has been found to be at 488 m below seafloor. Using the velocity analysis of Van de Beuque *et al.* [1998a] and Van de Beuque [2003], a correspondence is established between the unconformity and reflector RN on seismic line FAUST1-S206-2 (Figure 2-E). Reflector RN has two remarkable characteristics:

- it corresponds to a seismic phase inversion, because the presence of radiolarians in Eocene sediments induces lower seismic velocity in the lower layer (see above, section II). The phase inversion can be followed all the way along seismic line FAUST1-S206-

2, except in some places where the unconformity sits on basement (basement which corresponds to the top of the crust, confirmed by the wide angle seismic data), in which case RN swaps to a positive phase [Nouzé *et al.*, 2005; Nouzé *et al.*, 2007] (Figure 4)

- Sediments overlying RN have a characteristic, transparent seismic facies with a single intra-Miocene reflector (highlighted in pink in Figure 2-E) whereas sediments underlying RN present rather well bedded sequences.

These two well-marked characteristics are used hereafter to identify and follow reflector RN over the Lord Howe Rise, the Fairway Basin, the Fairway Ridge and the New Caledonia Basin. The path we use is highlighted in yellow on Figure 3.

IV.2 Link from LHR down into the Fairway Basin

Reflector RN is continuous along profile FAUST1-S206-2, from Site DSDP 208 to the junction with FAUST1-S206-1 (Figure 5). The reflector can then be followed along line FAUST1-S206-1 (Figure 6), up to the crest of Lord Howe Rise. On the eastern flank of Lord Howe Rise, a reflector is found that is related to a seismic phase inversion and located at the base of a seismically transparent layer. Hence, this reflector is identified as being RN. This interpretation is strongly supported by the presence of top laps truncated by this reflector (black arrows on Figure 6). The interpolation of RN from the crest of Lord Howe Rise down to the Fairway Basin is shown in Figure 6. Because of the presence of the Fairway Ridge, it is difficult to follow reflector RN all the way along line FAUST1-S206-1 down to the New Caledonia Basin. We thus use seismic profiles Z11-11, Z11-10 and Z11-09 (Figure 7, Figure 8, Figure 9) that link the northern and the central parts of the Fairway Basin (see location on Figure 3).

In the northern Fairway Basin, a specific seismic survey was conducted during the Zoneco-11 cruise near 23°15'S - 163°30'E, in order to confirm or deny the presence of gas hydrates in

the area. Five high resolution profiles, simultaneously recorded on a 3.5 km long 360 channel seismic streamer and on 12 Ocean Bottom Seismometers (OBSs), provided the detailed velocity structure of the substratum, in a narrow zone of 30x40 km². *Nouzé et al.* [2005; 2007] identified two distinct reflectors (RN and RP on Figure 7), both located at the base of a seismically transparent, upper sedimentary sequence: the first one (identified as RN) is located between 3 and 3.5 s two-way-time-travel (s-twt) below sea level and related to a seismic phase inversion, whereas the second one (identified as RP) is a bottom simulating reflector, as documented by previous authors [*Exon et al.*, 1998; *Auzende et al.*, 2000a; *Exon et al.*, 2004; *Pecher*, 2004]. The presence of a seismically transparent layer overlying reflector RN and the phase inversion suggests that RN likely corresponds to the regional Eocene Oligocene unconformity.

Profiles z11-11, z11-10 and z11-09 (Figure 7, Figure 8, Figure 9) allow us to follow reflector RN further south in the Fairway Basin, down to the intersection with profile NCP2-2 (Figure 10).

IV.3 Linking the Fairway Basin to the New Caledonia Basin across the Fairway Ridge

Profile NCP2-2 (Figure 10) is the only profile that enables us to continuously follow reflector RN, from the Fairway Basin to the New Caledonian Basin without interruption across the Fairway Ridge. Hence, it is the key profile for this new seismic interpretation. NCP2-2 shows a clear and continuous reflector RN, which runs over the Fairway Ridge and down in the New Caledonian Basin where all other eastward oriented profiles reveal RN as an onlap reflector truncated by the Fairway Ridge (see for example profile z11-01A on Figure 11) or lead to ambiguous interpretations using a seismic facies recognition method. The position of the Eocene-Oligocene unconformity in the New Caledonia Basin is the main new result of this

interpretation, positioning it deeper than in previous interpretations [*Van de Beuque*, 1999; 2003; *Vially et al.*, 2003].

Following RN along profile z11-04 (Figure 12) northward, we identify the position of RN in the northern New Caledonian Basin, and match it to the tilted reflector observed on profile z11-01A (Figure 13). Previous authors [*Willcox et al.*, 1980; *Exon et al.*, 1998; *Lafoy et al.*, 1998; *Van de Beuque*, 1999; *Auzende et al.*, 2000a; *Vially et al.*, 2003; *Lafoy et al.*, 2005b] interpreted the New Caledonia Basin as Cretaceous-Paleocene and placed the Eocene Oligocene unconformity U1 around 4 - 5 stwt below sealevel (bsl) in the New Caledonia Basin and around 4 stwt (bsl) in the Fairway Basin. U1's position is here confirmed in the Fairway Basin and revised in the NCB where we suggest it to be at 8 stwt at its deepest position.

IV.4 A rapid post-early Oligocene eastward-tilt of the NNCB

The stratigraphic analysis of the profile z11- 01A suggests that the depocenter of the NNCB migrated eastward before late Oligocene, as a response to a rapid eastward tilt of its basement:

(1): Well-stratified reflectors onlapping the pre-Oligocene series beneath RN on profile z11-01A (black arrows pointing to the right on Figure 13) indicate the existence of a tilted paleo-basin that had its depocenter beneath the western flank of the present-day NNCB (see reconstitution of profile z11-01A's history on Figure 14). This basin was tilted eastward during RN time as indicated by its eastward-dipping reflectors (dark blue reflector).

(2): The post-early Oligocene onlap series (black arrows pointing to the left on Figure 13) does not show evidence for syntectonic sedimentation, which would be distinguished by a fanning reflector pattern. Instead the entire post-early Oligocene fill is horizontal, with parallel-bedding. The geometrical relationship between unconformity RN and the basin stratification denotes a rapid tilt of the basin that terminated prior to massive post-early

Oligocene sedimentation. The thin post-RN sequence (top of which is marked by a dark green reflector on Figure 13) on the flank of the FR (SP3000) is draped sediment or, more likely, contourites.

These observations indicate that the NNCB's history evolved following two major phases: pre- and post-early Oligocene, with a major tectonic event as boundary.

V. Implications

V.1 Causal relationship with early Oligocene obduction in New Caledonia

Seismic reflector RN (the Eocene/Oligocene unconformity) dips eastward in the NNCB (profile z11-01A on Figure 13), indicating that the basin subsided, whereas the Fairway Ridge was uplifted some time in-between mid-Eocene and early Oligocene times (i.e. the upper and lower age boundaries of the tilt are ~ 45 and 34-28.5 Ma respectively, corresponding to the dating of U1 in the works of *Burns et al.* [1973d]). The tilting of FR appears to be small compared to the subsidence of the NNCB which denotes that the formation of a foredeep basin is the significant part of the event. This new result has several implications regarding the New Caledonia Basin in terms of age, tectonic and sedimentary evolution and links with its neighboring structural elements. Taking a close look at the New Caledonia onshore geology [*Avias*, 1967; *Brothers and Blake*, 1973; *Prinzhofer et al.*, 1980; *Paris*, 1981; *Collot et al.*, 1987; *Aitchison and Meffre*, 1992; *Aitchison*, 1995; *Cluzel et al.*, 1997; *Aitchison et al.*, 1998; *Cluzel*, 1998; *Cluzel et al.*, 1998; *Ali and Aitchison*, 2000; *Cluzel et al.*, 2001; *Ali and Aitchison*, 2002; *Spandler et al.*, 2004; *Fitzherbert et al.*, 2005; *Spandler et al.*, 2005; *Cluzel et al.*, 2006], we here propose to establish links between the marine tectonic event observed in the seismic data and the tectonic events observed onshore.

Rigolot and Pelletier [1988] and *Rigolot* [1989] have described from single channel seismic data (ZOE-400) compressional features, reverse faulting and duplexes along the Western margin of New Caledonia. They claimed that this compressive phase was active during New Caledonia's obduction episode from latest Eocene until late Miocene/Pliocene and resulted from displacements along a major boundary located west of New Caledonia.

The mafic allochthon that crops out over more than 250 km along the west coast of New Caledonia (the Poya Terrane), and extends further North over 150 km beneath the Belep Islands [*Collot et al.*, 1987] is evidence of the South Loyalty Basin. On micropaleontological and geochemical evidence, the South Loyalty Basin oceanic crust is known to have been formed from the Campanian to earliest Eocene in a back-arc basin setting [*Aitchison*, 1995; *Cluzel et al.*, 2001]. *Cluzel et al.* [2001] proposed three main steps regarding the evolution of the South Loyalty Basin (Figure 15-a,b,c):

1. Ca. 53Ma, Phase 1 (Figure 15-a): inception of east-dipping intra-oceanic subduction of the South Loyalty Basin and onset of the Loyalty Arc (LR). This phase is dated at 53 Ma by the arc related magmatism (dykes) that crosscut the New Caledonian ophiolites without piercing the underlying strata [*Cluzel et al.*, 2006]. These dykes testify that the peridotites, at present obducted upon New Caledonia, were part of the upper plate in the subduction system
2. 53-37 Ma, Phase 2: subduction of the western South Loyalty Basin. This event is recorded by the Pouebo Terrane, an eclogitised mélange that contains mafic rocks from the Poya Terrane. At ca. 45 Ma, elements of the oceanic crust from the South Loyalty Basin were brought into the subduction zone down to 70 km equivalent pressure depth [*Clarke et al.*, 1997; *Spandler et al.*, 2004; *Spandler et al.*, 2005]

3. 37-34 Ma, Phase 3: the arrival of the Norfolk Ridge near the trench at ca. 37 Ma, progressively blocked the subduction and initiated a compressive phase. This phase is expressed by three diachronous events observed onshore–New Caledonia:
 - i. the thrusting of mafic oceanic crust slices scraped off the South Loyalty Basin (the Poya Terrane) onto the Norfolk Ridge at 37 Ma, is recorded by the foreland deposits (the Bourail flysch of [Cluzel *et al.*, 2001]) (Figure 15-b)
 - ii. the exhumation (supposedly buoyancy driven, e.g. [Chemenda *et al.*, 1996]) between 44 and 34 Ma of eclogitized mafic mélange (the Pouebo Terrane) and Cretaceous metasediments (the Diahot Terrane) in northern New Caledonia [Spandler and Hermann, 2006; Baldwin *et al.*, 2007]
 - iii. the Late Eocene obduction of the ultramafic lower part of the upper plate (the Ophiolitic Nappe) onto the Norfolk Ridge (Figure 15-c,d). The end of this phase is poorly constrained by the late Priabonian sediments (ca. 34 Ma) that underlie the ophiolites [Cluzel, 1998], and mid Oligocene post-obduction granitoids at ca. 27 Ma [Cluzel *et al.*, 2005; Paquette and Cluzel, 2006]

We here hypothesize that the tilting of the NNCB (the age range of which is bounded between 45 Ma and 34-28.5 Ma) occurred during this compressive phase (37-34 Ma).

Prior to Oligocene times the depocenter was on the western side of the NNCB (observation 1 of section IV.4) implying a westward dip of the NCB basement.

As documented by the post-obduction tectonosedimentary records [Dubois *et al.*, 1974; Coudray, 1976; Chevillotte, 2005; Chardon and Chevillotte, 2006; Chevillotte *et al.*, 2006],

New Caledonia underwent a post-obduction uplift through erosional unloading of the high-density obducted allochthon. However, by contrast, based on the analysis of seismic data, no vertical movements were involved in the NNCB after Oligocene times (observation (2) of section IV-4). Together, these facts indicate that the NNCB and the NR did not react as a single block after Oligocene times but rather that the differential motion must have been controlled by a structural break (decoupling) between the NNCB and the NR. We suggest that this structural limit was inherited from the eastward thrusting during obduction described below (part 2 of the diachronic model).

In this context, we propose a diachronous model explaining the rise of the northern Fairway Ridge and the associated creation of a vast depot-center along New Caledonia's western margin :

1. As the NR was plunging into the subduction zone and subsiding, the NR and the NNCB reacted as a single rigid block. The NNCB initiated its eastward tilt by accompanying this downward motion and acted therefore as a subsiding flexural basin (Figure 15-C)
2. As the horizontal compressive stress increased within the obduction taking place, the NR and the NNCB decoupled into two separate blocks, the NNCB absorbing part of the shortening by underthrusting beneath the NR. This differential motion between the two blocks is simplified into a single thrust-fault on Figure 15-d, as documented by the analogue modelling of a horizontal lithospheric compression physical modelling by *Shemenda* [1992], Figure 15-d'). This hypothesis further places the Eocene/Oligocene New Caledonia Basin stage close to that of the initiation of a subduction. The present knowledge and available marine-data do not allow us to give an objective answer to this question of "subduction along New Caledonia's western margin" suggested by *Regnier* [1988] and *Cluzel et al.* [2005]. Discussing whether or not true-subduction

occurred and continued until late Oligocene times in the NNCB is beyond of the scope of this paper.

The positive topography of the Fairway Ridge is well expressed only between 20°S and 24°S, where it faces New Caledonia. This suggests that the FR and the NCB are tectonically related north of 24°S: the Fairway Ridge uplifted to form a prominent topographic high only where it faces New Caledonia, contributing to the tilt of the NNCB. The regional gravity map south of 24°S indicates that the Fairway Ridge extends southward down to 31°S in a N-S direction [Dupont *et al.*, 1975; Ravenne *et al.*, 1976; Eade, 1988; Lafoy *et al.*, 2005a]. The seismic profiles recorded from R.V. Franklin in 2001 [Exon *et al.*, 2004] show that an intermittent ridge extends southward from the NW-trending Fairway Ridge, to join the West Norfolk Ridge SSW of Norfolk Island. Exon *et al.* [2007] named this intermittent ridge the northern West Norfolk Ridge.

V.2 Causal relationship between early Oligocene obduction and LHR's return to marine environment

Because reflector RN coincides with the regional Eocene-Oligocene unconformity U1 observed at DSDP208, it is here suggested that the NNCB tilting is synchronous with LHR's return to a bathyal marine environment: while the NNCB was tilting and the Fairway Ridge rising, 300 km further west the LHR was subsiding. In the literature, this return to a bathyal marine environment of the LHR is either attributed to its subsidence associated with a major tectonic event [Launay *et al.*, 1977; Lafoy *et al.*, 1994; Van de Beuque *et al.*, 1998b] or to variations of the oceanic currents related to the opening between Australia-Antarctica in Eocene times [Andrews *et al.*, 1973]. Profile z05-12 (Figure 16) right on top of the LHR (see location on Figure 3) shows U1 as an angular unconformity with a clear erosional surface which favors a tectonic origin hypothesis. Variations of oceanic currents alone would show a

paraconformity corresponding to a lack of sedimentation with sub-parallel strata over and under the unconformity. The tilted strata underlying the hiatus being Eocene, we therefore reach the conclusion that the strata was tilted tectonically after the Eocene. LHR's return to a marine environment (marked by U1) is of post-Eocene tectonic origin.

The Eocene-Oligocene unconformity is region-wide, its base is roughly common to all wells but its top varies from a well to another (see chapter II and Figure 2). This means that the timing of the return in calm sedimentary deposit is different from a place to another. Looking at this fact on the LHR, it is remarkable to notice that the return in normal sedimentary deposit conditions is mid-Miocene in the southern LHR and mid-Oligocene in the northern LHR. This correlates very well with obduction in New Caledonia (37-34 Ma) and obduction in Northland - New Zealand (25-22 Ma) [*Ballance and Sporli, 1979; Brothers and Delaloye, 1982; Hayward et al., 1989; Herzer, 1992*] and shows the local uniqueness of the northern LHR's subsidence during Oligocene.

This leads us to propose that as the obduction event terminated in New Caledonia, the stress was simultaneously released further west which was manifested by the beginning of subsidence in the northern LHR. (Figure 15).

V.3 Age constrains on the sedimentary sequences within the NNCB

Using the refraction seismic velocities [*Klingelhoefer et al, 2006*], the absolute thickness of the sedimentary cover can be determined (non-sedimentary layers are discriminated as having velocities > 6 km/s).

The position of U1 (Figure 6 and Figure 13) indicates that the pre-Oligocene section is thinner than previously reported [*Van de Beuque, 1999; 2003; Vially et al., 2003*].

By comparing the sedimentary thicknesses of the Fairway (Figure 6) and New Caledonia (Figure 13) basins, it is noteworthy that the pre-Oligocene sediment of the Fairway Basin is much greater than that of New Caledonia (2.5 stwt against 1.1 stwt in the NNCB). On the

other hand the post Oligocene sediment thickness of the New Caledonia Basin is much greater than that of the Fairway Basin one (3.5 stwt against 0.5 stwt in the Fairway Basin). This observation is compatible with the fact that the LHR was in a sub-aerial state between 45-34/28.5 Ma and NC in a sub-aerial state after ca. 34 Ma.

We therefore propose that most of the thick pre-Oligocene sedimentary sequences of the Fairway Basin correspond to the product of erosion of the LHR and FR and most of the thick post-early Oligocene sedimentary sequences of the NNCB correspond to the product of erosion of the New Caledonian allochthon material. This corroborates the interpretation of *Jongsma and Mutter* [1978], who proposed that the Fairway Basin is the eastern equivalent of the Middleton Basin (west of LHR) and that their sedimentary infill corresponds with the erosion of the LHR.

CONCLUSIONS

The interpretation of reflector-RN and its trace without interruption from DSDP 208 to the NNCB enable to identify a post Eocene depocentre located adjacent to New Caledonia western margin which fits with the obduction and post-obduction story of New Caledonia. The uplifting of the NNCB's western margin (eg. the Fairway Ridge) and the deepening of its eastern margin is a direct result of New Caledonia's major obduction-compressive phase from 37 Ma to 34 Ma. The plunging of the Norfolk Ridge into the Loyalty subduction zone initiated the eastward subsidence of the NNCB, and compression continued as underthrusting under west of NC until the end of obduction to accommodate the shortening. As the shortening was taking place in New Caledonia and in the NNCB, stress was simultaneously released further west which was manifested by the beginning of subsidence in the LHR. As the LHR was subsiding and becoming more deeply submerged, NC was rising with the emplacement of the allochthon and becoming emergent. This is shown by the migration of the major sedimentary source from the LHR to NC, which is supported by our interpretation of

the position of the pre- and post-Oligocene sedimentary depocenters in the NNCB and the Fairway Basin.

ACKNOWLEDGEMENTS

The ZoNéco 11 cruise was conducted within the framework of the ZoNéCo program (Assessment of living and non-living resources of New Caledonia EEZ) funded by the ADECAL (Agency for the Economic development of New Caledonia). We thank the following people and organizations: IFREMER, IFP and the Service Géologique de Nouvelle Calédonie shipboard party of cruises ZoNéCo-11 and Noucaplac-2, and GENAVIR officers and crew of R/V l'Atalante. We are also grateful to W. Roest from IFREMER for allowing the Noucaplac-2 cruise (UNCLOS line) to be available for this study and to G. Bernardel from Geoscience Australia (former AGSO) for the co-owned FAUST I data New Caledonia / Australia. Special thanks to E. Cosquer for the data reprocessing of ZoNéCo 5, to Jean-Louis Olivet and François Bache for their great earth science knowledge and seismic interpretation tips. Maps were produced with the Generic Mapping Tool 4.2.1 (<http://gmt.soest.hawaii.edu>), seismic data processing with CGG-Geocluster and IFREMER-Sisbise; seismic interpretation with Seismic Micro Technology Inc.-The Kingdom Suite.

REFERENCES

- Aitchison, J., and S. Meffre (1992), New Caledonia a tectonic collage in the Southwest Pacific, *International Geological Congress, Abstracts*, 29.
- Aitchison, J. C. (1995), Eocene arc continent collision in New Caledonia and implications for regional southwest Pacific tectonic evolution, *Geology*, 23, 161-164.
- Aitchison, J. C., T. R. Ireland, G. L. Clarke, D. Cluzel, A. M. Davis, and S. Meffre (1998), Regional implications of U/Pb SHRIMP age constraints on the tectonic evolution of New Caledonia, *Tectonophysics*, 299, 333-343.
- Ali, J. R., and J. Aitchison (2002), Paleomagnetic-tectonic study of the New Caledonia Koh Ophiolite and the mid-Eocene obduction of the Poya Terrane, *New Zealand Journal of Geology and Geophysics*, 45, 313-322.
- Ali, J. R., and J. C. Aitchison (2000), Significance of palaeomagnetic data from the oceanic Poya Terrane, New Caledonia, for SW Pacific tectonic models, *Earth and Planetary Science Letters*, 177, 3-4.
- Andrews, J. E., R. E. Burns, M. Churkin, Jr., T. A. Davies, P. Dumitrica, A. R. Edwards, J. S. Galehouse, J. P. Kennett, G. H. Packham, and G. J. Van der Lingen (1973), Deep Sea Drilling Project: Leg 21; Tasman Sea-Coral Sea; (Preliminary Results); Oceanography of the South Pacific 1972, 185-199 pp.
- Auzende, J.-M., N. Exon, and S. Party (1999), Campagne ZoNéCo 5 (14 octobre - 7 novembre 1999) - Rapport de mission, 1-74 pp.
- Auzende, J.-M., S. Van de Beuque, G. Dickens, C. François, Y. Lafoy, O. Voutay, and N. Exon (2000a), Deep sea diapirs and bottom simulating reflector in Fairway basin (SW Pacific), *Marine Geophysical Researches*, 21, 579-587.
- Auzende, J.-M., S. Van de Beuque, M. Regnier, Y. Lafoy, and P. Symonds (2000b), Origin of the New Caledonian ophiolites based on a French-Australian Seismic Transect, *Marine Geology*, 162, 225-236.
- Avedik, F., V. Renard, J. P. Allenou, and B. Morvan (1993), "Single bubble" air-gun array for deep exploration, *Geophysics*, 58, 366-382.
- Avias, J. (1958), Sur l'existence d'une phase tectonique hercynienne tardive ayant affecté les formations antétriasiques de la cote ouest de la Nouvelle-Caledonie, *Comptes Rendus Hebdomadaires des Seances de l'Academie des Sciences*, 246, 136-137.
- Avias, J. (1967), Overthrust structure of the main ultrabasic New Caledonian massives, *Tectonophysics*, 4, 531-541.
- Baldwin, S., T. Rawling, and P. G. Fitzgerald (2007), Thermochronology of the New Caledonian high-pressure terrane: Implications for middle Tertiary plate boundary processes in the southwest Pacific, *Geological Society of America Special Paper*, 49.
- Ballance, P. F., and K. B. Sporli (1979), Northland Allochthon, *Journal of the Royal Society of New Zealand*, 9, 259-275.
- Brothers, R. N., and J. M. C. Blake (1973), Tertiary plate tectonics and high-pressure metamorphism in New Caledonia, *Tectonophysics*, 17, 337-358.
- Brothers, R. N., and M. Delaloye (1982), Obducted ophiolites of North Island, New Zealand: origin, age, emplacement and tectonic implications for Tertiary and Quaternary volcanicity, *New Zealand Journal of Geology and Geophysics*, 25, 257-274.
- Burns, R. E., J. E. Andrews, G. J. Van der Lingen, J. E. Andrews, M. Churkin, Jr., T. A. Davies, P. Dumitrica, A. R. Edwards, J. S. Galehouse, J. P. Kennett, and G. H. Packham (1973a), Regional aspects of deep sea drilling in the Southwest Pacific, Lithostratigraphy of

Eight Drill Sites in the South-west Pacific, *Preliminary Results of Leg 21 of the Deep Sea Drilling Project., Oceanography of the South Pacific* 1972.

Burns, R. E., J. E. Andrews, G. J. van der Lingen, M. Churkin, Jr., J. S. Galehouse, G. Packham, T. A. Davies, J. P. Kennett, P. Dumitrica, A. R. Edwards, and R. P. Von Herzen (1973b), Site 206, *Initial Reports of the Deep Sea Drilling Project - Leg 21*, 103-195.

Burns, R. E., J. E. Andrews, G. J. van der Lingen, M. Churkin, Jr., J. S. Galehouse, G. Packham, T. A. Davies, J. P. Kennett, P. Dumitrica, A. R. Edwards, and R. P. Von Herzen (1973c), Site 207, *Initial Reports of the Deep Sea Drilling Project - Leg 21*, 197-269.

Burns, R. E., J. E. Andrews, G. J. van der Lingen, M. Churkin, Jr., J. S. Galehouse, G. Packham, T. A. Davies, J. P. Kennett, P. Dumitrica, A. R. Edwards, and R. P. Von Herzen (1973d), Site 208, *Initial Reports of the Deep Sea Drilling Project - Leg 21*, 271-331.

Cabioch, G. (1988), Récifs frangeants de Nouvelle-Calédonie (Pacifique sud-ouest) - Structure interne et influences de l'eustatisme et de la néotectonique, PhD thesis, 291 pp, Université de Provence Aix Marseille I.

Cabioch, G., J. Recy, C. Jouannic, and L. Turpin (1996), Contrôle climatique et tectonique de l'édification récifale en Nouvelle Calédonie au cours du Quaternaire terminal, *Bulletin Société Géologique de France*, 167, 729-742.

Chardon, D., and V. Chevillotte (2006), Morphotectonic evolution of the New Caledonia ridge (Pacific Southwest) from post-obduction tectonosedimentary record, *Tectonophysics*, 420, 473-491.

Chemenda, A. I., M. Mattauer, and A. N. Bokun (1996), Continental subduction and a mechanism for exhumation of high-pressure metamorphic rocks: new modelling and field data from Oman, *Earth and Planetary Science Letters*, 143, 173-182.

Chevillotte, V. (2005), Morphogénèse tropicale en contexte epirogénique modéré, exemple de la Nouvelle Calédonie (Pacifique Sud Ouest), PhD thesis, 161 pp, University of Nouméa, Nouméa.

Chevillotte, V., D. Chardon, A. Beauvais, P. Maurizot, and F. Colin (2006), Long term tropical morphogenesis of New Caledonia (Pacific SW): importance of positive epeirogeny and climate change, *Geomorphology*, 81, 361-375.

Clarke, G., J. C. Aitchison, and D. Cluzel (1997), Eclogites and blueschists of the Pam Peninsula, NE New Caledonia: A reappraisal, *Journal of Petrology*, 38, 843-876.

Cluzel, D. (1998), Le flysch post-obduction de Nepoui, un bassin transporté? Consequences sur l'âge et les modalités de l'obduction tertiaire en Nouvelle-Calédonie (Pacifique sud-ouest), *Comptes Rendus de l'Académie des Sciences - Series IIA - Earth and Planetary Science*, 327, 419-424.

Cluzel, D., J. C. Aitchison, and C. Picard (2001), Tectonic accretion and underplating of mafic terranes in the late Eocene intraoceanic fore-arc of New Caledonia (Southwest Pacific): geodynamic implications, *Tectonophysics*, 340, 23-59.

Cluzel, D., D. Bosch, J. L. Paquette, Y. Lemennicier, P. Montjoie, and R. P. Menot (2005), Late Oligocene post-obduction granitoids of New Caledonia: A case for reactivated subduction and slab break-off, *Island Arc*, 14, 254-271.

Cluzel, D., D. Chiron, and M.-D. Courme (1998), Discordance de l'Eocène supérieur et événements pré-obduction en Nouvelle-Calédonie, *Comptes Rendus de l'Académie des Sciences - Series IIA - Sciences de la terre et des planètes*, 327, 485-491.

Cluzel, D., S. Meffre, P. Maurizot, and A. J. Crawford (2006), Earliest Eocene (53 Ma) convergence in the Southwest Pacific: evidence from pre-obduction dikes in the ophiolite of New Caledonia, *Terra Nova*, 18, 395-402.

Cluzel, D., C. Picard, J. C. Aitchison, C. Laporte, S. Meffre, and F. Parat (1997), La nappe de Poya (ex-formation des Basaltes) de Nouvelle-Calédonie (Pacifique Sud-Ouest) : un plateau

océanique Campanien-Paleocène supérieur obducté à l' Eocène supérieur, *Comptes rendus de l'Académie des sciences, Série II, Sciences de la terre et des planètes*, 324, 443-451.

Collot, J. Y., A. Malahoff, J. Recy, G. Latham, and F. Missegue (1987), Overthrust emplacement of New Caledonia ophiolite: geophysical evidence, *Tectonics*, 6, 215-232.

Coudray, J. (1976), Recherche sur le Néogène et le Quarternaire marin de la Nouvelle Calédonie. Contribution de l'étude sédimentologique à la connaissance de l'histoire géologique post-Eocène.

Crawford, A. J., S. Meffre, and P. A. Symonds (2002), 120 to 0 Ma tectonic evolution of the southwest Pacific and analogous geological evolution of the 600 to 220 Ma Tasman Fold Belt System, *Geological Society of Australia Special Publication*, 22, 377-397.

Dubois, J., J. Launay, and J. Recy (1974), Uplift movements in New Caledonia-Loyalty Islands area and their plate tectonics interpretation, *Tectonophysics*, 24, 133-150.

Dubois, J., J. Launay, J. Recy, and J. F. Marshall (1977), New Hebrides trench: subduction rate from associated lithospheric buldge, *Can. J. Earth. Sci.*, 14, 250-255.

Dupont, J., J. Launay, C. Ravenne, and C. E. De Broin (1975), Données nouvelles sur la ride de Norfolk (Sud Ouest Pacifique), *Comptes Rendus de l'Académie des Sciences, Série D*, 281.

Eade, J. V. (1988), The Norfolk Ridge System and its margins, in *The Ocean basin and margins*, edited by A. E. M. Nairn, Stehli, F.G and Uyeda, S. (eds), 7th ed., Plenum Press, New York & London, pp. 303-324.

Exon, N., G. Dickens, J.-M. Auzende, Y. Lafoy, P. Symonds, and S. van de Beuque (1998), Gaz hydrates and free gas on the Lord Howe Rise, Tasman Sea, *Pesa Journal*, 148-158.

Exon, N., P. Hill, Y. Lafoy, M. Fellows, K. Perry, P. Mitts, R. Howe, G. C. H. Chaproniere, G. Dickens, B. Ussler, and C. K. Paull (2004), Geology of the Fairway and New Caledonia basins in the Tasman Sea: sediment, pore water, diapirs and bottom simulating reflectors (Franklin cruise FR9/01 and Geoscience Australia Survey 232), *Geoscience Australia Records*, 2004, 1-112.

Exon, N., Y. Lafoy, P. J. Hill, G. Dickens, and I. Pecher (2007), Geology and petroleum potential of the Fairway Basin in the Tasman Sea, *Australian Journal of Earth Sciences*, 54, 629-645.

Fitzherbert, J. A., G. L. Clarke, and R. Powell (2005), Preferential retrogression of high-P metasediments and the preservation of blueschist to eclogite facies metabasite during exhumation, Diahot terrane, NE New Caledonia, *Lithos*, 83, 67-96.

Gaina, C., D. Müller, J. Royer, J. Stock, J. Hardebeck, and P. Symonds (1998), The tectonic history of the Tasman Sea: A puzzle with 13 pieces, *Journal of geophysical research*, 103, 12.413-412.433.

Guyomard, T. S. A., D.M; McNeill, D.F. (1996), Magnetostratigraphic dating of the uplifted atoll of Maré: Geodynamics of the loyalty Ridge, SW Pacific, *Journal of Geophysical Research*, 101, 601-612.

Hayward, B. W., F. J. Brook, and M. J. Isaac (1989), Cretaceous to middle Tertiary stratigraphy, paleogeography and tectonic history of Northland, New Zealand.; Geology of Northland; accretion, allochthons and arcs at the edge of the New Zealand micro-continent, *Bulletin Royal Society of New Zealand*, 47-64.

Herzer, R. H. (1992), The Northland Basin, New Zealand, from rift to active margin; tectonic evolution and petroleum potential, *AAPG Bulletin*, 76, 1107.

Jongsma, D., and J. C. Mutter (1978), Non-axial breaching of a rift valley: evidence from the Lord Howe Rise and the southeastern Australian margin, *Earth and Planetary Science Letters*, 39, 226-234.

Kennett James, P., C. von der Borch Christopher, A. Baker Paul, E. Barton Charles, A. Boersma, P. Caulet Jean, C. Dudley Walter, Jr., V. Gardner James, D. G. Jenkins, H. Lohman William, E. Martini, B. Merrill Russell, H. Morin Roger, S. Nelson Campbell, C. Robert, M.

S. Srinivasan, R. Stein, and A. Takeuchi (1986a), Site 590, *Initial Reports of the Deep Sea Drilling Project - Leg 90*, 263-376.

Kennett James, P., C. von der Borch Christopher, A. Baker Paul, E. Barton Charles, A. Boersma, P. Caulet Jean, C. Dudley Walter, Jr., V. Gardner James, D. G. Jenkins, H. Lohman William, E. Martini, B. Merrill Russell, H. Morin Roger, S. Nelson Campbell, C. Robert, M. S. Srinivasan, R. Stein, and A. Takeuchi (1986b), Site 591, *Initial Reports of the Deep Sea Drilling Project - Leg 90*, 377-486.

Kennett James, P., C. von der Borch Christopher, A. Baker Paul, E. Barton Charles, A. Boersma, P. Caulet Jean, C. Dudley Walter, Jr., V. Gardner James, D. G. Jenkins, H. Lohman William, E. Martini, B. Merrill Russell, H. Morin Roger, S. Nelson Campbell, C. Robert, M. S. Srinivasan, R. Stein, and A. Takeuchi (1986c), Site 593, *Initial Reports of the Deep Sea Drilling Project - Leg 90*, 551-651.

Kennett James, P., C. von der Borch Christopher, A. Baker Paul, E. Barton Charles, A. Boersma, P. Caulet Jean, C. Dudley Walter, Jr., V. Gardner James, D. G. Jenkins, H. Lohman William, E. Martini, B. Merrill Russell, H. Morin Roger, S. Nelson Campbell, C. Robert, M. S. Srinivasan, R. Stein, A. Takeuchi, and H. Blakeslee Jan (1986d), Site 587, *Initial Reports of the Deep Sea Drilling Project - Leg 90*, 115-138.

Kennett, J. P., R. E. Houtz, P. B. Andrews, A. R. Edwards, V. A. Gostin, M. Hajos, M. Hampton, D. G. Jenkins, S. V. Margolis, A. T. Ovenshine, and K. Perch Nielsen (1975), Site 284, *Initial Reports of the Deep Sea Drilling Project*, 29, 403-445.

Klingelhoefer, F., Y. Lafoy, J. Collot, E. Cosquer, L. Géli, H. Nouzé, and R. Vially (2007), Crustal structure of the basin and ridge system west of New Caledonia (southwest Pacific) from wide-angle and reflection seismic data *Journal of Geophysical Research*, 112.

Lafoy, Y., I. Brodien, R. Vially, and N. Exon (2005a), Structure of the basin and ridge system west of New Caledonia (Southwest Pacific): A synthesis, *Marine Geophysical Researches*, 00, 1-13.

Lafoy, Y., L. Géli, F. Klingelhoefer, R. Vially, B. Sichler, and H. Nouzé (2005b), Discovery of continental stretching and oceanic spreading in the Tasman sea, *Eos Trans Am Geophys Union*, 86, 101+104-105.

Lafoy, Y., L. Géli, R. Vially, F. Klingelhoefer, H. Nouzé, Y. Auffret, J. Bégot, D. Buisson, J. Collot, E. Cosquer, J. Crozon, A. Nercessian, S. Rouzo, S. Serbutoviez, B. Sichler, E. Théreau, C. Tzimeas, and J. Yi_II (2004), Rapport de mission de la campagne ZoNéco 11 de sismique lourde (multitrace, réfraction, haute résolution) à bord du N/O L'Atalante (8 sept. - 5 oct. 2004) - Volume texte, 1-147 pp.

Lafoy, Y., B. Pelletier, J. M. Auzende, F. Missègue, and L. Mollard (1994), Tectonique compressive cénozoïque sur les rides de Fairway et Lord Howe, entre Nouvelle Calédonie et Australie, *Comptes rendus de l'Académie des sciences, Série II, Sciences de la terre et des planètes*, 319, 1063-1069.

Lafoy, Y., S. van de Beuque, F. Missegue, A. Nercessian, and G. Bernadel (1998), Campagne de sismique multitrace entre la marge Est Australienne et le Sud de l'arc des Nouvelles-Hébrides - Rapport de la campagne RIG SEISMIC 206 (21 avril - 24 mai 1998) - Programme FAUST, 1-40 pp.

Lagabrielle, Y., P. Maurizot, Y. Lafoy, G. Cabioch, B. Pelletier, M. Régnier, I. Wabete, and S. Calmant (2005), Post-Eocene extensional tectonics in Southern New Caledonia (SW Pacific): Insights from onshore fault analysis and offshore seismic data, *Tectonophysics*, 403, 1-28.

Launay, J. (1985), Paléoniveaux marins et néotectonique à l'île des Pins (Nouvelle Calédonie), *Bulletin Société Géologique de France*, 1, 77-81.

Launay, J., J. Dupont, A. Lapouille, C. Ravenne, and C. E. de Broin (1977), Seismic traverses across the northern Lord Howe Rise and comparison with the southern part (South West

- Pacific), in *International symposium on geodynamics in South-West Pacific*, edited, pp. 155-164, Technip, Nouméa (New Caledonia).
- Launay, J., and J. Recy (1972), Variations relatives du niveau de la mer et néotectonique en Nouvelle Calédonie au Pléistocène Supérieur et à l'Holocène, *Rev. Géogr. phys. Géol. dyn.*, *14*, 47-65.
- Loubrieu, B., W. Roest, and S. Party (2004), Rapport de mission de la campagne Noucaplac-2 de sismique lourde à bord du N/O L'Atalante - Programme EXTRAPLAC, IFREMER.
- Mignot, A. (1984), Sismo-stratigraphie de la terminaison nord de la ride de Lord Howe. Evolution géodynamique du Sud-Ouest Pacifique entre l'Australie et la Nouvelle-Calédonie, PhD thesis, 205 pp, Université Pierre et Marie Curie, Paris.
- Nouzé, H., E. Cosquer, Y. Lafoy, L. Géli, F. Klingelhoefer, J. Collot, and J. P. Foucher (2007), Seismic characterization of bottom simulating reflectors in the Fairway basin, offshore New Caledonia, *Journal of Geophysical Research*, in prep.
- Nouzé, H., Y. Lafoy, L. Géli, F. Klingelhoefer, and Z. c. party (2005), First results of a high resolution seismic study of a Bottom Simulating Reflector in the Fairway Basin, offshore New Caledonia, paper presented at Fifth International Conference on Gas Hydrates (ICGH 5), Trondheim, June 13-16.
- Paquette, J. L., and D. Cluzel (2006), U-Pb zircon dating of post-obduction volcanic-arc granitoids and a granulite-facies xenolith from New Caledonia. Inference on Southwest Pacific geodynamic models, *Int J Earth Science*, *96*, 613-622.
- Paris, J. P. (1981), *Géologie de la Nouvelle Calédonie: un essai de synthèse*, 279 pp.
- Pecher, I. A. (2004), Waveform inversion applied to a bottom simulating reflection on the eastern Lord Howe Rise, 1-24 pp.
- Prinzhofer, A., A. Nicolas, D. Cassard, J. Moutte, M. Leblanc, J. P. Paris, and M. Rabinovitch (1980), Structures in the new caledonia peridotites-gabbros: Implications for oceanic mantle and crust, *Tectonophysics*, *69*, 85-112.
- Ravenne, C., C. E. De Broin, J. Dupont, A. Lapouille, and J. Launay (1976), New Caledonia basin-Fairway Ridge: structural and sedimentary study, *International symposium on geodynamics in south-west pacific*, 145-154.
- Regnier, M. (1988), Lateral variation of upper mantle structure beneath New Caledonia determined from p-wave receiver function: evidence for a fossil subduction zone, *Geophysical journal Oxford*, *95*, 561-577.
- Rigolot, P. (1989), Evolution morphologique et structurale de la marge occidentale de la ride de Nouvelle-Calédonie (SW Pacifique), *Bulletin des Centres de recherches exploration production Elf Aquitaine*, *13*, 319-344.
- Rigolot, P., and B. Pelletier (1988), Tectonique compressive recente le long de la marge Ouest de la Nouvelle-Caledonie: resultats de la campagne ZOE 400 du N/O Vauban (mars 1987), *Comptes rendus de l'Académie des sciences, Série 2, Mécanique, physique, chimie, sciences de l'univers, sciences de la terre*, *307*, 179-184.
- Shemenda, A. I. (1992), Horizontal Lithosphere Compression and Subduction: Constraints Provided by Physical Modelling, *Journal of Geophysical Research*, *97*, 11,097-011,116.
- Smith, W. H. F., and D. T. Sandwell (1997), Global sea floor topography from satellite altimetry and ship depth soundings, *Science*, *277*, 1956-1962.
- Spandler, C., and J. Hermann (2006), High-pressure veins in eclogite from New Caledonia and their significance for fluid migration in subduction zones, *Lithos*, In Press, Corrected Proof.
- Spandler, C., J. Hermann, R. Arculus, and J. Mavrogenes (2004), Geochemical heterogeneity and element mobility in deeply subducted oceanic crust; insights from high-pressure mafic rocks from New Caledonia, *Chemical Geology*, *206*, 21-42.

- Spandler, C., D. Rubatto, and J. Hermann (2005), Late Cretaceous-Tertiary tectonics of the southwest Pacific: Insight from U-Pb sensitive, high-resolution ion microprobe (SHRIMP) dating of eclogite facies from New Caledonia, *Tectonics*, 24.
- Symonds, P. A., J. B. Colwell, H. I. Struckmeyer, J. B. Willcox, and P. J. Hill (1996), Mesozoic rift basin development off eastern Australia, *Geological Society of Australia*, 43, 528-542.
- Uruski, C., and R. Wood (1991), A new look at the New Caledonia Basin, an extension of the Taranaki Basin, offshore North Island, New Zealand, *Marine and Petroleum Geology*, 8, 379-391.
- Van de Beuque, S. (1999), Evolution géologique du domaine péri-calédonien (Sud-Ouest Pacifique), PhD thesis, 270 pp, Université de Bretagne Occidentale, Brest.
- Van de Beuque, S. (2003), Geological framework of the Northern Lord Howe Rise and adjacent areas, *Geoscience Australia Records*, 2003.
- Van de Beuque, S., J.-M. Auzende, Y. Lafoy, G. Bernardel, A. Nercessian, M. Regnier, P. Symonds, and N. Exon (1998a), Transect sismique continu entre l'arc des Nouvelles-Hébrides et la marge orientale de l'Australie: programme FAUST (French Australian Seismic Transect), *Comptes Rendus de l'Académie des Sciences - Series IIA - Sciences de la terre et des planètes*, 327, 761-768.
- Van de Beuque, S. V., J.-M. Auzende, Y. Lafoy, and F. Missegue (1998b), Tectonique et volcanisme tertiaire sur la ride de Lord Howe (Sud-Ouest Pacifique), *Comptes Rendus de l'Académie des Sciences - Series IIA - Earth and Planetary Science*, 326, 663-669.
- Vially, R., Y. Lafoy, J.-M. Auzende, and R. France (2003), Petroleum potential of New-Caledonia and its offshore basins, in *AAPG International conference*, edited, Barcelona (Spain).
- Willcox, J. B., J. Sayers, H. M. J. Stagg, and S. van de Beuque (2001), Geological framework of the Lord Howe Rise and adjacent ocean basins, Eastern Australasian basins symposium 2001: A refocused energy perspective for the future, *Petroleum Exploration Society of Australia Special Publication*, 1, 211-225.
- Willcox, J. B., P. Symonds, K. Hinz, and D. Bennett (1980), Lord How Rise, Tasman Sea - Preliminary geophysical results and petroleum prospects, *BMR Journal of Australian Geology and Geophysics*, 5, 225-236.

FIGURE CAPTIONS

Figure 1

SW Pacific bathymetric location map with main regional Deep Sea Drilling Program boreholes (blue circles) and names of the main structural elements (DR: Dampier ridge, MB: Middleton Basin, FR: Fairway Ridge, NC: New Caledonia, NR: Norfolk Ridge, LR: Loyalty Ridge, NB: Norfolk Basin, TKR: Three Kings Ridge). Sea floor topography from [*Smith and Sandwell, 1997*]. Black dashed rectangle locates the study zone.

Figure 2

Diagram synthesizing the regional geological data (boreholes and tectonic onshore events)

- A- Tectonic events related to the history of the South Loyalty Basin observed onshore New Caledonia [*Cluzel et al., 1997; Cluzel, 1998; Cluzel et al., 1998; Cluzel et al., 2001; Cluzel et al., 2006*]
- B- Geological time scale
- C- Seafloor spreading of main SW Pacific oceanic basins in literature
- D- SW Pacific regional boreholes reaching the Eocene/Oligocene unconformity U1 [*Burns et al., 1973d; c; b*]
- E- Correlation of borehole DSDP208 with seismic line FAUST1-S206-2 using the velocity analysis of CDP 27328 [*Van de Beuque, 1999*]. Reflector RN marked in violet is identified with U1, and is characterized by a transparent overlying seismic facies and a negative polarity. In pink, the characteristic intra-Miocene reflector.

Figure 3

Bathymetric location map of all available seismic lines in the study zone. This zone corresponds with the black dashed rectangle on Figure 1.

Highlighted in yellow is the path we use to identify and follow continuously reflector RN from DSDP 208 over the Lord Howe Rise, the Fairway Basin, the Fairway Ridge and down into the New Caledonia Basin. Red dashed line is the morpho-tectonic feature described by *Lafoy et al.* [2005a] which separates the North from South New Caledonia Basin. Sea floor topography from *Smith and Sandwell* [1997].

Figure 4

Phase inversion of the reflector associated to the Eocene/Oligocene unconformity [*Nouzé et al.*, 2005; *Nouzé et al.*, 2007] on profile z05-11 from the ZoNéCo-5 survey.

- where the unconformity sits on basement (ellipse) RN has a positive phase
- where the unconformity is in the basin (rectangle) RN has a negative phase

Figure 5

Seismic profile FAUST1-S206-2 modified from [*Lafoy et al.*, 1998]

The yellow line in the horizontal scale indicates the seismic path we use to follow the Eocene/Oligocene unconformity from well DSDP 208 to line FAUST1-S206-1. This path is also underlined in yellow on the location map (Figure 3).

Acoustic basement is orange, reflector RN associated to the Eocene/Oligocene unconformity is violet and intra-Miocene reflector is pink (see Figure 2-E for more details). Other reflectors picked on the following figures 6-7-8-9-10-11-12-13-14 are not dated and are not the focus of this paper. Their presence is for information only.

Figure 6

Seismic profile FAUST1-S206-1 modified from [*Lafoy et al.*, 1998]

The yellow line in the horizontal scale indicates the seismic path we use to follow the Eocene/Oligocene unconformity from line FAUST1-S206-2 to line z11-11. This path is also

underlined in yellow on the location map (Figure 3). See Figure 5 for color code of seismic reflectors. Black arrows show toplaps beneath RN.

Figure 7

Seismic profile z11-11 modified from [Nouzé *et al.*, 2005; Nouzé *et al.*, 2007]

The yellow line in the horizontal scale indicates the seismic path we use to follow the Eocene/Oligocene unconformity from line FAUST1-S206-1 to line z11-10. This path is also underlined in yellow on the location map (Figure 3).

RN and RP identified by [Nouzé *et al.*, 2005; Nouzé *et al.*, 2007]: respectively the Eocene/Oligocene unconformity and the regional BSR which confirms our position of U1.

See Figure 5 for color code of seismic reflectors

Figure 8

Seismic profile z11-10 modified from [Nouzé *et al.*, 2005; Nouzé *et al.*, 2007]

The yellow line in the horizontal scale indicates the seismic path we use to follow the Eocene/Oligocene unconformity from line z11-11 to line z11-09. This path is also underlined in yellow on the location map (Figure 3). See Figure 5 for color code of seismic reflectors

Figure 9

Seismic profile z11-09 modified from [Lafey *et al.*, 2004]

The yellow line in the horizontal scale indicates the seismic path we use to follow the Eocene/Oligocene unconformity from line z11-10 to line NCP2-2. This path is also underlined in yellow on the location map (Figure 3). See Figure 5 for color code of seismic reflectors

Figure 10

Seismic profile NCP2-2 modified from [Loubrieu *et al.*, 2004]

The yellow line in the horizontal scale indicates the seismic path we use to follow the Eocene/Oligocene unconformity from line z11-09 to line z11-04. This path is also underlined in yellow on the location map (Figure 3).

Reflector RN is continuous over the Fairway Ridge. See Figure 5 for color code of seismic reflectors

Figure 11

Onlaps along Fairway Ridge (seismic profile z11-01 A)

RN is here truncated by the Fairway Ridge (red arrow) and no correlation between the Fairway Basin and the New Caledonia Basin is therefore possible. See Figure 5 for color code of seismic reflectors

Figure 12

Seismic profile z11-04 modified from [Lafey *et al.*, 2004]

The yellow line in the horizontal scale indicates the seismic path we use to follow the Eocene/Oligocene unconformity from line NCP2-2 to line z11-01A. This path is also underlined in yellow on the location map (Figure 3).

RN deepens as it gets closer to the New Caledonia main land. See Figure 5 for color code of seismic reflectors

Figure 13

Seismic profile z11-01A modified from [Lafey *et al.*, 2004]

The yellow line in the horizontal scale indicates the seismic path we use to follow the Eocene/Oligocene down in the New Caledonia Basin. This path is also underlined in yellow on the location map (Figure 3).

RN coincides with a main tilted reflector. Black arrows pointing to the right show toplaps beneath RN and testify to the existence of a pre-Oligocene basin with its deep part on the

western side. Black arrows pointing to the left show onlaps on RN, and the absence of fanning reflectors testifies to the rapidity of the tilting. The deep part of the basin is on the Eastern side after Oligocene time. See Figure 5 for color code of seismic reflectors

Figure 14

Reconstitution of New Caledonia Basin history using z11-01A's reflectors (Figure 13) from Eocene times to the Present by flattening RN:

A/ The deep part of the basin is on the Western side and the obduction event in New Caledonia creates an overloading (represented by the blue arrows) and a compressive field

B/ Uplift of the Fairway Ridge and eastward deepening of the NNCB during the obduction episode

C/ Sedimentary filling of the newly created depression in a quasi-static tectonic setting

Figure 15

Mid-Eocene to Upper-Oligocene geodynamic evolution of the basin and ridge system effected by the New Caledonian obduction :

(LHR : Lord Howe Rise, FB : Fairway Basin, FR: Fairway Ridge, NNCB : northern New Caledonia Basin, NR: Norfolk Ridge, LR: Loyalty Ridge, SZ : Subduction Zone. For other acronyms see Figure 1)

a) 53 Ma, eastward inception of subduction of the South Loyalty Basin [*Cluzel et al.*, 2006]

b) 37 Ma, arrival of the NR in the subduction. Regional compression begins, the LHR is in a subaerial position and subject to erosion, as evidenced by the unconformity U1 documented at drilling site DSDP 208.

c) between 37 and 34 Ma, overriding of the SLB's oceanic crust onto the NR creates subsidence of the NR (marked by the vertical arrow). This subsidence triggers the tilting of the NNCB while compression continues. The tilting absorbs part of the shortening (dy) which releases strain further west at the LHR and makes the LHR subside.

d) As the allochthon's progressively overrides towards the west onto New Caledonia, the deformation front migrates towards the west and the NNCB is under-thrust beneath New Caledonia. At 34 Ma, the end of obduction and general release of strain, the LHR continues to subside.

d') Physical modelling of a horizontal lithosphere compression from the work of *Shemenda* [1992]

Then isostatic uplift of NC begins, related to its unloading by erosion of the allochthon. By contrast no vertical movements are observed in the NNCB, confirming the hypothesis of a dissociation between the 2 blocks.

The question marks indicate that the role played by the FB in this system is unknown.

Figure 16

Seismic profile z05-12 on the top of the LHR modified from [*Auzende et al.*, 1999]

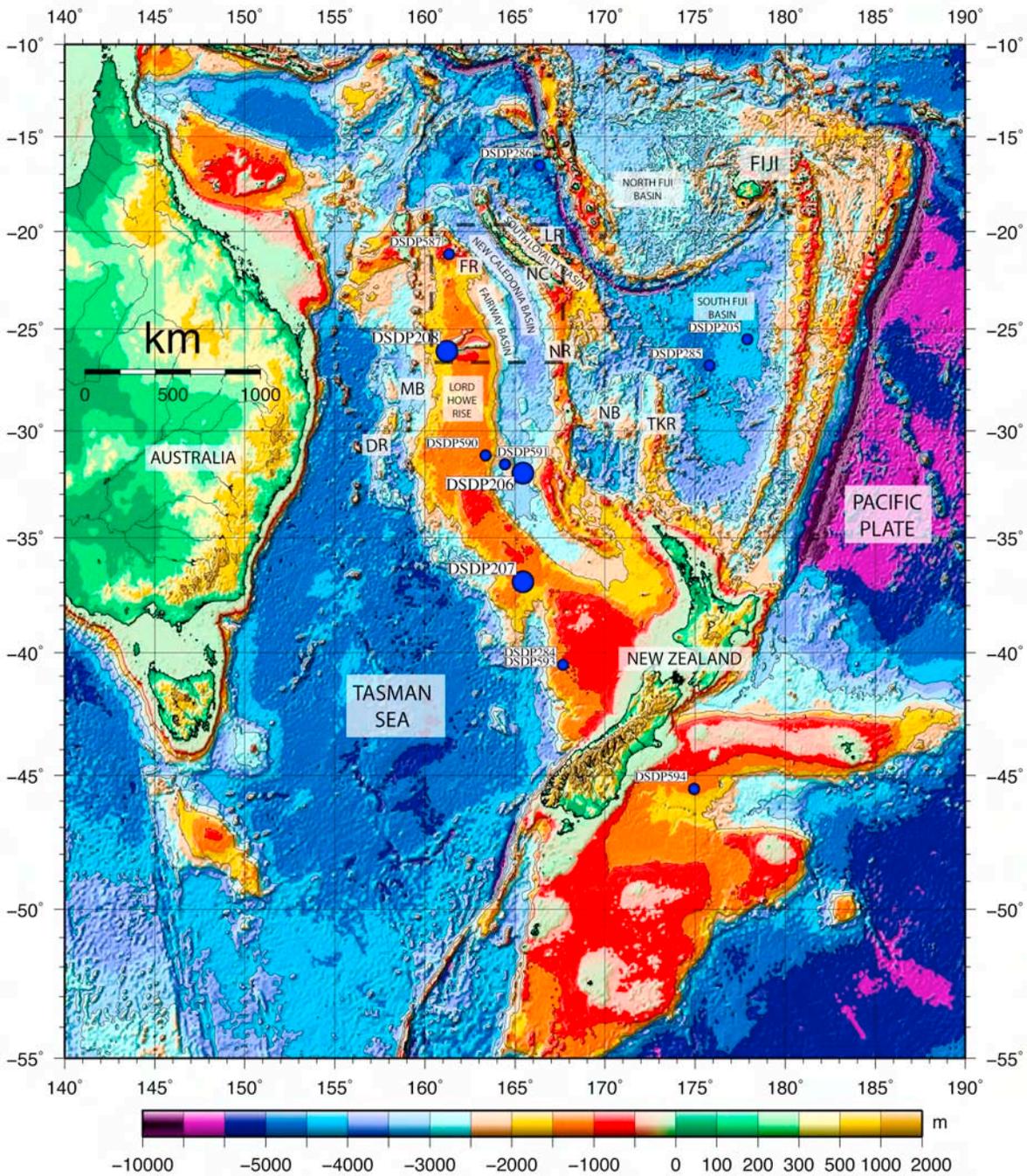
Red arrows show toplaps revealing U1 as an angular unconformity over the LHR.

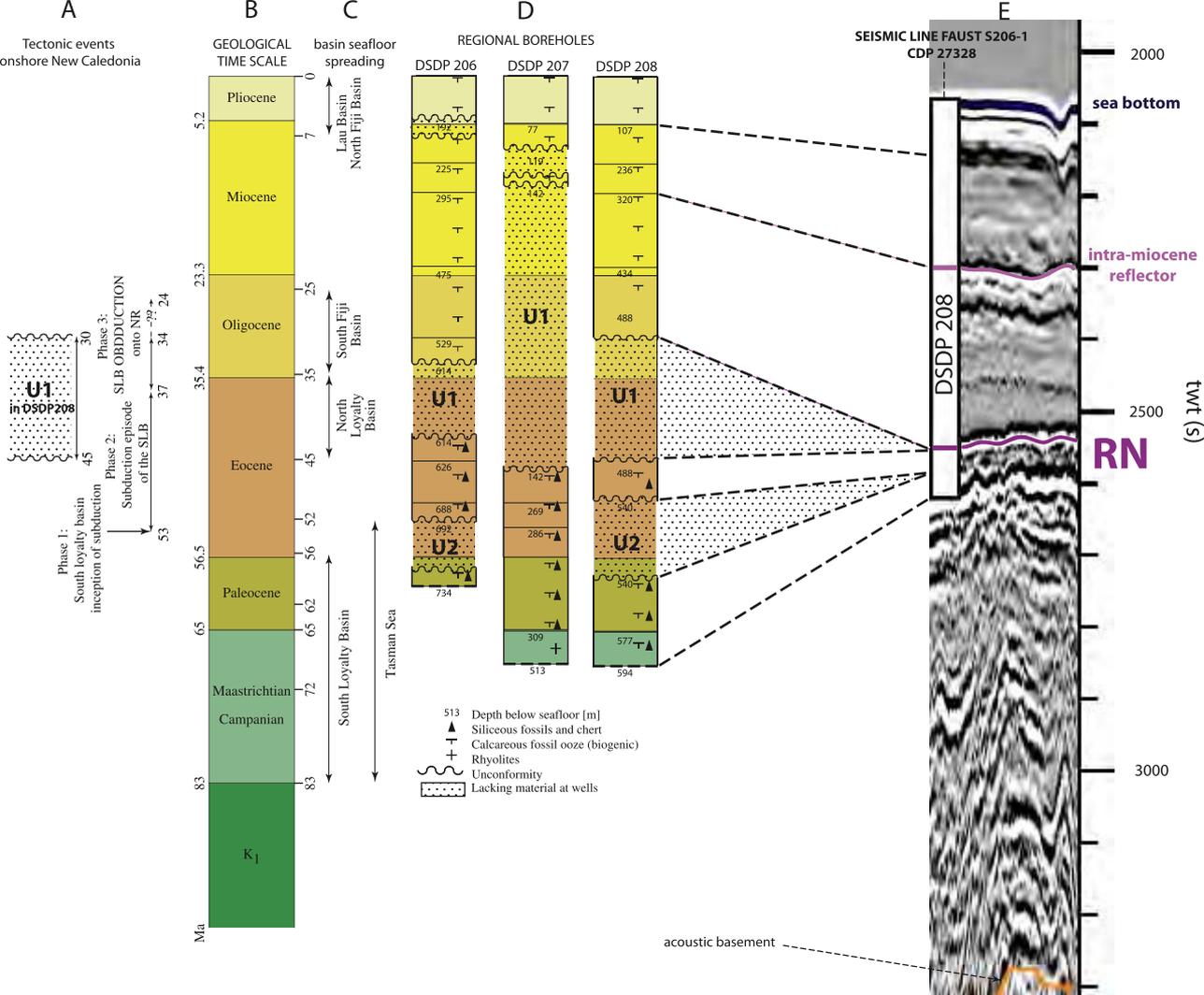
Table 1 Characteristics of the seismic acquisition devices. The single-bubble source [*Avedik et al.*, 1993] used during the ZoNéCo-11 [*Lafoy et al.*, 2004] and Noucaplac-2 [*Loubrieu et al.*, 2004] campaigns explains the low frequency content of profiles z11-01A (Figure 13), z11-04 (Figure 12) and NCP2-2 (Figure 10).

TABLES

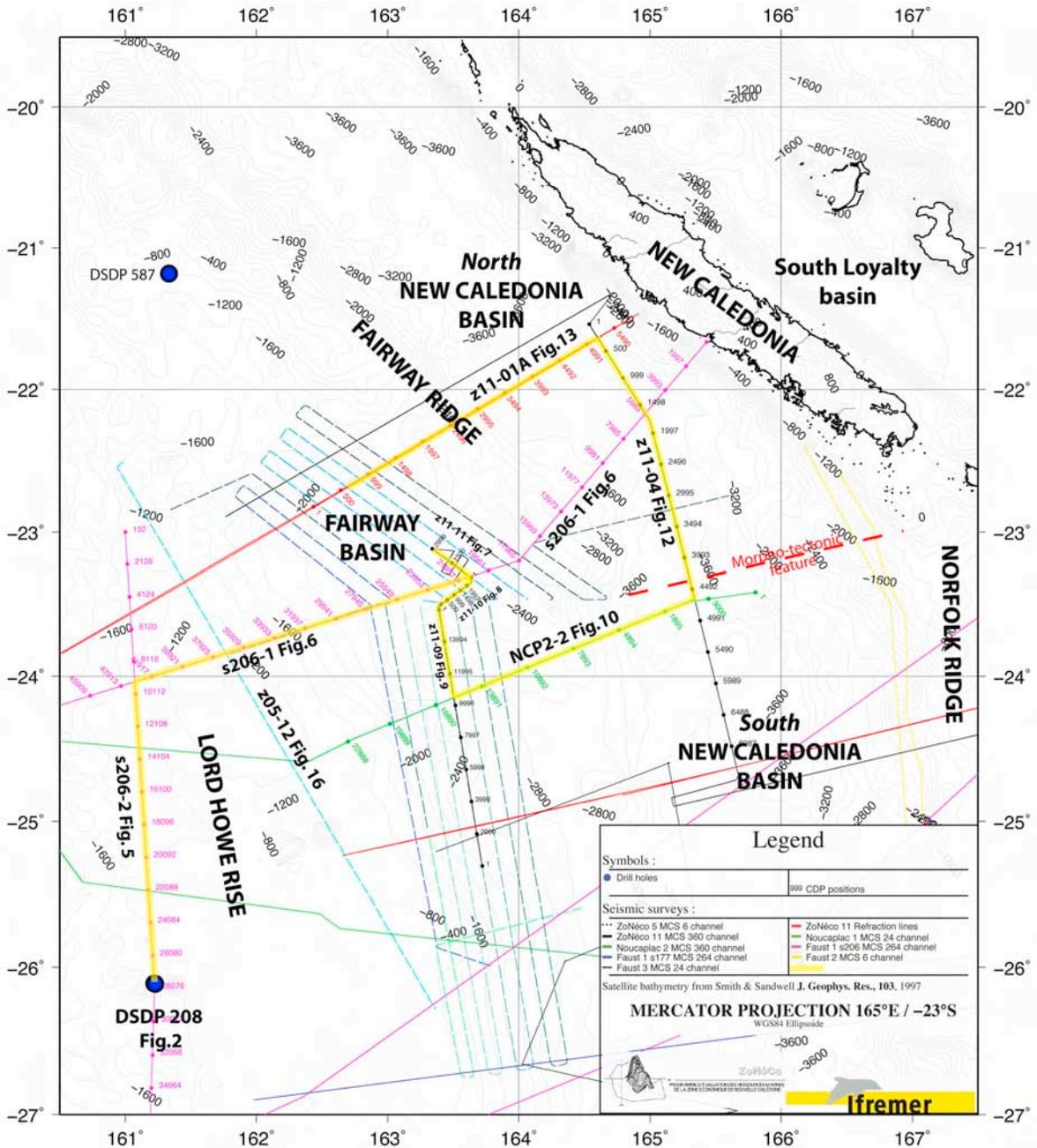
Cruises or profiles number	Streamer length (km)	Number of channels	Receiver immersion (m)	Source type	Source immersion (m)	Source band width (Hz)	Shot interval (m)	Interpreted Seismic data
FAUST FAUST1-S206 (LHRNC)	3.3	264	10	First peak	10	50-60	50	Time migrated
Noucaplac-2	4.5	360	15	Single bubble	15	4-80	75	Time migrated
Z11-01, z11-04 and z11-07	4.5	360	15	Single bubble	20	4-60	150	Time migrated
ZoNéCo 5	0.6	6	7	First peak	10	50-130	50	Stacked
Z11-09, z11-10 and z11-11	3.3	264	3	First peak	3	50-130	25	Time migrated

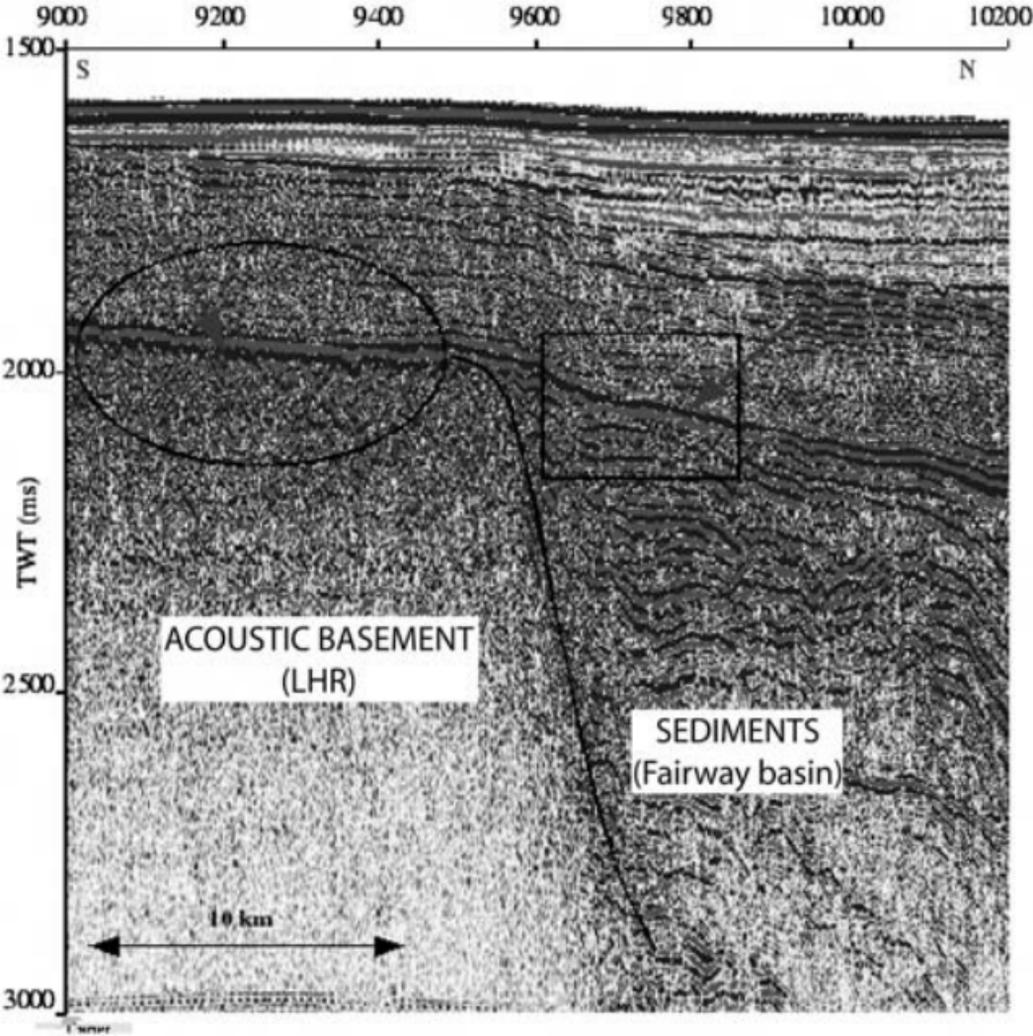
SOUTH WEST PACIFIC OCEAN BATHYMETRY MAP

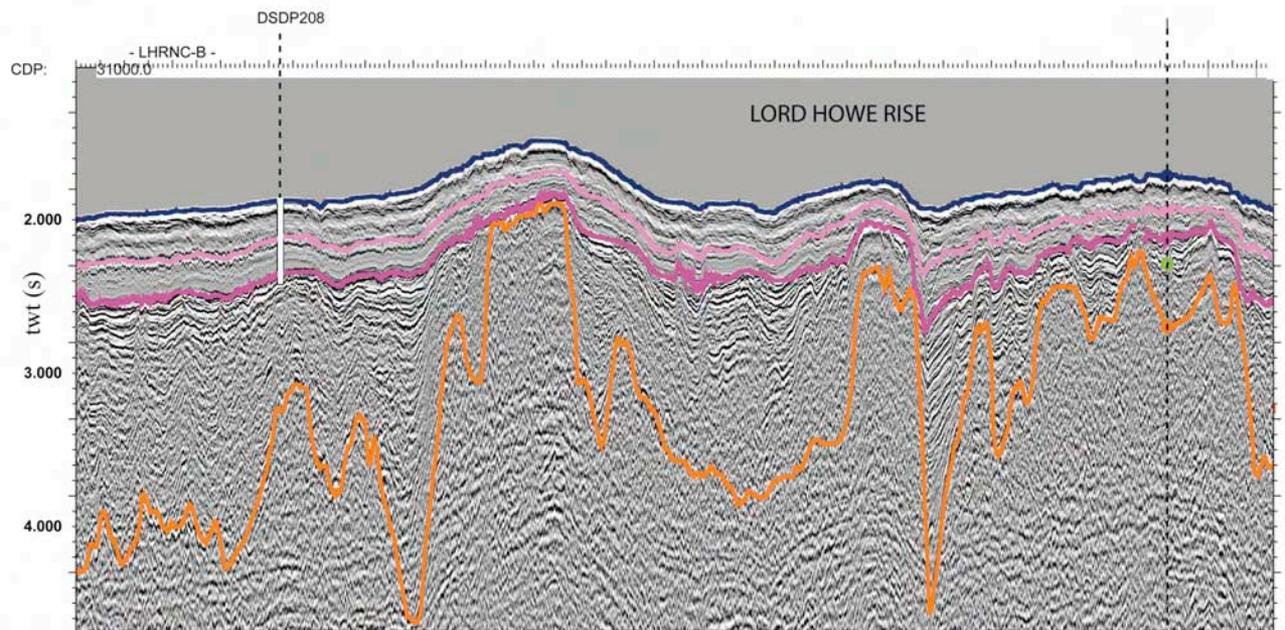
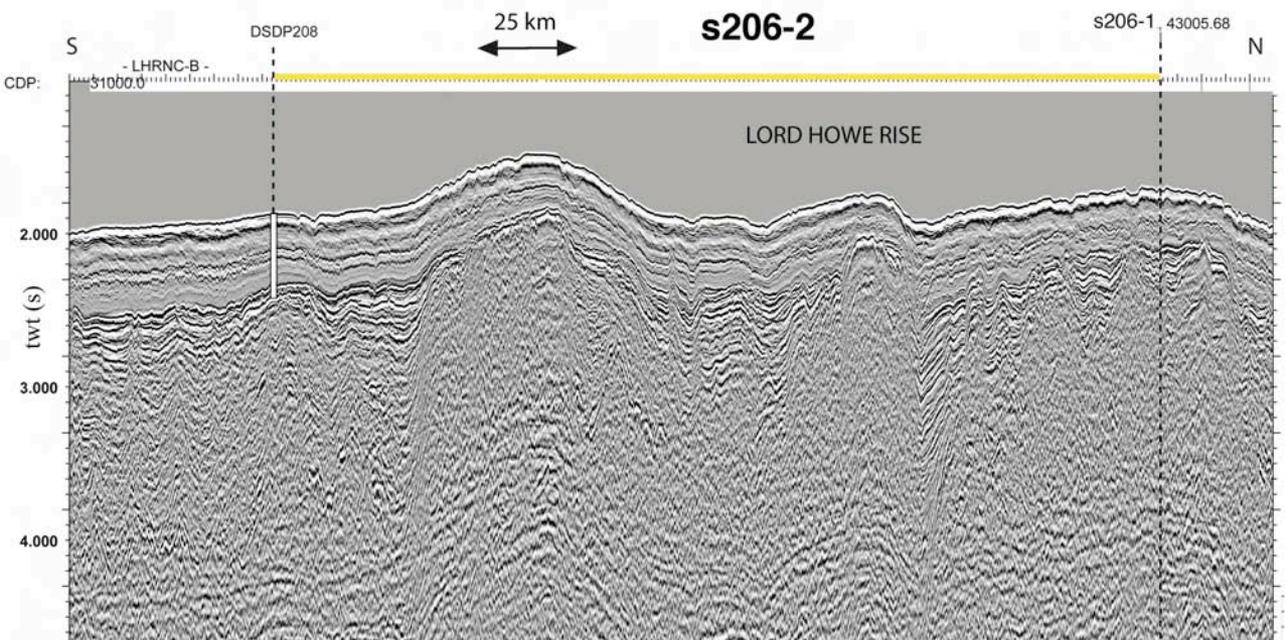


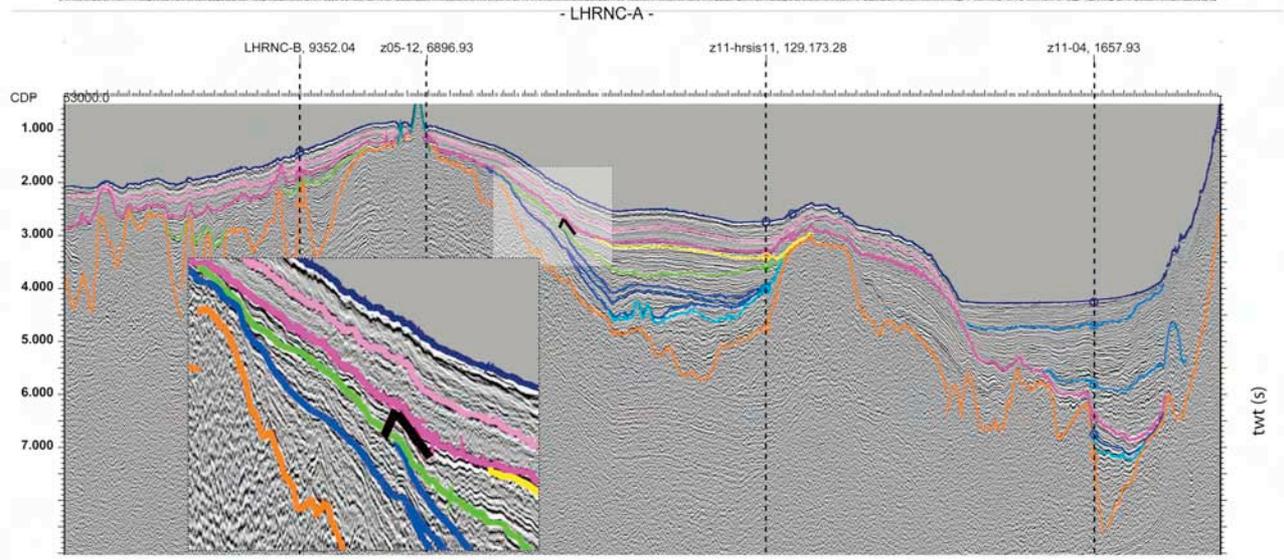
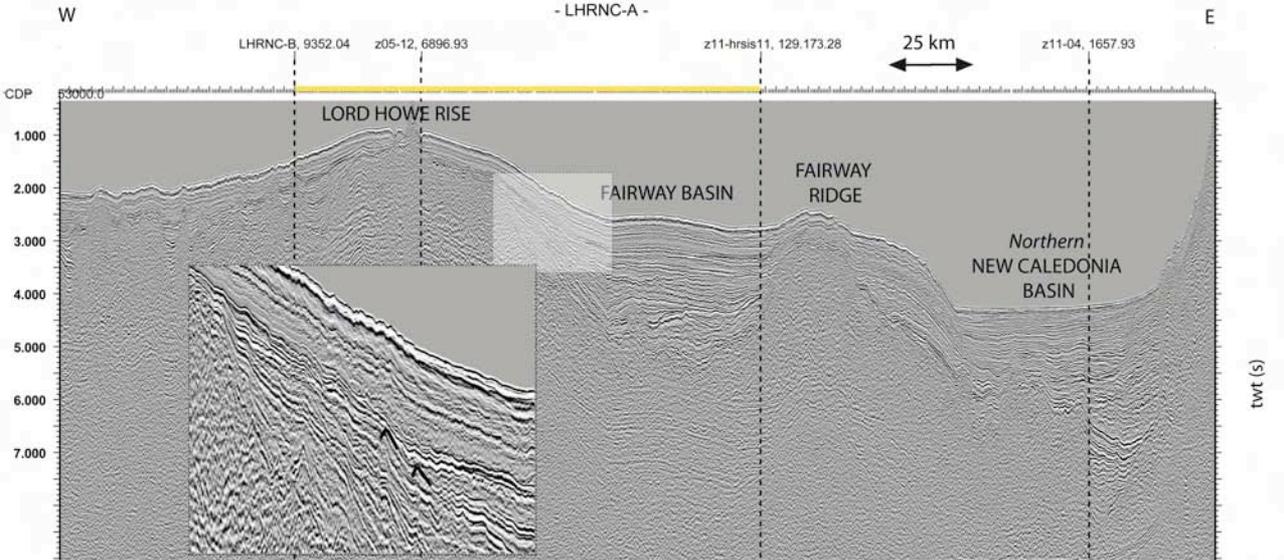


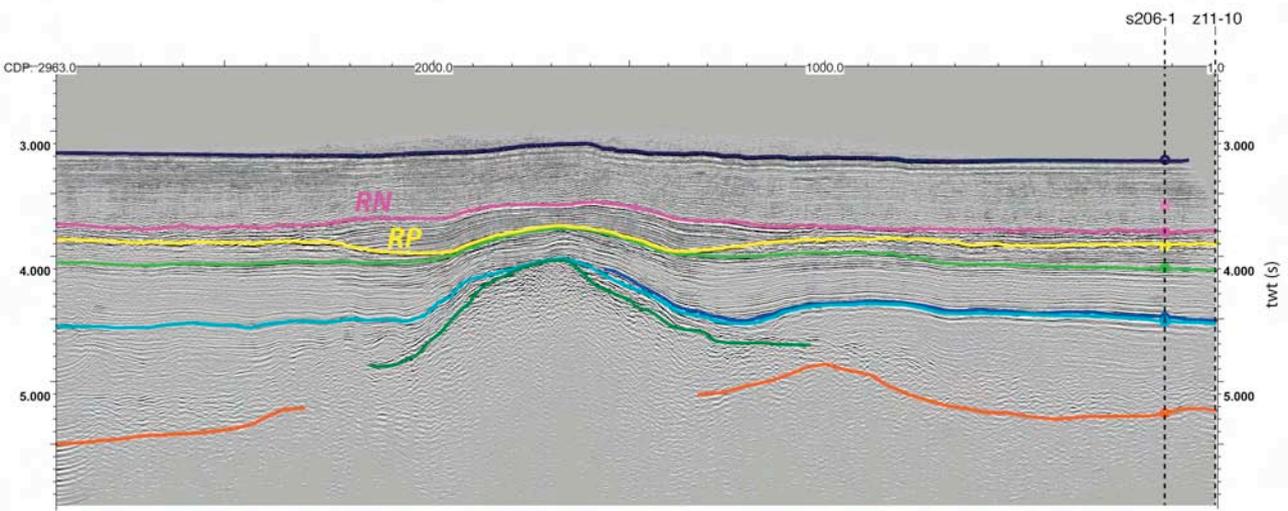
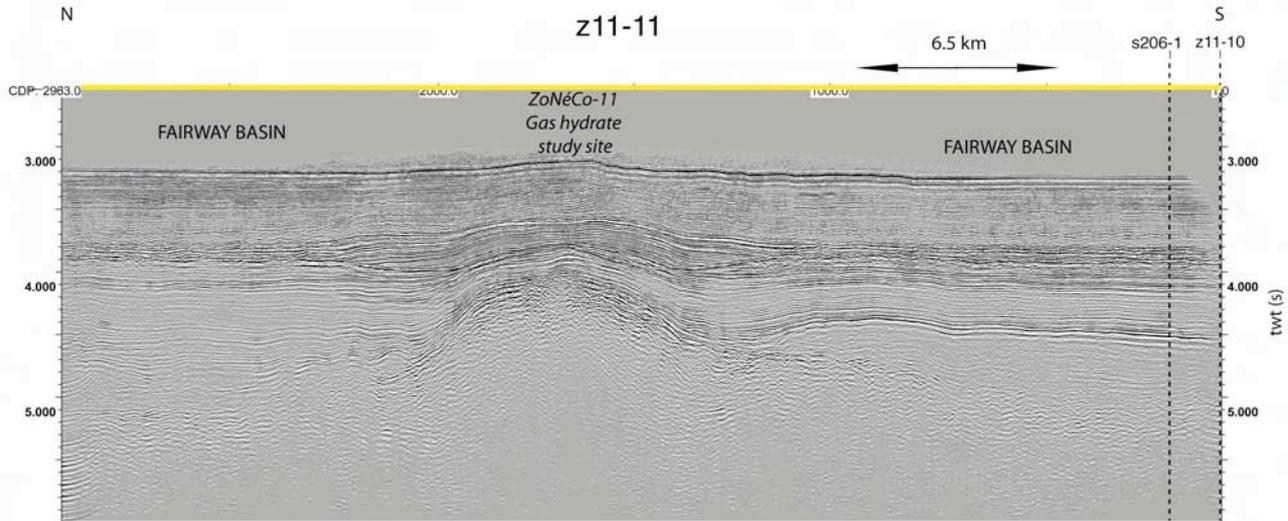
BATHYMETRIC LOCATION MAP



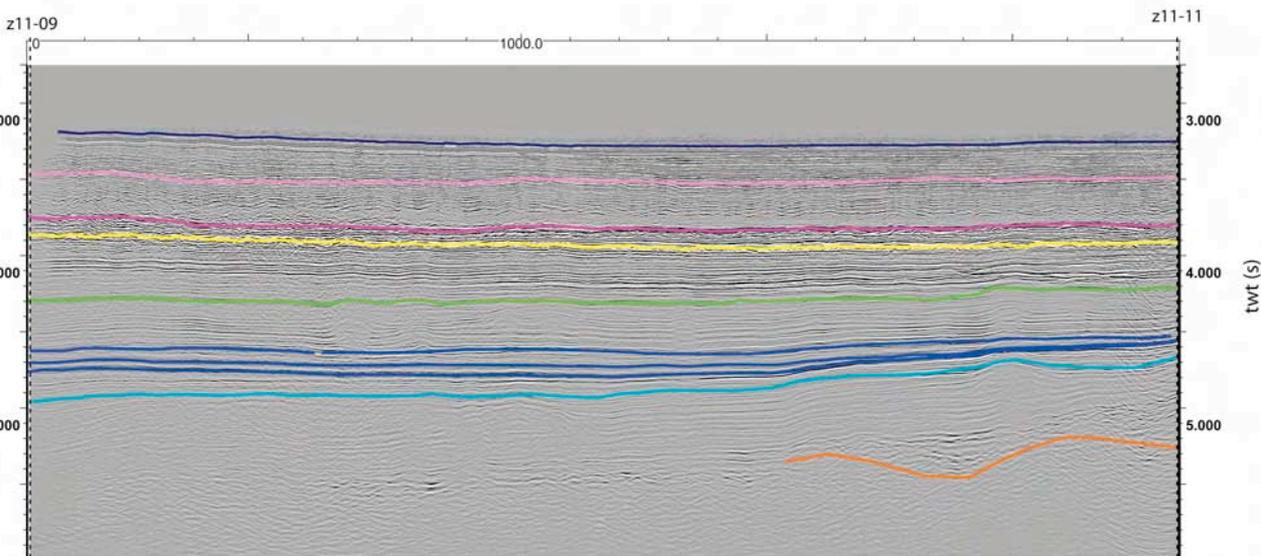
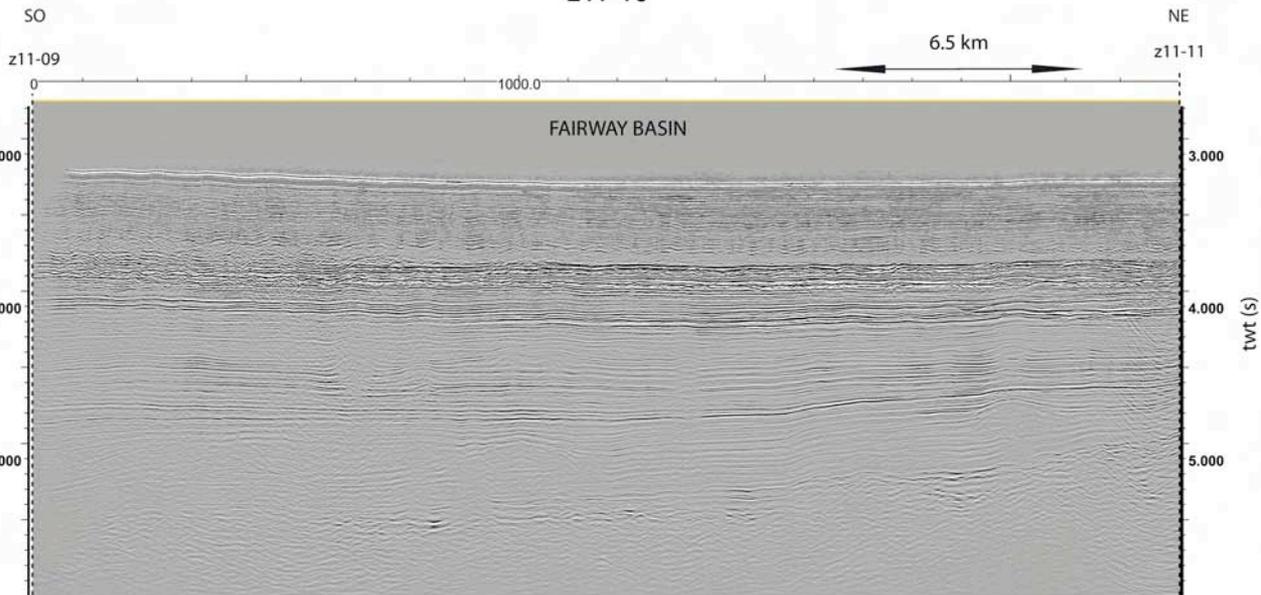




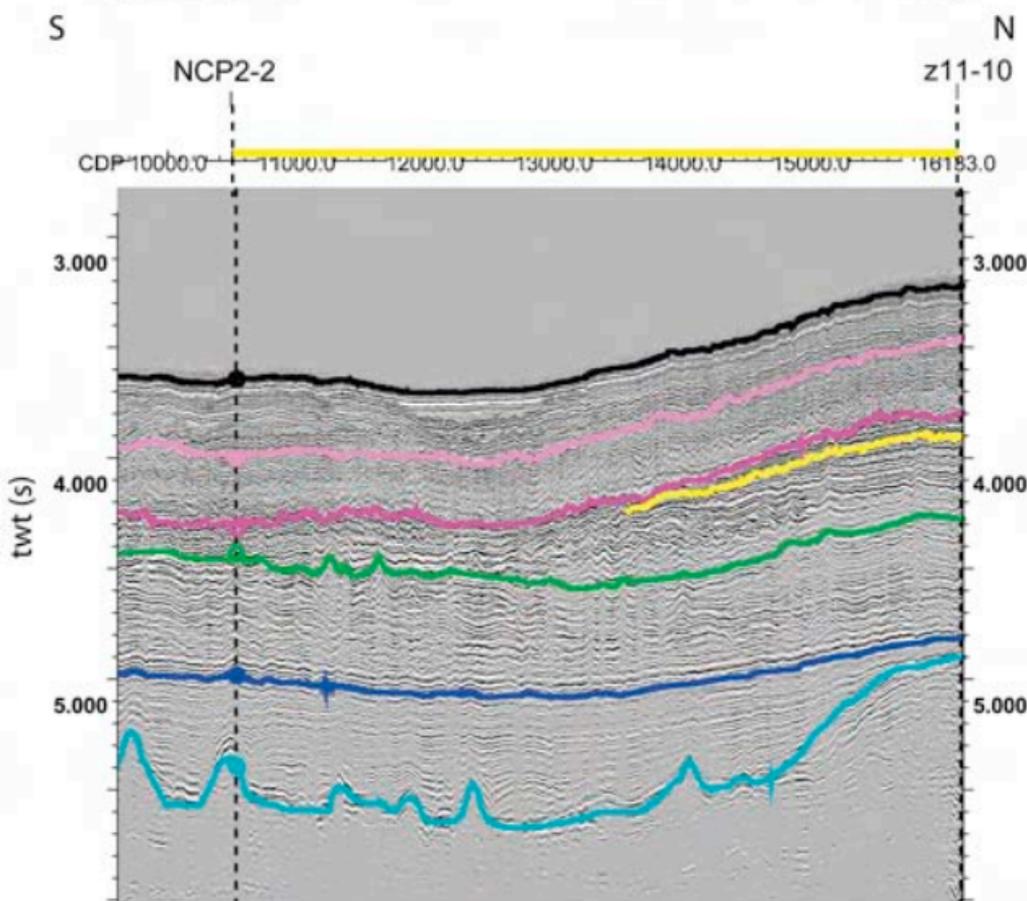
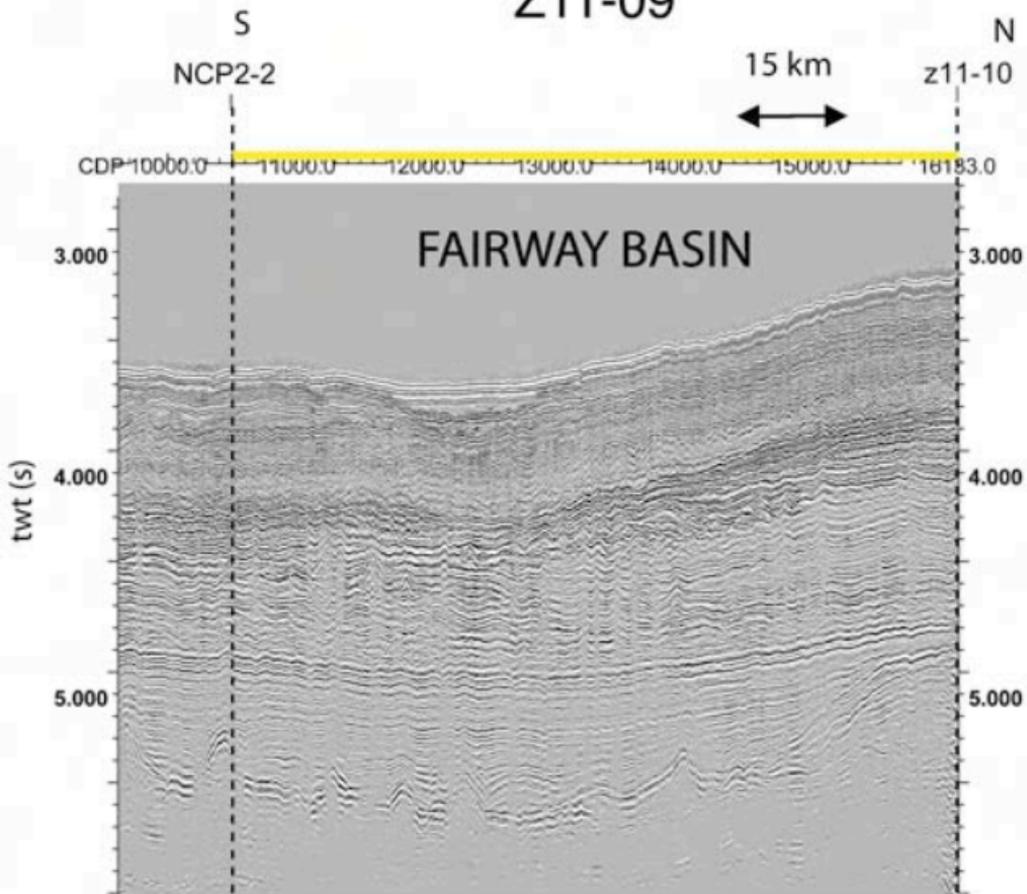


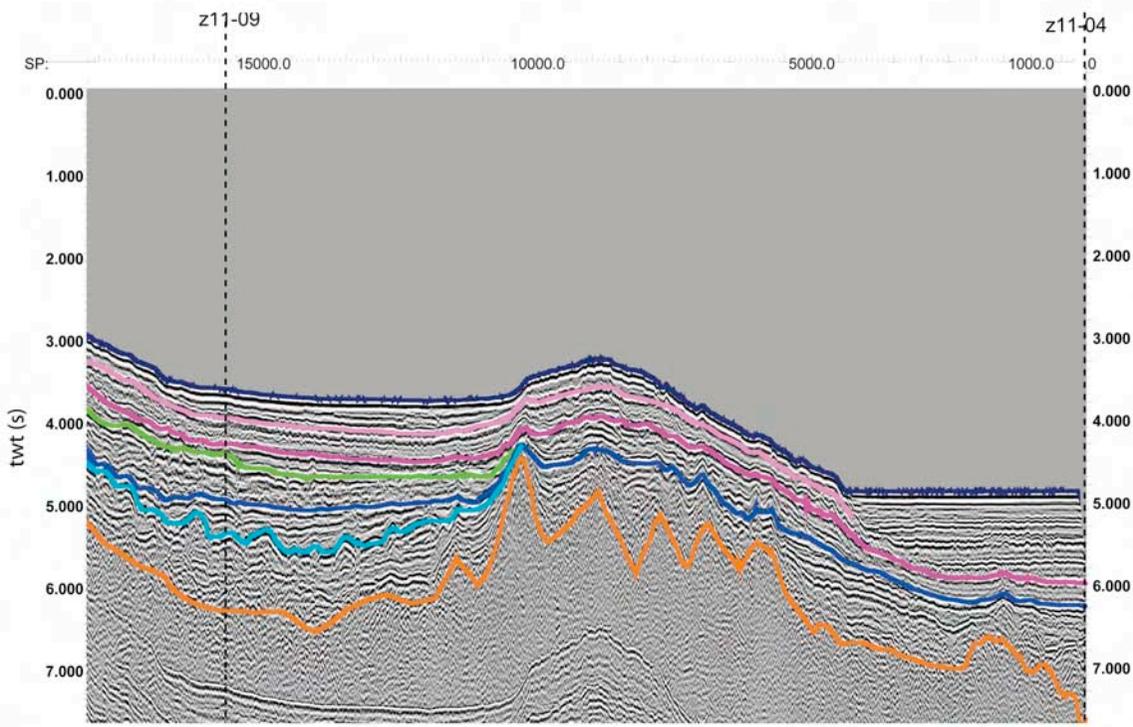
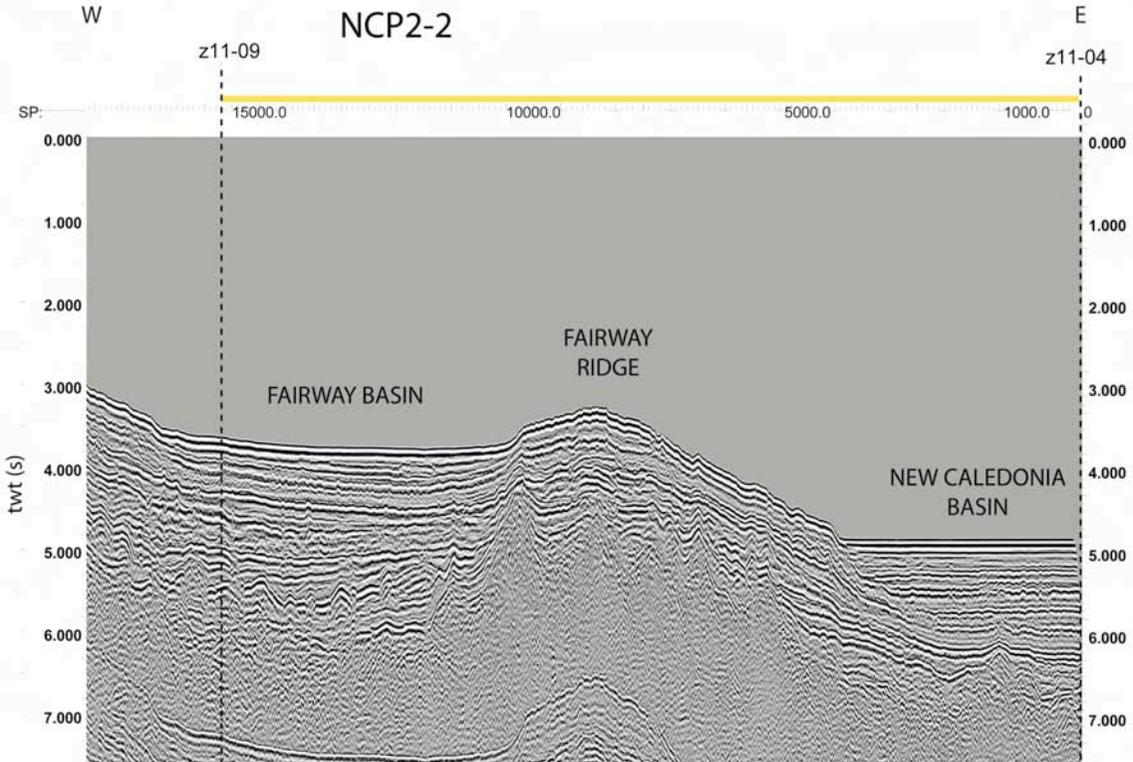


z11-10



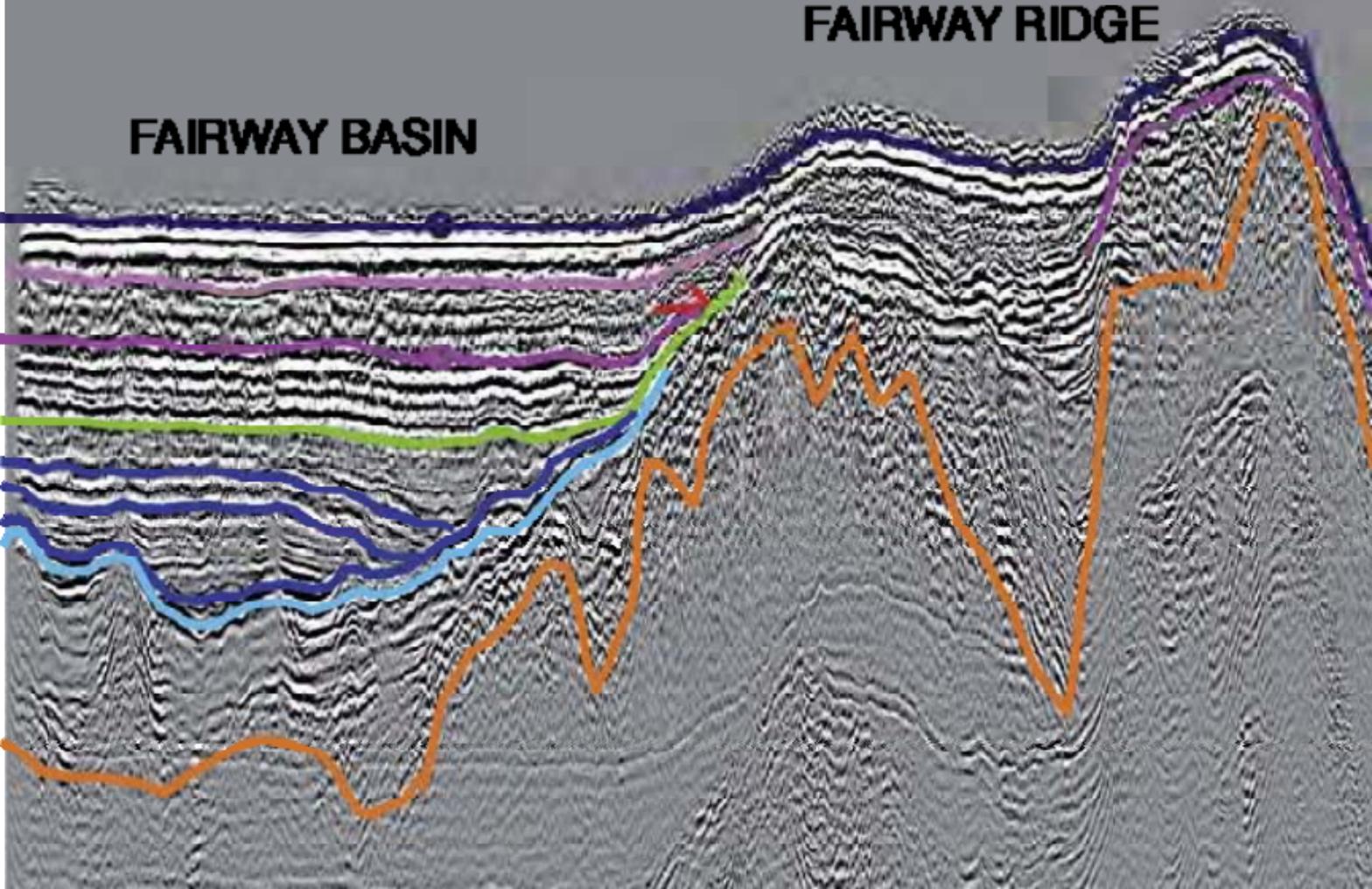
Z11-09

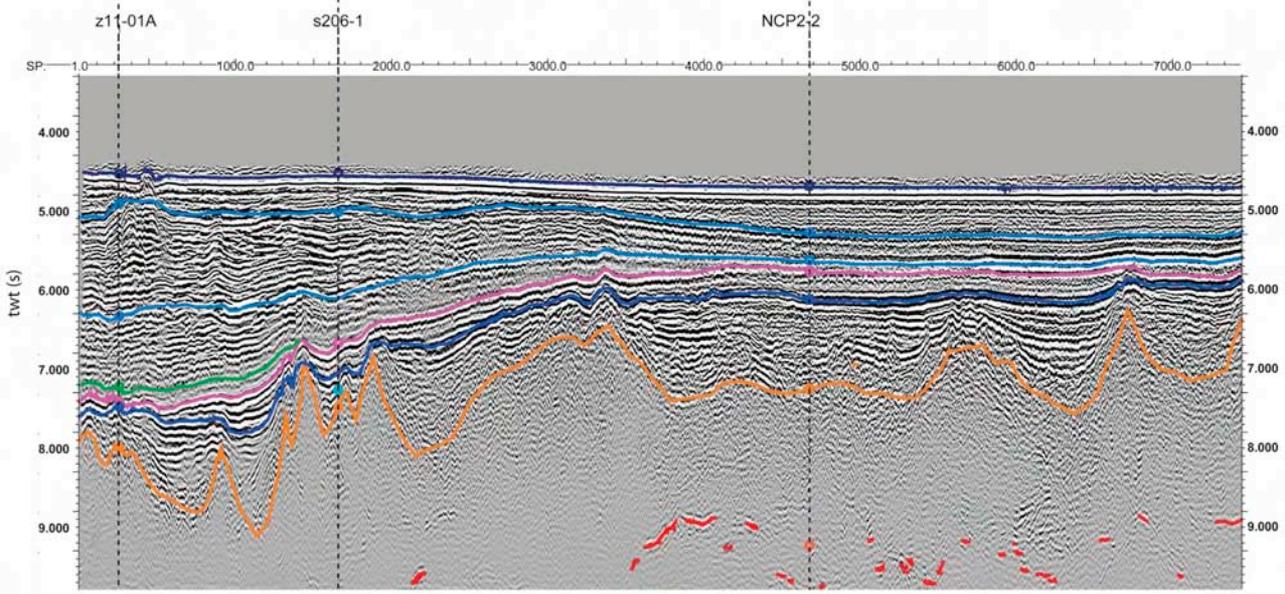
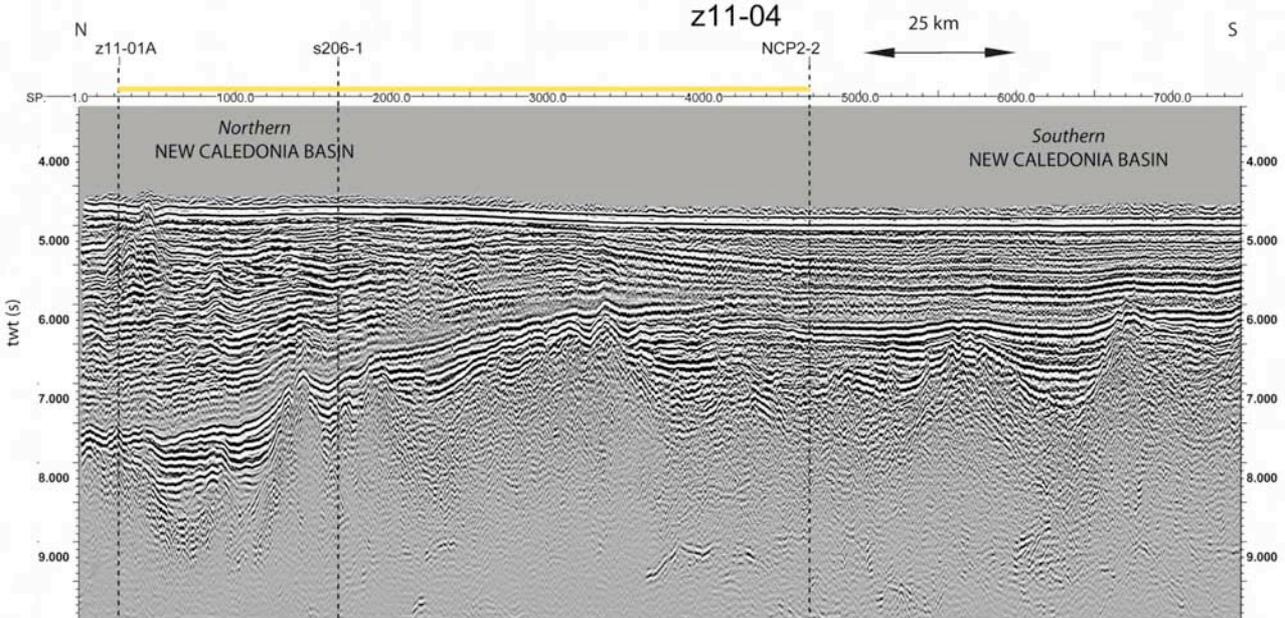


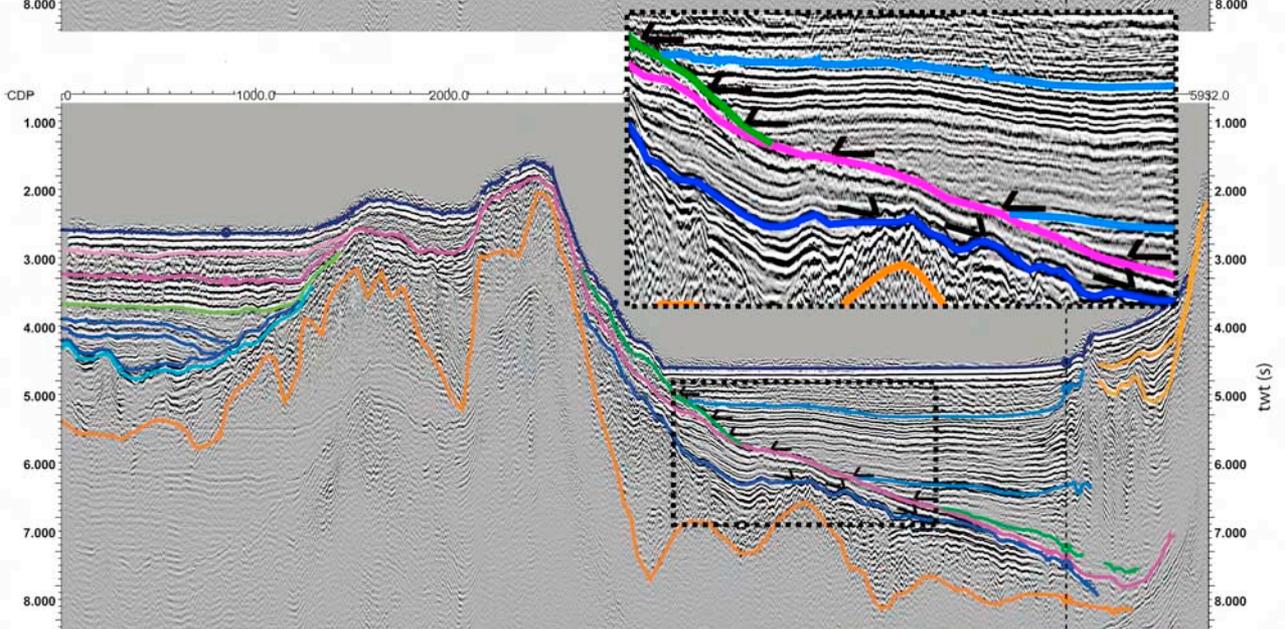
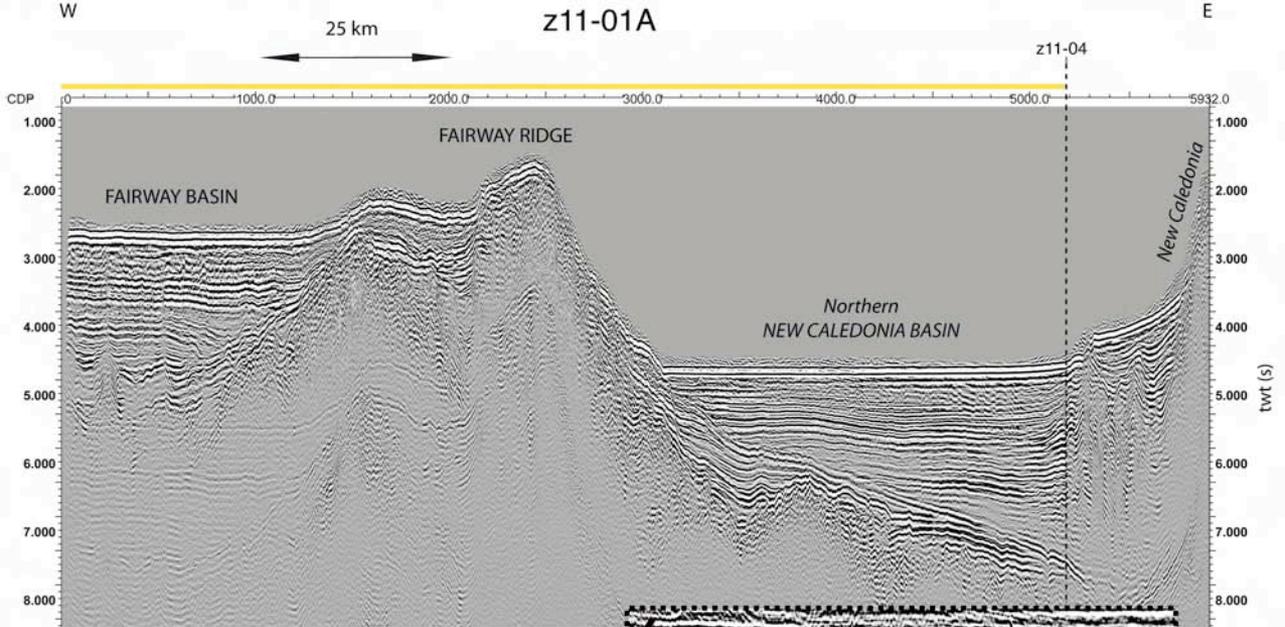


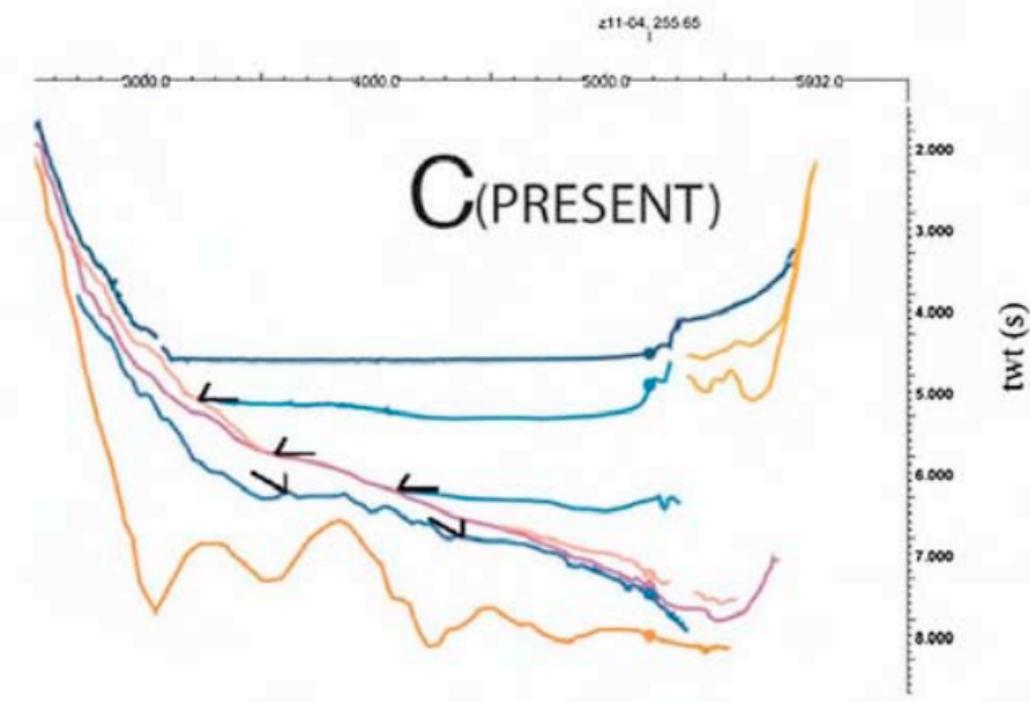
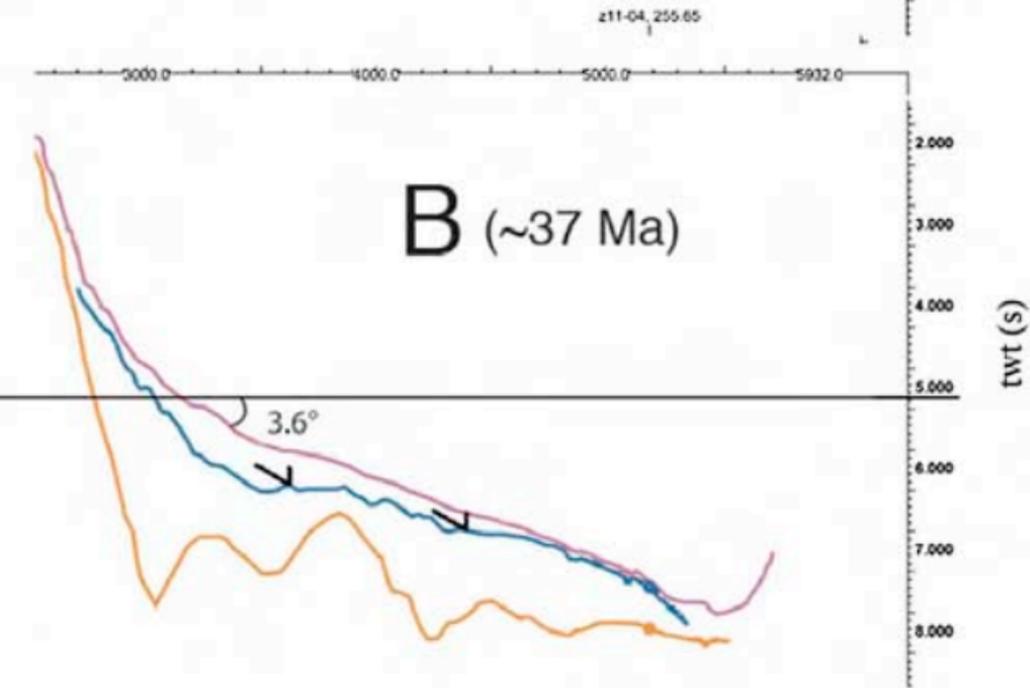
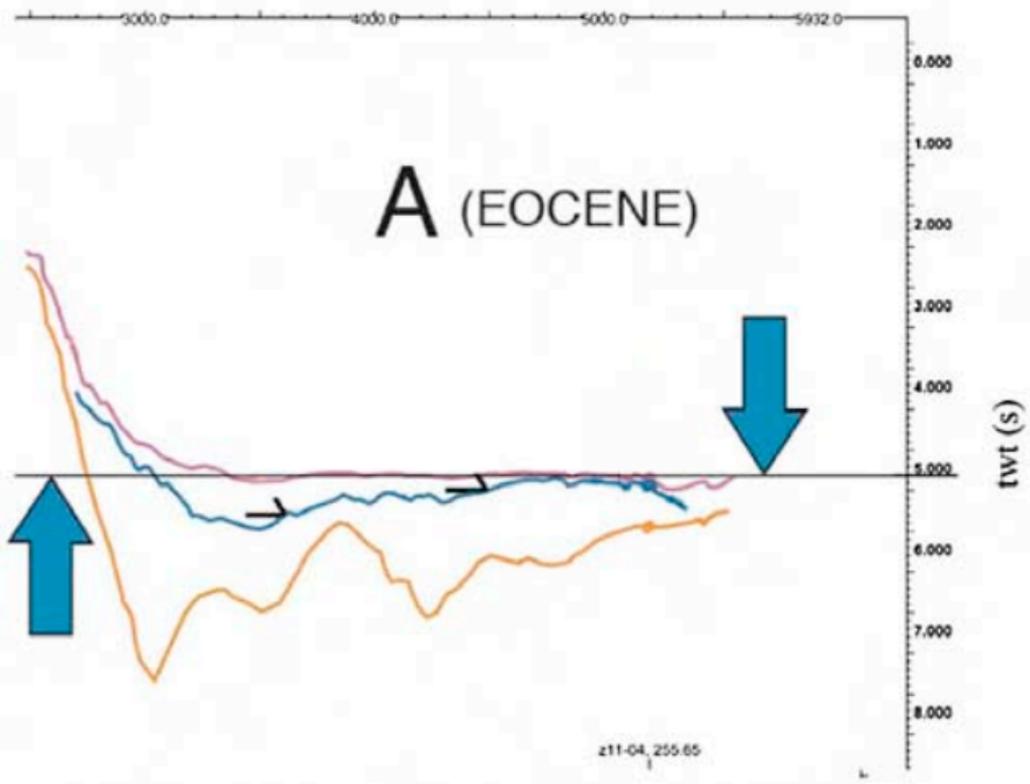
FAIRWAY RIDGE

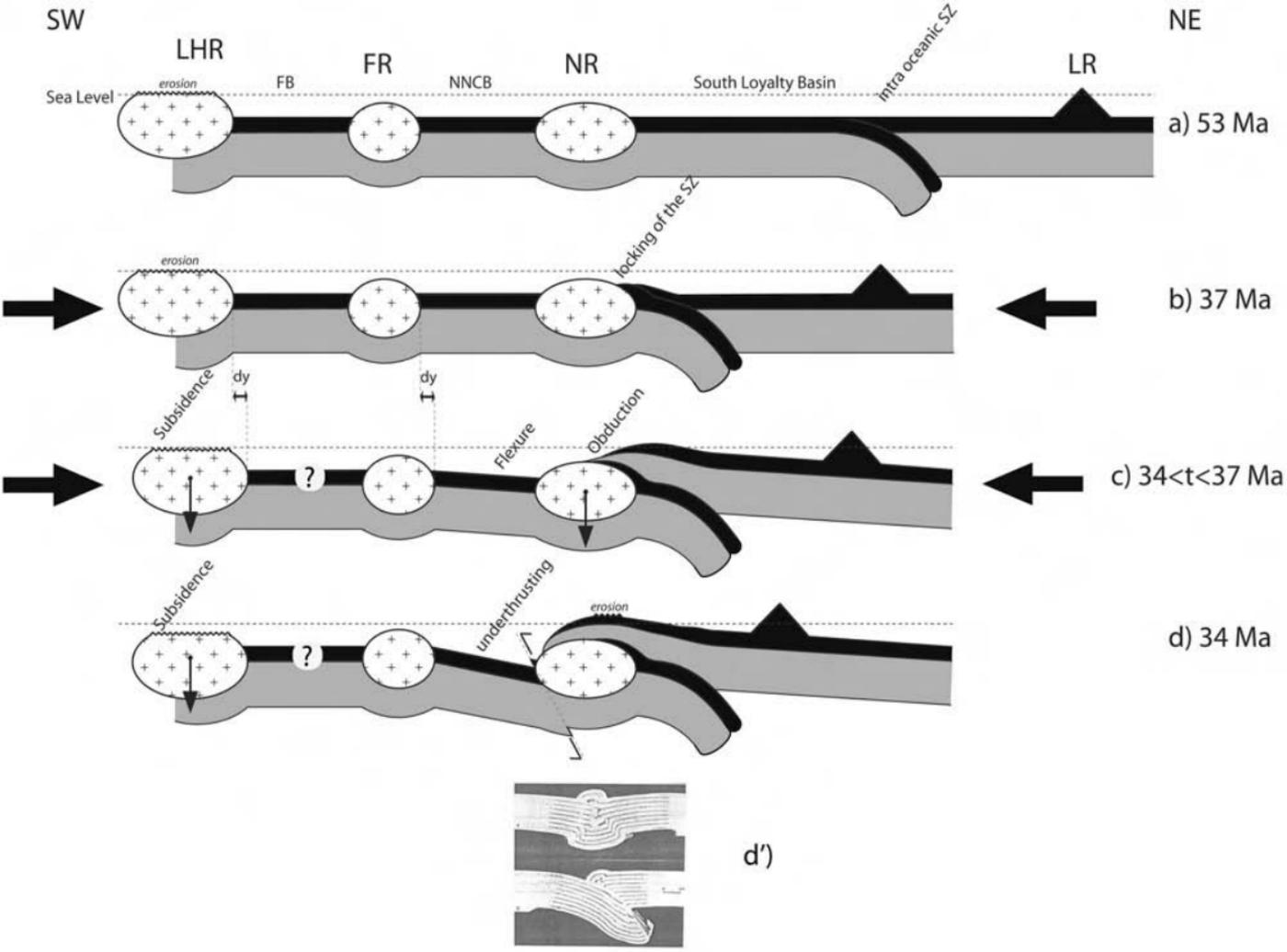
FAIRWAY BASIN











N

z05-12

5 km

S

7400 7600 7800 8000 8200 8400 8600 8800 9000 9200 9400 9600 9800 10000 10200 10400 10600 10800 11000

1.2

LORD HOWE RISE

1.4

1.6

1.8
twt(s)

2.0

2.2

2.4

